

## **Stomatal and non-stomatal responses of typical temperate C3 crops to soil water stress**

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**Stomatal and non-stomatal responses of typical temperate  
C3 crops to soil water stress**

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## Abstract

Photosynthesis plays a key role in the carbon cycle, being the primary mechanism through which carbon dioxide ( $\text{CO}_2$ ) is converted into organic compounds by plants, algae, and certain bacteria. On average, terrestrial ecosystems uptake around one-third of anthropogenic  $\text{CO}_2$  emissions and participate to the mitigation of climate change. Temperature, solar irradiance or soil water availability are important factors mediating the strength of terrestrial ecosystems. Therefore, drought poses a major threat on the atmospheric  $\text{CO}_2$  balance due to its profound implications for plant survival, ecosystem dynamics and agricultural productivity.

Water stress induces a cascade of physiological responses in plant functioning. During hot and dry weather, plants often close their stomata to save water and prevent dehydration at the expense of a decrease in  $\text{CO}_2$  supply to the chloroplasts, which results in a decrease in both carbon assimilation and transpiration. In addition, photosynthesis and stomatal closure can be impacted by non-stomatal factors, which complexifies the mechanisms regulating plant responses to water stress. Models of photosynthesis and stomatal conductance still lack from a detailed characterization of the intertwined physiological processes under soil water-limiting conditions. Understanding, disentangling and quantifying the importance of photosynthesis limiting factors is crucial for selecting drought-tolerant varieties and for improving model predictions.

The gold standard method for estimating photosynthesis is the measurement of the  $\text{CO}_2$  net assimilation rate by gas exchange techniques. While providing key information on the factors influencing the temporal variability of photosynthesis, gas exchanges alone do not allow to fully characterize the limitations on carbon assimilation under water stress. Additional measurements of chlorophyll fluorescence by active methods are needed to improve the knowledge about these constraints and to quantify the importance of each limiting factors. In particular, the response of mesophyll conductance to water stress remains a pivotal uncertainty in models of photosynthesis.

The interpretation of actively-induced chlorophyll fluorescence measurements is however limited to the leaf scale and does not allow to elucidate the impact of water stress on large patches of vegetation. The recent emergence of passive techniques has allowed the monitoring of fluorescence emission induced by solar irradiance (SIF) at different temporal scales and spatial resolutions. These measurements provide a promising indicator of vegetation physiological processes which can be used to calibrate empirical models to estimate carbon assimilation at local and large scales. However, such relationships typically fail in reproducing photosynthesis temporal evolution at short timescales or during climate extremes such as droughts or heatwaves. More mechanistic approaches are needed to fully exploit the physiological message carried by SIF measurements.

In this thesis, we first aimed to unravel the origins of photosynthesis limitations under water stress. The first chapters are dedicated to providing a calibration of the

response of stomatal and non-stomatal factors to the decrease in soil water availability. This method, first applied on forest ecosystems (chapter 3) and potato (chapter 4) with eddy covariance (EC) data, emphasizes the key role played by non-stomatal factors in regulating canopy-scale photosynthesis for these two very different plant functional types. By performing measurements of leaf-level chlorophyll fluorescence by active methods, and by using a new partitioning method, we also showed (chapter 5) that mesophyll conductance is a pivotal trait which regulates both carbon assimilation and stomatal conductance of potato. These chapters provide functions of soil water availability for stomatal and non-stomatal factors which can be used to reduce uncertainties in climate models. Finally, we evaluated the capability of a new mechanistic approach to model CO<sub>2</sub> assimilation and transpiration of a winter wheat crop from proximal sensing of SIF (chapter 6). A very strong correlation between model estimates and EC measurements was found across a wide range of environmental conditions including edaphic drought. A sensitivity analysis revealed that the fraction of open photosystem II centers is a key parameter affecting model robustness. This last chapter paves the way towards an improvement of the understanding of the interactions between the water and carbon cycles by using the physiological information carried by SIF.

## Résumé

La photosynthèse joue un rôle essentiel dans le cycle du carbone, en tant que mécanisme principal par lequel le dioxyde de carbone ( $\text{CO}_2$ ) est converti en composés organiques par les plantes, les algues et certaines bactéries. En moyenne, les écosystèmes terrestres absorbent environ un tiers des émissions anthropiques de  $\text{CO}_2$ , contribuant ainsi à la lutte contre le changement climatique. La température, l'irradiance solaire ou la disponibilité en eau du sol sont des facteurs importants qui influencent la force des puits de carbone terrestre. Par conséquent, la sécheresse constitue une menace majeure pour le cycle du carbone, en raison de ses implications profondes pour la survie des plantes, la dynamique des écosystèmes et la productivité agricole.

Le stress hydrique induit une cascade de réponses physiologiques et biochimiques chez les plantes. Par temps chaud et sec, les plantes ont tendance à fermer leurs stomates pour économiser de l'eau et éviter la déshydratation, au détriment d'une diminution de l'approvisionnement en  $\text{CO}_2$  vers les chloroplastes, entraînant ainsi une réduction à la fois de l'assimilation du carbone et de la transpiration. De plus, la photosynthèse et la fermeture stomatique peuvent être affectées par des facteurs non-stomatiques comme les capacités photosynthétiques ou la conductance mésophyllienne, ce qui complexifie les mécanismes régulant les réponses des plantes au stress hydrique. Les modèles de photosynthèse et de conductance stomatique manquent encore d'une caractérisation détaillée des processus physiologiques interconnectés dans des conditions de sécheresse édaphique. Comprendre, quantifier et clarifier l'importance des facteurs limitants la photosynthèse est crucial pour sélectionner des variétés tolérantes à la sécheresse et améliorer les prédictions des modèles climatiques.

La méthode de référence pour estimer la photosynthèse est la mesure du taux d'assimilation nette en  $\text{CO}_2$  par des techniques d'échange gazeux. Bien que fournissant des informations clés sur les facteurs influençant la variabilité temporelle de la photosynthèse, les échanges gazeux seuls ne permettent pas une caractérisation complète de la limitation de la photosynthèse en cas de stress hydrique. Des mesures supplémentaires de fluorescence chlorophyllienne par des méthodes actives sont nécessaires afin de quantifier l'importance de chaque facteur limitant. En particulier, la réponse de la conductance mésophyllienne au stress hydrique reste une importante source d'incertitude dans les modèles de photosynthèse.

L'interprétation des mesures de fluorescence chlorophyllienne par des méthodes actives est néanmoins limitée à l'échelle de la feuille et ne permet pas de caractériser l'impact du stress hydrique sur des grandes étendues de végétation. L'émergence récente des méthodes passives pour mesurer la fluorescence chlorophyllienne a permis de suivre l'émission de fluorescence naturellement induite par le soleil (SIF) à différentes échelles spatiales et temporelles. Ces mesures constituent un indicateur prometteur des processus physiologiques de la végétation et peuvent être utilisées pour calibrer des modèles empiriques qui fournissent une estimation directe de l'assimilation en

CO<sub>2</sub> à des échelles locales et plus larges. Cependant, de tels modèles ne permettent généralement pas de reproduire l'évolution temporelle de la photosynthèse à court terme ou pendant des extrêmes climatiques tels que des sécheresses ou des vagues de chaleur. Des approches plus mécanistes sont nécessaires pour exploiter pleinement le message physiologique transmis par la SIF.

Dans cette thèse, nous avons d'abord cherché à identifier les origines des limitations de la photosynthèse pendant des épisodes de sécheresse édaphique. Les premiers chapitres sont consacrés à la caractérisation complète de la réponse des facteurs stomatiques et non-stomatiques à la diminution de la disponibilité en eau du sol. Cette méthode, d'abord appliquée aux écosystèmes forestiers (chapitre 3) et aux pommes de terre (chapitre 4) avec des données d'eddy covariance (EC), a mis en évidence le rôle clé des facteurs non-stomatiques dans la régulation de la photosynthèse à l'échelle de la canopée pour ces deux différents types d'écosystèmes. En effectuant des mesures de fluorescence chlorophyllienne actives et en utilisant une nouvelle méthode de partitionnement, nous avons également montré que la conductance mésophyllienne était un facteur essentiel qui régulait à la fois l'assimilation du carbone et la conductance stomatique de la pomme de terre (chapitre 5). Ces chapitres fournissent des fonctions de disponibilité en eau dans le sol pour l'influence des facteurs stomatiques et non-stomatiques sur la photosynthèse et la transpiration, qui peuvent être utilisés pour réduire les incertitudes dans les modèles climatiques. Dans le dernier chapitre, nous avons testé une nouvelle approche mécaniste visant à modéliser les échanges de carbone et d'eau d'une culture de blé d'hiver à partir de mesures de SIF à l'échelle de l'écosystème. Une corrélation très forte entre les estimations du modèle et les mesures d'EC a été trouvée, y compris pendant les périodes de sécheresse édaphique. Une analyse de sensibilité a révélé que la fraction de centres ouverts du photosystème II était un paramètre clé impactant les performances du modèle. Ce dernier chapitre ouvre la voie à une amélioration de notre compréhension des interactions entre les cycles de l'eau et du carbone en utilisant les informations physiologiques contenues dans la SIF.



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# Acronyms

**ABA** ABscissic Acid. 41, 43, 178

**ADP** Adenosine Di-Phosphate. 7

**AGB** AboveGround Biomass. 203

**AIC** Akaike Information Criterion. 116, 143, 145, 147, 154, 169

**APAR** Absorbed Photosynthetically Active Radiation. 159, 160, 162

**ASCAT** Advanced SCATterometer. 202

**ASWC** Available Soil Water Content. 43

**ATP** Adenosine Tri-Phosphate. 7–9, 136, 137, 160

**BPGA** 1,3-Biphosphoglycerate. 8, 9

**CABLE** Community Atmosphere Biosphere Land Exchange. 46, 178

**CBB** Calvin-Benson-Bassham. 5, 8, 9, 13, 16, 29, 30, 34, 44

**ChlF** Chlorophyll Fluorescence. i, 12, 13, 16–18, 24–26, 32, 33, 35–37, 51, 52, 177, 178, 180, 181, 188

**CLM5** Community Land Model v5. 35, 192

**cytb6f** Cytochrome b6f. 7

**DHAP** Dihydroxyacetone Phosphate. 8

**DT** DayTime. 22, 23, 181–183, 196

**EC** Eddy Covariance. vi, 18–23, 27, 29, 35, 42, 51, 57–59, 70–72, 77, 178, 181, 183, 184, 189, 191, 196, 205

- ECOSTRESS** ECOSystem Spaceborne Thermal Radiometer Experiment on Space Station. 201
- EnMAP** Environmental Mapping and Analysis Program. 202
- ENVISAT** ENVIronmental SATellite. 27
- FD** Ferredoxine. 7
- FLD** Fraunhofer Line Depth. 26, 27, 180
- FLEX** FLUorescence EXplorer. 27, 29, 138, 199
- FLORIS** FLuORescence Imaging Spectrometer. 27, 29
- FloX** Fluorescence boX. 29, 138, 140, 146
- FNR** Ferredoxin-NADP+ reductase. 7
- FTSW** Fraction of Transpirable Soil Water. 43
- FvCB** Farquhar-von Caemmerer-Berry. iv, 29–32, 35–37, 45–47, 52, 58, 78, 110–112, 126, 137, 138, 177, 179, 193, 202
- FWHM** Full Width of the absorption band at Half of the Maximum depth. 25
- G3P** Glyceraldehyde 3-phosphate. 8, 9
- GAI** Green Area Index. 81
- GAPDH** 3-phosphate dehydrogenase. 8
- GOME-2** Global Ozone Monitoring Experiment-2. 27
- GOSAT** Greenhouse gases Observing SATellite. 27
- GPP** Gross Primary Production. ii, v–vii, x, xii, 21–24, 27, 29, 30, 35–37, 51, 52, 57–65, 67, 69–73, 77–81, 85, 86, 91, 93, 95, 96, 99, 135–141, 143, 156, 159, 162–165, 171, 178–183, 187, 189, 195–199, 201, 202
- HyPlant** High-performance airborne imaging spectrometer for vegetation monitoring. 27
- ICOS** Integrated Carbon Observation System. 59, 77, 80, 109, 135, 139, 149, 181, 189, 205
- IPCC** Intergovernmental Panel on Climate Change. 35

- IRGA** Infrared Gas Analyser. 14, 21
- JULES** Joint UK Land Environment Simulator. 45, 178, 191
- LAI** Leaf Area Index. 58, 60, 70, 81, 197, 198
- LE** Latent heat. 20, 139, 140, 144, 145
- LHCs** Light Harvesting Centers. 5, 6, 11
- LIFT** Light-Induced Fluorescence Transients. 17, 18
- LRC** Light-Response Curve. 9, 12
- LSMs** Land Surface Models. 30, 32, 34, 43, 45, 46, 51, 77, 79, 111, 116, 123, 136, 138, 177, 178, 187, 189, 192, 195, 196
- LST** Land Surface Temperature. 200, 201
- LUE** Light-Use Efficiency. xi, 29, 35, 36, 137, 145, 150, 154, 159, 160, 163, 168, 201
- LWP** Leaf Water Potential. ii, 40–42, 192, 202, 204
- MERIS** Medium Resolution Imaging Spectrometer. 27
- MLR** Mechanistic Light Response. v–vii, xi–xiii, 36, 37, 52, 135, 137–139, 141, 143, 145, 147, 148, 156–165, 169, 171, 172, 180–182, 195–200, 202, 204, 205
- MODIS** Moderate Resolution Imaging Spectroradiometer. 202
- MTG-S** Meteosat Third Generation-Sounder. 199
- NADP+** oxidized form of Nicotinamide Adenine Dinucleotide Phosphate. 7, 44
- NADPH** reduced form of Nicotinamide Adenine Dinucleotide Phosphate. 7–9, 136, 137, 160
- NDVI** Normalized Difference Vegetation Index. 142, 149, 150, 198
- NEE** Net Ecosystem Exchange. vi, 21–23, 81, 136, 139, 164, 181–183, 196
- NIR** Near-Infrared. 5, 11, 12, 24, 140–142, 146
- NPQ** Non-Photochemical Quenching. 12, 13, 36, 137, 147, 148, 158, 165, 201

- NSOL** Non-Stomatal Origin Limitation of photosynthesis. iii, vi, x, 44–46, 51, 52, 57–59, 63, 64, 67, 69–73, 77, 79, 80, 83, 86, 87, 91, 95–99, 111, 115, 117, 125, 178–180, 183, 186, 187, 189–191, 194, 196
- NT** NightTime. 22, 181–183, 196
- OCO-2** Orbiting Carbon Observatory-2. 138, 199
- OCO-3** Orbiting Carbon Observatory-3. 199
- OEC** Oxygen-Evolving Complex. 7
- PAM** Pulse Amplitude Modulation. 16, 17, 37
- PAR** Photosynthetically Active Radiation. 5, 16, 22, 29, 32, 33, 35, 113, 114, 136, 137, 140, 141, 146, 149, 151, 154, 155, 158–160, 162–165, 171, 197, 198, 201
- PASWC** Plant Available Soil Water Content. 43
- PC** Platocyanine. 7
- PFT** Plant Functional Type. 42, 46, 51, 79, 110–112, 123, 137, 158, 163, 181, 184, 187, 189, 190, 194, 195, 197, 199, 205
- PG** 2-phosphoglycolate. 9
- PGA** 3-Phosphoglycerate. 8, 9
- PGK** Phosphoglycerate kinase. 8
- PheoD1** Pheophytin D1. 7
- Pi** Inorganic phosphate. 7
- PM** Penman-Monteith. 20, 144
- PPFD** Photosynthetic Photon Flux Density. 9, 14, 15, 22, 23, 113–115, 180
- PQ** Photochemical Quenching. 36, 136, 147, 160
- PQH2** Plastoquinol. 7
- PQi** Plastoquinone. 7
- PR** Photorespiration. 9, 13
- PRI** Photochemical Reflectance Index. 201, 202

- PSI** Photosystem I. 6, 7, 12, 16, 30, 32, 37, 141, 146, 148, 158
- PSII** Photosystem II. v, 6, 7, 12, 16, 17, 30, 32, 36, 37, 131, 137, 141–143, 146, 148, 155, 159, 160, 165, 169–172, 174, 180
- R<sub>ECO</sub>** Ecosystem Respiration. vi, 21–23, 181, 183, 196
- RBD** Root Biomass Density. xi, 45, 46, 166, 184, 185, 192
- RCs** Reaction Centers. 5, 6, 9, 12, 16
- REW** Relative Extractable Water. iv, vi, x–xii, 42, 43, 45, 46, 59–68, 70–73, 77, 79, 80, 82, 83, 86–88, 90–99, 109, 111, 112, 115, 117–119, 121–125, 129, 130, 145, 147–149, 152, 154–160, 163–166, 168, 173, 177–181, 183–187, 189–191, 196
- ROS** Reactive Organic Species. 12, 43, 44
- RS** Remote Sensing. vi, vii, 23, 27, 42, 177, 197–200, 202–204
- Ru5P** Ribulose 5-P. 9
- Rubisco** Ribulose 1,5-biphosphate carboxylase. 8
- RuBP** Ribulose-1,5-biphosphate. 8, 9, 31
- RWU** Root Water Uptake. 184, 192
- SCOPE** Soil Canopy Observation, Photochemistry and Energy fluxes. 137, 142
- SFM** Spectral Fitting Methods. 26, 140, 141, 180
- SIF** Sun-Induced chlorophyll Fluorescence. v, vi, xi, xii, 24–27, 29, 35–37, 52, 135–143, 145–149, 154, 158–160, 162–165, 171, 180, 181, 183, 188, 195–199, 201, 202, 204, 205
- SMAP** Soil Moisture Active Passive. 202, 203
- SMOS** Soil Moisture and Ocean Salinity. 202–204
- SOL** Stomatal Origin Limitation of photosynthesis. iii, vi, x, 44–46, 51, 52, 57–59, 62–64, 67, 69–71, 78–80, 83, 86, 87, 91, 96, 111, 115–118, 178, 180, 183, 187, 189, 190, 194
- SPAC** Soil-Plant-Atmosphere Continuum. ii, xii, 39–41, 187, 192, 195, 204
- SPEI** Standardized Precipitation Evaporation Index. 38

- SVD** Singular Vector Decomposition. 142
- SWA** Soil Water Availability. 43
- SWC** Soil Water Content. ii, 38, 39, 42, 43, 45, 60, 64, 67, 70, 79, 111–113, 130, 140, 145, 166, 177, 183–185, 189–191, 194, 195
- SWP** Soil Water Potential. vi, xii, 40–43, 183, 186, 187, 193–195
- SWRC** Soil Water Retention Curves. xii, 42, 43, 130, 183, 186, 194
- TAW** Total Available Water. 189, 190
- TEMPO** Tropospheric Emission Monitoring of POLLution. 199
- TIR** Thermal InfraRed. 200
- TPI** Triose phosphate isomerase. 8
- TPU** Triose Phosphate Utilization. 30, 31
- Tr** Transpiration. 20, 37, 52, 135, 138, 139, 145, 156, 163, 165, 179, 180, 192
- TROPOMI** TROPOspheric Monitoring Instrument. 27, 138, 198
- USO** Unified Stomatal Optimality. iv–vi, xi, 34, 35, 44–47, 52, 58, 59, 61, 62, 71–73, 78, 111, 112, 115, 117, 118, 123, 126, 128, 135, 138, 139, 144, 145, 148, 156–159, 161–165, 167, 169, 171, 177, 179–182, 195, 196, 204, 205
- UV** Ultraviolet. 5, 11
- VG** van Genuchten. 42, 113, 186, 193
- VOD** Vegetation Optical Depth. 203–205
- VPD** Vapor Pressure Deficit. x, 3, 9, 14, 22, 29, 33, 39, 40, 58, 78, 110–113, 117, 118, 121–123, 126, 128–130, 138, 144, 179, 182, 195, 198
- VRWC** Vegetation Relative Water Content. 203
- VWC** Vegetation Water Content. 202–204
- VWRC** Vegetation Water Retention Curves. 204
- WUE** Water-Use Efficiency. 34, 35, 58, 62, 70–72, 115, 124, 125, 179

**1**

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**Introduction**





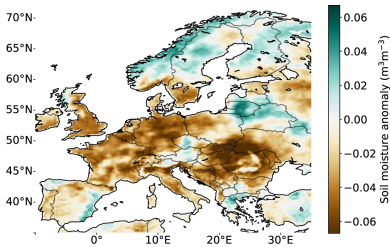
## 1. Climate change

Since the middle of the 18<sup>th</sup> century, the concentration of carbon dioxide (CO<sub>2</sub>) in the Earth's atmosphere has increased by about 52% up to a level that has never been reached since the mid-Pliocene, 3 million years ago (Keeling & Tans, 2023; Intergovernmental Panel On Climate Change, 2023). Anthropogenic fossil fuel emissions are responsible for the increase in CO<sub>2</sub> concentration, where coal burning still represented the largest contribution in annual carbon emissions in 2021 (Friedlingstein *et al.*, 2022). CO<sub>2</sub> emissions will cause irreversible changes in the climate, shifting conditions on Earth outside of the 'safe operating space' for humanity for at least the next millennia (Solomon *et al.*, 2009; Richardson *et al.*, 2023).

