



# **CO<sub>2</sub> fluxes in the Coastal Ocean : a short synthesis**

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# Why the Coastal Ocean ?

Facts and figures from Gattuso et al. (1998)

- < covers 7% of surface of global ocean
- < massive inputs of terrestrial organic matter & nutrients
- < intense exchange of energy and matter with open ocean
- < one of most biogeochemically active areas of the biosphere
  
- < 14-30% of oceanic primary production
- < 80% of oceanic organic matter burial
- < 90% of oceanic sedimentary mineralization
- < 75-90% of oceanic sink of suspended river load
- < 50% of oceanic  $\text{CaCO}_3$  deposition
  
- < 37% of human population live within 100km of the coastline

# Gas transfer velocity in estuaries

$$F = a k dp\text{CO}_2$$

$F$  = air-water flux of  $p\text{CO}_2$

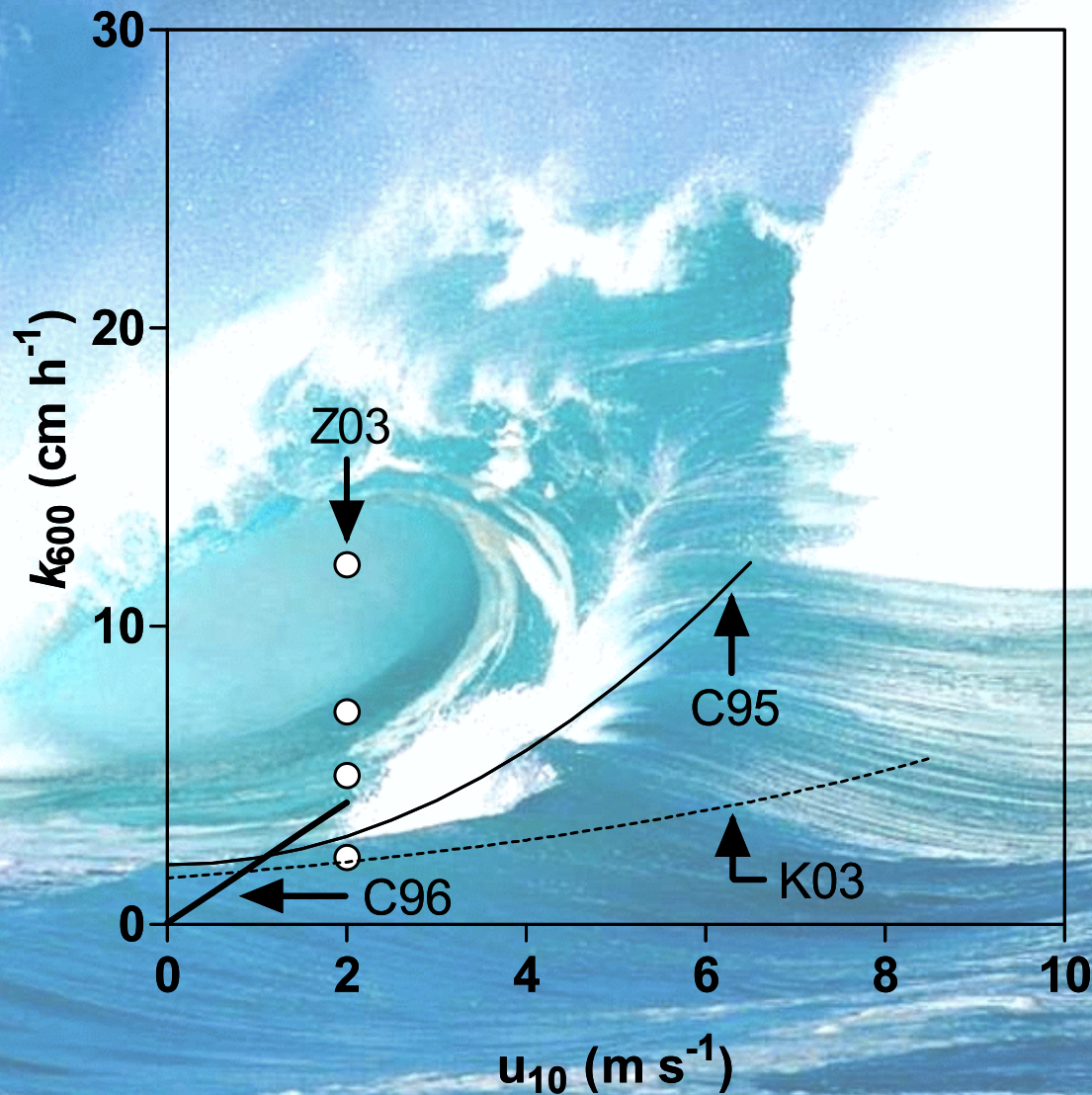
$a$  =  $\text{CO}_2$  solubility coefficient

$dp\text{CO}_2$  = air-water gradient of  $\text{CO}_2$

$k$  = gas transfer velocity of  $\text{CO}_2$

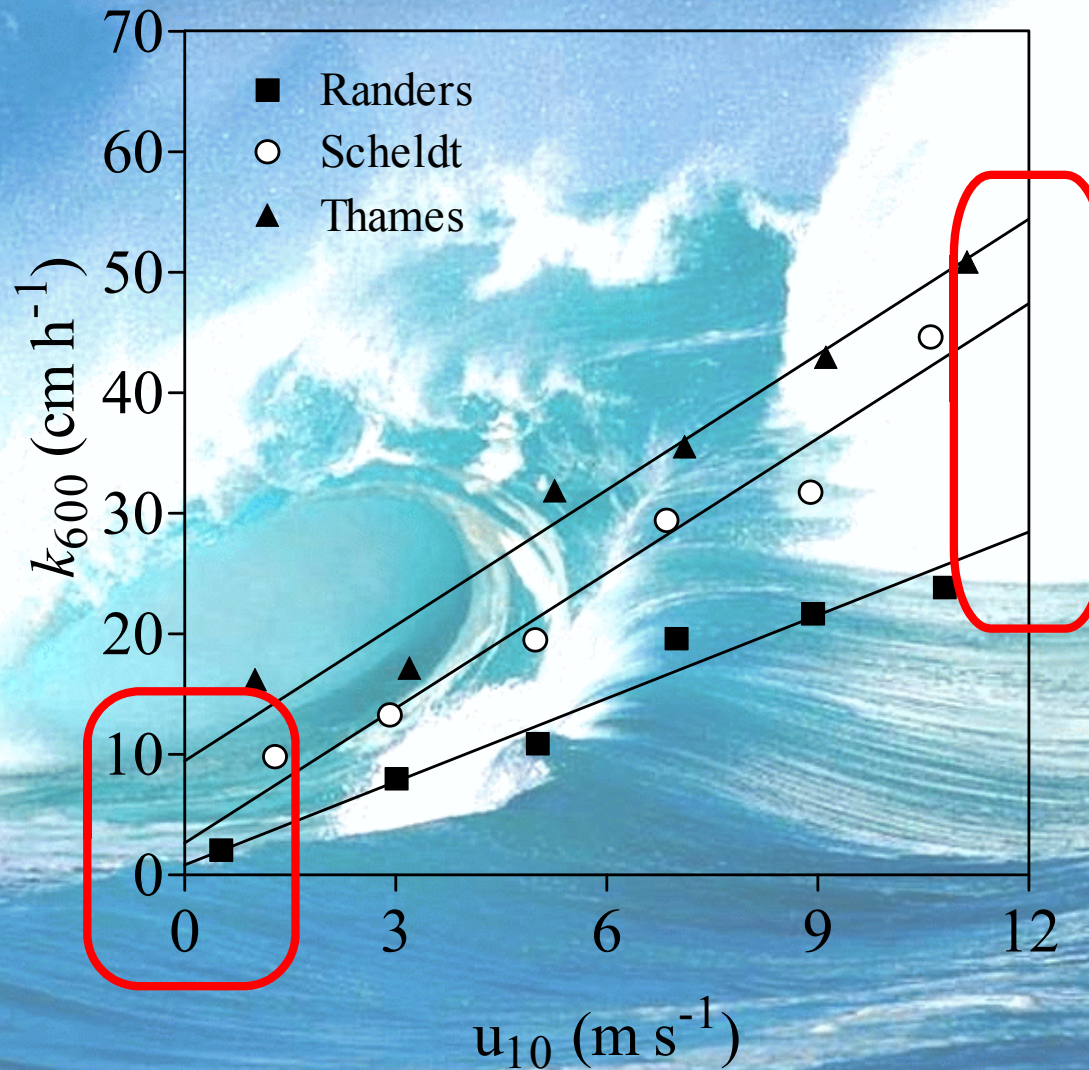
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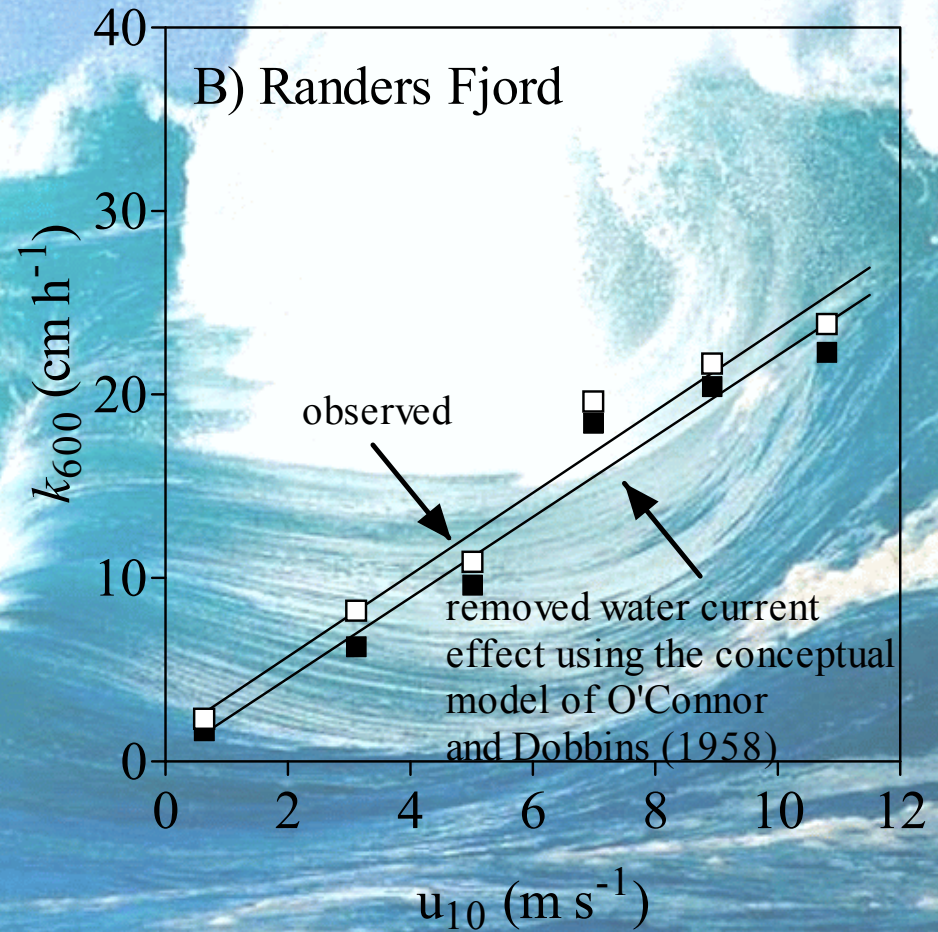
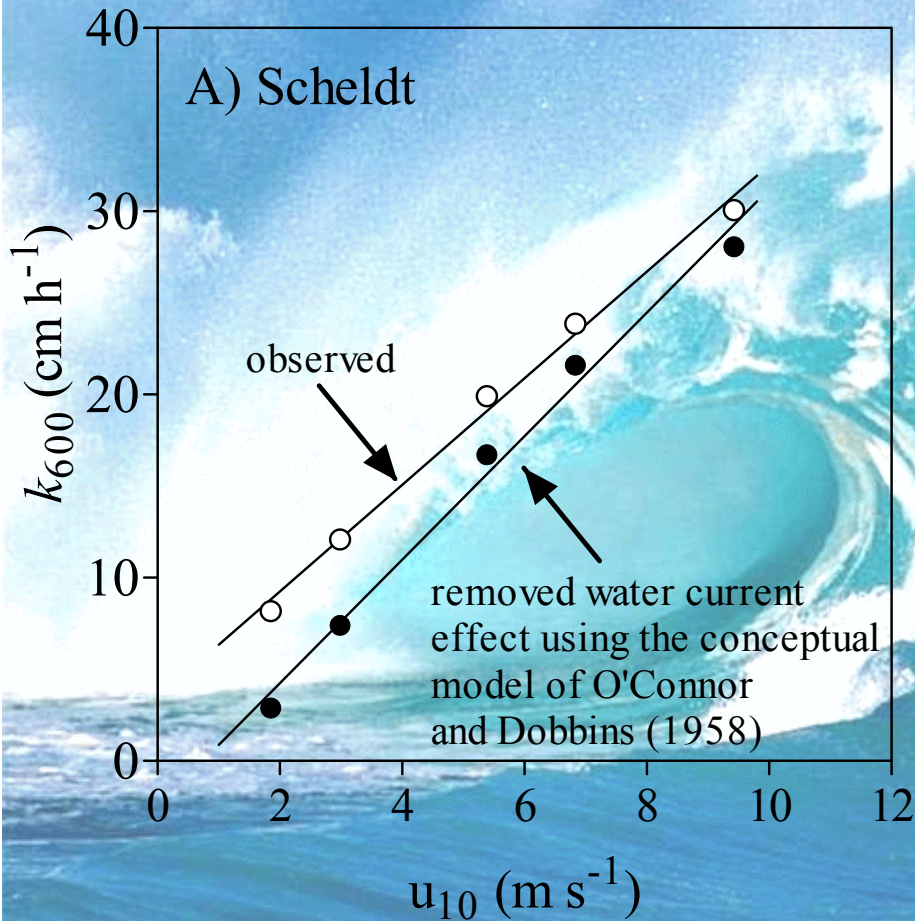
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$$F = a k dp\text{CO}_2$$



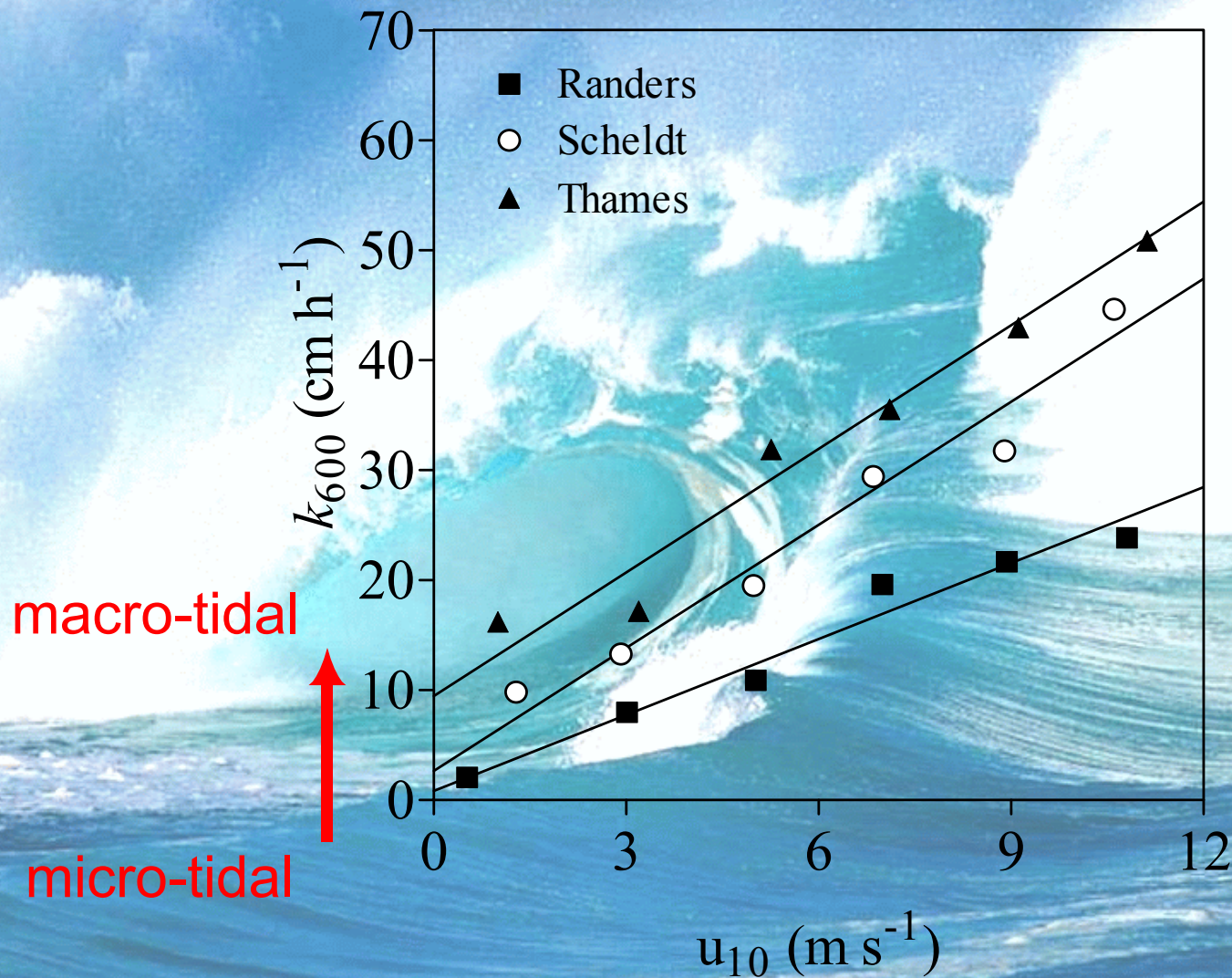
# Gas transfer velocity in estuaries

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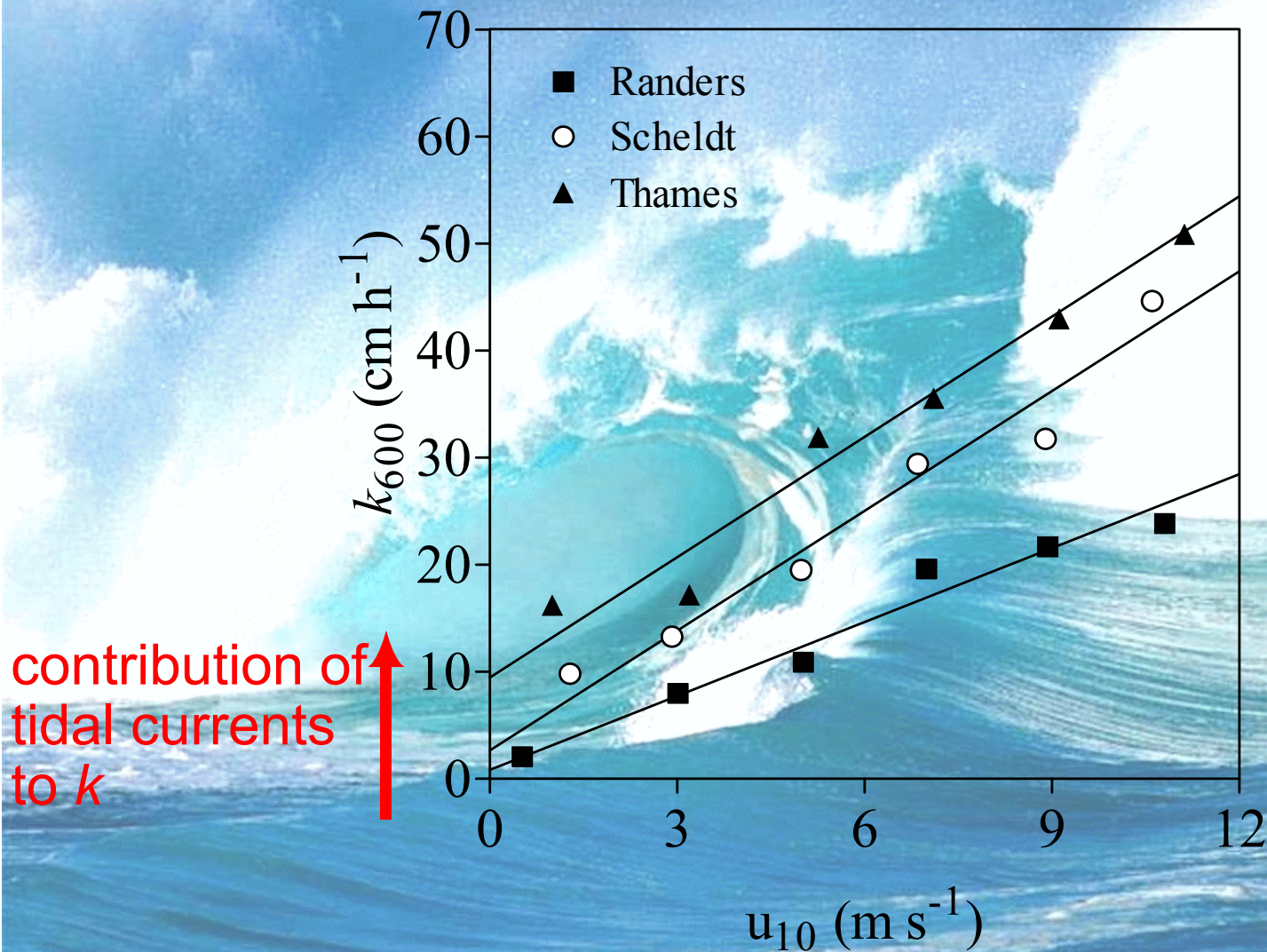
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# Gas transfer velocity in estuaries