Once released in the atmosphere, CO<sub>2</sub> acts as a greenhouse: it absorbs and radiates heat in all directions, creating an insulating shell in the Earth's atmosphere. The fraction of solar energy radiated by the Earth's surface as heat towards the atmosphere is trapped and contributes to the 'greenhouse effect'. Climate model scenarios predict that the global surface temperature will increase within a range of 1.0°C to 5.7°C by the end of the 21<sup>st</sup> century compared to the 1850-1900 reference period (Masson-Delmotte *et al.*, 2021). A simulated change at 2°C global warming will cause an increase in 5°C to 7°C above the Arctic while Europe will warm up by 2°C to 3°C (Masson-Delmotte *et al.*, 2021).

The increase in the global surface temperature causes perturbations on the water cycle. One dominating thermodynamic constraint is the Clausius-Clapeyron (CC) equation, which predicts an increase in the water-holding capacity of the atmosphere (i.e., the saturated vapor pressure) of approximately 7% per degree Celsius of warming. In other words, the moisture content in the air must increase by 7% per degree Celsius to keep a constant relative humidity. Land surfaces warm up faster than oceans, mostly because of the predominance of sensible heat and longwave upward fluxes in the energy budget (Sutton *et al.*, 2007). As a result, the increase in saturated vapor pressure is higher over lands compared to oceans. However, the actual vapor pressure does not follow the CC rate and increases more slowly because of water atmospheric circulation and interactions between land surfaces and the atmosphere (Ficklin & Novick, 2017; Novick *et al.*, 2024). In particular, climate change is intensifying the strength of land-atmosphere feedbacks which limit evapotranspiration through, for instance, a lack of soil water, or by interacting with ecosystem carbon cycling (Ficklin & Novick, 2017). The difference between the saturation vapor pressure and the actual vapor pressure (i.e., the Vapor Pressure Deficit (VPD)) is a measurement of the evaporating force of the atmosphere, and acts as a major driver of the water cycle and plant-water relations (Grossiord *et al.*, 2020). Overall, a global increase in VPD has been observed in the past years and will continue to increase in a near future (Intergovernmental Panel On Climate Change, 2023).

The intensification of the hydrological cycle with climate change leads to very contrasting patterns of precipitation regimes. Although the increase in global mean precipitation is following the CC equation rate, major geographical heterogeneities in precipitation regime and distribution are observed because of the complex effects of climate change on the water cycle (Byrne & O’Gorman, 2015; Allan *et al.*, 2020). In particular, the occurrence of floods and extreme precipitation events is modulated by the availability of the water resource at the regional scale, also impacted by the seasonality of water availability (Tabari, 2020). Therefore, some regions are already facing an intensification of atmospheric water transport and precipitation frequency, while others are experiencing an increase in precipitation shortage episodes (Intergovernmental Panel On Climate Change, 2023). In Europe, major disruptions of seasonal precipitation regimes are expected with less precipitation on average but more frequent extreme events (Huo *et al.*, 2021). As a result, model simulations converge towards an increase in temperature and soil moisture anomalies in spring and summer across Europe, regardless of global warming scenarios (Spinoni *et al.*, 2018).



**Figure 1.1:** Soil moisture anomalies during the 2022 summer drought relative to the 1980–2022 reference period (not including 2022) in Europe. Soil moisture is taken from ERA5-Land (Muñoz-Sabater *et al.*, 2021). Adapted from Van Der Woude *et al.* (2023).

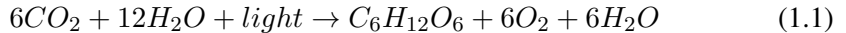
The increasing number of soil moisture and temperature anomalies over the past two decades in the Northern Hemisphere have shown that climate change is not only a model prediction, but is also becoming the new normal. These recent climate extremes include the 2003, 2018 and 2022 droughts (Ciais *et al.* (2005); Aalbers *et al.* (2023); Van Der Woude *et al.* (2023), Fig. 1.1), the 2010 heatwave in Russia, the 2020 warm anomaly over Siberia, or the combined droughts and heatwaves in Central Europe in summers 2015, 2018 and 2019 (Rousi *et al.*, 2022). Every year is a potential candidate for breaking previous-standing records (Fischer *et al.*, 2021). Summer droughts are also caused by a combination of lack of rainfall in winter and increasing evapotranspiration in early spring due to vegeta-

tion onset and seasonally increasing temperatures (Ionita *et al.*, 2020). Not only having considerable societal and economic impacts (Ionita & Nagavciuc, 2021; Yin *et al.*, 2023), water stress poses a major threat on terrestrial ecosystems. For instance, the 2003 drought caused a reduction in vegetation growth by  $\sim 30\%$  (Ciais *et al.*, 2005), while a reduction of net carbon uptake ( $\sim 60$  TgC) by terrestrial ecosystems during the 2022 drought in Europe was observed (Van Der Woude *et al.*, 2023). Over the last five decades in Europe, crop losses have tripled (Beillouin *et al.*, 2020; Brás *et al.*, 2021). The combination of soil water stress and heatwave is the worst abiotic stress factor for global development (Yin *et al.*, 2023).

The net amount of CO<sub>2</sub> remaining in the atmosphere depends on the balance between CO<sub>2</sub> sources and natural sinks (i.e., land and ocean sinks). Each year, the land biosphere removes from the atmosphere around one third of the total carbon emitted by human activities (Friedlingstein *et al.*, 2022). The carbon uptake rate of land biosphere strongly depends on the ability of terrestrial ecosystems to keep removing carbon in the context of global climate change. Understanding the processes at the basis of CO<sub>2</sub> uptake by vegetation is pivotal for predicting ecosystem capacities to mitigate climate change in a near future (Rogger *et al.*, 2024).

## 2. Photosynthesis

The strength of the CO<sub>2</sub> land sink is regulated by the capacity of plants to convert atmospheric CO<sub>2</sub> into carbohydrates by photosynthesis as follows:

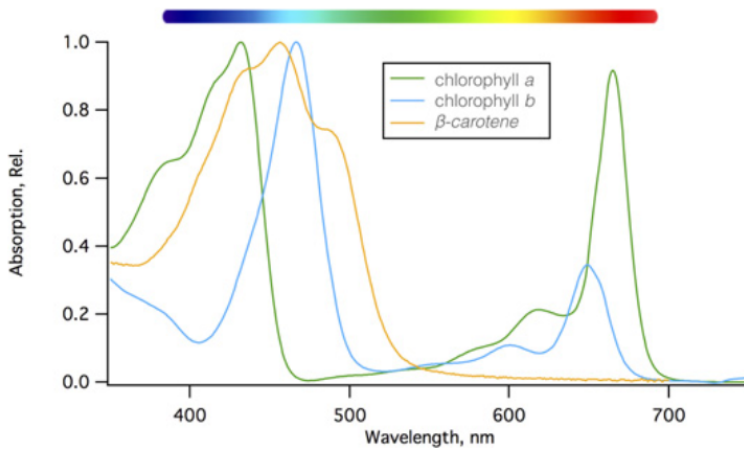


where the dioxygen (O<sub>2</sub>) molecule released to the atmosphere originates from the oxidation of water (H<sub>2</sub>O) and not from CO<sub>2</sub> (Blankenship, 2002). Not only photosynthesis uptake CO<sub>2</sub> from the atmosphere, but it also contributes to maintain O<sub>2</sub> concentration at a suitable level for sustaining life on Earth.

Photosynthesis is divided into four phases: (i) light is absorbed and transformed into chemical energy, (ii) electrons are transferred to reaction centers, (iii) solar irradiance energy is stabilized by the production of high energy molecules, and (iv) carbon is fixed, and final products are exported. In C3 plants (i.e., plants which produce a 3-carbon molecule from CO<sub>2</sub> fixation, see section 2.2), the first three phases are commonly referred to as the light reactions while the last phase is known as the Calvin-Benson-Bassham (CBB) cycle or dark reactions.

### 2.1. Absorption of light and electron transport

Light is absorbed by a complex structure combining proteins and pigments located in the thylakoids membrane within the chloroplasts, also known as Light Harvesting Centers (LHCs). Such structures regroup chlorophylls, carotenes and xanthophylls pigments. For example, one LHCs II monomer binds six chlorophylls *b*, eight chlorophylls *a* and four carotene molecules (Liguori *et al.*, 2015). LHCs are responsible for photon absorption and collection, acting like antennas within specific spectral regions. In particular, plant pigments can absorb solar irradiance between 400 nm and 700 nm (know as Photosynthetically Active Radiation (PAR)), with different peaks of absorbance. Chlorophyll pigments contain two major absorption bands in the Ultraviolet (UV) and Near-Infrared (NIR) regions, while carotenoids absorb most of the light in the near-UV (Fig. 1.2). LHCs concentrate sunlight energy towards Reaction Centers (RCs) where a cascade of redox reactions is initiated.



**Figure 1.2:** Absorption spectra of chlorophylls (chlorophyll a and chlorophyll b) and carotenoid ( $\beta$  carotene) pigments in plants (Johnson, 2016)

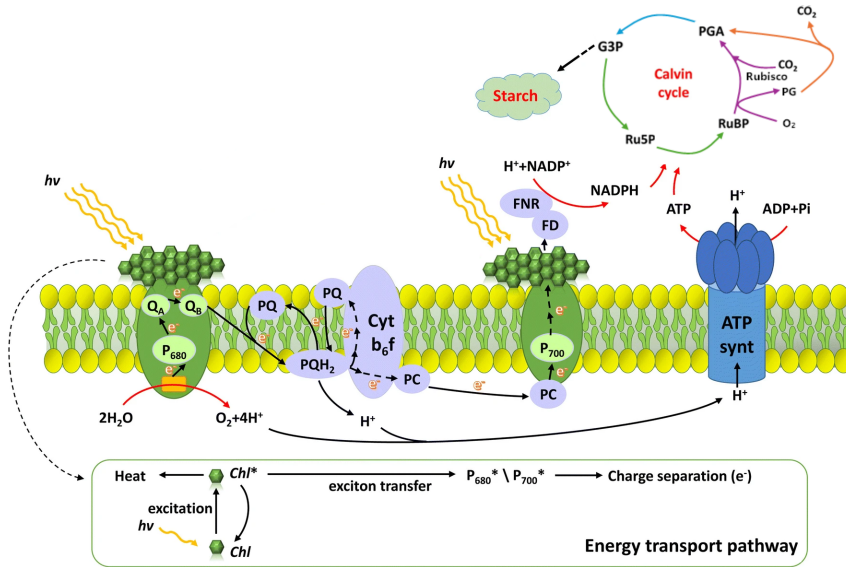
The light absorbed by LHCs excites pigments from ground to high electronic states. Before initiating photochemistry, this energy must be transferred to one of the RCs. This transfer of energy can be explained by the Förster theory (Förster, 1948), which is a very fast process (within femtoseconds) characterizing the nonradiative resonance transfer of energy between two chromophores (i.e., light-sensitive molecules). The Förster theory is based on several assumptions, including a weak coupling between two chromophores (distance superior to 10 nm) and an absence of quantum superpositions of excited electronic states. If these assumptions are met, the excitation can theoretically hop in a ‘random walk’ between two delocalized excitation states. As pigments are densely packed in LHCs, the hypothesis of the Förster theory are often challenged. Alternative mechanisms involving quantum coherence and excitons have been developed to describe the energy transfer dynamics in the antennas (Fassioli *et al.*, 2014; Ullah & Dral, 2022).

Photochemistry is initiated when the excitation energy is trapped by RCs, which are pigment-protein complexes, composed of chlorophylls, electron cofactors and hydrophobic peptides (Blankenship, 2002). RCs can lose or acquire an electron and therefore be oxidized or reduced. Once one of the RCs is promoted to a higher energy state by the electronic energy transferred by LHCs, it loses one electron to a primary acceptor, forming altogether an ion-pair state. This primary charge separation allows to transform electronic energy from solar irradiance into chemical energy. LHCs and RCs are regrouped within structural units known as Photosystem I (PSI) and Photosystem II (PSII), which differ by the main absorption bands of their primary acceptors, either at 680 nm (P680) or at 700 nm (P700).

PSII is the first complex in the light-dependant reactions of photosynthesis. Once excited, P680 is brought to an excited state ( $P680^*$ ) and transfers one electron to the primary acceptor, a nearby Pheophytin D1 (PheoD1) molecule. P680 becomes photo-oxidized ( $P680^{\bullet}$ ) and the electron hole must be replenished by the oxidation of water. This chemical reaction, which consists in breaking two  $H_2O$  molecule into one  $O_2$  while releasing four electrons and protons, is hard to initiate due to the high redox potential of  $H_2O$  (+820 mV).  $P680^{\bullet}$  is the strongest biological oxidizing agent and has a higher redox potential than  $H_2O$  (+1200 mV), which is sufficient to oxidize water (Johnson, 2016). However, more electrons are released from water oxidation than what is needed for reducing  $P680^{\bullet}$ . Therefore, these highly reactive particles are stored in the Oxygen-Evolving Complex (OEC), a cluster of manganese, calcium and oxygen which has two main functions: binding water molecules and storing electrons from water oxidation. As the OEC can store up to four electrons, two water molecules are oxidized after four turnovers of primary charge separation. The OEC is connected to the reaction center by the active tyrosine and transfers, one at the time, the electrons needed to reduce  $P680^{\bullet}$ . The OEC enzyme acts as a catalyst for water oxidation and is a key component of the light reactions of photosynthesis (Lubitz *et al.*, 2019). The electron carried by PheoD1 is then transferred to Plastoquinone (PQi), and ultimately reduced in Plastoquinol (PQH2) after the addition of two protons from the cytoplasm. As water oxidation releases four protons into the lumen while two protons are taken for forming PQH2, an electrochemical potential between the two sides of the thylakoid membrane is created, which contributes to the production of Adenosine Tri-Phosphate (ATP). PSII is also known as water-plastoquinone oxidoreductase for its ability to use light energy to oxidize water and to metabolically bound hydrogen.

Similarly, PSI utilizes light energy to bring the reaction center P700 to an excited state of energy ( $P700^*$ ). One electron is lost by  $P700^*$  (becoming  $P700^{\bullet}$ ) and is carried along a chain of electron carriers until ultimately being transferred to the final acceptor Ferredoxine (FD) (Grotjohann & Fromme, 2013). FD is one of the three substrates of Ferredoxin-NADP+ reductase (FNR) enzyme which catalyzes the production in the cytoplasm of the reduced form of Nicotinamide Adenine Dinucleotide Phosphate (NADPH) from the oxidized form of Nicotinamide Adenine Dinucleotide Phosphate (NADP+) by consuming one proton. The reduction of  $P700^{\bullet}$  is achieved by Platocyanine (PC) which transfers electrons from the Cytochrome b6f (cytb6f) to PSI. Not only providing a bridge for electrons between PSII and PSI, cytb6f and the Q-cycle oxidizes PQH2 and reduces PC (Malone *et al.*, 2021). The oxidation of PQH2 results in the release of protons in the lumen, thus increasing the  $H^+$  concentration and supplying the proton gradient across the thylakoid membrane. The ATP synthase enzyme uses the energy provided by the electrochemical proton gradient to make ATP from Adenosine Di-Phosphate (ADP) and Inorganic phosphate (Pi). PSI, PSII and cytb6f are connected in series and allow the circulation of electrons in the thylakoid membrane from water oxidation to the production of NADPH by FNR. The final prod-

ucts of the light reactions and the electron transport chain are ATP and NADPH which provide chemical energy for CO<sub>2</sub> fixation by the CBB cycle (Fig. 1.3).



**Figure 1.3:** Schematic representation of the light and dark reactions of photosynthesis. The definition of acronyms can be found in the text. *hν* is sunlight energy. In the CBB cycle, the purple, blue and green arrows represent respectively the fixation, reduction and regeneration steps. The orange arrow represents the photorespiratory cycle. The fatty acid synthesis is not represented. Adapted from Yang *et al.* (2020b) and Raines (2022).

## 2.2. CO<sub>2</sub> fixation

The CBB cycle (named after the American chemists Melvin Calvin, Andy Benson and James Bassham) is one of the most important biochemical cycles on Earth which fixes CO<sub>2</sub> into carbohydrates. It is a complex series of chemical reactions taking place in the chloroplast and divided into three phases: carboxylation, reduction, and regeneration (Blankenship, 2002). The carboxylation phase consists in the fixation of CO<sub>2</sub> into Ribulose-1,5-biphosphate (RuBP) which is catalyzed by Ribulose 1,5-biphosphate carboxylase (Rubisco). The fixation of CO<sub>2</sub> into RuBP gives two molecules of 3-Phosphoglycerate (PGA), which is an unstable carbon intermediate and is further phosphorylated by ATP into 1,3-Biphosphoglycerate (BPGA). NADPH is then needed for the reduction of BPGA into Glyceraldehyde 3-phosphate (G3P). The reduction of PGA in G3P consumes two molecules of ATP and NADPH and is catalyzed by Phosphoglycerate kinase (PGK) and 3-phosphate dehydrogenase (GAPDH). Triose phosphate isomerase (TPI) catalyzes the production of Dihydroxyacetone Phosphate (DHAP) from BPGA. BPGA and DHAP (also known as triose phosphates) are the starting point of the regeneration of RuBP. The triose phosphate pool is needed for

the regeneration of RuBP by the Ribulose 5-P (Ru5P) enzyme (Raines, 2022). These reactions consume one ATP for each CO<sub>2</sub> fixed. Since one G3P molecule has three carbon atoms, three CO<sub>2</sub> molecules must undergo the fixation and reduction steps to synthesize one GAP. In total, three ATP and two NADPH are needed for each CO<sub>2</sub> molecule fixed. Following the cycle stoichiometry, three CO<sub>2</sub> gives six BPGA and G3P. Among the six trioses phosphates produced by the fixation of three CO<sub>2</sub>, five are used for RuBP regeneration and one is exported as carbohydrates.

Rubisco can also react with O<sub>2</sub> by photorespiration. This process releases one molecule of 2-phosphoglycolate (PG), which is a potential inhibitor of several enzymes involved in the CBB cycle and must be metabolized. The Photorespiration (PR) cycle converts PG into PGA which is fed back into the CBB cycle at the beginning of the reduction phase. The PR cycle also releases CO<sub>2</sub> by the conversion of glycine into serine in the mitochondria and produces hydrogen peroxide in the peroxisome (Busch, 2020; Shi & Bloom, 2021). The CO<sub>2</sub> concentration at which the rate of photorespiration equals the rate of photosynthesis is known as the CO<sub>2</sub> compensation point. Photorespiration (also known as ‘respiration in the light’) affects the net balance of CO<sub>2</sub> assimilated by plants. An overview of the light reactions and the CBB cycle is given below (Fig. 1.3).

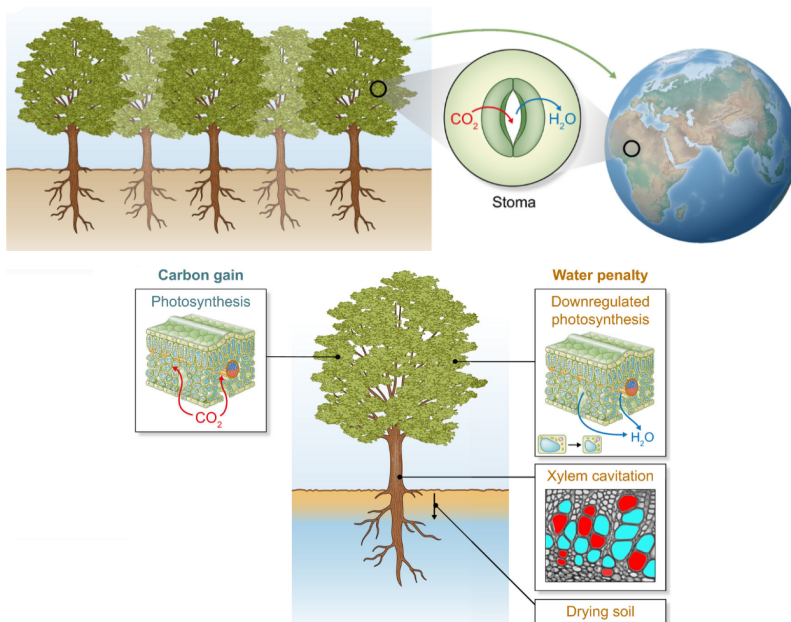
Photosynthesis is strongly limited by the supply of CO<sub>2</sub> to chloroplast and the light intensity which fuels ATP and NADPH to the CBB cycle by the light reactions. The tight link between dark and light reactions leads to a nonlinear relationship between net CO<sub>2</sub> assimilation and solar irradiance, as illustrated by Light-Response Curve (LRC)s describing how net photosynthesis reacts to changes in light intensity (measured by the Photosynthetic Photon Flux Density (PPFD)). These curves show a strong increase in photosynthesis at low PPFD, before a gradual decrease of the slope and a saturation from a specific PPFD threshold where all RCs are closed and all primary acceptors oxidized. Very similar trends are observed when characterizing the response of photosynthesis to the increase in CO<sub>2</sub> supply. These two curves provide key information on the maximum capacities of plants for absorbing light and assimilating CO<sub>2</sub> (Kubiske & Pregitzer, 1996; Zi-Piao, 2010). Combining these measurements with models of photosynthesis provide key information on the response of plants to a changing climate (see section 4, (Sharkey *et al.*, 2007)).

### **2.3. Diffusion of gases from the leaf surface to chloroplasts**

CO<sub>2</sub> diffusion to chloroplasts is a key factor limiting photosynthesis. Stomata are pores on the epidermis of leaf cells, which are bounded by guard cells and provide a direct pathway for CO<sub>2</sub> diffusion (Fig. 1.4). Many factors control stomatal opening, such as light intensity, VPD or molecular signaling (Buckley & Mott, 2013). For instance, stomata open as light intensity increases due to photoreceptors of blue light located on the guard cells (Shimazaki *et al.*, 2007). Under high VPD, water is being

evaporated from guard cells which causes membrane depolarization and stomatal closure (i.e., hydropassive mechanism). Stomata may also close in response to molecular signaling under drought by sensing of soil water status (i.e., hydroactive mechanism (Cowan, 1978)). Whether one contribution is predominant over the other is however often difficult to assess (Franks, 2013). Due to the difference in  $\text{CO}_2$  and  $\text{H}_2\text{O}$  concentration between the inside of the leaf and the outer surrounding atmosphere, the opening of stomata leads to a diffusion of  $\text{CO}_2$  to the mesophyll and an unavoidable loss of water by transpiration.

Stomatal response is coordinated to use the direct benefit of stomatal opening (i.e., fueling photosynthesis by facilitating  $\text{CO}_2$  diffusion) while avoiding the corresponding penalty (i.e., losing water and increasing the risk of cavitation, Fig. 1.4, (Cowan & Farquhar, 1977)). This hypothesis has led to numerous optimization models which aimed at predicting stomatal conductance over a wide range of environmental conditions (Wang *et al.*, 2020b).

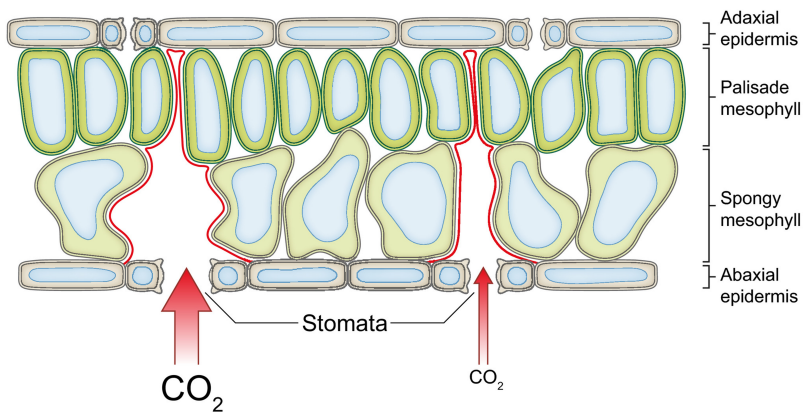


**Figure 1.4:** Stomata are tiny pores on leaf epidermis allowing  $\text{CO}_2$  diffusion to chloroplasts where carbon fixation occurs. This carbon gain through photosynthesis comes at the expense of a loss of water by transpiration, which increases the risk of cavitation under dry soil conditions. Adapted from Wang *et al.* (2020b).

The diffusion of  $\text{CO}_2$  in the plant cells is also mediated by the size of mesophyll airspaces which forms a continuum for  $\text{CO}_2$  diffusion to chloroplasts. A series of aqueous (i.e., diffusion through cell walls, lipid membranes, liquid cytoplasm and stroma)



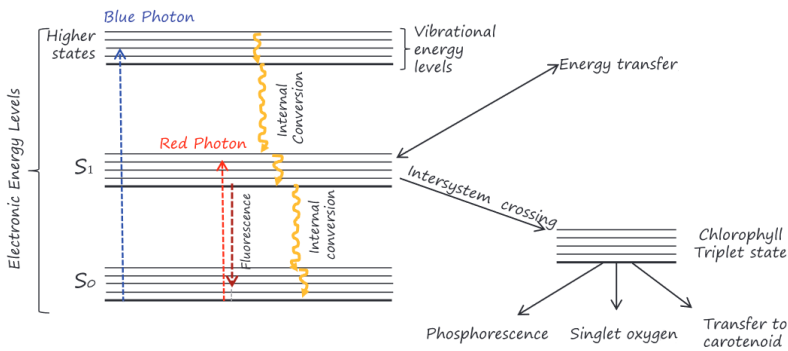
and gaseous (i.e., diffusion through leaf mesophyll airspaces) barriers affect photosynthesis by limiting  $\text{CO}_2$  diffusion. These diffusive limitations on photosynthesis are represented as two distinct conductances (i.e., the stomatal and mesophyll conductance) which determine the  $\text{CO}_2$  flux through the cell by the Fick's diffusive law. Mesophyll conductance varies accordingly to anatomic traits (e.g., cell wall thickness or chloroplast surface area exposed to intercellular airspaces per unit of leaf area) or biochemical components (e.g., aquaporins or changes in carbonic anhydrase activity) (Flexas *et al.*, 2018a; Baillie & Fleming, 2020). It is a key ecophysiological parameter which allows to identify the diffusive constraints on photosynthesis others than stomatal closure and also characterizes the ability of  $\text{CO}_2$  to diffuse through airspaces to the fixation sites in the chloroplasts (Fig. 1.5). Stomatal and mesophyll conductance are pivotal plant traits which limit the  $\text{CO}_2$  diffusion to chloroplasts. Identifying the response of these diffusive conductances to environmental drivers is crucial for predicting the effects of climate change on plant gas exchanges.



**Figure 1.5:** Leaf structure can increase  $\text{CO}_2$  diffusion to chloroplasts through airspaces by rearranging mesophyll architecture. The red area represents the  $\text{CO}_2$  diffusive continuum including substomatal cavities and mesophyll airspaces. Adapted from Baillie & Fleming (2020).

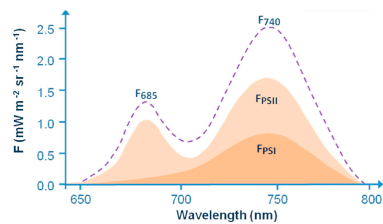
## 2.4. Fluorescence

As detailed in section 2.1, chlorophyll pigments have two absorption peaks in the NIR and UV regions, resulting in the absorption of blue photons ( $\sim 400\text{-}500\text{ nm}$ ) and red photons ( $\sim 600\text{-}700\text{ nm}$ ) by LHCS. Photon absorption by chlorophyll triggers the electron transfer from a ground state orbital ( $S_0$ ) to a higher excited state orbital ( $S_1$  or  $S_2$ ) (Gouterman, 1961). Hopping from  $S_0$  to  $S_1$  requires less energy than hopping from  $S_0$  to  $S_2$ . Therefore, the absorption of a blue photon brings an electron from  $S_0$  to  $S_2$  and a red photon from  $S_0$  to  $S_1$ . The electron transferred in  $S_2$  is highly unstable, and rapidly decays by molecular vibration (internal conservation) to  $S_1$  (Fig. 1.6).



**Figure 1.6:** Idealized Jablonski diagram of the fates of electrons on chlorophyll excitation states ( $S_0$ ,  $S_1$ ). Higher states correspond to  $S_2$ . Adapted from Porcar-Castell *et al.* (2014)

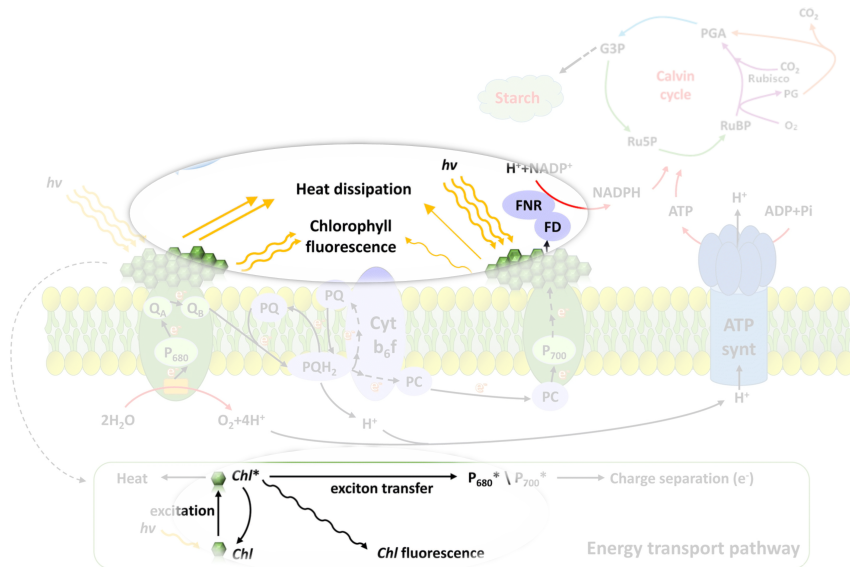
The electron on  $S_1$  has three possible fates: either returning to  $S_0$ , being lost as heat by molecular vibration, or being transferred to other chlorophyll pigments of the light harvesting complex through resonant energy or exciton transfer. This transfer (known as intersystem crossing) can lead to the formation of chlorophyll triplets, which can relax by transferring the energy to carotenoids, by phosphorescence or by reacting with oxygen producing reactive singlet oxygen (Clegg, 2004). These Reactive Organic Species (ROS) are highly reactive and can lead to damages in the cell by absorbing electrons from stable molecules or by disrupting cell functioning (Tripathy & Oelmüller, 2012). Returning to  $S_0$  from  $S_1$  can be done by molecular vibration or by emitting a photon at a longer wavelength than the absorbed photon. This process, known as Chlorophyll Fluorescence (ChlF), characterizes the emission of a photon resulting from an electron hopping from  $S_1$  to  $S_0$  (Fig. 1.6). The emission spectra of ChlF has two peaks, one at  $\sim 685$  nm in the red, and one at  $\sim 740$  nm in the NIR. While PSII contributes to these two peaks, PSI only emits at 740 nm (Fig. 1.7).



**Figure 1.7:** Chlorophyll fluorescence emission spectrum under steady state conditions. Two peaks can be observed at 685 nm ( $F_{685}$ ) and 740 nm ( $F_{740}$ ) resulting from the contribution of PSI ( $F_{PSI}$ ) and PSII ( $F_{PSII}$ ) emission spectra (Mohammed *et al.*, 2019)

Once on the saturation plateau of the LRC (i.e., RCs fully closed) or under stressed conditions (i.e., high temperature or low soil moisture), plants can no longer use all the incoming energy carried by solar irradiance for fixing  $\text{CO}_2$ . This excessive energy must be dissipated by alternative processes which are pivotal for avoiding stress on plants. Balancing the utilization and absorption of light energy to minimize the potential photo-oxidative damage is therefore crucial. Non-Photochemical Quenching

(NPQ) regroups the different mechanisms utilized by plants to dissipate the excess of energy mainly through heat, including internal conservation and intersystem crossing (Demmig-Adams *et al.*, 2014). Photosynthesis, fluorescence and heat dissipation are in competition for the incoming light energy. As a result, ChlF measurements can provide key information on carbon assimilation and on the physiological barriers on CO<sub>2</sub> diffusion (Porcar-Castell *et al.*, 2021). An overview of the light and dark reactions of photosynthesis including ChlF and NPQ is given below (Fig. 1.8).



**Figure 1.8:** Heat dissipation and chlorophyll fluorescence in the light reactions of photosynthesis. The definition of acronyms can be found in the text.  $h\nu$  is sunlight energy. In the CBB cycle, the purple, blue and green arrows represent respectively the fixation, reduction and regeneration steps. The orange arrow represents the PR cycle. The fatty acid synthesis is not represented. Adapted from Yang *et al.* (2020b) and Raines (2022).

### 3. Estimating photosynthesis and fluorescence at different spatial scales

Measuring photosynthesis is crucial for understanding how plants respond to changing environmental conditions. Moreover, ChlF provides additional information that can help in identifying the constraints of drought on the carbon cycle. The following section details the methods for measuring photosynthesis and fluorescence at the leaf level and at the ecosystem level.

## 3.1. Leaf-level measurements

### 3.1.1. Gas exchanges

Gas exchange systems measure the gas concentration (either CO<sub>2</sub> or H<sub>2</sub>O) with an Infrared Gas Analyser (IRGA) and the flow rate at which the air interacts with leaf tissues. The technique used in IRGAs exploits the absorption properties of CO<sub>2</sub> and H<sub>2</sub>O in the infrared wavelengths. In particular, IRGAs use an infrared source and detector to measure the difference of infrared radiation which depends on the concentration of the tracer. Most of commercially available gas exchange systems are designed to measure the response of carbon assimilation and transpiration to a wide variety of environmental conditions which can be set by the user (e.g., H<sub>2</sub>O, CO<sub>2</sub> concentrations or light intensity). Gas exchange measurements are conducted by enclosing a part of a leaf (or the entirety) within a cuvette. Although many cuvette types exist to cover the wide variety of leaves anatomy, the characteristics of the cuvette must maintain uniform climate conditions across the leaf. For instance, cuvettes are constructed from a metal with high thermal conductivity and painted to reflect most of the solar radiation (Bernacchi *et al.*, 2012). The following paragraph details the functioning of the LI-COR LI-6400 device (LI-COR Inc., Lincoln, NE, USA), which has been and continues to be, used in numerous studies conducted in photosynthesis research (Savvides & Fotopoulos, 2018). This device is available at the Biosystem Dynamics and Exchanges (BIODYNE) research department of Gembloux Agro-Bio Tech. It was used in this thesis in chapter 5 and 6.

In an open system such as the LI-6400, the air flows into the console, and passes through a CO<sub>2</sub> scrubber and a desiccant, allowing to set specific concentrations in H<sub>2</sub>O and CO<sub>2</sub> in the air surrounding the leaf sample by adjusting bypass valves. When the system is equipped with a CO<sub>2</sub> mixer, the CO<sub>2</sub> concentration in the air flow can be increased up to more than 2000 ppm to build CO<sub>2</sub> response curves. The flow rate is measured and adjusted by a flow meter and a pump before circulating through the cuvette in the sensor head. Two IRGAs are used to measure independently the gas concentration within the cuvette (i.e., the 'sample') and in a parallel branch of the circuit that does not interact with the cuvette (i.e., the 'reference'). In addition, the incident PPFD is measured by a quantum sensor on the sensor head. The PPFD inside the chamber can be set to ambient level or to a specific value by a parameterization of the LED light source to build light-response curves. Finally, leaf temperature is measured by a thermocouple in contact with the leaf sample while air temperature is measured by a thermistor located beneath the IRGA circulation fan. Many air humidity variables are also computed from the H<sub>2</sub>O concentration and temperature measurement such as relative humidity, VPD or dewpoint temperature.

Open IRGA systems use measurements of sample and reference H<sub>2</sub>O and CO<sub>2</sub> concentrations, as well as the flow rate to calculate the transpiration and the net carbon assimilation rate ( $A_n$ ). This latter corresponds to photosynthesis ( $A$ ) minus leaf respi-

ration ( $R_d$ ):

$$A_n = A - R_d \quad (1.2)$$

$R_d$  can be estimated by measuring the net CO<sub>2</sub> flux on a dark-adapted sample ( $A = 0$  in Eq. 1.2), by being set as a fraction of photosynthetic capacities (see 4), or by being neglected when  $A \gg R_d$  under high irradiance. In addition, the degree of stomatal opening can be estimated by the calculation of the stomatal conductance from transpiration measurements and molar concentration of water vapor within the leaf by Fick's law of diffusion (Farquhar & Sharkey, 1982). Such system provides a complete characterization of leaf gas exchanges in a wide variety of environmental conditions through, for instance, light and CO<sub>2</sub> response curves. In these cases, the net assimilation rate is measured in response of changes in PPFD and reference CO<sub>2</sub> concentration. Instantaneous measurements are also very useful to measure plant responses to changes in ambient climate conditions over time, or to compare values between species.

Gas exchange measurements rely on steady state conditions at the leaf surface. Equilibrium must be reached before each measurement, which means that the variation through time of sample H<sub>2</sub>O and CO<sub>2</sub> concentration must be minimum. The LI-6400 software provides stability variables to help in logging the measurements at the right time, such as the variation coefficient or the rate of change over time. Waiting for stability can take up to 1 hour, which makes leaf level measurements time-consuming and limits the maximum number of measurements per day. Problems can also be encountered, such as cuvette leakages, lateral fluxes through leaves or calibration issues (Bernacchi *et al.*, 2012), which may alter the interpretation of measurements. Despite these issues, gas exchange measurements have become the standard benchmark for measuring photosynthesis and transpiration at the leaf level when compared to other techniques (Siebers *et al.*, 2021).

Upscaling leaf-level measurements over space and time is challenging when using gas exchange systems. Not only time and labor-consuming, it also requires a representative sampling throughout the canopy and the study site for integrating the complexity of the ecosystem structure. The fraction of sunlit and shaded leaves, the leaf angle orientation to the sun, or the micro-climate gradients within the canopy are issues that are difficult to assess when upscaling local measurements over space. It is practically impossible to make enough measurements without sampling artefacts (Baldocchi, 2014). Another type of gas exchange measurement device is canopy chambers, which are whole plant enclosure systems allowing to measure plant-scale processes. Such systems are typically used to study the impact of increasing CO<sub>2</sub> concentration on vegetation processes, the exchanges processes of reactive trace gases, or continuous measurements of photosynthesis and transpiration (Perez-Priego, 2021). Although very suitable for measuring net fluxes of uniform and short vegetation, this technique suffers from several issues (Baldocchi, 2014; Perez-Priego, 2021). For instance, cham-

bers affect plant's microclimate by damping turbulent mixing or increasing humidity and heat. Moreover, many chambers are needed to sample the canopy (i.e., between 10 and 20). Such density is rarely seen in the literature (Baldocchi, 2014).

### 3.1.2. Measurements of chlorophyll fluorescence with active techniques

ChlF has been used for many years to monitor the photosynthetic activity of plants. ChlF intensity measurements can be used to infer many physiological parameters such as the proportion of open PSII RCs, the PSII maximum efficiency, or the PSII operating efficiency (Baker, 2008). In particular, this last characterizes the efficiency at which light absorbed by PSII is used for primary acceptor reduction, and is directly related to the electron transport rate and to carbon assimilation. These parameters are calculated from ChlF intensities in different light conditions. ChlF measurements consists in measuring the peak of ChlF induced by a pulse from a light source at a specific wavelength (usually at  $\sim 650$  nm to drive electron hopping from  $S_0$  to  $S_1$ , Fig. 1.6). Such measurements are 'active' because the ChlF peak results from the application of a saturating pulse on the leaf. One of the challenges for measuring ChlF intensities is the identification of the ChlF peak among the natural sunlight reflected by the plant and the incident sunlight in the red and infrared bands. While not being necessary for the determination of photosynthetic parameters relying on dark-adapted sample measurements such as the PSII maximum efficiency, this is essential for deriving key parameters of the electron transport chain and the CBB cycle.

The key feature of active ChlF measurement devices is the selective capacity of the detection system to separate the ChlF signal from the ambient and reflected light within the chamber. Pulse Amplitude Modulation (PAM) fluorometers use lock-in amplifiers to measure the ChlF changes induced by the measuring light (i.e., the light signal which drives the ChlF peak) while ignoring the ChlF changes induced by the actinic light (i.e., the light signal which drives photo-chemistry - a full review of PAM fluorometry is given in Schreiber (2004) and in Cornic & Massacci (2004)). A typical ChlF quenching analysis from PAM fluorometry is described in Fig. 1.9. The analysis begins with a dark-adapted leaf, which has primary acceptors fully oxidized and RCs open. The measuring light is turned on, and the leaf is exposed to a very low PAR, resulting in a minimum level of ChlF ( $F_0$ ) while primary acceptors remain fully oxidized. A saturating pulse is then applied (several thousand  $\mu\text{molm}^{-2}\text{s}^{-1}$  in less than 1 s) to drive the closure of RCs and allow the measurement of the maximum ChlF intensity in the dark ( $F_m$ ). The actinic light (the light driving the partial reduction of primary acceptors in PSII) is then set to a specific value, and the ChlF emission of a light-adapted sample ( $F_s$  or  $F$ ) is measured. A saturating pulse is once again applied to fully reduce PSII primary acceptors and gives the maximum ChlF intensity in the light ( $F'_m$ ). Finally, the actinic light is turned off, and a weak far-red pulse of light is applied simultaneously to excite PSI RCs and to remove electrons on the primary acceptors, giving the minimum ChlF intensity in the dark ( $F'_0$ ) of a light-adapted sample.