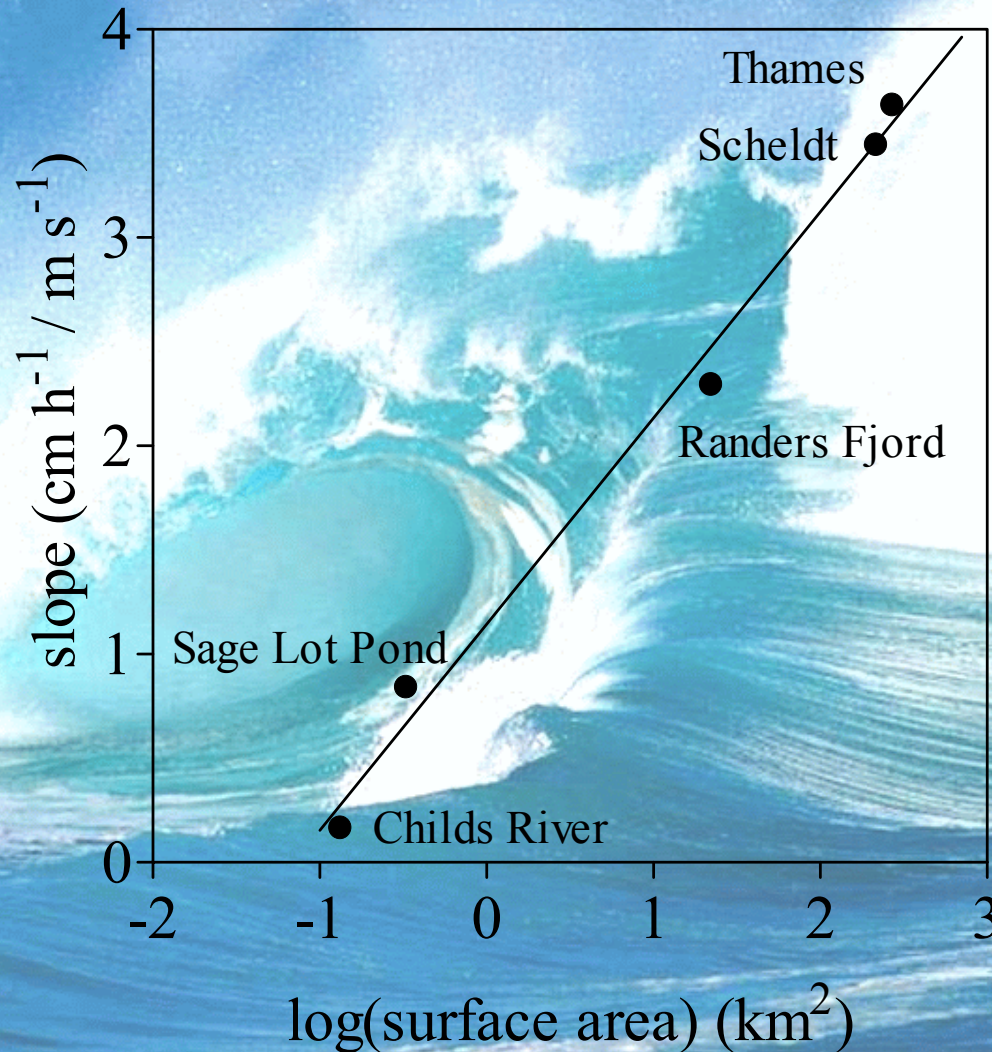
$$F = a k dp\text{CO}_2$$





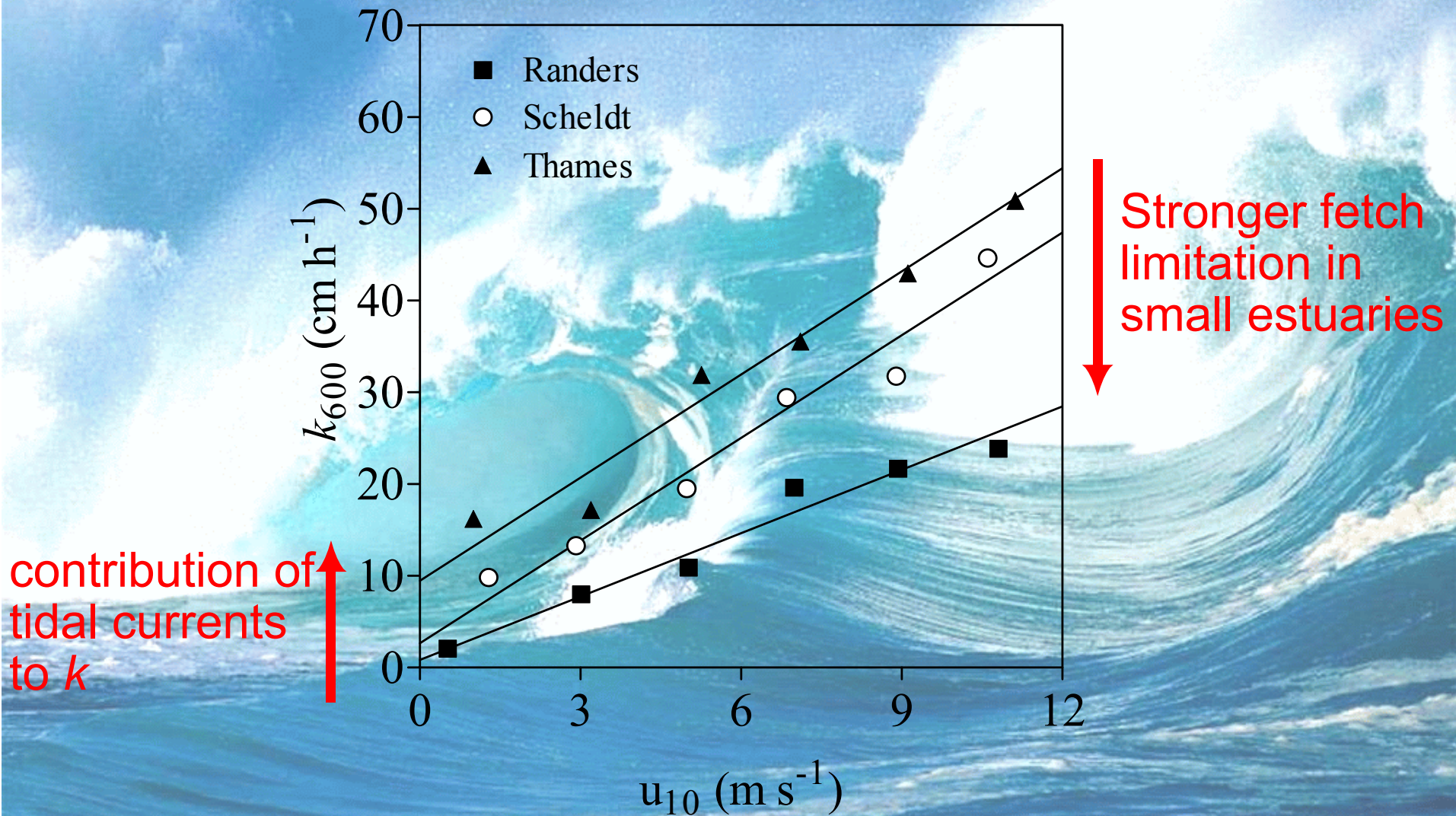
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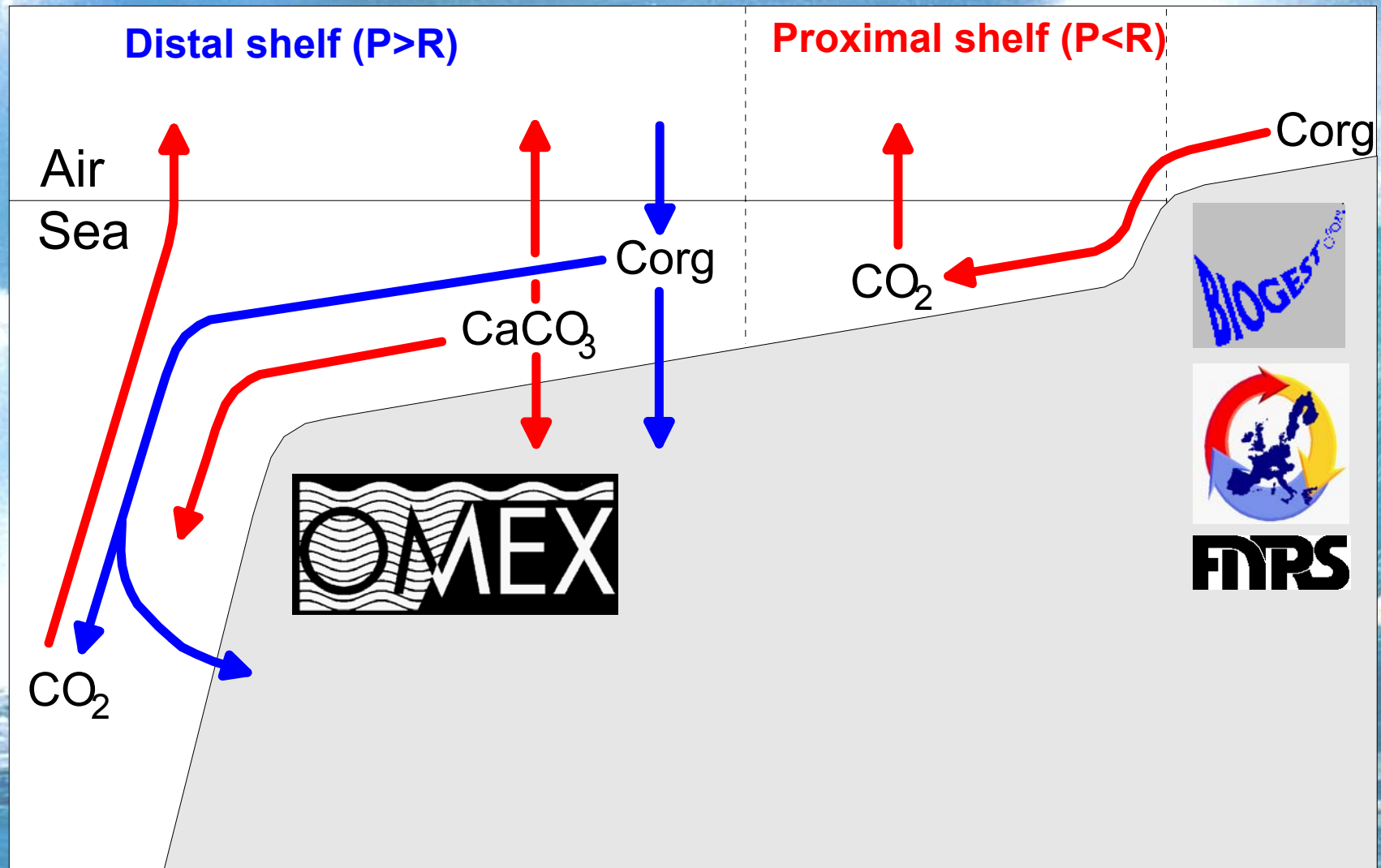