Assuming a 'lake' model describing the structure of RC and antennas, the maximum quantum yield (i.e., the ratio of photons used for a process to the photons absorbed by the leaf) of PSII in the dark ( $\phi_{PSII,max}$ ) or the quantum yield of PSII in the light ( $\phi_{PSII}$ ) and the fraction of open PSII centers ( $q_L$ ) are directly calculated from ChlF intensities as (Baker, 2008):

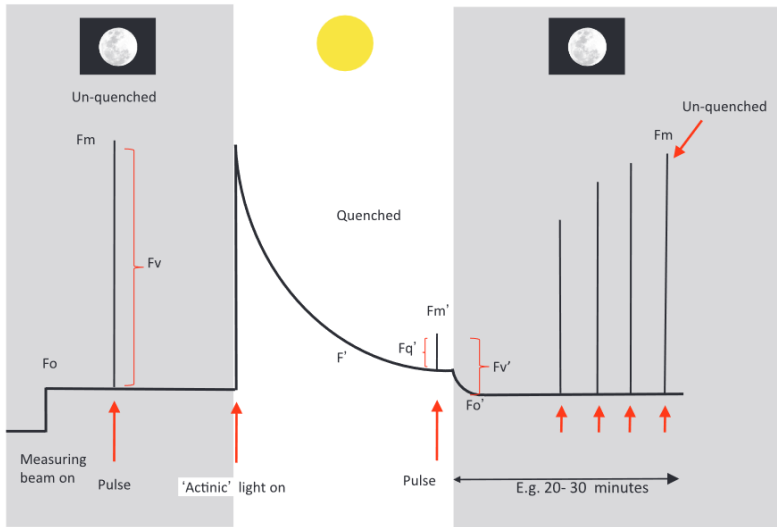
$$\phi_{PSII} = \frac{F'_m - F'_0}{F'_m} \quad (1.3)$$

$$\phi_{PSII,max} = \frac{F_m - F_0}{F_m} \quad (1.4)$$

$$q_L = \frac{F'_m - F_s F'_0}{F'_m - F'_0 F_s} \quad (1.5)$$

ChlF quenching analysis is a key tool for estimating the photosynthetic capacities of leaves. Actively-induced ChlF measurements can be coupled with gas exchange measurements, giving a complete set of parameters for characterizing photosynthetic capacities of a plant and limiting factors of photosynthesis. For example, the mesophyll conductance to CO<sub>2</sub> diffusion can be determined from  $\phi_{PSII}$  and gas exchange measurements of the substomatal CO<sub>2</sub> concentration (detailed in section 4). These measurements can be performed using a LI-6400 device with a LI-6400-40 leaf chamber fluorometer which has similar technical properties as PAM fluorometers. The fluorescence chamber is available at the BIODYNE research department.

ChlF measurements with active techniques can only be performed at the leaf scale, so the same spatial extrapolation issues as gas exchange measurements are encountered (Amoros-Lopez *et al.*, 2008). Moreover, the physiological interpretation of ChlF measurements resulting from the application of saturating pulses with an intensity much greater than what plants experience in natural conditions may be challenged. The recent development of the Light-Induced Fluorescence Transients (LIFT) device aimed at overcoming these spatial upscaling issues (Raesch *et al.*, 2014; Kolber *et al.*, 2005; Pieruschka *et al.*, 2014). The first versions of LIFT devices used a laser beam to project the excited signal from a distance up to 50 m away onto a target area of about 100 cm<sup>2</sup> (Raesch *et al.*, 2014; Moya *et al.*, 2019). Because of the raising concerns for eye safety involving the use of lasers in terrestrial environments, Osmond *et al.* (2017) developed an eye-safe LIFT prototype using light emitting diodes rather than laser, which is suited for measuring photosynthesis parameters at a distance inferior to 2 m (Osmond *et al.*, 2017; Keller *et al.*, 2019). However, several aspects of the LIFT technique complicate the interpretation of ChlF measurements and the comparison with PAM fluorometry such as the influence of complex dense canopies or changes in leaf angle due to wind which impacts scattering and absorption properties of the target



**Figure 1.9:** Example of ChlF quenching analysis using PAM fluorometry (Murchie & Lawson, 2013).

areas (Pieruschka *et al.*, 2014). Although promising for measuring photosynthetic parameters at larger scales, data analysis and processing with the LIFT technique needs further development (Pérez-Bueno *et al.*, 2019).

## 3.2. Ecosystem-level measurements

### 3.2.1. Eddy covariance

Undoubtedly, measuring gas exchanges at the interface between ecosystems and the atmosphere would significantly increase understanding of the functioning of ecosystems and the interactions with environmental drivers. The theoretical framework for measuring gas exchanges at the ecosystem surface was proposed by Montgomery (1948), Obukhov (1951) and Swinbank (1951). Sensors providing continuous measurements of flux densities at a high frequency were developed few decades later in the 90's. Since then, the Eddy Covariance (EC) technique has become widely used in climatology and ecophysiology, as it integrates the complexity of plants and canopies into a single high frequency measurement of gas exchanges.

#### 3.2.1.1 Physical principle and main equations

The net flux exchanged at the ecosystem-atmosphere interface is calculated from the covariance between turbulent fluctuations of vertical wind and the dry mole fraction of the gas of interest (Foken *et al.*, 2012):



$$F_{EC} = \overline{\rho_d w s} \quad (1.6)$$

with  $F_{EC}$  the EC flux,  $\rho_d$  the dry molar air density,  $w$  the vertical wind speed and  $s$  the dry mole fraction of the gas of interest. Eq. 1.6 can be decomposed into the sum of the time-mean part and the fluctuation part (i.e., the Reynolds decomposition):

$$F_{EC} = \overline{(\bar{\rho}_d + \rho'_d)(\bar{w} + w')(\bar{s} + s')} \quad (1.7)$$

with  $'$  denoting the fluctuation term and  $\bar{w}$ ,  $\bar{\rho}_d$ ,  $\bar{s}$  the time-mean terms. Eq. 1.7 is further simplified into:

$$F_{EC} = \bar{\rho}_d \bar{w} \bar{s} + \bar{\rho}_d \overline{w' s'} + \bar{s} \overline{\rho'_d w'} + \overline{\rho'_d w' s'} \quad (1.8)$$

As the air density fluctuations are assumed to be negligible (which is the case over flat and vast spaces) as well as the mean vertical flow (for homogeneous horizontal terrains), Eq. 1.8 can be written as:

$$F_{EC} \simeq \bar{\rho}_d \overline{w' s'} \quad (1.9)$$

Eq. 1.9 is valid only if most of the vertical transfer is assumed to be driven by eddies. Integrating the one-point equation conservation of dry mole fractions in the control volume (i.e., a parcel of air above the surface and below the height of measurements  $h_m$ ) gives the total ecosystem flux ( $F_{ECO}$ ) (Foken *et al.*, 2012):

$$F_{ECO} = F_{EC} + \int_0^{h_m} \bar{\rho}_d \frac{\delta \bar{s}}{\delta t} dz = F_{EC} + F_S \quad (1.10)$$

Eq. 1.10 shows that the total flux exchanged between the ecosystem surface and the atmosphere is the sum of the EC flux ( $F_{EC}$ ) and the storage flux ( $F_S$ ) which is calculated from the integration of temporal changes over a time interval of the dry mole fraction with height ( $\int_0^{h_m} \bar{\rho}_d \frac{\delta \bar{s}}{\delta t} dz$ ). Storage calculations are important when turbulence is not the main governing transport, in case of low wind, stable stratification or decoupling between the surface and the atmosphere. Nighttime is often a critical period as most of these conditions are observed. Above homogeneous terrains, for small canopies and when turbulence is present, the storage flux may be neglected. Eq. 1.10 can be therefore simplified as (Foken *et al.*, 2012):

$$F_{ECO} = F_{EC} \simeq \bar{\rho}_d \overline{w' s'} \quad (1.11)$$

which is the equation for EC fluxes calculation (Burba, 2021). Eq. 1.11 is valid for any gas and corresponds to the product of the mean air density and the mean covari-

ance between instantaneous deviations in vertical wind speed and dry mole fraction. Sensible heat ( $H$ ) and Latent heat (LE) fluxes are calculated in a similar fashion:

$$H = \bar{\rho} c_p \overline{wT'_a} \quad (1.12)$$

$$LE = \lambda \frac{M_w/M_a}{P} \overline{\rho_d w e'_a} \quad (1.13)$$

where  $\rho$  is wet molar air density,  $c_p$  is heat capacity,  $T_a$  is air temperature,  $M_w/M_a$  is the ratio of molar masses of water and air,  $P$  is air pressure and  $e_a$  is air vapor pressure. The latent heat flux characterizes a flux of water changing from liquid to gaseous phase.

LE flux measurements can be used to derive canopy-scale stomatal conductance. In particular, LE is related to the water vapor pressure gradient between the ecosystem surface and the atmosphere, and the conductances opposing to the water vapor transfer to the atmosphere:

$$LE = \frac{\rho_a c_p (e^*(T_s) - e_a)}{\gamma} \frac{G_{aw} G_{sw}}{G_{aw} + G_{sw}} \quad (1.14)$$

with  $e^*(T_s)$  the saturated air vapor pressure calculated at surface temperature ( $T_s$ ),  $\rho_a$  the air density,  $c_p$  the heat capacity of dry air,  $\gamma$  the psychrometric constant,  $G_{aw}$  the aerodynamic resistance and  $G_{sw}$  the canopy-scale stomatal conductance to water vapor. When  $T_s$  is not measured, Eq. 1.14 can be written as:

$$LE = \frac{\Delta (R_n - G) + \rho_a c_p (e^*(T_a) - e_a) G_{aw}}{\Delta + \gamma(1 + \frac{G_{sw}}{G_{aw}})} \quad (1.15)$$

with  $G_{sw}$  the canopy-scale stomatal conductance to water vapor transfer,  $R_n$  the net radiation,  $G$  the ground heat flux and  $T_a$  the air temperature. Eq. 1.15 is also known as the Penman-Monteith (PM) equation (Monteith, 1965) and is widely used for (i) estimating the latent heat flux at the ecosystem scale from meteorological measurements (Maes *et al.*, 2019), and (ii) estimating canopy conductance to water vapor from LE measurements, under the assumption that evaporation is negligible compared to Transpiration (Tr) (i.e.,  $LE \sim Tr$  (Knauer *et al.*, 2018a)). Partitioning LE into evaporation and Tr is a currently active research topic (Nelson *et al.*, 2020). One of the most straightforward approaches for estimating Tr from LE measurements is to filter EC data for selecting active vegetation periods when the canopy surface is dry. Canopy dryness can be evaluated by calculating an index from a water balance model (Nelson *et al.*, 2018) or by excluding a certain number of days after a precipitation event (at least 2 days (Knauer *et al.*, 2018b)).

### 3.2.1.2 Eddy covariance setup

EC systems are composed of two key components which measure the wind speed in three dimensions, and the concentration of the gas of interest. The IRGA and the high-frequency three-dimensional (3D) sonic anemometer are both placed on a tower or on a mast. The wind speed is determined by knowing the forward and return transit time of ultrasonic pulses sent between two 3D sonic anemometers transducers. The air temperature can also be deduced from the speed of sound, calculated by the average of the inverse of the return and forward transit times. 3D sonic anemometers used in EC are made of three pairs of transducers, which are orientated to minimize wind perturbation and to reduce the impact of rain. A conversion to orthogonal 3D is needed to determine the wind speed in three dimensions (Aubinet, 2023).



**Figure 1.10:** Example of an eddy covariance system configuration with a three-directional sonic anemometer (Gill HS-50) and a closed-path infra-red gas analyzer (LICOR 7200). Photo taken in BE-Lon.

Gas concentrations are measured by the far-red absorption spectrometry. Most recent IRGAs used in EC have an enclosed path (Rebmann *et al.*, 2018), which allows to measure gas concentration during precipitation and icing, and do not require a filter for eliminating sunlight interference (Burba *et al.*, 2010). Enclosed analyzers use short intake tubes to pump the air to the measurement chamber, which minimizes frequency losses and sorption-desorption of water molecules (Burba *et al.*, 2010). Both IRGA and 3D sonic anemometer must be separated from the smallest distance possible to minimize frequency losses (Aubinet, 2023). A typical setup is shown in Fig. 1.10. More details about data

post-processing steps are given in Aubinet *et al.* (2012) and Aubinet (2023).

### 3.2.1.3 Partitioning of the net CO<sub>2</sub> flux

The EC technique measures the Net Ecosystem Exchange (NEE) at the ecosystem-atmosphere interface, which is composed of Gross Primary Production (GPP) (i.e., the total amount of carbon fixed by all plants) and Ecosystem Respiration ( $R_{ECO}$ ) (i.e., autotrophic and heterotrophic respiration):

$$F_{ECO} = NEE = R_{ECO} - GPP \quad (1.16)$$

GPP (or  $R_{ECO}$ ) corresponds to a downward (or inward) carbon flux and is counted negative (or positive) by convention. When  $|GPP| > |R_{ECO}|$ , the ecosystem fixes more carbon through photosynthesis than it emits through respiration. A carbon sink is

therefore characterized by  $NEE < 0$  and a carbon source by  $NEE > 0$ . Eq. 1.16 is the equivalent form of Eq. 1.2 written for a net  $CO_2$  flux measured at the ecosystem scale.

An important step for interpreting EC measurements in physiological and modeling studies is the separation of NEE into GPP and  $R_{ECO}$ . This is carried out by partitioning methods, which are commonly classified into two categories. Either  $R_{ECO}$  is determined from nighttime  $CO_2$  fluxes measurements (i.e.,  $GPP = 0$  in Eq. 1.16) and extrapolated with temperature (i.e., the NightTime (NT) partitioning method (Reichstein *et al.*, 2005)), or  $R_{ECO}$  is determined from daytime  $CO_2$  fluxes measurements using the intercept of light response curves of NEE (Falge *et al.*, 2001). One important bias of the NT method is that it relies on data which can be significantly affected by low turbulence or advection, resulting in a limited dataset for extrapolating  $R_{ECO}$  with temperature. Lasslop *et al.* (2010) combined the two approaches by proposing a modified partitioning method using both the temperature-dependance of  $R_{ECO}$  and the light response curve of NEE. In this method, nighttime  $CO_2$  measurements are first extrapolated with air temperature using an Arrhenius-type model:

$$R_{ECO} = R_b \exp \left[ E_0 \left( \frac{1}{T_{ref} - T_0} - \frac{1}{T_a - T_0} \right) \right] \quad (1.17)$$

where  $R_b$  is the base respiration at the reference temperature  $T_{ref}$  (set to 15 °C),  $E_0$  is the temperature sensitivity,  $T_0$  is set to -46.02 °C and  $T_a$  is air temperature (which can be replaced by soil temperature - the choice of driving temperature for  $R_{ECO}$  is discussed in Lasslop *et al.* (2012)). The rectangular hyperbolic light response curve (Falge *et al.*, 2001) is then modified to account for the temperature effect on  $R_{ECO}$  to model NEE using daytime data:

$$NEE = \frac{\alpha_{NEE} \beta_{NEE} R_g}{\alpha R_g + \beta_{NEE}} - R_b \exp \left[ E_0 \left( \frac{1}{T_{ref} - T_0} - \frac{1}{T_a - T_0} \right) \right] \quad (1.18)$$

where  $R_g$  is global radiation (can be replaced by PPFD, Gilmanov *et al.* (2013)),  $\alpha_{NEE}$  is the canopy light utilization efficiency (representing the initial slope of the light response curve) and  $\beta_{NEE}$  the maximum NEE at light saturation. Modeling solely NEE from PAR and temperature without accounting for stomatal closure effects on GPP can also generate uncertainties in the results. Lasslop *et al.* (2010) modified  $\beta_{NEE}$  to account for the effect of stomatal closure on GPP through VPD as:

$$\beta_{NEE} = \begin{cases} \beta_{NEE,0} \exp^{-k(VPD - VPD_0)} & VPD > VPD_0 \\ \beta_{NEE,0} & VPD < VPD_0 \end{cases} \quad (1.19)$$

with  $VPD_0$  set to 10 hPa and  $\beta_{NEE,0}$  a fitted parameter. Lasslop *et al.* (2010) partitioning method is referred to as the DayTime (DT) partitioning method. The

multi-step process of model parameterization of the DT method is described in Lasslop *et al.* (2010) and in Wutzler *et al.* (2018).

### 3.2.1.4 Footprints and upscaling

EC fluxes originate from a target area (i.e., a footprint), which should be representative of the ecosystem of interest. An analysis of the turbulent field around the measurements point is required for identifying the contribution of all sources and sinks to EC fluxes. Footprint models use wind direction, height of the measurement point, surface roughness and stability conditions to provide a mapping of contributing areas around the EC tower (Schmid, 2002). Assessing the contribution of the target area to the global EC footprint is crucial prior to fluxes analysis (Burba, 2021).

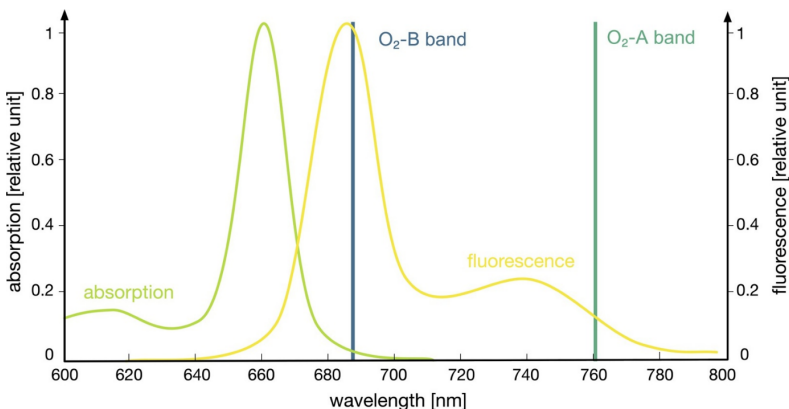
The unique features of the EC technique (i.e., a non-destructive, automatic and high frequency method for measuring gas exchanges and energy fluxes greenhouse gases and energy fluxes at the ecosystem-atmosphere interface) have made it popular for understanding the impacts of climate on ecosystems (Baldocchi, 2014). EC measurements have been conducted on 2155 locations, and more than 900 sites are currently part of the FLUXNET network, covering all ecosystem types (Baldocchi *et al.*, 2001; Chu *et al.*, 2017; Burba, 2019). EC measurements have been regrouped within global datasets such as the 2007 LaThuile dataset with 965 site-years of data and the FLUXNET2015 dataset with 1532 site-years of data. These datasets provide high quality data using a common data processing pipeline (i.e., the ONEFLUX pipeline) for all stations and site-years (Pastorello *et al.*, 2020). Notably, this standardized data treatment provides gap-filling of missing data and partitioning of NEE between GPP and  $R_{ECO}$ . The joint measurements of EC fluxes along with meteorological variables (e.g., air vapor pressure, air temperature or PPFD), vegetation variables (e.g., leaf area index, crop yield or canopy cover), and edaphic variables (e.g., soil moisture or soil temperature) have made possible to study the response of ecosystems to climate drivers (see Baldocchi (2020) for several references).

Notwithstanding the undeniable contribution of EC to understanding of ecophysiological processes in ecosystems, EC stations only provide a partial spatial coverage of fluxes measurements. EC fluxes can only be interpreted within the footprint of the station, and providing fluxes "everywhere and all the time" (Chu *et al.*, 2021) is actually not really feasible, especially near the tropics where climate conditions and ecosystem structure complexifies the installation of EC stations. Despite the growing number of EC under the tropics (e.g., Sibret *et al.* (2022)), these regions still remain poorly monitored, even though they are the most important carbon sinks (Chu *et al.*, 2017). Alternatively, Remote Sensing (RS) of vegetation greenness provide a proxy of vegetation health which can be used to build machine learning algorithm to understand the response of fluxes to climate drivers at large scales (Jung *et al.*, 2020). Although GPP and vegetation greenness are undeniably linked, this relationship is typically weakened

under water or heat stress and it does not allow to detect early stress on vegetation (Zhao *et al.*, 2024; Maleki *et al.*, 2022). In particular, changes in biochemistry and photosynthetic activity cannot be related to greenness-based vegetation indexes. ChlF has the potential to capture the early response of ecosystems to climate extremes by being directly emitted from the core of photosynthesis reaction. One technical limitation remains : ChlF retrieved by active methods cannot be related to GPP because of the spatial and temporal mismatch of measurements. This has been overcome by the recent emergence of ChlF measurements with passive methods, which provide automatic ChlF measurements at a high frequency at larger spatial scale.

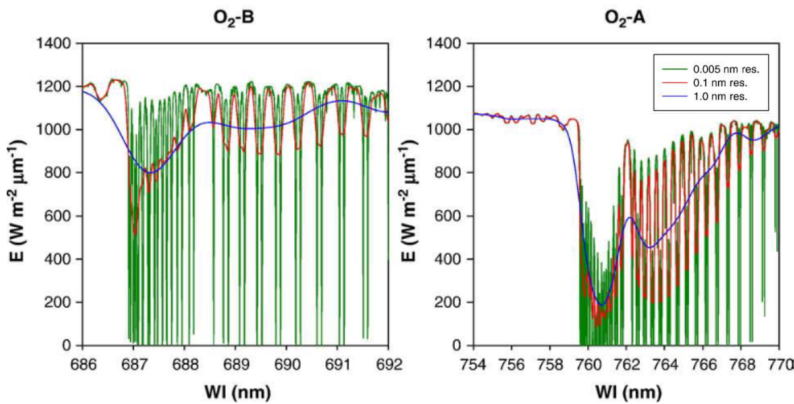
### 3.2.2. Measurements of chlorophyll fluorescence with passive techniques

Light is affected by absorption and scattering properties of molecules. In particular, solar irradiance is attenuated at specific wavelengths either by the Sun's atmosphere (i.e., known as Fraunhofer lines) or by the Earth's atmosphere (i.e., known as telluric absorption bands). More than 25 000 Fraunhofer lines have been identified such as the hydrogen H $\alpha$  line (656 nm) or the potassium KD1 line (770 nm). Another example is O<sub>2</sub> which has several absorption lines regrouped as bands (i.e., O<sub>2</sub>-A band 759-770 nm and O<sub>2</sub>-B band 687-692 nm). When induced by natural sunlight, ChlF is known as Sun-Induced chlorophyll Fluorescence (SIF), at the opposite of ChlF measured by active methods which is based on the excitation of ChlF pigments by an artificial light. As mentioned before (see section 2.4), SIF has two peaks in the red (~685-690 nm) and in the NIR (~730-740 nm) region (Mohammed *et al.*, 2019). The peak in the red is very close to the O<sub>2</sub>-B telluric absorption band, while the one in the NIR is slightly offset (Fig. 1.11).



**Figure 1.11:** Absorption spectra, fluorescence emission spectra (yellow) and oxygen absorption bands of chlorophyll (light green). Only the near-infrared and red domains are shown (Jonard *et al.*, 2020)

The shape of telluric bands and absorption lines is strongly influenced by the spectral resolution of the spectrometer. One important parameter for assessing the precision of spectral measurements is the Full Width of the absorption band at Half of the Maximum depth (FWHM); the depth being the difference between maximum and minimum reflectance within the band (Kruse & Lefkoff, 1999). Since the FWHM of spectral features is measured by fitting a Gaussian curve on the dark well, the precision of FWHM measurements is directly dependent on the spectral resolution of the spectrometer (i.e., the ability to detect observable wavelength intervals). For example, the O<sub>2</sub>-B band has a FWHM of 1.54 nm at 1 nm spectral resolution and a 0.03 FWHM at 0.005 nm spectral resolution (Meroni *et al.*, 2009). In other words, information on O<sub>2</sub>-A and O<sub>2</sub>-B bands is typically lost when spectral resolution increases (Fig. 1.12). One should note that the term FWHM can be used for both characterizing a spectral feature and the precision of a spectrometer. This is because the spectral resolution of a spectrometer is currently determined by measuring the FWHM of a reference light source.



**Figure 1.12:** Incident irradiance ( $E$ ) at ground levels and the two O<sub>2</sub> absorption bands at three different spectral resolutions (0.005, 0.1 and 2 nm). Adapted from Meroni *et al.* (2009).

Passive measurements of SIF use the similarity between dark lines and ChlF peaks to separate the SIF signal from the ambient light. More precisely, SIF contributes to fill-in the dark lines by adding a signal to the reflected radiance (the signal is directly proportional to the SIF emission spectrum). The upwelling radiance from vegetation ( $L(\lambda)$  - in blue on Fig. 1.13-B) is therefore composed of a reflected component ( $R(\lambda)$  - in green on Fig. 1.13-B) and an emitted component ( $F(\lambda)$  - in red on Fig. 1.13-B), which is related to SIF. Therefore, following Lambert's law, one can write:

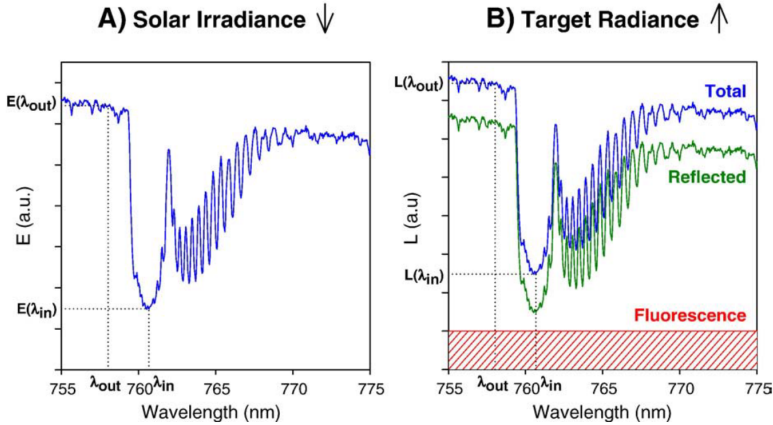
$$L(\lambda) = R(\lambda) + F(\lambda) = \frac{r(\lambda) E(\lambda)}{\pi} + F(\lambda) \quad (1.20)$$

with  $E(\lambda)$  the incident solar irradiance on the target (in blue on Fig. 1.13-A) and

$r(\lambda)$  the reflectance. The utilization the SIF contribution to the reflected radiance is at the basis of the Fraunhofer Line Depth (FLD) – based methods (Plascyk & Gabriel, 1975). The FLD method uses radiance and irradiance measurements outside ( $L(\lambda_{out})$ ,  $E(\lambda_{out})$ ) and inside ( $L(\lambda_{in})$ ,  $E(\lambda_{in})$ ) the well. Assuming that  $F(\lambda)$  and  $r(\lambda)$  remain constant, the FLD method consists in solving Eq. 1.20 outside and inside the well, which gives (Meroni *et al.*, 2009):

$$F = \frac{E(\lambda_{out}) L(\lambda_{in}) - E(\lambda_{in}) L(\lambda_{out})}{E(\lambda_{out}) - E(\lambda_{in})} \quad (1.21)$$

The main underlying hypothesis of the FLD method is that ChlF and reflectance do not vary on both sides of a dark line. This hypothesis is actually often challenged, which has led to the development of alternative FLD-based methods which use correction factors or spectral curve fitting to account for variability in ChlF and reflectance (Meroni *et al.*, 2009). Finally the Spectral Fitting Methods (SFM) uses spectral curve fitting on all available hyperspectral bands to determine variations in both reflectance and ChlF (Cogliati *et al.*, 2015). Additional information on SIF retrieval methods can be found in Meroni *et al.* (2009), Cendrero-Mateo *et al.* (2019) and Chang *et al.* (2020).



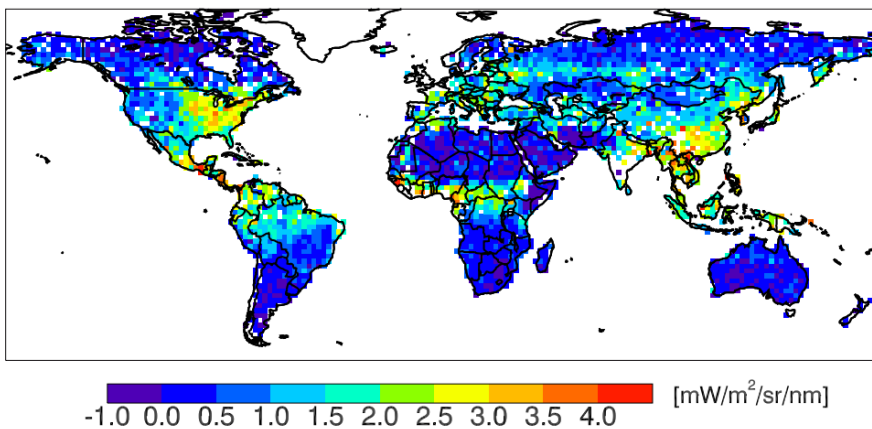
**Figure 1.13:** Solar irradiance measurements showing the  $O_2$ -A absorption band with a FWHM of 0.13 nm (A). Radiance measurements over vegetation (B) where the total signal (blue) is the sum of the reflected signal (green) and the fluorescence emission (red).  $\lambda_{out}$  and  $\lambda_{in}$  are the wavelengths nearby and inside the well (Meroni *et al.*, 2009).

### 3.3. Measuring SIF at large scales

SIF can be retrieved from spaceborne, airborne or field spectrometers with a high spectral resolution (Jonard *et al.*, 2020). First SIF observations were retrieved from satellite measurements which had a general purpose of studying the Earth's atmo-



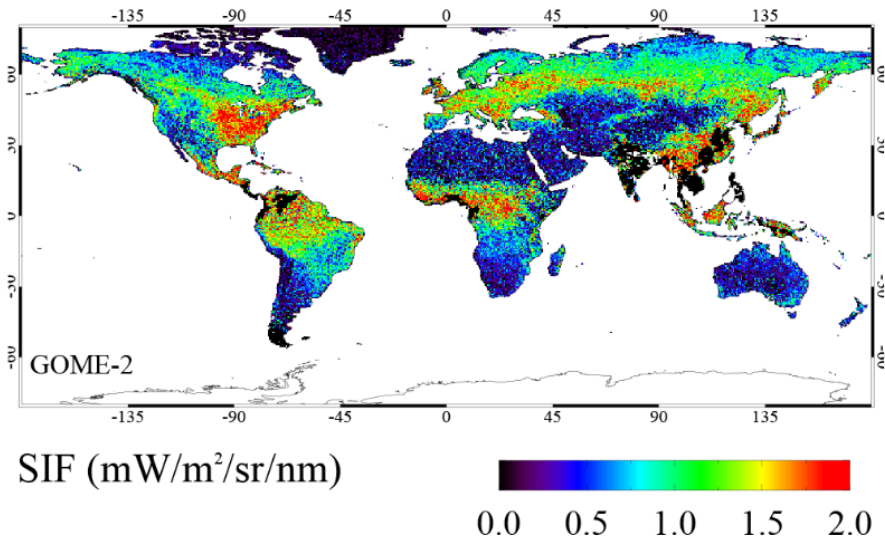
sphere. More than 15 years ago, the data acquired by the Medium Resolution Imaging Spectrometer (MERIS) on board of the ENVironmental SATellite (ENVISAT) was used by Guanter *et al.* (2007) to retrieve SIF in the O<sub>2</sub>-A band using the FLD method. Four years later, Joiner *et al.* (2011) used Greenhouse gases Observing SATellite (GOSAT) spectral measurements within the potassium Fraunhofer line KD1 to retrieve SIF (Fig. 1.14). The first comparison between SIF from GOSAT and GPP from EC towers on the ground was achieved the same year (Frankenberg *et al.*, 2011). A strong correlation between SIF and GPP was found, suggesting that SIF could be used to estimate GPP. As the technical characteristics of the following satellites improved (such as the Global Ozone Monitoring Experiment-2 (GOME-2) and the TROPospheric Monitoring Instrument (TROPOMI)), spatial resolution increased and global SIF maps became more precise (Fig. 1.14 to 1.16). Current TROPOMI SIF product has a pixel resolution of 19 km<sup>2</sup> at nadir (i.e., the direction pointing directly below a particular location, orthogonal to a horizontal flat surface) (Guanter *et al.*, 2021).



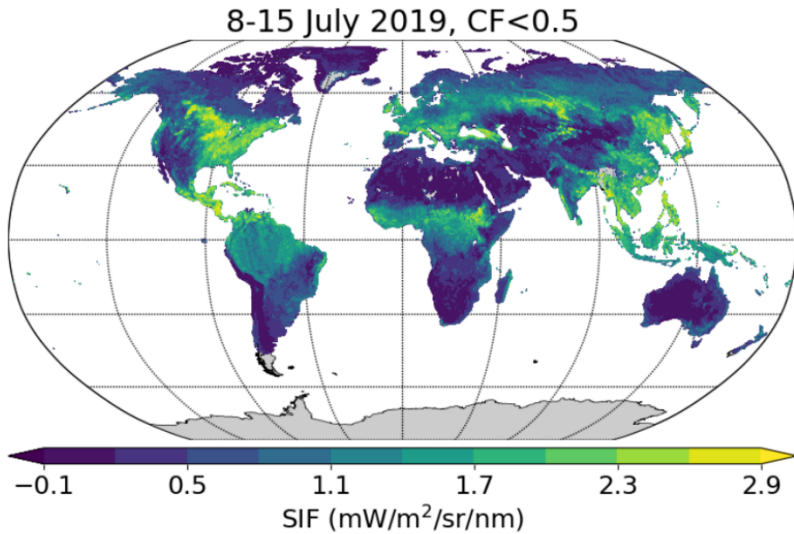
**Figure 1.14:** Global map of SIF monthly averages for July 2009 using GOSAT measurements. Data gridded in 2° grid boxes. Adapted from Joiner *et al.* (2011).

The first SIF-dedicated satellite will be launched by the FLUorescence EXplorer (FLEX) mission. FLEX will provide global maps of SIF emission of vegetation as primary measurement (Drusch *et al.*, 2017). The main instrument embedded on the FLEX space segment is the FLUOREscence Imaging Spectrometer (FLORIS) payload, a high precision spectrometer with a spectral resolution of 0.1 nm in the O<sub>2</sub> telluric bands and a pixel size of 300x300 m every 10 to 25 days, which is very similar to most of EC tower footprints. FLEX will orbit in tandem with Sentinel-3, providing joint measurements of SIF and Sentinel-3 data products to unravel complex ecosystem processes.

The development of FLEX has led to many ground campaigns for direct validation of future RS SIF products with the development of the High-performance airborne



**Figure 1.15:** Global map of SIF monthly averages for July 2013 using GOME-2 measurements. Data gridded in  $0.5^\circ$  grid boxes. Adapted from Guanter *et al.* (2015).



**Figure 1.16:** Global map of SIF averages between 8 and 15 July 2019 from TROPOMI measurements. Data gridded in  $0.2^\circ$  grid boxes. CF means cloud fraction. Adapted from Guanter *et al.* (2021).

imaging spectrometer for vegetation monitoring (HyPlant) airborne high-resolution spectrometer (Rascher *et al.*, 2015) or the field-based high-resolution Fluorescence boX (FloX) spectrometer (Buman *et al.*, 2022), which have similar technical features as FLORIS. While awaiting the launch of FLEX (expected for 2025), more than 40 FloX systems have been installed on EC sites, covering a wide variety of ecosystems (Buman *et al.*, 2022). Joint SIF and EC measurements have helped in understanding the relationship between SIF and GPP at the ecosystem scale (e.g., Helm *et al.* (2020); Magney *et al.* (2020); Marrs *et al.* (2020); Martini *et al.* (2022)), opening the way for building models estimating GPP from SIF (more details in section 4).

## 4. Photosynthesis modeling

Gas exchange measurements provide an important source of data describing the dynamics of net assimilation in a broad range of environmental conditions. For instance, gas exchange devices allow to change light intensity and CO<sub>2</sub> concentrations to build light-response or CO<sub>2</sub>-response curves under varying abiotic factors (e.g., temperature, soil moisture or relative humidity). While the response of CO<sub>2</sub> exchange is quite predictable (e.g., a saturation under high irradiance and CO<sub>2</sub> concentration, or a decrease under increasing VPD and lack of soil water), and the underlying mechanisms were well understood even before the 1980s, a global scheme for representing these processes and their interactions with the environment was lacking. With the feeling of "one doesn't really understand something until one can describe it mathematically" (Farquhar *et al.*, 2001), Farquhar, Berry and von Caemmerer brought biochemistry and gas exchanges together into the Farquhar-von Caemmerer-Berry (FvCB) model (Farquhar *et al.*, 1980). The FvCB model is a biochemical model for C<sub>3</sub> photosynthesis which is based on a mechanistic representation of the processes associated to carbon assimilation by the CBB cycle. It has been used for more than four decades to interpret gas exchange measurements and predict the impacts of environmental drivers on photosynthesis.

Biochemical processes in the FvCB model are dependant on the instantaneous supply of CO<sub>2</sub> and the effect of temperature on photosynthetic capacities, which leads to a nonlinear relationship between light and net assimilation (Farquhar *et al.*, 1980). When integrated over large spatial and time scales, ecosystem productivity is linearly proportional to the fraction of absorbed PAR by vegetation (Medlyn, 1998; Monteith *et al.*, 1997). These empirical relationships, also known as Light-Use Efficiency (LUE) models, can be generalized to various environmental conditions by adding a multiplier factor which accounts for the effects of environmental variables (other than PAR) on GPP (Pei *et al.*, 2022). While this straightforward approach can be used to model GPP at large scales, it does not allow an identification of the physiological factors limiting photosynthesis. In this perspective, the FvCB model provides a mechanistic framework for assessing the impacts of water stress on ecosystem productivity. Yet,

the response of model parameters on soil water availability remains a key uncertainty when modeling photosynthesis under water stress (Rogers *et al.*, 2017).

### 4.1. The Farquhar-von Caemmerer-Berry (FvCB) model

The mathematical representation of photosynthetic physiological processes in the FvCB model is at the basis of the modeling of GPP in many Land Surface Models (LSMs) (Rogers *et al.*, 2017). Net CO<sub>2</sub> assimilation rate ( $A_n$ ) is calculated as the minimum of three different processes: either by Rubisco activity ( $A_c$ ), electron transport ( $A_j$ ), or Triose Phosphate Utilization (TPU) ( $A_p$ ) (see section 2.2 for the role of triose phosphate in the CBB cycle). The equations for these limiting rates are:

$$A_c = \frac{C_c - \Gamma^*}{C_c + K_m} V_{cmax} \quad (1.22)$$

$$A_j = \frac{C_c - \Gamma^*}{4C_c + 8\Gamma^*} J \quad (1.23)$$

$$A_p = 3T_p \quad (1.24)$$

with  $C_c$  the CO<sub>2</sub> concentration in the chloroplast,  $\Gamma^*$  the CO<sub>2</sub> compensation point,  $K_m$  the Michaelis-Menten constant for Rubisco kinetics,  $T_p$  the rate of inorganic phosphate supply to the chloroplast,  $V_{cmax}$  the maximum carboxylation rate of Rubisco and  $J$  the potential electron transport rate.  $A_n$  is estimated as the minimum of the potential rates of these three processes minus the respiration in the light  $R_d$ :

$$A_n = \min(A_c, A_j, A_p) - R_d = A - R_d \quad (1.25)$$

While the modeling of the CBB cycle in the FvCB model is strictly mechanistic (i.e., based on Rubisco kinetics), the representation of light reactions is empirical. Indeed, the FvCB model calculates a potential electron transport rate between PSII and PSI instead of an actual rate.  $J$  is calculated using a rectangular hyperbola function of irradiance:

$$J = \frac{I_2 + J_{max} - \sqrt{(I_2 + J_{max})^2 - 4\theta_J I_2 J_{max}}}{2\theta_J} \quad (1.26)$$

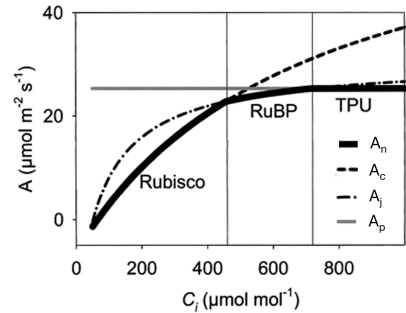
with  $I_2$  the actual light absorbed by PSII calculated from incident irradiance and leaf absorptance,  $J_{max}$  the maximum electron transport rate and  $\theta_J$  a curvature factor. One should note that  $C_c$  in Eq. 1.22 and Eq. 1.23 is the CO<sub>2</sub> concentration in the chloroplast, which is different from the CO<sub>2</sub> concentration in substomatal cavities. CO<sub>2</sub> pathways in the cell can be described using the Fick's law of diffusion (i.e., the electricity analogy of resistances - Farquhar & Sharkey (1982)), with stomata and mesophyll being the two limiting factors to CO<sub>2</sub> diffusion from the surrounding atmosphere to chloroplasts connected in series:

$$C_c = C_i - \frac{A_n}{g_m} = C_s - \frac{A_n}{g_s} - \frac{A_n}{g_m} \quad (1.27)$$

where  $C_i$  is the  $\text{CO}_2$  concentration in the substomatal cavities,  $C_s$  is the  $\text{CO}_2$  concentration at the leaf surface,  $g_m$  is the mesophyll conductance and  $g_s$  is the stomatal conductance to  $\text{CO}_2$  transfer. Note that  $g_s = g_{sw}/1.6$  with 1.6 being the ratio of  $\text{H}_2\text{O}$  to  $\text{CO}_2$  diffusivity. Most of the time,  $g_m$  is considered as infinite and Eq. 1.22 and Eq. 1.23 are used with  $C_i$  instead of  $C_c$ , and apparent  $V_{cmax}$  ( $V_{cmax,app}$  instead of  $V_{cmax}$ ). While kinetics constants of Rubisco ( $\Gamma^*$ ,  $K_m$ ) are regulated by leaf temperature in the same way for all C3 plants (Bernacchi *et al.*, 2001),  $J_{max}$  and  $V_{cmax}$  follow a species-specific relationship to temperature and are impacted by soil water status. The effect of these variables is commonly implemented using respectively a peak-Arrhenius model and an empirical factor considering the change in soil water availability (see section 5.2) (Medlyn *et al.*, 2002; Rogers *et al.*, 2017). These relationships must be calibrated on field data.