# Gas transfer velocity in estuaries

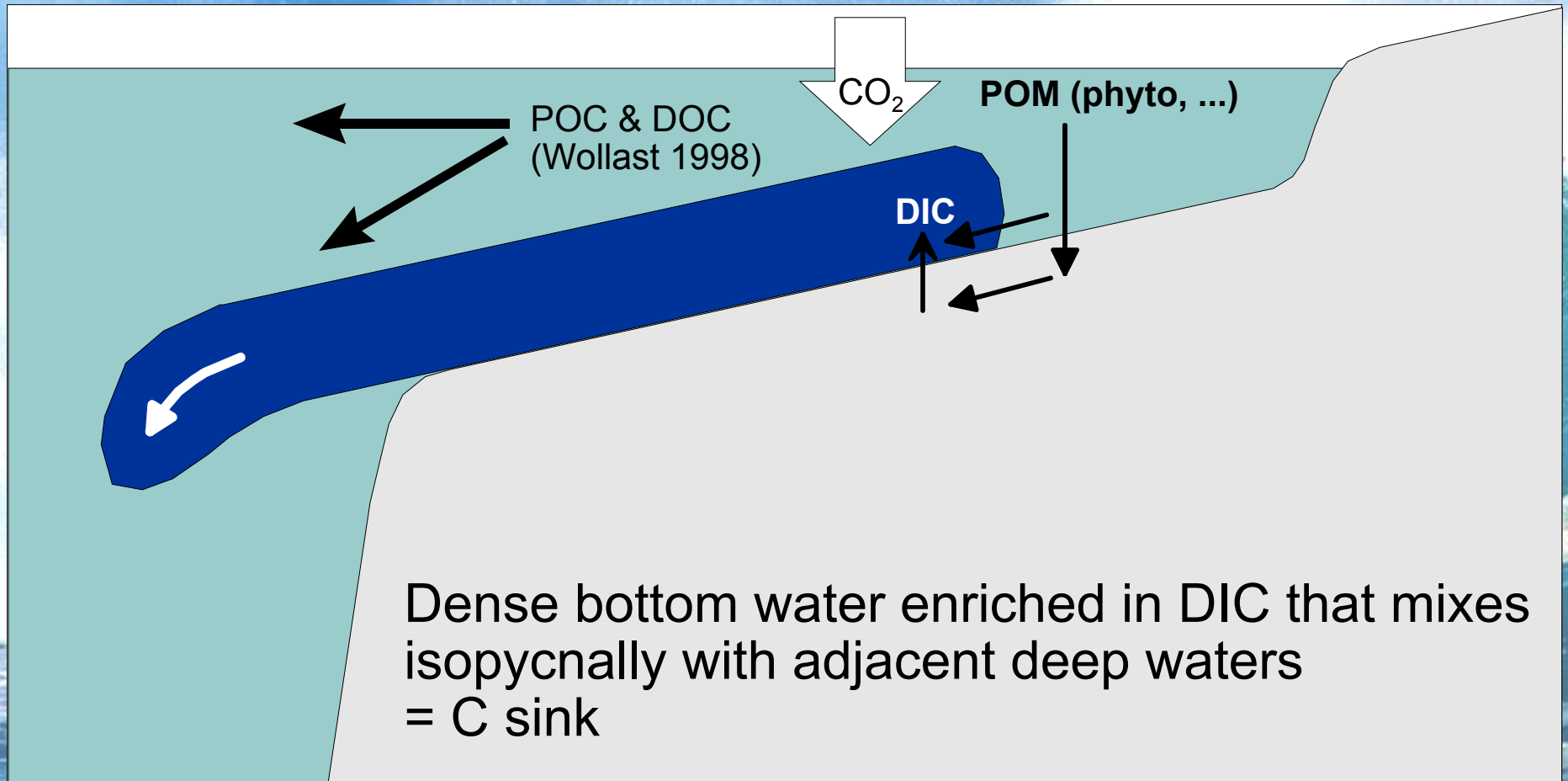
$$F = a k dp\text{CO}_2$$



# Conceptual model of C cycling in shelves



# Continental shelf pump hypothesis Tsunogai et al. (1999)



# Continental shelf pump hypothesis Tsunogai et al. (1999)

5 cruises in the East China Sea

pCO<sub>2</sub> measurements

< ) pCO<sub>2</sub> = -55 ppm

< CO<sub>2</sub> flux of - 8 mmol m<sup>-2</sup> d<sup>-1</sup>

< global shelf surface area = 26 10<sup>6</sup> km<sup>2</sup>

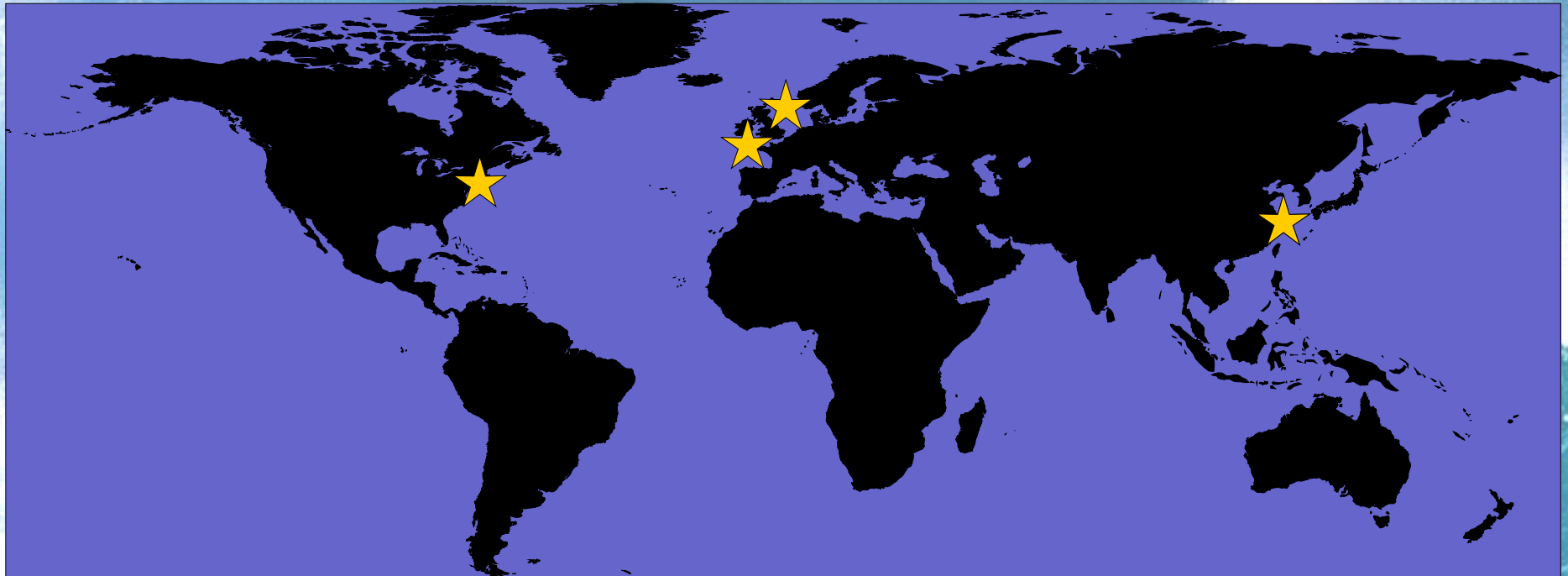
< 1 PgC yr<sup>-1</sup> (10<sup>15</sup> gC yr<sup>-1</sup>)

Open oceanic waters (Takahashi et al. 2002)

< 2.2 PgC yr<sup>-1</sup> (10<sup>15</sup> gC yr<sup>-1</sup>)

< Major revision of oceanic CO<sub>2</sub> pump !!

# Continental shelf pump hypothesis



East China Sea	$-8 \text{ mmol m}^{-2} \text{ d}^{-1}$	Tsunogai et al. (1999)
Gulf of Biscay	$-5 \text{ to } -8 \text{ mmol m}^{-2} \text{ d}^{-1}$	Frankignoulle & Borges (2001)
US MAB	$-2 \text{ to } -3 \text{ mmol m}^{-2} \text{ d}^{-1}$	DeGranpre et al. (2002)
North Sea	$-4 \text{ mmol m}^{-2} \text{ d}^{-1}$	Thomas al. (2004)

**However :**

- < Temperate systems
- < Non-upwelling
- < Little influence of terrestrial inputs

CO<sub>2</sub> fluxes in :