#### 4.1.1. Determination of FvCB model parameters from $\text{CO}_2$ and light-response curves

$\text{CO}_2$ -response curves are built by collecting net assimilation measurements under varying  $\text{CO}_2$  concentration with a LI-6400 equipped with a  $\text{CO}_2$  mixer. These curves typically show three phases illustrating the limiting rates of photosynthesis (Fig. 1.17). When  $\text{CO}_2$  concentration is low, the slope of the curve ( $dA_n/dC_i$ ) is high as  $\text{CO}_2$  is rapidly fixed by Rubisco which induces a sharp increase in  $A_n$ . As Rubisco progressively becomes saturated in  $\text{CO}_2$ ,  $dA_n/dC_i$  approaches zero and photosynthesis shifts from Rubisco-limiting to RuBP-limiting stage (Fig. 1.17). A second transition point may be observed when shifting to TPU-limiting stage with  $dA_n/dC_i$  slightly negative. This third limiting stage is not usually detected and therefore often neglected. FvCB model parameters (i.e.,  $V_{cmax}$ ,  $g_m$ ,  $T_p$ ,  $J$  and  $R_d$ ) can be determined by fitting the FvCB model on gas exchange measurements with, for instance, the R package 'Plantecophys' (Duursma, 2015; R Core Team, 2020). However, estimating model parameters from fitting methods may be challenging (Long, 2003; Sharkey, 2016). The fitting procedure requires to find the minimum of two non-linear equations (either Eq. 1.23, or 1.22 and 1.27) while adjusting five parameters (or four if  $A_p$  is neglected).



**Figure 1.17:** Idealized  $A_n$  vs  $C_i$  response curve simulated using the following FvCB model parameters:  $V_{cmax} = 70 \mu\text{mol m}^{-2} \text{s}^{-1}$ ,  $J = 130 \mu\text{mol m}^{-2} \text{s}^{-1}$ ,  $T_p = 9.1 \mu\text{mol m}^{-2} \text{s}^{-1}$  and  $R_d = 2 \mu\text{mol m}^{-2} \text{s}^{-1}$  (Long, 2003).

Several issues have been identified (Sharkey, 2016) such as strong variability of parameters according to the statistical fitting method (Miao *et al.*, 2009), or divergence of algorithms when data is collected on plants under water or heat stress as the saturation of  $A_n$  is not even reached (De Kauwe *et al.*, 2015b). The use of alternative methods for estimating parameters independently of gas exchanges is advised to reduce these uncertainties.

#### 4.1.2. Determination of FvCB model parameters from a single measurement of gas exchanges

An alternative method for estimating  $V_{cmax}$  consists in measuring  $A_n$  and  $C_c$  under ambient CO<sub>2</sub> concentration and saturated irradiance ( $R_d \ll A$ ,  $A = A_c$ ,  $A_n = A_c = A_{sat}$ ) by inverting Eq. 1.22, which gives:

$$V_{cmax} = A_{sat} \frac{C_c + K_m}{C_c - \Gamma^*} \quad (1.28)$$

In Eq. 1.28,  $g_m$  is considered as finite. The same equation can be applied considering a infinite mesophyll conductance (i.e.,  $C_c = C_i$ ). In these conditions, the apparent  $V_{cmax}$  parameter ( $V_{cmax,app}$ ) can be calculated by Eq. 1.28. In a global analysis using CO<sub>2</sub>-response curves from 594 species across various ecosystems, De Kauwe *et al.* (2015b) showed that  $V_{cmax,app}$  estimated from these single measurements correlated well with  $V_{cmax,app}$  estimated from  $A_n/C_i$  response curves (De Kauwe *et al.*, 2015b). These results support the use of one-point measurements of  $A_{sat}$  and  $C_i$  to estimate  $V_{cmax,app}$  instead of  $A_n - C_i$  response curves. Because of the challenges faced when estimating  $g_m$  and  $C_c$ ,  $V_{cmax,app}$  is generally reported in ecophysiological studies and used in LSMs (De Kauwe *et al.*, 2015b).

#### 4.1.3. Estimation of the CO<sub>2</sub> concentration in the chloroplasts from gas exchange and ChlF measurements

Estimating  $C_c$  requires first the determination of the actual electron transport rate ( $J_F$ ) which can be estimated from ChlF measurements of the quantum yield of PSII in the light ( $\phi_{PSII}$ ) as Genty *et al.* (1989) :

$$J_F = \alpha \beta_{PSII} \phi_{PSII} PPFD \quad (1.29)$$

with  $\alpha$  the leaf absorptance and  $\beta_{PSII}$  the distribution of electrons between PSII and PSI.  $\alpha$  and  $\beta_{PSII}$  should be ideally measured under non-photorespiratory conditions, or can also be set respectively to 0.83 and 0.5 (Genty *et al.*, 1989; Von Caemmerer, 2013).  $\alpha$  and  $\beta_{PSII}$  measurement procedure is detailed in chapter 5. Eq. 1.29 hypothesizes that a certain fraction of PAR can be directly used to induce electron transport between PSII and PSI.  $J_F$  can be used in the FvCB model as an actual electron transport rate under two conditions (Pons *et al.*, 2009): if alternative electron sinks are negligible, and if the leaf cross-section sampled for ChlF measurements can be con-

sidered as similar to the one for gas exchange measurements. This can be challenged as ChlF measurements use a red light to induce electron hopping between  $S_0$  and  $S_1$  orbitals, which has different penetration properties within the leaves compared to sunlight. This will be further discussed in chapter 5. Eq. 1.23 can be solved for  $C_c$  using  $J_F$ , which gives:

$$C_c = \frac{\Gamma^*(J_F + 8(A_n + R_d))}{(J_F - 4(A_n + R_d))} \quad (1.30)$$

and  $g_m$  can be determined by the Fick's law using measured  $C_i$  and  $C_c$  from Eq. 1.30:

$$g_m = \frac{A_n}{C_i - C_c} \quad (1.31)$$

## 4.2. Stomatal conductance modeling

Stomatal apertures are the gates of photosynthesis (Gago *et al.*, 2020). As detailed in section 2.3, stomatal opening is sensitive to multiple environmental factors. Early approaches for modeling stomatal conductance used empirical multiplicative functions integrating the effects of PAR, temperature, VPD and  $C_s$  on  $g_s$  (Jarvis *et al.*, 1976). With experimental growing evidences that stomatal conductance is strongly correlated with net carbon assimilation (Wong *et al.*, 1979), Ball *et al.* (1987) and Leuning (1995) integrated the dependence of  $g_s$  on  $A_n$  as (Eq. 1.32 and Eq. 1.33):

$$g_s = g_0 + g_1 RH \frac{A_n}{C_s} \quad (1.32)$$

$$g_s = g_0 + g_1 \frac{A_n}{(1 + \frac{VPD}{VPD_0}(C_s - \Gamma^*))} \quad (1.33)$$

where  $VPD_0$  is a constant. Note that  $g_0$  represents the minimum leaf conductance. It corresponds either to a stomatal conductance to water vapor during nighttime when photosynthesis does not occur, or to a cuticle conductance facilitating water vapor diffusion which is typically non-negligible under water stress (Boyer, 2015a).  $g_0$  is usually set to a constant value (Duursma *et al.*, 2019). In Eq. 1.32 and Eq. 1.33, the effect of temperature on  $g_s$  and  $A_n$  is implemented through  $RH$  or  $VPD$  (see Damour *et al.* (2010) for a review of empirical models of stomatal conductance).

An alternative to empirical approaches is to model  $g_s$  using the optimality framework of Cowan & Farquhar (1977). This approach assumes that stomata should be opened such as carbon assimilation is maximized while water losses are minimized, resulting in maximizing the following integral over a constant time interval (from  $t$  to  $t'$ , Fig. 1.18):

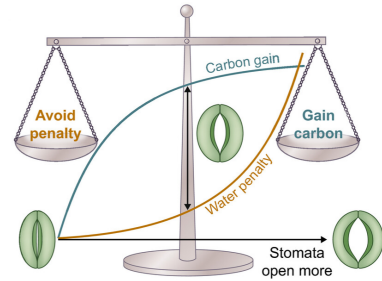
$$\int_t^{t'} (A_n(t) - \lambda_{opt} T_r(t)) dt \quad (1.34)$$

where  $\lambda_{opt}$  is the marginal carbon cost of water use. Medlyn *et al.* (2011a) used the optimality theory to derive a formulation (i.e., the Unified Stomatal Optimality (USO) model) that is closely analogous to the empirical models of Ball *et al.* (1987) and Leuning (1995):

$$g_s = g_0 + \left( 1 + \frac{g_1}{\sqrt{VPD}} \right) \frac{A_n}{C_s} \quad (1.35)$$

In the past 10 years, numerous stomatal conductance models have been developed based on the optimality theory (Mrad *et al.*, 2019). These models mainly differ from the formulation of the cost of stomatal opening, either attributed to maintaining a constant CO<sub>2</sub> concentration in the leaves (Prentice *et al.*, 2014), an impairment of the water flow through xylem cavitation and soil drying (Sperry *et al.*, 2016, 2017), or an impairment of photosynthetic apparatus through biochemical effects on the CBB cycle (Dewar *et al.*, 2018). The implementation of more mechanistic approaches for modeling  $g_s$  in LSMs is currently explored. While it is recognized that the representation of physiological processes is improved by such models, the calibration remains challenging because as it comes with a very high computational cost. So far, it does not seem that any stomatal conductance model rules all the others (Sabot *et al.*, 2022) and the USO model has become a reference for representing stomatal behavior in LSMs.

A key parameter of the USO model (Eq. 1.35) is the slope parameter  $g_1$ , which is inversely related to the Water-Use Efficiency (WUE) (i.e., the ratio of carbon gained to water use (Medlyn *et al.*, 2011a; Lin *et al.*, 2015a)). The concept of WUE is often used to characterize the productivity of ecosystems in relation to their water use and for understanding the response of photosynthesis to changes in water availability (Hatfield & Dold, 2019). The advantage of the USO model is that  $g_1$  can be linked to plant physiology through WUE. A decrease in  $g_1$  (i.e., increase in WUE) is expected when the optimality theory is verified (i.e., decreasing  $g_s$  while maximizing  $A_n$  (Mäkelä, 1996)). On the other hand, if  $A_n$  is limited by biochemical factors such as  $V_{cmax}$  or  $g_m$  in addition of the effect of decreasing CO<sub>2</sub> supply by stomatal closure,  $g_1$  is expected to



**Figure 1.18:** Stomatal optimization models hypothesize that the actual stomatal opening is what maximizes the difference between carbon gain and water penalty. Adapted from Wang *et al.* (2020b)



increase (i.e., decrease of WUE and of the ratio  $A_n/g_s$ ). Although the response of  $g_1$  to decreasing soil water availability is still highly uncertain (Zhou *et al.*, 2013), the USO model is currently implemented in climate models such as the Australian Community Climate Earth Systems Simulator ACCESS1.3b (Kala *et al.*, 2015a) or the Community Land Model v5 (CLM5) which is the main resource for Intergovernmental Panel on Climate Change (IPCC) assessment reports (Lawrence *et al.*, 2019).

### 4.3. Scaling from leaves to ecosystems with 'big leaf' models

The FvCB and USO models describe gas exchanges and their interdependencies with biotic and abiotic factors at the leaf level. One important question is how these models can be used to interpret and model fluxes at the canopy level. The issue of spatial extrapolation of gas exchange measurements from leaves to canopy has been investigated in the late 70's by Sinclair *et al.* (1976) who compared a complex micro-meteorological model to simpler models which assumed either that environmental conditions are constant within the canopy, or that the canopy can be simplified as a single leaf (i.e., a 'big leaf'). 'big leaf' models differs from multi-layer canopy models which resolve vertical micro-climate profiles within the canopy. Those are especially relevant for dense forest canopies where within-canopy gradient of environmental conditions can be observed (Bonan *et al.*, 2021; Luo *et al.*, 2018). However, single-layer canopy models are preferred to multi-layer canopy models due to the simplicity of the approach which avoids a heavy computational cost while estimating GPP with a good accuracy (Bonan *et al.*, 2021; Lawrence *et al.*, 2019).

### 4.4. Photosynthesis modeling from passive measurements of chlorophyll fluorescence

Using SIF to estimate GPP has several advantages. First, measurements of ChlF with passive techniques can be collected at various spatial scales and do not require the installation of a complex structure on the ground such as an EC flux tower. Second, satellite measurements of SIF have the potential to unravel the spatial and temporal dynamics of vegetation growth at regional and global scales, thus providing a key tool to upscale GPP measurements (Pei *et al.*, 2022).

#### 4.4.1. The Light-Use Efficiency approach

Similarly to GPP with the LUE model, a straightforward approach for estimating ground-based GPP from proximal sensing of SIF consists in assuming a proportionality to the absorbed PAR ( $APAR$ ) and ChlF quantum yield ( $\phi_F$ ) (Joiner *et al.*, 2011):

$$SIF = APAR \phi_F f_{esc} \quad (1.36)$$

where  $\tau_{atm}$  is the atmosphere transmittance which characterizes the scattering effect of the atmosphere (Köhler *et al.*, 2015) and  $f_{esc}$  is the probability that SIF photons

escape the canopy and reach the sensor (similar to Eq. 6.1). Combining the LUE approaches for SIF and GPP gives:

$$GPP = \frac{\phi_{PSII}}{\phi_F f_{esc}} SIF \quad (1.37)$$

Note that  $\tau_{atm}$  can be neglected when SIF is measured at the ecosystem scale. Eq. 1.37 translates the close coupling between SIF and GPP. The LUE approach provides a rationale for establishing empirical relationship between SIF and GPP at large spatial and temporal scales by using machine learning techniques (Guanter *et al.*, 2014; Pickering *et al.*, 2022; Bai *et al.*, 2022; Pei *et al.*, 2022).

#### 4.4.2. The Mechanistic Light-Response model

Empirical relationships do not always capture the variability of SIF and GPP. For instance, it has been shown that SIF may remain constant while GPP decreased during stomatal closure in several temperate tree species (Marrs *et al.*, 2020). A breakdown of the SIF-GPP relationship was also observed for forests during a heatwave (Martini *et al.*, 2022) or during fall transition (Kim *et al.*, 2021), which can be generalized at the canopy scale for several biomes (Liu *et al.*, 2022b). These evidences highlight a strong control of physiological and structural factors on the SIF-GPP relationship that are not capture by empirical relationships. Recently, Gu *et al.* (2019) combined the FvCB model and the first-order rate constants for the de-excitation of chlorophyll to link GPP to SIF using key physiological factors (i.e., the Mechanistic Light Response (MLR) model). The starting point of MLR model equations is the combination of Eq. 1.29 and Eq. 1.36, which gives:

$$J_{SIF} = \frac{\phi_{PSII} SIF}{\phi_F f_{esc}} \quad (1.38)$$

with  $J_{SIF}$  the electron transport rate estimated from SIF.  $\phi_{PSII}$  and  $\phi_F$  can be expressed as a function of the first-order rate constants for the de-excitation of chlorophyll pigments after being excited by sunlight:

$$\phi_{PSII} = \frac{k_p}{k_D + k_F + k_{NPQ} + k_p} \quad (1.39)$$

$$\phi_F = \frac{k_F}{k_D + k_F + k_{NPQ} + k_P} \quad (1.40)$$

where  $k_D$ ,  $k_F$ ,  $k_{NPQ}$  and  $k_P$  are respectively the first-order rate constants for the de-excitation of chlorophyll pigments by constitutive thermal dissipation, ChlF emission, NPQ and Photochemical Quenching (PQ). The ratio  $\phi_{PSII}/\phi_F$  can be linked to the fraction of open PSII centers ( $q_L$ ) and the maximum photochemical quantum yield

( $\phi_{PSII,max}$ ), which can be measured by ChlF measurements with active techniques (Eq. 1.5 and Eq. 1.4):

$$\frac{\phi_{PSII}}{\phi_F} = q_L \frac{\phi_{PSII,max}(1 + k_{DF})}{1 - \phi_{PSII,max}} \quad (1.41)$$

with  $k_{DF} = k_D/k_F$ . Combining Eq. 1.41 with Eq. 1.38 gives:

$$J_{SIF} = q_L \frac{\phi_{PSII,max}(1 + k_{DF})}{1 - \phi_{PSII,max}} \frac{SIF}{f_{esc}} \quad (1.42)$$

Combining Eq. 1.42 with the FvCB model (Eq. 1.23) gives:

$$GPP = q_L \frac{\phi_{PSII,max}(1 + k_{DF})}{1 - \phi_{PSII,max}} \frac{SIF}{f_{esc}} \frac{C_c - \Gamma^*}{4C_c + 8\Gamma^*} \quad (1.43)$$

The main advantage of representing  $J_{SIF}$  as a function of SIF is to provide an instantaneous linkage between the light and dark reactions of photosynthesis. By calculating the actual linear electron transport rate between PSII and PSI, the MLR model integrates the complexity of the energy partitioning processes through  $q_L$  and  $SIF$ . Therefore, if SIF is observed, " (...) the complexity of modeling photosynthesis is greatly reduced because it integrates over this complexity" (Gu *et al.*, 2019).

In the MLR model,  $\phi_{PSII,max}$  and  $q_L$  must be determined from leaf-level actively-induced ChlF measurements. The LeafWeb database (Han *et al.*, 2022a) regroups measurements of leaf gas exchanges and PAM fluorometry of light, CO<sub>2</sub>, O<sub>2</sub>, and temperature responses from several C3 and C4 species. Researchers can upload their measurements in the database, which contributes to the effort initiated by Gu *et al.* (2019) and Han *et al.* (2022a) to provide model parameters for different species.  $\phi_{PSII,max}$  and  $q_L$  can also be found in the literature as numerous studies have already independently measured these parameters (e.g., Cendrero-Mateo *et al.* (2015); Zivcak *et al.* (2014)). Existing studies have tested the MLR model at the leaf scale under controlled conditions (Shi *et al.*, 2022), or at the ecosystem scale for an irrigated winter wheat crop (Liu *et al.*, 2022c). Large uncertainties remain on the capability of the MLR model for estimating GPP under water stress. In addition, coupling the MLR model to a model of stomatal conductance could allow the estimation of GPP and Tr from a single SIF measurements. These research questions should be addressed before considering an application of the model at larger scales with remote sensing data.

## 5. Drought impacts on photosynthesis

Unraveling the impacts of drought on carbon and water fluxes is pivotal for understanding ecosystem functioning and improving model predictions in the context of climate change. Photosynthesis model parameters must account for environmental effects on physiological processes to predict the future strength of carbon land sinks (Fisher *et al.*, 2014; Friedlingstein *et al.*, 2022).

### 5.1. What is drought ?

The World Meteorological Organization (1992) defines drought as a “period of abnormally dry weather sufficiently prolonged for the lack of precipitation to cause a serious hydrological imbalance”. The diversity of temporal distribution, spatial scale and disciplines affected by drought explains that “...we cannot reasonably expect the existence of any workable generalized objective definition of drought.” (Lloyd-Hughes, 2014). As a consequence, different drought definitions coexist. Among them, physical-based droughts are commonly classified into three categories (Belal *et al.*, 2014): meteorological drought (lack of precipitation compared to a reference period), agricultural drought (crop failure caused by a decrease in soil moisture), and hydrological drought (limiting period when surface water resources are no longer available for water management systems). These definitions must be associated to the calculation of parameters which characterizes the status of the physical variable of interest. Such parameters, also known as drought indexes, can be a single variable (e.g, precipitation) or a combination of different variables (e.g, the Standardized Precipitation Evaporation Index (SPEI) which uses both precipitation and potential evapotranspiration). A flurry of drought indexes has been defined over the past decades, which constantly evolves with the availability of new data or techniques to monitor the impact of climate change on vegetation (Gebrechorkos *et al.*, 2023).

A more relevant definition of drought when considering vegetation is ecological drought, which is defined as an “episodic deficit in water availability that drives ecosystems beyond thresholds of vulnerability, impacts ecosystem services, and triggers feedbacks in natural and/or human systems” (Crausbay *et al.*, 2017). This definition puts plant hydraulics and physiological processes at the core of the analysis of water stress effects on ecosystems. Assessing ecological droughts requires the joint analysis of the relationship between a plant functional trait and a physical variable characterizing the water status in the soil (e.g., the relationship between stomatal conductance, photosynthesis at light saturation, or the hydraulic conductivity with Soil Water Content (SWC) (Anderegg *et al.*, 2019; Maynard *et al.*, 2022; Quetin *et al.*, 2023)). For instance, the drought resistance of an ecosystem can be quantified by analyzing the differences in plant functional traits under similar edaphic conditions. A key aspect for characterizing ecological droughts is the separation of the effect of soil water status among all the other meteorological conditions which impact plant processes. This

requires a detailed knowledge of the effects of temperature or atmospheric dryness on photosynthesis-related variables such as stomata or Rubisco (see section 4).