< Upwelling systems } DIC inputs

< Estuaries

< River plumes

} DIC and OM inputs

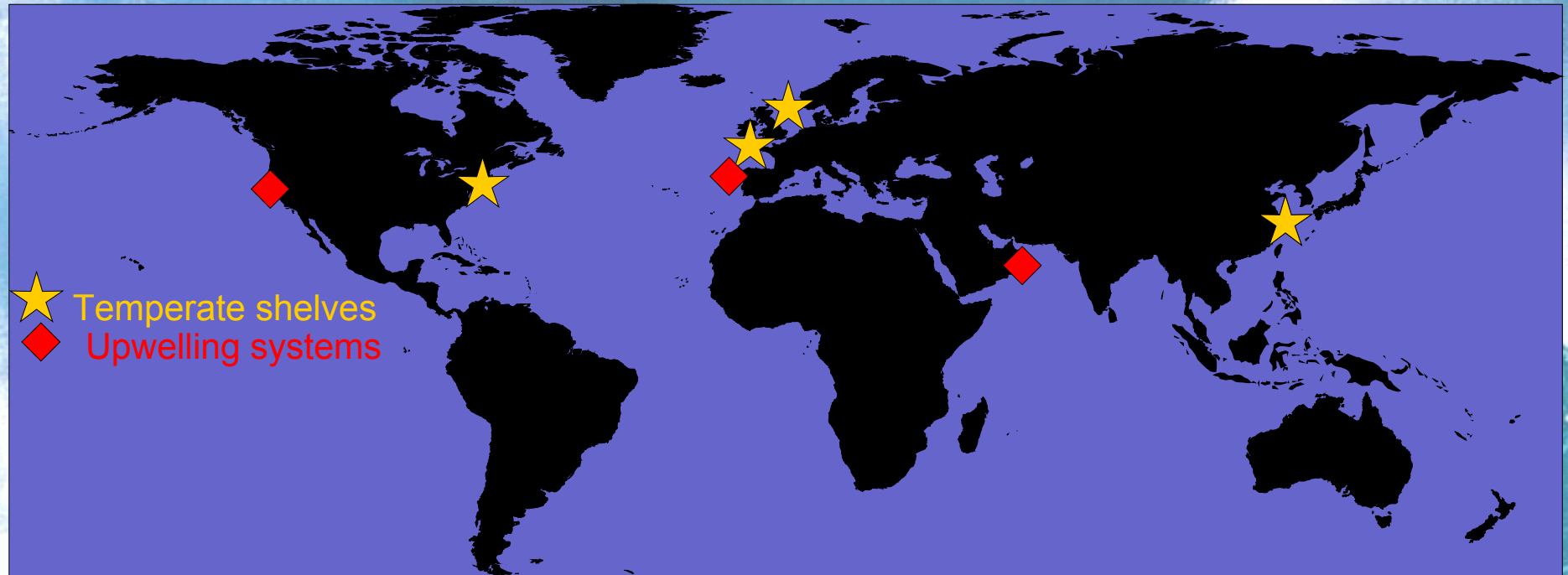
< Sub-tropical shelves

< Coral reefs

< Mangroves

} Non-temperate systems

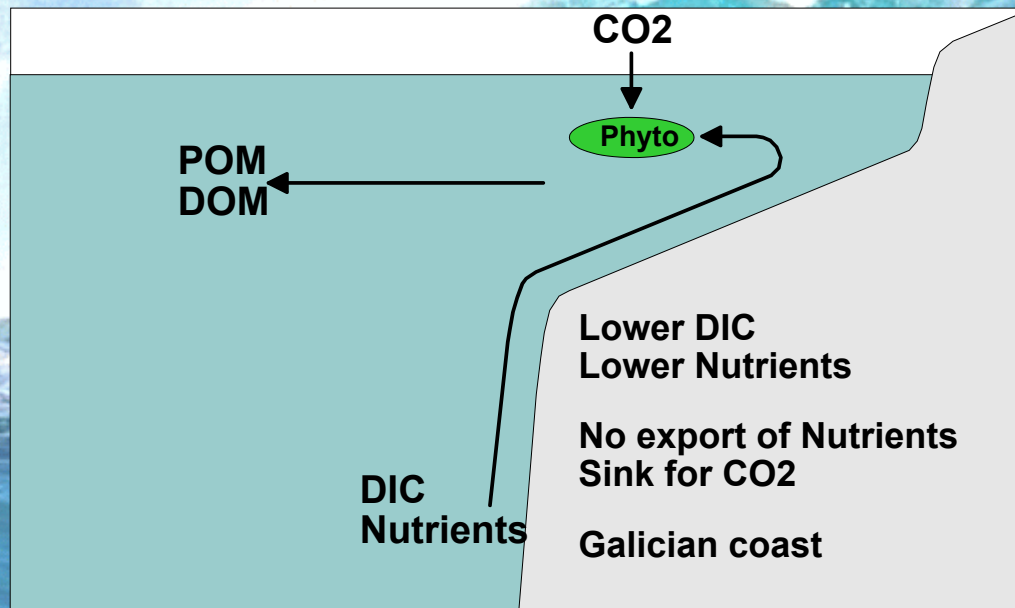
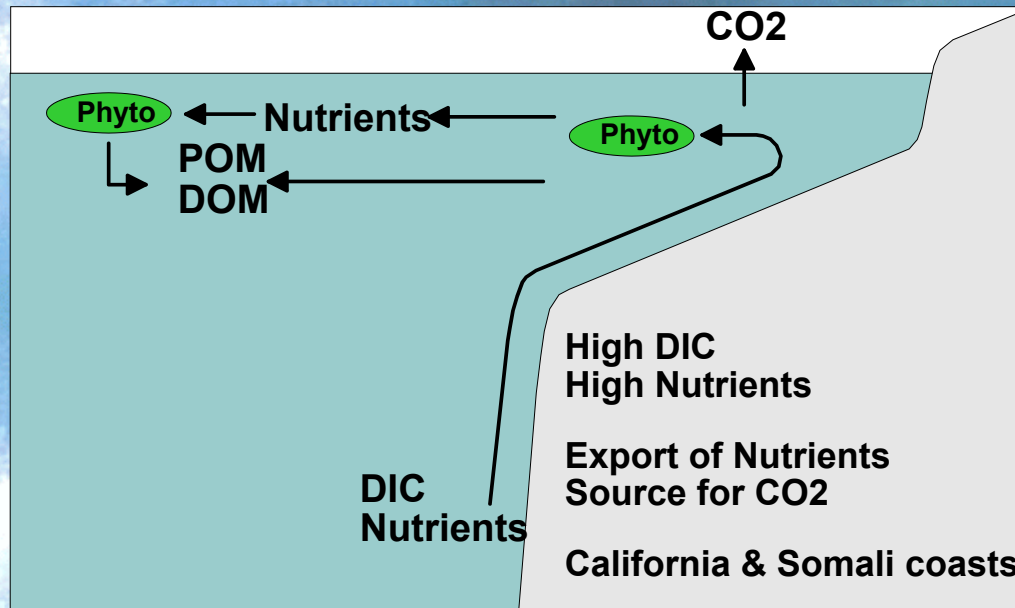
# Upwelling systems



Somali coast	+2.5 mmol m <sup>-2</sup> d <sup>-1</sup>	Goyet et al. (1998)
Galician coast	-3.5 to -7.0 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges & Frankignoulle (2002)
California coast	+4.1 to +6.0 mmol m <sup>-2</sup> d <sup>-1</sup>	Friederich et al. (2002)



# Upwelling systems



California coast  
(Friederich et al. 2002)

El Nina  
< 4.1 to 6.0 mmol m<sup>-2</sup> d<sup>-1</sup>

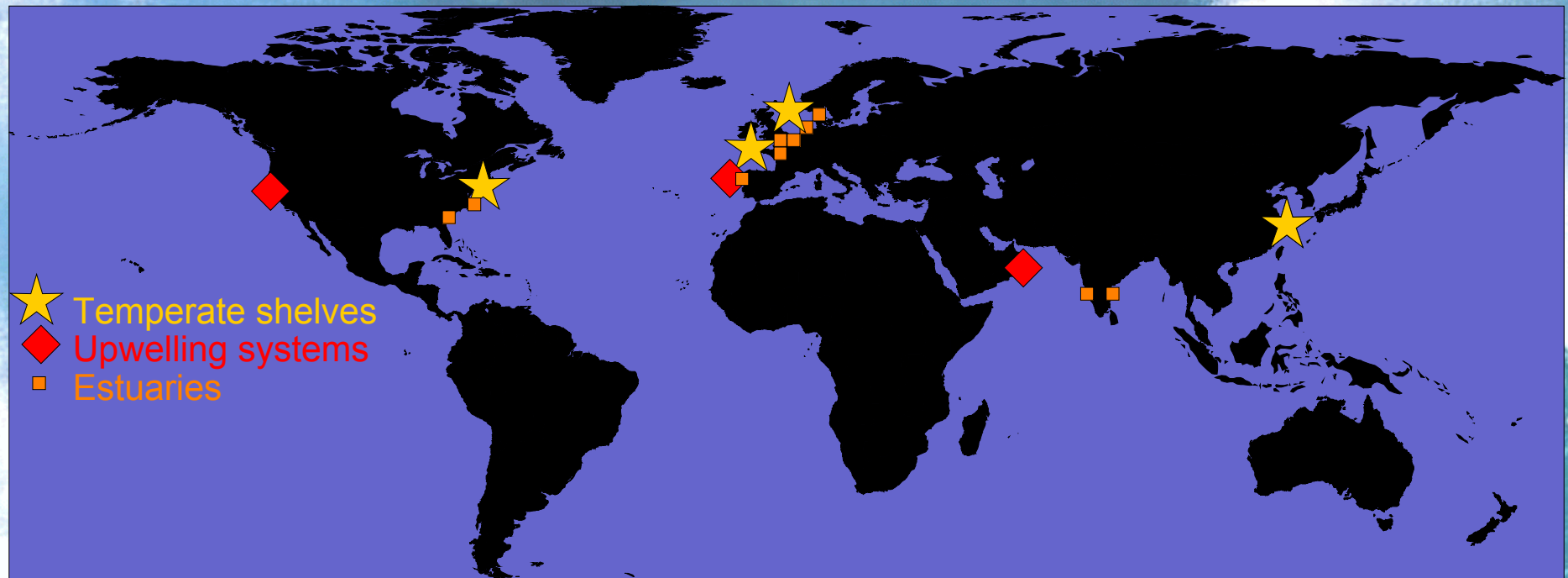
El Nino  
< -1 to -2.0 mmol m<sup>-2</sup> d<sup>-1</sup>



# Estuaries

- < Receive massive inputs (riverine and lateral) of dissolved inorganic carbon, organic matter and nutrients
- < Long residence time promotes biogeochemical transformations during estuarine transit
- < Residence time is highly variable (days to months) !

# Estuaries



European estuaries (11):

Randers Fjord, Elbe, Ems, Rhine, Scheldt, Thames, Tamar, Loire, Gironde, Douro & Sado = +40 to +210  $\text{mmol m}^{-2} \text{d}^{-1}$

Frankignoulle et al. (1998), Abril et al. (1998) & Borges (unpublished)

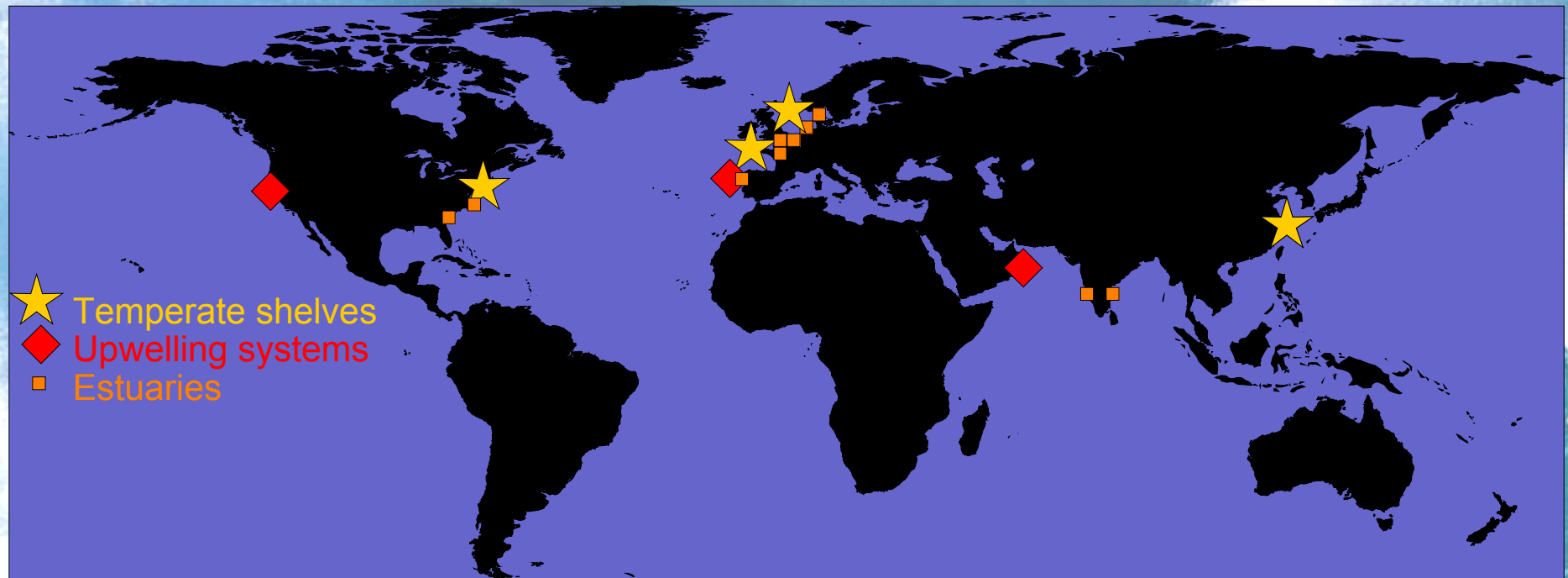
York & Satilla Rivers +17 to 116  $\text{mmol m}^{-2} \text{d}^{-1}$

Raymond et al. (2000) and Cai & Wang (1998)

Mandovi-Zuari & Godavari +15 to +39  $\text{mmol m}^{-2} \text{d}^{-1}$

Sarma et al. (2001) & Bouillon et al. (2003)

# Estuaries



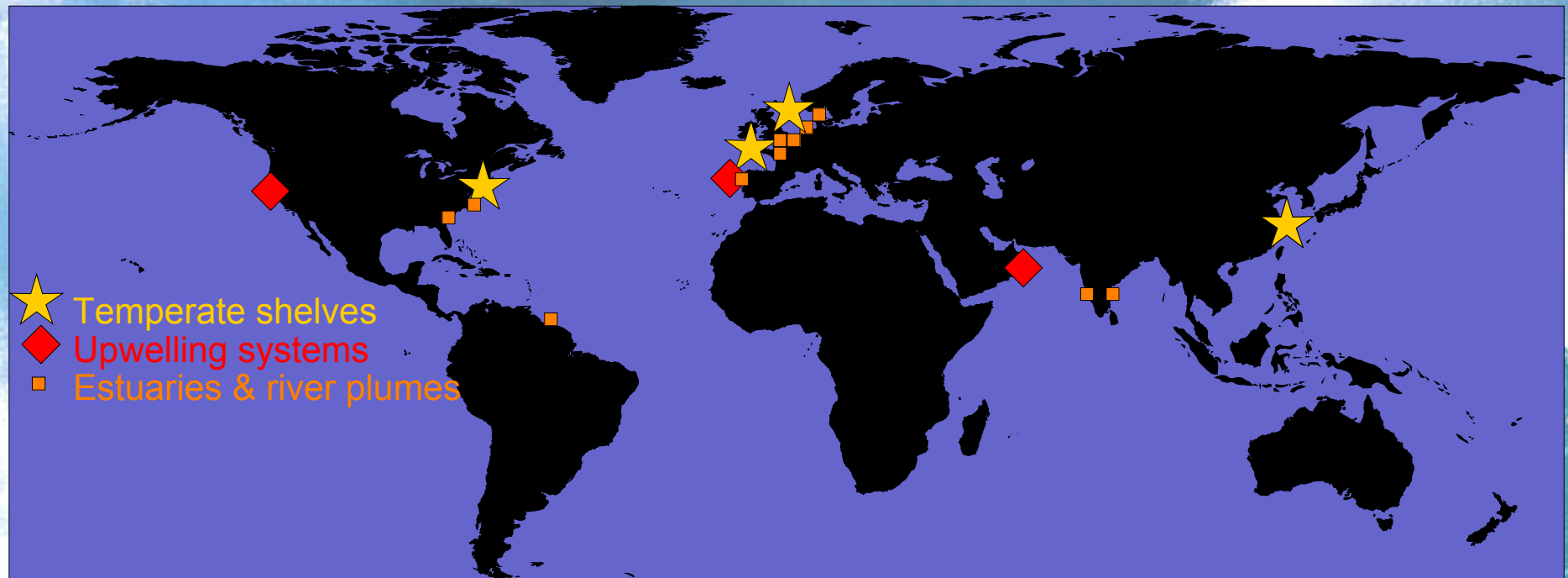
CO<sub>2</sub> source (temperate sub-tropical and tropical)  
CO<sub>2</sub> fluxes +20 to +200 mmol m<sup>-2</sup> d<sup>-1</sup>  
CO<sub>2</sub> fluxes one to two orders of magnitude  
higher than open shelves



# River plumes

- < Receive inputs of dissolved inorganic carbon, organic matter and nutrients from inner estuaries
- < Lower end of the mixing with seawater end-member
- < Decrease of light limitation in relation to inner estuaries
- < Biogeochemical processes less intense than inner estuaries

# River Plumes



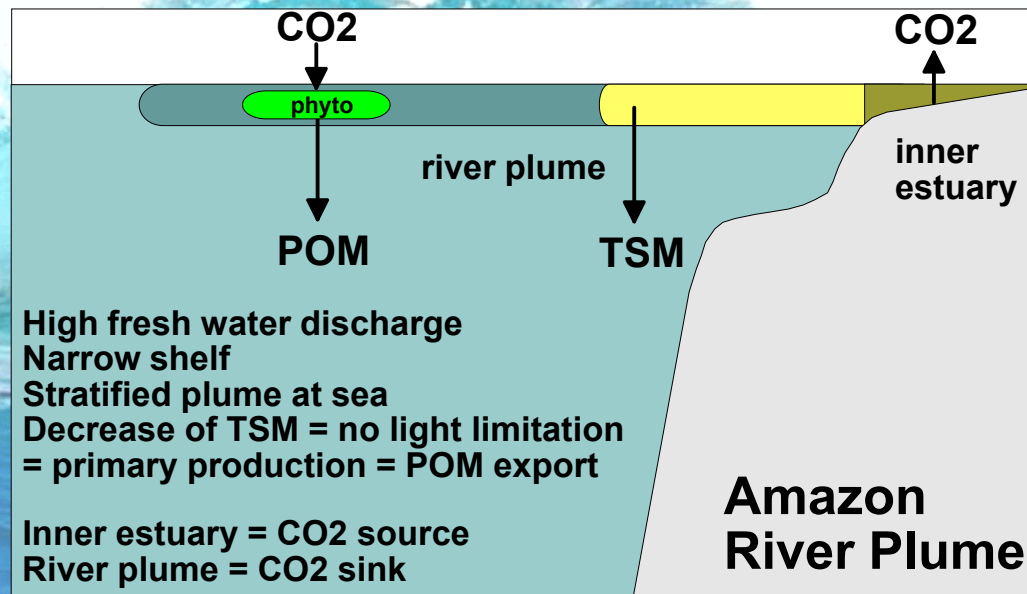
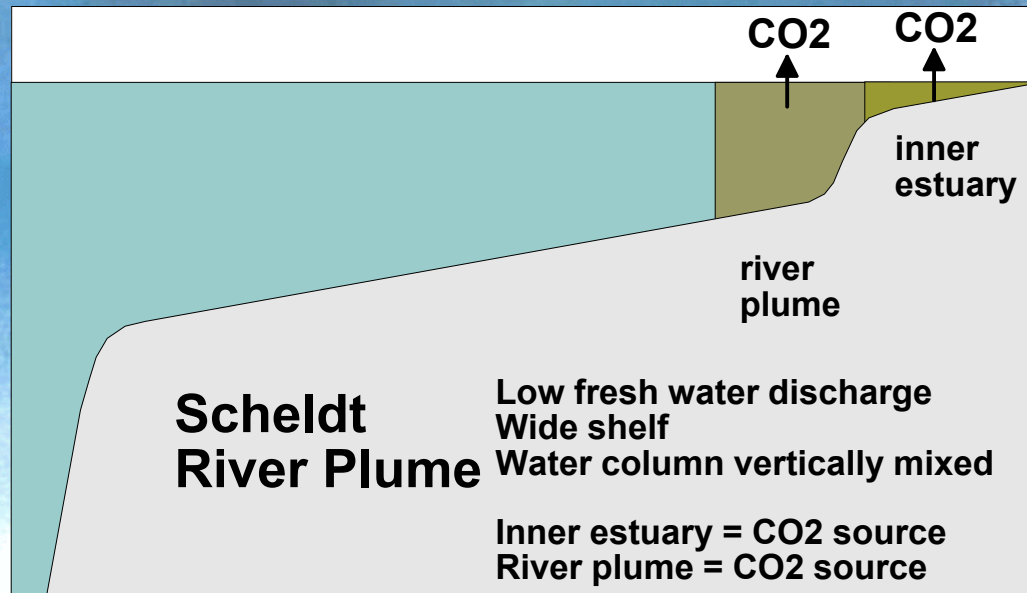
Scheldt river plume  
 $+3$  to  $+5 \text{ mmol m}^{-2} \text{ d}^{-1}$

Borges & Frankignoulle (2002)

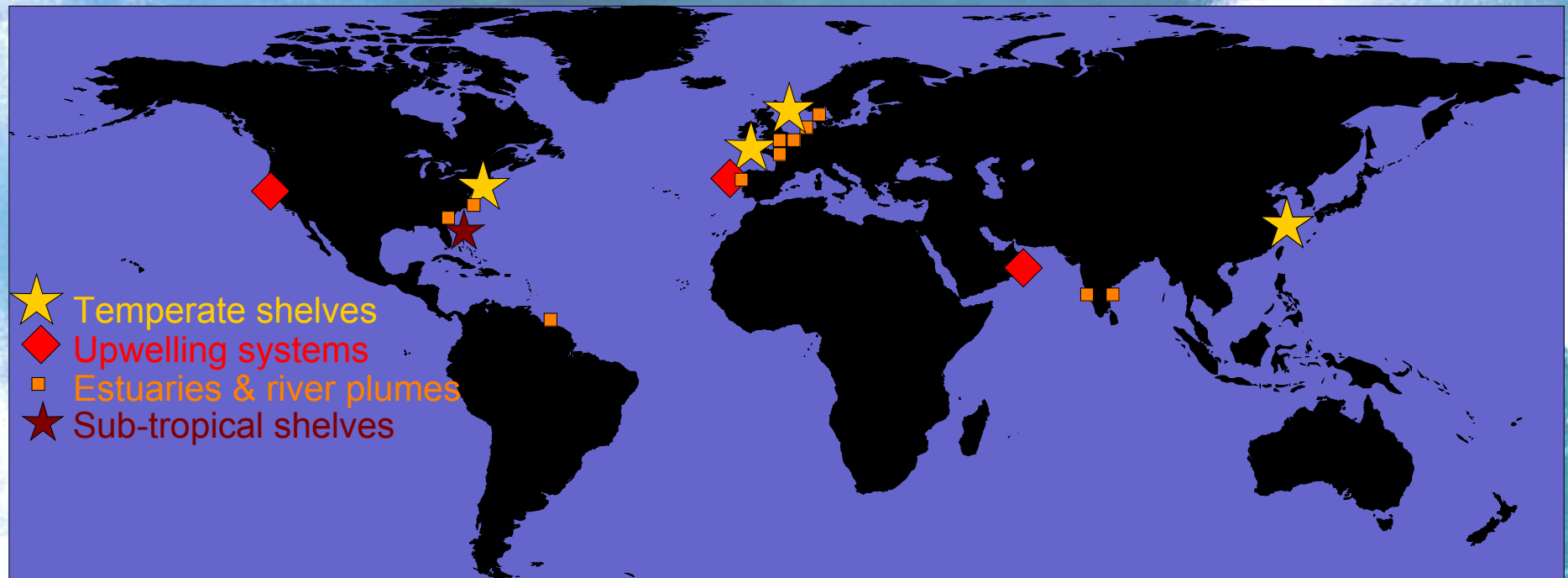
Amazon river plume  
 $-1.4 \text{ mmol m}^{-2} \text{ d}^{-1}$

Körtzinger (2003)