## 5.2. Drought limitations on the water flow in plants

### 5.2.1. The transpiration stream from soil to atmosphere

Lack of soil water poses major perturbations on water, carbon and energy fluxes in the plant. The Soil-Plant-Atmosphere Continuum (SPAC) refers to a physically integrated and dynamic system which describes the water pathway from soil via plants to atmosphere. Water flows within a SPAC component ( $X$ ) following Darcy's law (Darcy, 1856) which establishes a direct proportionality between a gradient of water potentials and the hydraulic conductance following:

$$q = K_X(\psi_X)(\psi_{X_{i-1}} - \psi_X) \quad (1.44)$$

with  $\psi_{X_{i-1}} - \psi_X$  the difference in water potential between two consecutive SPAC components (with water flowing from more positive to more negative values),  $q$  the water flux and  $K_X(\psi_X)$  the hydraulic conductance in the  $X$  SPAC component which is expressed as a function of  $\psi_X$ . For example, the flow equation in the leaf is represented by:

$$q = K_{leaf}(\psi_{leaf})(\psi_{xylem} - \psi_{leaf}) \quad (1.45)$$

Assuming a single component in the SPAC, a constant hydraulic conductance in the plant, no water storage and no cavitation, Eq. 1.44 gives:

$$q = K_{plant}(\psi_{soil} - \psi_{atm}) \quad (1.46)$$

This very simple representation of water flow in plants has served as a base for the development of more complex approaches where hydraulic conductances and potentials are broken down into different components as a function of plant tissues and organ proprieties (Carminati & Javaux, 2020). It is clearly evidenced in Eq. 1.46 that transpiration occurs only if the water potential of the atmosphere is lower than the water potential of the soil (counted negative). These water potentials are mediated respectively by the VPD in the atmosphere and the SWC in the soil, both related to the moisture of the media. This gradient of potentials between the soil and the atmosphere allows water to be transported by the cohesion-tension mechanism through the SPAC from roots to leaves (Wolfe *et al.*, 2023). Realistic representations of the SPAC includes many different structures, connected through a dense network of conductances and capacitances following the electric analogy (Hunt *et al.*, 1991). Recent models represent the SPAC as a complex network composed of numerous components including dozens of conductances and capacitances (Cochard *et al.*, 2020). Representing plant hydraulics

yields to complex models that will be described elsewhere in this manuscript (section 3.4). Among the different components of the SPAC, the Soil Water Potential (SWP) and Leaf Water Potential (LWP) are key regulators of the plant water status and its response to edaphic drought (De Swaef *et al.*, 2022; Carminati & Javaux, 2020).

### 5.2.2. Impact of drought on water potentials in the SPAC

When stomata open to sustain CO<sub>2</sub> diffusion to chloroplasts for photosynthesis, it inevitably leads to a loss of water through transpiration. Under high VPD, the water in the leaves is evaporated, the suction forces between water molecules and leaf cells increase, and so does the LWP towards negative values. Water flows from roots to leaves by either osmotic absorption (when the atmospheric demand is low) or passive absorption (when the atmospheric demand is high) through the SPAC. The tension in the xylem vessels can be modulated by aquaporins which facilitates water transport from the roots to the plant vascular system (Chavarria & dos Santos, 2012). The transpiration flux is maintained as long as roots keep supplying water from the soil to meet the evaporative demand of the atmosphere.

In any aqueous SPAC component, the water potential  $\psi_{X,aq}$  can be decomposed as the sum of three components: the matric potential  $\psi_{X,mat}$  (due to water-matrix interactions), the gravitational potential  $\psi_{X,gra}$  (due to the gravitational force field) and the osmotic potential  $\psi_{X,osm}$  (due to electrolytes in the water) (Or *et al.*, 2005):

$$\psi_{X,aq} = \psi_{X,gra} + \psi_{X,osm} + \psi_{X,mat} \quad (1.47)$$

In saturated soils (i.e., all soil pores are filled with water), the gravitational potential is dominant and the matric potential is negligible. In unsaturated soils, when gravitational water has drained out, soil pores exert a force on water molecule high enough to hold water against gravity. The SWP is therefore strongly constrained by the matric potential, which is determined by the surface tension ( $T_{cap}$ ), the radius of the curvature of the air-water interface ( $r_{cap}$ ), the soil-water contact angle ( $\alpha_{cap}$ ) and a adsorptive component directly related to the Van der Waals interaction forces, hydrogen bonding, and electrostatic forces between water molecules and soil pores ( $A_{dv}(z)$  at a depth  $z$ ), following (Tuller *et al.*, 1999):

$$\psi_{soil,mat} = -\frac{2T_{cap} \cos(\alpha_{cap})}{r_{cap}} + A_{dv}(z) = \underbrace{-\frac{2T_{cap} \cos(\alpha_{cap})}{r_{cap}}}_{\text{Capillary water}} + \underbrace{\frac{A_H}{6\pi z^3}}_{\text{Adsorptive water}} \quad (1.48)$$

where  $A_H$  is the Hamaker constant between soil pores and water molecules, which varies with the mineral type of the soil particle (Tuller *et al.*, 1999). In unsaturated soils, both capillary and adsorptive water co-exist above a typical value of  $-10^4$  kPa. Under soil drying, capillary water is the being extracted by the plants or being evaporated at the soil surface. Soil water enters into the adsorptive film water regime, which

is the transition zone between capillary and adsorptive soil water. Capillary water disappears from  $-10^5$  kPa for most soils (Luo *et al.*, 2022). It is assumed that water uptake can no longer occur from a matric potential of  $\sim -1.5 \times 10^3$  kPa, known as the permanent wilting point located near the adsorbed film region (Richards & Weaver, 1943). From this point onwards, it is hypothesized that soil water cannot be extracted by the roots due to the increasing predominance of adsorptive forces over capillary forces. In the definition of the wilting point itself lies a strong assumption regarding the root water uptake capacities of plants. Is it impeded from the wilting point or from higher soil potential values? The capacity SPAC components to sustain the water flow is characterized by vulnerability curves, which are pivotal for predicting the distribution of water potential in the plant (see chapter 7, section 2.2.1). Water transport in the plant may cease when SWP is significantly higher than the wilting point. This is one of the main uncertainties when modeling the response of ecosystem to water stress (Rogers *et al.* (2017), see section 5.4).

Critical water potentials trigger hormonal and molecular responses at the whole plant level to avoid desiccation (Aroca *et al.*, 2012). For instance, closure of aquaporins, changes in morphological features (e.g., root depth, root length, or root angles) or tissue dehydration (Aroca *et al.*, 2012; Kou *et al.*, 2022; Jalakas *et al.*, 2021) disrupts the transpiration flow along the SPAC. In particular, aquaporins can facilitate the transport of solutes across the cell membrane and act as a regulator of root water uptake (Aroca *et al.*, 2012; Gambetta *et al.*, 2017; Javot & Maurel, 2002). ABA is also known to be a major hormone regulating the architecture of the root system and the overall plant response to drought (Harris, 2015; Sun *et al.*, 2018). Early sensing of root-induced ABA by leaves allows a quick reaction of plant to the increasing suction forces in the soil by closing stomata to prevent water losses even when LWP remained constant (Davies *et al.*, 2005). Moreover, ABA is known to affect aquaporins gene expression which results in a regulation of root hydraulic conductivity (Kapilan *et al.*, 2018). While it is well established that critical levels of turgor pressure in stomata guard cells triggers the synthesis of ABA (McAdam & Brodribb, 2016), the mechanisms inducing ABA production in roots during water are still poorly understood (Thompson *et al.*, 2007). The increase in ABA biosynthesis by the roots and the facilitation of ABA transport through xylem vessels to stomata is a key adaptive feature for regulating growth and gas exchanges (Davies *et al.*, 2005; Kou *et al.*, 2022; Liu *et al.*, 2022a). As a result, LWP is tightly linked to soil water status and is one of the most important ecophysiological variable for monitoring plant water stress (Ratzmann *et al.*, 2019). Consequently, stomatal dynamics operate on a thin safety margin by dynamically regulating the water flow to protect the vascular system from hydraulic failure while still facilitating CO<sub>2</sub> diffusion to chloroplasts for photosynthesis (Martínez-Vilalta & Garcia-Forner, 2017).

### 5.2.3. SWC as a proxy of LWP

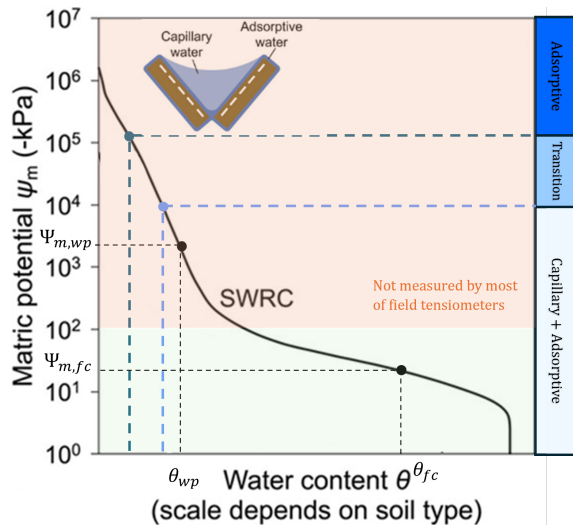
While being directly related to the forces holding water to plant cells or soil pores, measuring water potentials presents some key limitations. First, LWP is typically measured by labor-consuming and destructive sampling techniques, which complicates the monitoring of plant water status at sub-daily and seasonal temporal scales, or across an entire ecosystem (García-Tejera *et al.*, 2021). Second, in-situ automatic measurements of SWP by tensiometers only provide limited data for studying the response of plants to water stress by (i) not covering the entire range of possible SWP values (Fig. 1.19) with an upper limit of SWP typically below 100 kPa due to cavitation (Menne *et al.*, 2022; Dasgupta *et al.*, 2015; Durner & Or, 2006), (ii) showing low sensitivity to drying-rewetting repeated cycles (noz Carpena *et al.*, 2005), (iii) showing lags in SWP readings when equilibrium is particularly slow to reach and (iv) being sensitive to cavitation of water bubbles at very high pressure in water-filled tensiometers (Bittelli, 2010). The recent development of microtensiometers might represent an interesting technical solution for continuous measurements of LWP (Pagay, 2022; Christenson *et al.*, 2024).

Field automatic SWC sensors are very common and do not suffer from the same shortcomings. SWC is routinely measured on EC sites and provides an estimation of the water content in the soil as a percentage of the control volume of the soil sampled by the sensor. Global in situ soil moisture measurements are also regrouped within unique datasets spanning across many Plant Functional Type (PFT) and climate zones (Dorigo *et al.*, 2011). In addition, SWC in the topsoil can be retrieved at large scales from RS of microwaves from satellites (Njoku & Entekhabi (1996); Zhang *et al.* (2019a), also detailed in chapter 7, section 3.9). The bridge between soil moisture and water potentials is provided by Soil Water Retention Curves (SWRC)s, which describe the relationship between SWC and SWP (Fig. 1.19). These curves are built from (i) joint measurements of SWP and SWC, and (ii) models for describing their relationship such as the van Genuchten (VG) (van Genuchten, 1980) and the Brooks-Corey (Brooks & Corey, 1966) models (Tuller & Or, 2005). SWRCs can also be built by using neural network predictions if the soil type is known (Pham *et al.*, 2019; Novick *et al.*, 2022). These curves are a major soil hydraulic propriety and can be used to illustrate the limiting soil water ranges for plant uptake (Fig. 1.19, (Pan *et al.*, 2019)).

For instance, SWRCs can be used to retrieve the lower and upper limits of accessible water for plant uptake (i.e., the SWC at a SWP of  $\sim$ -1500 kPa and  $\sim$ 33 kPa or  $\sim$ 10 kPa respectively) respectively known as the wilting point and the field capacity. The difference between the field capacity and the wilting point provides an estimation of the available water for plant uptake in the root zone. The ratio of the actual extractable water (i.e., the difference between the actual soil water content and the wilting point) to the available water (i.e., the difference between the field capacity and the wilting point) is known as Relative Extractable Water (REW) (Eq. 1.49), which is often used



for characterizing the soil water status (Granier *et al.*, 1999a, 2007a), or for defining thresholds from which water should be added to crops to avoid yield losses (Juenger & Verslues, 2022). Note that other names can be found in the literature such as Fraction of Transpirable Soil Water (FTSW), Plant Available Soil Water Content (PASWC), Available Soil Water Content (ASWC) or Soil Water Availability (SWA). The wide availability of SWC data, and the numerous evidences of a strong control of soil moisture on photosynthesis (e.g., Lv *et al.* (2023); Keenan *et al.* (2010b); Trugman *et al.* (2018)) have supported the use of SWC as a key variable controlling water and carbon exchanges. As a result, this variable is used in many LSMs to reproduce drought effects on photosynthesis with the empirical approach, which consists in downregulating model parameters from a threshold of REW. This is detailed in section 5.4 and further discussed in chapter 7, section 5.4.



**Figure 1.19:** Conceptual illustrations of soil water retention regimes at different ranges of matric potential. Most of field tensiometers do not capture the full range of the SWRC. The field capacity and wilting points are set to the SWC at a SWP of -33 kPa and -1500 kPa. Adapted from Luo *et al.* (2022).

### 5.3. How drought affects photosynthesis ?

Soil drying triggers complex mechanisms in the plant to avoid hydraulic failure and cavitation. One of these mechanisms is stomatal closure, which can be induced by either passive (i.e., a decrease in leaf water potential and a loss in turgor pressure induced by hydraulic signaling) or active factors (i.e., a sensing of molecules such as ABA,  $\text{Ca}^{2+}$ , or ROS) (Merilo *et al.*, 2018; Jalakas *et al.*, 2021; Nardini & Salleo, 2000). The closure of stomata limits the diffusion of  $\text{CO}_2$  to chloroplasts, which in turn induces a decrease in photosynthesis. As being one of the three pathways for dissipating the

energy carried by sun irradiance, the limitation on photosynthesis triggers a cascade of effects which affects whole plant metabolism. One of the consequences of stomatal closure is the reduction of NADP+ regeneration in the CBB cycle, thus limiting light reactions and electron transport. The limitation of photosynthesis by constraining CO<sub>2</sub> supply also leads to an accumulation of ROS and photoinhibition damages (Cruz De Carvalho, 2008). Severe water shortage episodes also directly impacts the CBB cycle by regulating Rubisco synthesis through the production of tight-binding inhibitors (Parry, 2002). In addition, a decrease in mesophyll conductance can also be observed under water stress, which increase the constraint on CO<sub>2</sub> supply to chloroplasts (Flexas *et al.*, 2012). Additional information regarding drought effects on the photosynthetic apparatus will be described in the next chapters.

The origins of photosynthesis limitations have been historically defined along the stomatal versus non-stomatal categorization, first described by Jones (1985) and later mathematically formulated by Grassi & Magnani (2005). The Jones' definition is based on the assumption that photosynthesis limitations either originate from a closure of stomata through a decrease in stomatal conductance ( $g_s$ ), and by non-stomatal factors through a decrease in the apparent maximum carboxylation rate of Rubisco ( $V_{cmax,app}$ ). These origins are refereed in this thesis to as respectively Stomatal Origin Limitation of photosynthesis (SOL) and Non-Stomatal Origin Limitation of photosynthesis (NSOL). This framework has been widely used to determine the importance of each physiological factor in the reduction of carbon assimilation under water stress (e.g., Grassi & Magnani (2005); Grassi *et al.* (2009); Zait & Schwartz (2018)). One main drawback of this approach is the strong correlation between stomatal conductance and photosynthesis (Wong *et al.*, 1979). In particular, in most of stomatal conductance models such as the USO model, both variables are linked to coordinate the opening of stomata with the supply of CO<sub>2</sub> to the fixation sites in the chloroplasts (Lemonnier & Lawson, 2023). As a result, a decrease in  $g_s$  can also originate from a decrease in  $A_n$ ,  $V_{cmax}$  or  $g_m$ , potentially leading to an underestimation of the contribution of NSOL in limiting photosynthesis (also corresponding to an overestimation of SOL). An alternative approach consists in using the USO model to quantitatively decompose the effects of stomatal closure drivers (i.e.,  $VPD$ ,  $A_n$ ,  $g_1$  or  $C_s$ , Eq. 1.35), and to redefine the stomatal control on photosynthesis. Instead of using a decrease in  $g_s$  as an evidence of SOL, Zhou *et al.* (2013) suggested to associate the stomatal effect on  $A_n$  to a change in the stomatal sensitivity to photosynthesis  $g_1$ , which varies according to the water cost of carbon gain (Medlyn *et al.*, 2011a). When photosynthesis limitations originate from NSOL,  $A_n$  is affected by additional internal factors ( $g_m$ , or  $V_{cmax}$ ) which decouple the relationship between  $A_n$  and  $g_s$ , and is evidenced by an increase in  $g_1$ . On the other hand, SOL on photosynthesis is associated to an increase in the ratio  $A_n/g_s$  (i.e., maximizing carbon gains while minimizing water losses, decrease in  $g_1$ ) or a constant  $A_n/g_s$  (i.e., no maximization of carbon gains : both  $A_n$  and  $g_s$  decrease with the same magnitude,  $g_1$  constant). With this framework, variation of

$g_1$  can be used to clearly identify the role played by non-stomatal factors and avoids any underestimation of NSOL in limiting photosynthesis (Zhou *et al.*, 2013).

While the response of SOL and NSOL to water stress provides an indication of the dominant mechanism constraining photosynthesis, it does not allow to quantify the importance of these processes. To that end, partial derivative calculus can be used to link the total derivative of photosynthesis to the total derivative of the explanatory variables according to a selected model. With this approach, the relative variation of  $A_n$  compared to a reference value can be decomposed as a sum of the relative variations of the explanatory variables. The reference value can be typically associated to the soil water availability conditions that are not limiting for physiological processes. This framework can be adapted to any model for quantifying the role of model parameters in limiting photosynthesis under water stress if (i) the effects of other environmental factors are considered (e.g., by normalizing to 25°C and by collecting measurements at light saturation), (ii) the explanatory variables are measured and (iii) their relationship to soil water availability are characterized. It is therefore possible to combine the FvCB and USO models in the limitation analysis to quantitatively assess the importance of the constrains of  $VPD$ ,  $g_1$ ,  $C_s$ ,  $g_m$  and  $V_{cmax}$  on  $A_n$  and  $g_s$ . See chapter 5 for more details.

## 5.4. How to implement drought effects on photosynthesis in modeling

Investigating the response of physiological processes to the decrease in soil water availability is pivotal for understanding drought effects on ecosystems. This can be used for studying ecosystem resilience to drought or for improving the representation of water stress impact on physiological processes in LSMs. Most LSMs use an empirical factor (i.e., the  $\beta$  factor) which acts as a modifier on model parameters to reproduce the effects of soil moisture on photosynthesis (Rogers *et al.*, 2017).  $\beta$  is directly linked to REW by (Egea *et al.*, 2011; Harper *et al.*, 2021):

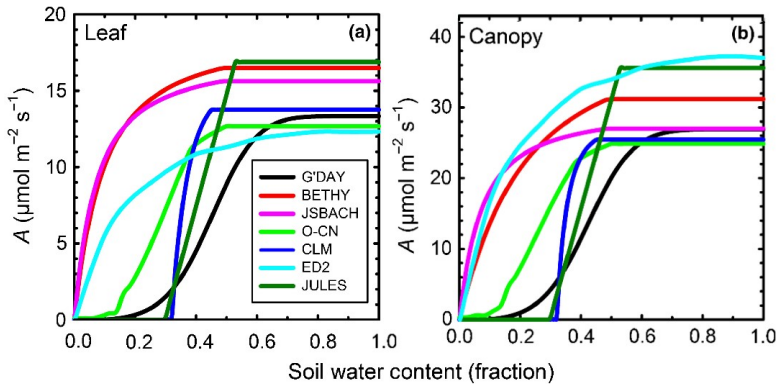
$$\beta = \begin{cases} 1 & \theta \geq \theta_{sup} \\ \left( \frac{\theta - \theta_{wp}}{\theta_{fc} - \theta_{wp}} \right)^p = REW^p & \theta_{inf} < \theta < \theta_{sup} \\ 0 & \theta \leq \theta_{inf} \end{cases} \quad (1.49)$$

with  $p$  allowing to implement a nonlinear dependence of model parameters on REW,  $\theta_{fc}$  the field capacity,  $\theta_{wp}$  the wilting point and  $\theta$  the actual SWC of the soil layer. Eq. 1.49 can be applied to the different soil layers in the root zone. A weighting parameter is used to consider the proportion of Root Biomass Density (RBD) or rooting depth within each soil layer. For instance, in the Joint UK Land Environment Simulator (JULES) model, this weighting factor is expressed as a function of the root fraction (Harper *et al.*, 2021):

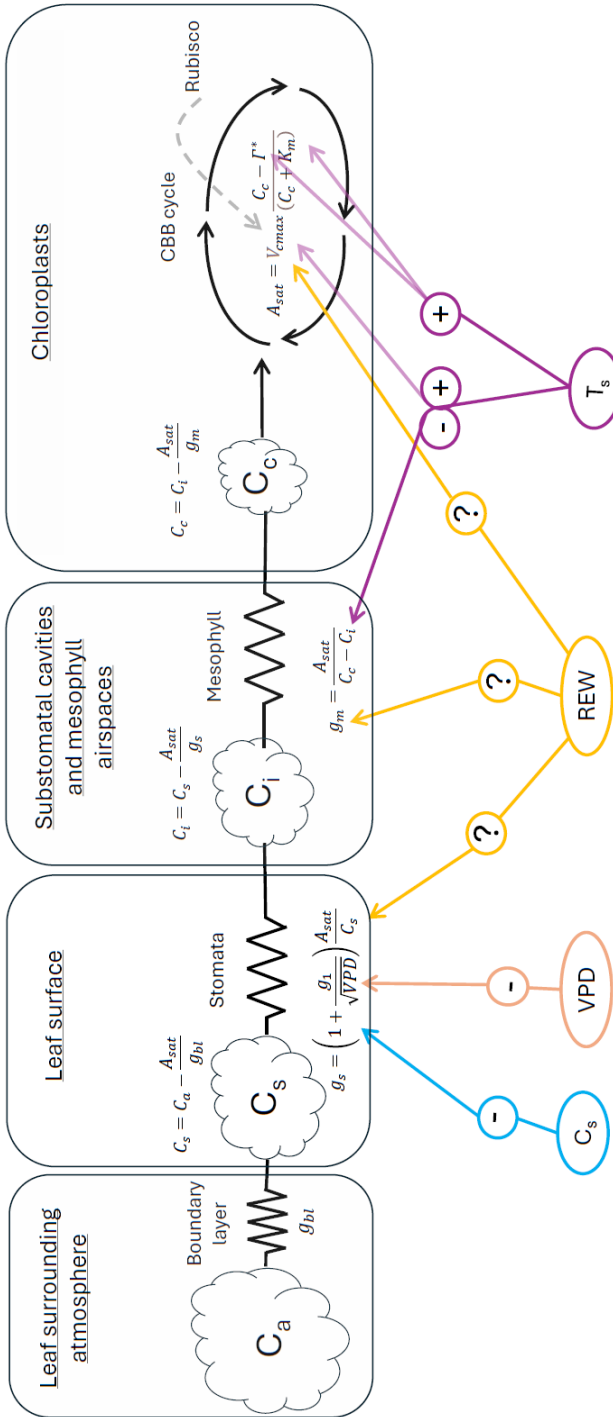
$$\beta = \sum_{k=1}^{n_{soil}} e^{z_k/d_r} \beta_k \quad (1.50)$$

where  $k$  is the soil layer index,  $n_{soil}$  is the total number of soil layers,  $z$  is the depth of the  $k$  soil layer,  $\beta_k$  is  $\beta$  within each soil layer  $k$ , and  $d_r$  a PFT-specific parameter which consider the root mass distribution in the soil (i.e., the smaller the  $d_r$ , the more important root biomass in the shallow soil layers). Other approaches consist in weighting  $\beta$  by a factor proportional to the RBD as implemented in several LSMs such as in Community Atmosphere Biosphere Land Exchange (CABLE) or in ORCHIDEE (Haverd *et al.*, 2016; Best *et al.*, 2011; Haverd *et al.*, 2016).