# River plumes



# Sub-tropical shelves



US South Atlantic Bight

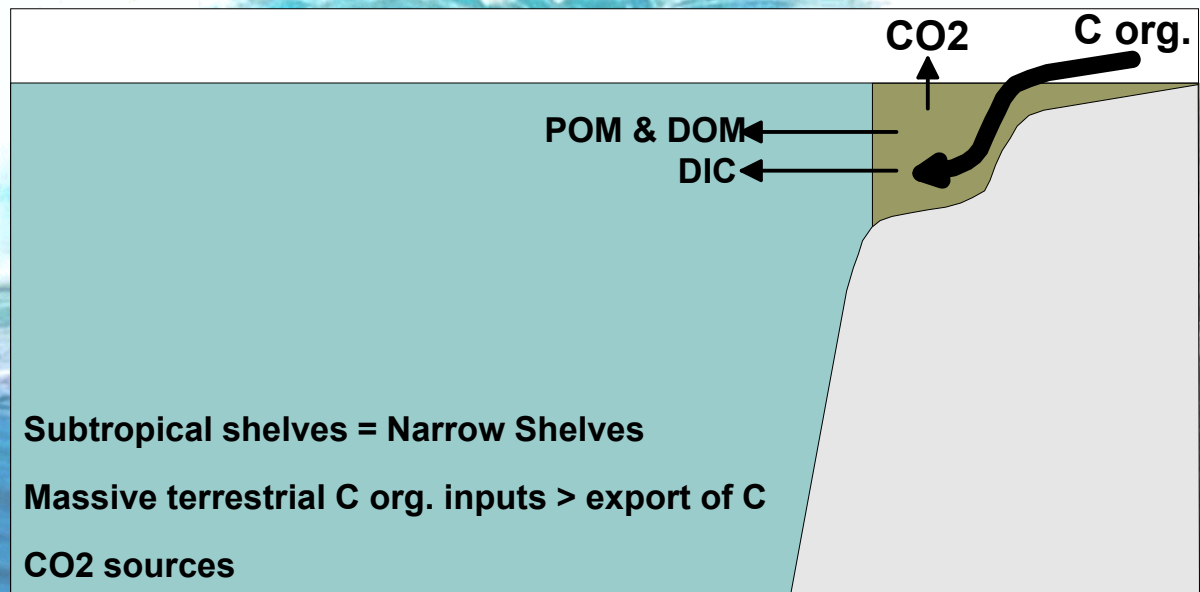
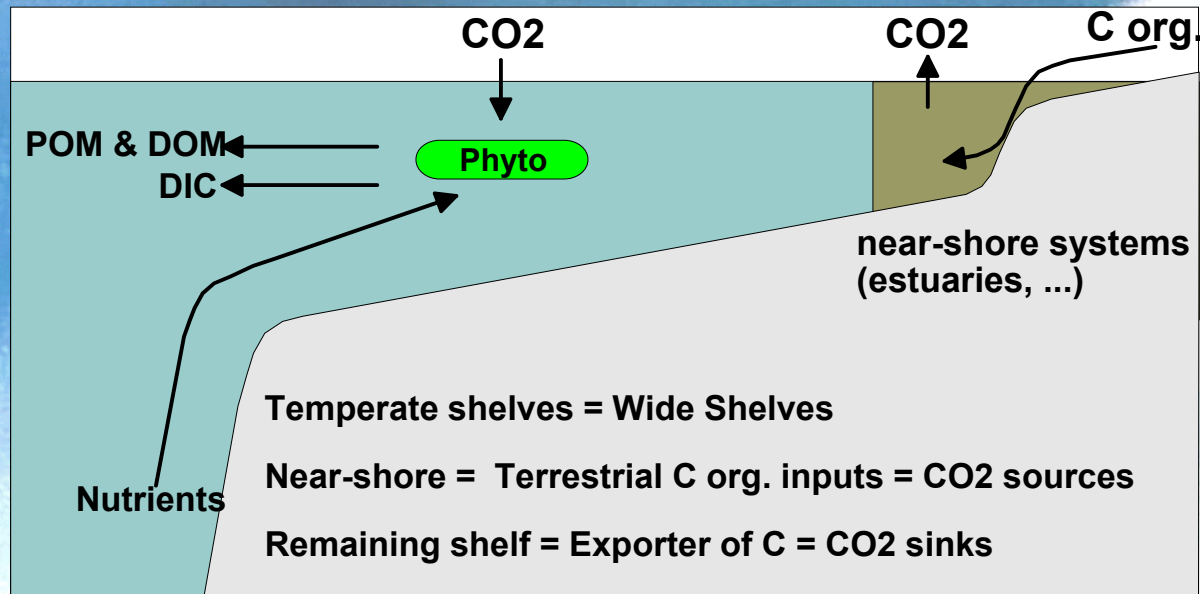
+7 mmol m<sup>-2</sup> d<sup>-1</sup>

Cai et al. (2003)

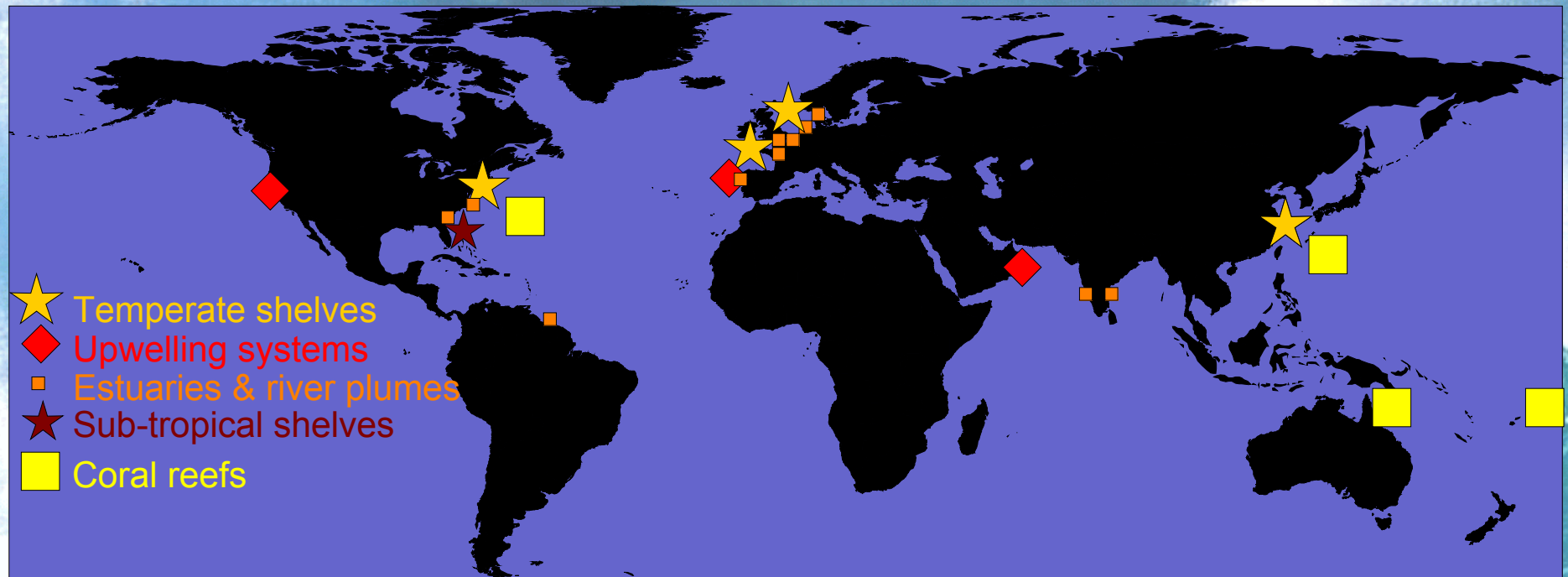
CO<sub>2</sub> fluxes of same order of magnitude as temperate shelves but OPPOSITE sign



# Sub-tropical shelves

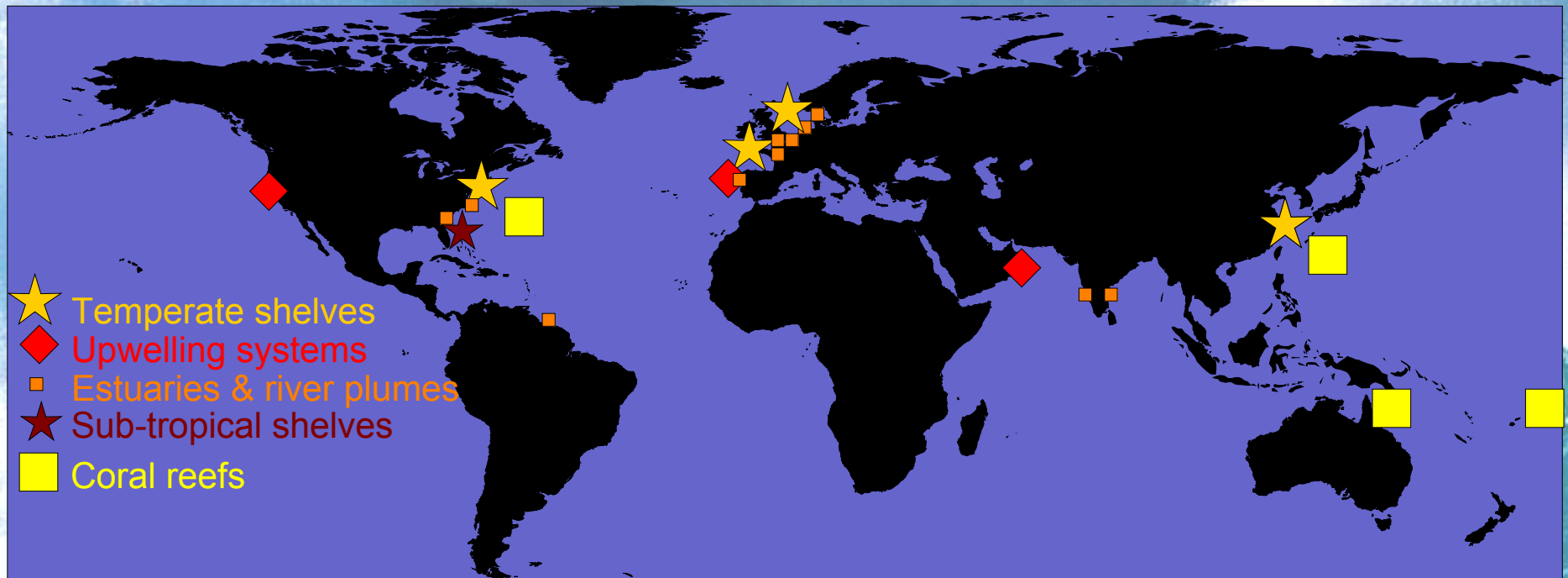


# Coral reefs



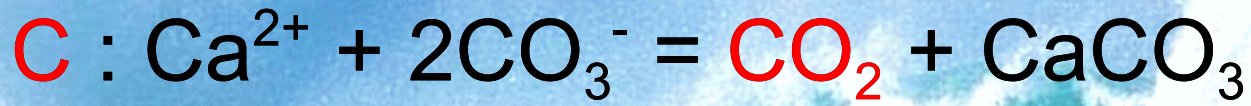
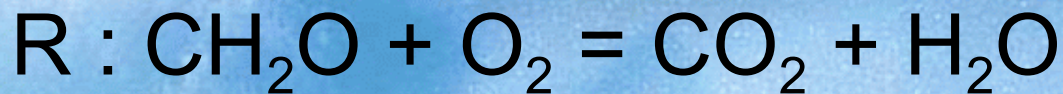
Moorea	+0.4 mmol m <sup>-2</sup> d <sup>-1</sup>	Gattuso et al. (1993; 1997)
Yonge Reef	+4.1 mmol m <sup>-2</sup> d <sup>-1</sup>	Frankignoulle (1996)
Okinawa	+5.0 mmol m <sup>-2</sup> d <sup>-1</sup>	Ohde & van Woesik (1999)
Hog Reef	+3.3 mmol m <sup>-2</sup> d <sup>-1</sup>	Bates et al. (2001)

# Coral reefs



Coral reefs are a moderate source of CO<sub>2</sub>

# Coral Reefs



$$P-R = 0$$

$$C > 0 \quad (2.2 \text{ gCaCO}_3 \text{ m}^{-2} \text{ d}^{-1})$$

## English Channel

$$) \text{ pCO}_2 = 0 \quad \text{Borges \& Frankignoulle (2003)}$$

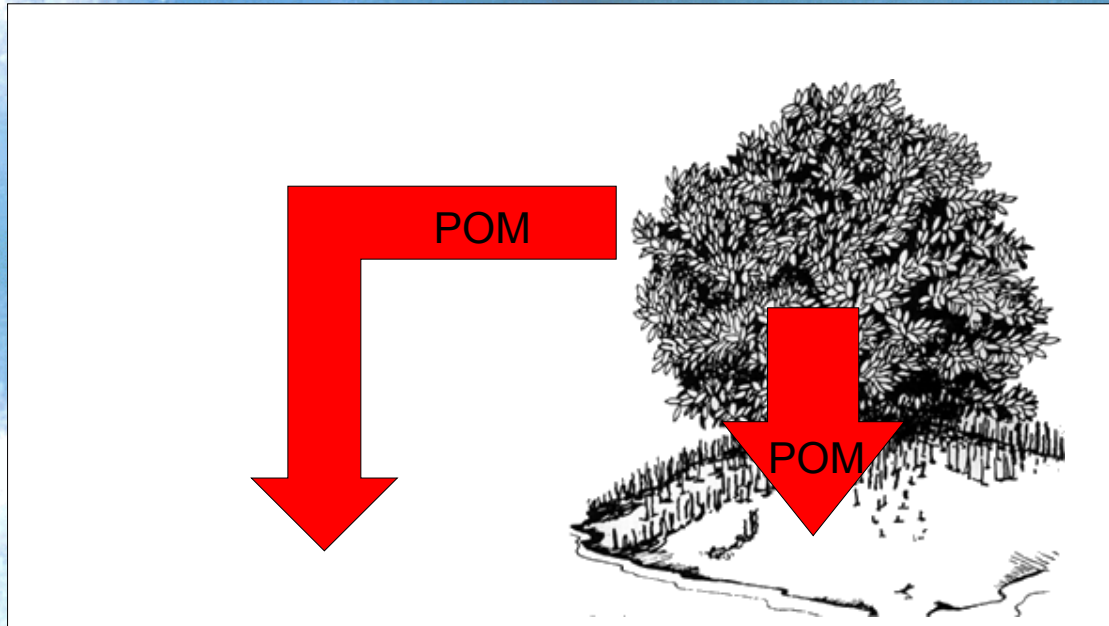
$$P-R = 1.0 \text{ mmol m}^{-2} \text{ d}^{-1}$$

Le Corre et al. (1996) L'Helguen et al. (1996)

$$\text{CO}_2 \text{ release} = 0.9 \text{ mmol m}^{-2} \text{ d}^{-1}$$

by brittle star populations (Migné et al. 1998)

# Mangrove surrounding waters

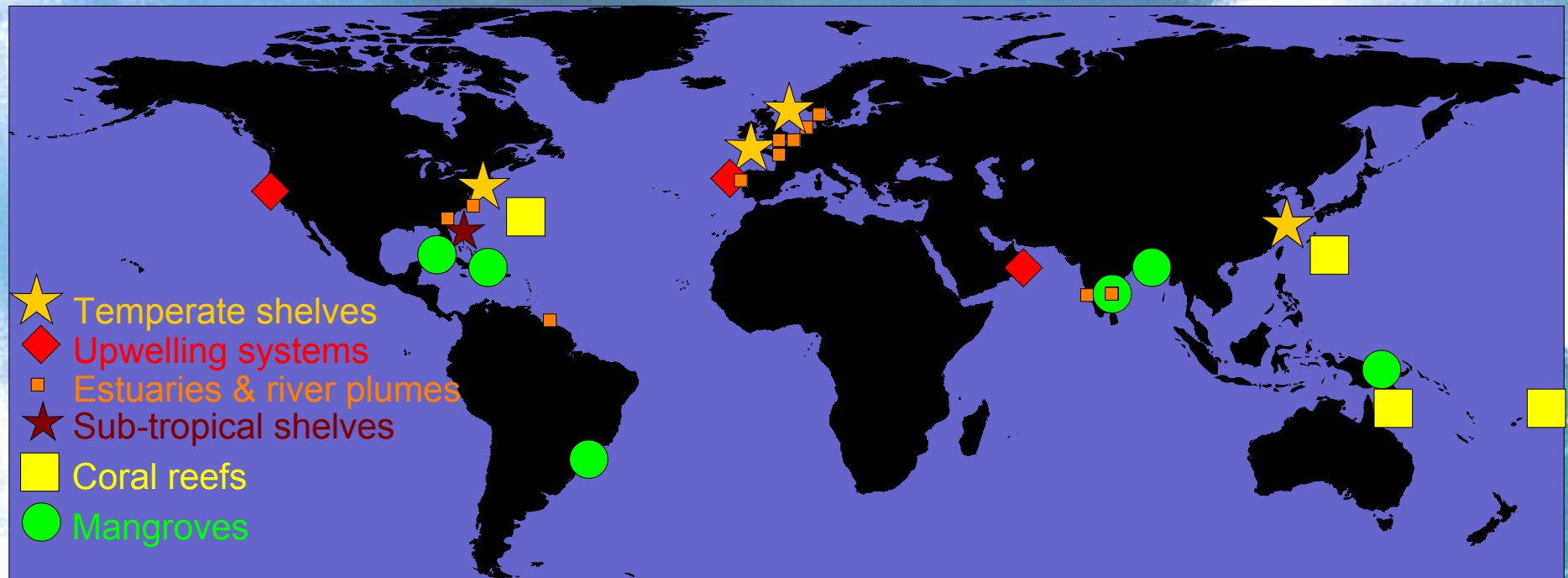


Leaf and wood litter transfer to water and sediments

Thus, water column and sediments metabolisms are net heterotrophic

Thus, water column should be a  $\text{CO}_2$  source

# Mangrove surrounding waters



Nagada Creek	+44 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003)
Saptamukhi Creek	+57 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003) based on Ghosh et al.(1987)
Mooringanga Creek	+23 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003) based on Ghosh et al.(1987)
Gaderu Creek	+56 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003)
Itacuraça Creek	+23 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003) based on Ovalle et al.(1991)
Norman's Pond	+14 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003)
Florida Bay	+3 mmol m <sup>-2</sup> d <sup>-1</sup>	Borges et al.(2003) based on Millero et al.(2001)

# Mangrove surrounding waters

< +50 mmol m<sup>-2</sup> d<sup>-1</sup>

< 0.2 10<sup>6</sup> km<sup>2</sup>

< CO<sub>2</sub> emission of 0.05 PgC yr<sup>-1</sup>

CO<sub>2</sub> emission in tropical sub-tropical oceanic waters

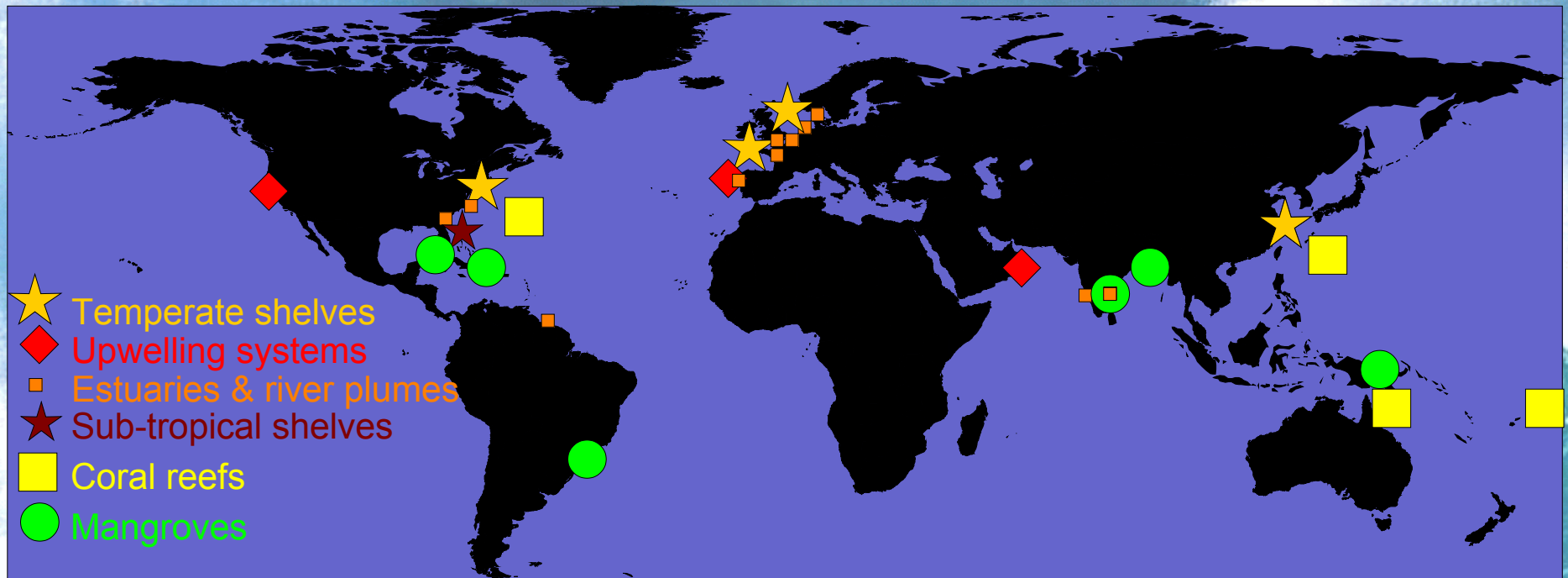
< 0.43 PgC yr<sup>-1</sup> (Takahashi et al. 1997)

< Mangrove surrounding waters provide an extra CO<sub>2</sub> emission of 12% at sub-tropical and tropical latitudes for a surface area 1000 times smaller than open oceanic waters !

# Conclusions

- < Coastal ocean as sink of  $\text{CO}_2$  of  $1.0 \text{ PgC yr}^{-1}$  as formulated by Tsunogai et al. (1999) would imply a major revision of the oceanic  $\text{CO}_2$  pump
- < Temperate wide shelves = sink of  $\text{CO}_2$  ( $-2$  to  $-5 \text{ mmol m}^{-2} \text{ d}^{-1}$ )
- < Sub-tropical (and tropical ?) shelves = sources of  $\text{CO}_2$  ( $+7 \text{ mmol m}^{-2} \text{ d}^{-1}$ )
- < Temperate, sub-tropical and tropical estuaries = sources of  $\text{CO}_2$  ( $+20$  to  $+200 \text{ mmol m}^{-2} \text{ d}^{-1}$ )
- < River plumes = ?
- < Coral reefs = sources of  $\text{CO}_2$  ( $+1$  to  $+5 \text{ mmol m}^{-2} \text{ d}^{-1}$ )
- < Mangrove surrounding waters = sources of  $\text{CO}_2$  ( $+50 \text{ mmol m}^{-2} \text{ d}^{-1}$ )
- < The geographic and ecological **diversity** of the Coastal Ocean **must** be accounted when integrating at global scale





Thank you