In addition to the root distribution profile, other important aspects of  $\beta$  calculation should be highlighted. First, it is unclear which parameter should be affected by  $\beta$  to reproduce realistic responses of photosynthesis and transpiration to water stress (Trugman *et al.*, 2018; Rogers *et al.*, 2017). Second, in Eq. 1.49,  $\beta$  affects SOL or NSOL when REW drops below  $\theta_{fc}$ . However, it has been shown that plants can maintain high photosynthetic capacities and sustain transpiration even when  $\theta \leq \theta_{fc}$  (e.g., Granier *et al.* (2007a)). As a result, the shape and thresholds of  $\beta$  are PFTs-specific and induce major divergences in the response of net photosynthesis to water stress (e.g., Fig. 1.20). Calibrating  $\beta$  on field data is therefore needed (Rogers *et al.*, 2017; Vidale *et al.*, 2021). An overview of the interdependences between stomatal conductance, photosynthesis at light saturation and abiotic drivers as formulated in the FvCB and USO models is shown in Fig. 1.21.



**Figure 1.20:** Examples of photosynthesis response to REW in several LSMs (Rogers *et al.*, 2017).



**Figure 1.21:** Diagram of the effects of abiotic factors on stomatal conductance and photosynthesis at light saturation. The control exerted by one variable over another (either a positive (+) or negative (-) relationship) is determined from calculating the partial derivatives using the FvCB and USO models. The equations for  $g_s$ ,  $g_m$ ,  $A_{sat}$ ,  $C_i$  and  $C_c$  can be found in section 4.  $g_{bl}$  is boundary layer conductance. The double sign for the surface temperature  $T_s$  refers to the Arrhenius-peak model: either a positive relationship below the optimum ( $25^\circ\text{C}$ ), or negative above the optimum.



## **Objectives and structure of the thesis**





## 1. Objectives of the thesis

Characterizing the response of photosynthesis model parameters under water stress is crucial for identifying the factors limiting GPP in the context of climate change. Modeling these constraints requires the calibration of  $\beta$ , an empirical factor responsible for implementing drought effects on photosynthesis in LSMs. In this thesis, we specifically focused on crops, which are key components of terrestrial ecosystems by participating to the regulation of the carbon and water cycles, and by providing pivotal ecosystem services such as food production for humanity and livestock. Large uncertainties still remain for identifying the physiological constraints on photosynthesis and transpiration of croplands under soil water stress (Yuan *et al.*, 2015; Xia *et al.*, 2021; Jiang *et al.*, 2023). The BIODYNE research department has a long-term expertise in analyzing gas exchanges of crops by managing the EC site of Loncée (site code BE-Lon) for more than 20 years. This 4-year rotation cropland is typical of central Belgium and is equipped with a collection of instruments for monitoring fluxes and their environmental drivers at the ecosystem-atmosphere interface (Delandmeter *et al.*, 2023; MOUREAUX *et al.*, 2008; Lognoul *et al.*, 2019; Moureaux *et al.*, 2012; Buysse *et al.*, 2017; Moureaux *et al.*, 2006; Bachy *et al.*, 2013; Suleau *et al.*, 2011; Kutsch *et al.*, 2010).

This thesis started in 2018, which was a year of severe drought in western Europe. This year, potato (*Solanum tuberosum*) was cultivated at BE-Lon, as part of the 4-year rotation with winter wheat (*Triticum aestivum*) and sugar beet (*Beta vulgaris*). At the same time, Louis Gourlez de la Motte (former BIODYNE researcher) started his post-doc on the effects of drought on forest ecosystems. The objectives of his work were very similar to those of this thesis: understanding the constraints of drought on forests by calibrating  $\beta$  on SOL and NSOL and by using EC data during the 2018 drought. It was a natural step for us to work together. This study, to which I participated as co-author for helping in designing the method (chapter 3), laid the foundations of the following chapters dedicated to potato (chapter 4 and chapter 5).

While providing large datasets of gas exchanges which can be used to calibrate the response of photosynthetic processes to water stress, the EC technique does not allow to fully unravel the physiological mechanisms involved in fluxes limitations. In particular, the response of  $V_{cmax,app}$  (i.e., considered as NSOL) can be driven by either the mesophyll conductance  $g_m$  or the true carboxylation rate  $V_{cmax}$ . Therefore, EC data only provides a limited understanding of the mechanisms regulating photosynthesis under water stress. Further information on photosynthesis constraints are provided by leaf-level measurements of ChlF, which allow to determine the mesophyll conductance, a critical ecophysiological parameter causing major restrictions in CO<sub>2</sub> diffusion to chloroplasts. While numerous studies have already highlighted that mesophyll conductance may decrease under water stress, strong uncertainties remain on the response of  $g_m$  to soil water availability, which is likely regulated by PFT and

specie-specific adaptive characteristics to drought. To reduce these uncertainties, we characterized the response of  $g_m$  and  $V_{cmax}$  to water stress by performing joint gas exchange and actively-induced chlorophyll fluorescence measurements at the leaf level during a drought experiment conducted on potato. In the perspective of improving crops yield, quantitatively assessing the most important constraints on photosynthesis under water stress is crucial. To that end, partitioning analysis can be used to decompose the relative variation of net photosynthesis into the sum of relative variations of SOL and NSOL (Grassi & Magnani, 2005). We aimed to modify this partitioning scheme by coupling the FvCB and USO models. The main advantage of this approach is to link the variability of  $A$  to  $g_m$ ,  $VPD$  and  $g_1$ , which therefore allows to identify the feedback effect of non-stomatal factors on stomatal closure. Such method would help in unravelling the role played by the mesophyll conductance in constraining stomatal conductance and photosynthesis responses to soil water availability.

Leaf-level ChlF measurements are however limited for estimating  $g_m$  of large patches of vegetation and do not allow to track sub-daily variations of NSOL by being time and labor-consuming. The recent emergence of SIF has allowed the collection of automatic measurements of ChlF at the ecosystem scale. SIF provides a proxy of vegetation physiological processes, which allows a direct estimation of photosynthesis using a small set of mechanistic equations if few key physiological parameters are known. The last objective of this thesis consisted in testing the capability of the MLR-USO model for estimating GPP and Tr under natural conditions, including soil water availability-limiting conditions. This approach will help in investigating the response of physiological processes to drought and improving the understanding of carbon and water fluxes temporal variability in the context of climate change. The first SIF measurements campaign started in 2022 at BE-Lon, which was the year of winter wheat cultivation. As a result, this thesis exclusively focused on winter wheat in chapter 6. Their importance for the agricultural production system in Belgium and Western Europe is outlined in chapter 4, chapter 5 and chapter 6.

The objectives of this thesis can be summarized as (Fig. 2.1):

**Objective 1:** Calibrating the response of stomatal and non-stomatal limitations on photosynthesis to soil water availability (chapters 3, 4 and 5).

**Objective 2:** Using a new partitioning approach to quantify the contributions of stomatal and non-stomatal limitations to the decrease in photosynthesis under water stress (chapter 5).

**Objective 3:** Testing a novel mechanistic approach to model photosynthesis and transpiration from sun-induced chlorophyll fluorescence at the ecosystem scale (chapter 6).

## 2. List of publications

- Gourlez de la Motte L., **Beauclaire, Q.**, Heinesch, B., Cuntz, M., Foltýnová, L., Šigut, L., Kowalska, N., Manca, G., Goded Ballarin, I., Vincke, C., Roland, M., Ibrom, A., Lousteau, D., Siebicke, L., Neiryink J., and Longdoz B., 2020. Non-stomatal processes reduce gross primary productivity in temperate forest ecosystems during severe edaphic drought. *Philosophical Transactions of the Royal Society B-Biological Sciences*.  
<https://doi.org/10.1098/rstb.2019.0527>
- **Beauclaire, Q.**, Heinesch, B., Longdoz, B., 2023. Non-stomatal processes are responsible for the decrease in gross primary production of a potato crop during edaphic drought. *Agricultural and Forest Meteorology*.  
<https://doi.org/10.1016/j.agrformet.2023.109782>
- **Beauclaire, Q.**, Vanden Brande F., Longdoz B. New method to partition the origins of photosynthesis limitations of potato under soil water availability-limiting conditions (*in prep*)
- **Beauclaire, Q.**, De Cannière, S., Jonard, F., Delhez, L., Pezzetti, N., Longdoz, B., 2023. Modeling gross primary production and transpiration from sun-induced chlorophyll fluorescence using a mechanistic light-response approach. *Remote Sensing of Environment*.  
<https://doi.org/10.1016/j.rse.2024.114150>

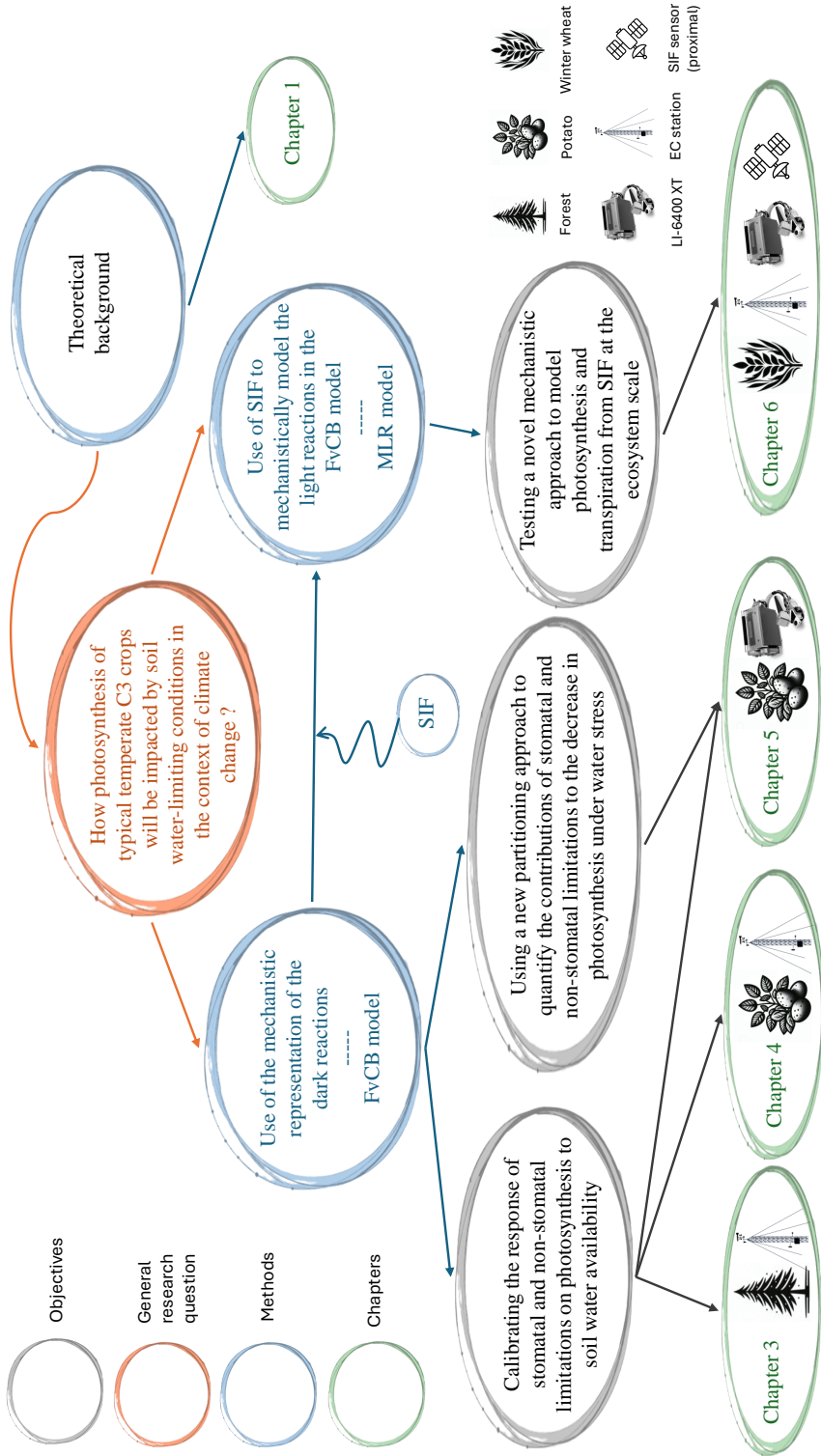


Figure 2.1: Flowchart of the thesis

**Non-stomatal processes reduce gross primary productivity in temperate forest ecosystems during severe edaphic drought**



This paper has been published as:

Gourlez de la Motte L., Beauclair, Q., Heinesch, B., Cuntz, M., Foltýnová, L., Šigut, L., Kowalska, N., Manca, G., Goded Ballarin, I., Vincke, C., Roland, M., Ibrom, A., Lousteau, D., Siebicke, L., Neiryink J., and Longdoz B., 2020. Non-stomatal processes reduce gross primary productivity in temperate forest ecosystems during severe edaphic drought. *Philosophical Transactions of the Royal Society B-Biological Sciences*.  
<https://doi.org/10.1098/rstb.2019.0527>

## Abstract

Severe drought events are known to cause important reductions of GPP in forest ecosystems. However, it is still unclear whether this reduction originates from SOL and/or NSOL. In this study, we investigated the impact of edaphic drought in 2018 on GPP and its origin (SOL, NSOL) using a dataset of 10 European forest ecosystem flux towers. In all stations where GPP reductions were observed during the drought, these were largely explained by declines in the maximum apparent canopy scale carboxylation rate  $V_{cmax,app}$  (NSOL) when the soil relative extractable water content dropped below around 0.4. Concurrently, we found that the stomatal slope parameter ( $G_1$ , related to SOL) of the Medlyn et al. unified optimization model linking vegetation conductance and GPP remained relatively constant. These results strengthen the increasing evidence that NSOL should be included in stomatal conductance/photosynthesis models to faithfully simulate both GPP and water fluxes in forest ecosystems during severe drought.

This article is part of the special issue ‘Impacts of the 2018 severe drought and heatwave in Europe: from site to continental scale’.

## 1. Introduction

With global climate change, droughts are likely to be more intense (Trenberth *et al.*, 2014; Vogel *et al.*, 2019). In 2018, a severe drought event occurred in Northern and central Europe causing forest fires and crop yield losses (Toreti *et al.*, 2019). Europe experienced a major reduction of GPP and transpiration (E) similarly to previous extreme events such as the 2003 Europe drought-heatwave (Ciais *et al.*, 2005) mostly because of soil water limitation (Granier *et al.*, 2007a; REICHSTEIN *et al.*, 2007). Continuous measurements of ecosystem CO<sub>2</sub> and water fluxes captured throughout Europe at EC flux tower stations thus provide a great large-scale ‘natural experiment’ to study the impact of drought on GPP and E (Sippel *et al.*, 2018).

There is increasing evidence that GPP reductions due to droughts could originate from both changes in stomatal behaviour (SOL) and non-stomatal traits (NSOL) (MIGLI-  
AVACCA *et al.*, 2011; Egea *et al.*, 2011; Zhou *et al.*, 2013; Xu & Baldocchi, 2003; reichstein *et al.*, 2002). Proposed NSOL mechanisms are reduced Rubisco activity

(carboxylation rate) and/or electron transport activity (Reichstein, 2003), reduced active Leaf Area Index (LAI) (Reichstein *et al.*, 2003), reduced mesophyll conductance (including the intercellular airspace, cell walls, plasma membranes, cytoplasm and the chloroplast envelopes (Knauer *et al.*, 2019a),  $g_m$ ) (Keenan *et al.*, 2010a) or a combination of those (Egea *et al.*, 2011; Zhou *et al.*, 2019b). The cause of GPP reduction is still subject to debate (Rogers *et al.*, 2017; Chaves *et al.*, 2009) and the modelling of SOL and NSOL and their (de)coupling is still poorly constrained by data. As a result, there is a strong need to examine whether different mechanisms are relevant and if models could be improved by developing more evidence-based functions for the impact of drought stress (Drake *et al.*, 2017; Medlyn *et al.*, 2016).

In leaf/canopy photosynthesis models, gross primary assimilation ( $A$ ) is very often modelled using the FvCB photosynthesis model for C3 species (Zhou *et al.*, 2013; Xu & Baldocchi, 2003; Reichstein, 2003). In this model, Rubisco-limited photosynthesis (usually close to light saturation) is a function of the maximum carboxylation rate ( $V_{cmax}$ ) and the internal CO<sub>2</sub> leaf concentration ( $C_i$ ) which implicitly considers that  $C_i$  is equal to the CO<sub>2</sub> concentration in the chloroplasts ( $C_c$ ). As  $C_i$  cannot be measured directly, it is usually approximated by employing Fick's diffusion law through the stomata using a stomatal conductance ( $g_s$ ). This representation requires the determination of stomatal conductance by modelling. In this study, following (Zhou *et al.*, 2013), we use the concept of apparent  $V_{cmax}$  ( $V_{cmax,app}$ ) recognizing that variations in  $V_{cmax,app}$  can result either from changes in the actual maximum rate of carboxylation or from changes in  $g_m$  which are not explicitly represented in this diffusion model. Consequently, when drought occurs, it impacts directly stomatal behaviour (closure) and then photosynthesis by limiting the diffusion of CO<sub>2</sub> into the leaf which results in reduced  $C_i$  (SOL) or/and it impacts non-stomatal mechanisms (NSOL) which result in decreases of  $V_{cmax,app}$  (Zhou *et al.*, 2013; Drake *et al.*, 2017).

A long standing stomatal conductance model from Cowan & Farquhar (1977) states that stomata should act to maximize carbon gains while minimizing water losses (transpiration,  $E$ ), that is to maximize the integrated sum of  $A-E$  where  $\lambda$  (mol C mol<sup>-1</sup> H<sub>2</sub>O) is the carbon cost of water gain  $\rho A/\rho E$  or marginal water-use efficiency (Manzoni *et al.*, 2011) (note that we inverted the original expression). Medlyn *et al.* (2011a) proposed a reconciliation of the optimal stomatal behaviour theory (Cowan & Farquhar, 1977) with empirical stomatal models linking  $g_s$  and  $A$ . Their work resulted in a USO model with a form similar to former empirical expressions (see equation (2.3)) where the slope between  $g_s$  and  $A * f(g_1)$  is a key parameter (called the stomatal slope parameter).  $g_1$  is directly interpretable as inversely related to  $\lambda$  and to intrinsic WUE ( $A/g_s$ ) normalized by VPD and CO<sub>2</sub> air concentration ( $C_a$ ) (Knauer *et al.*, 2018b).

The USO model has been used both at the leaf level using leaf gas exchange data (Lin *et al.*, 2015b) and at the ecosystem level using EC flux observations (Medlyn *et al.*, 2017a) during non-water limited periods. During water limited periods, various



responses of  $g_1$  (leaf level, SOL) to soil moisture were found (Zhou *et al.*, 2013) for a large range of species while a more consistent pattern of decreasing  $V_{cmax,app}$  was found. In a recent work, a good correlation between leaf scale and ecosystem scale  $g_1$  (or  $G_1$ ) response to soil moisture was found in a woodland dominated by Acacia trees thereby demonstrating the ability of both leaf and ecosystem scale approaches to quantify drought effect (Tarin *et al.*, 2020b).

In this study, we used the USO model combined with the Farquhar C3 model (considering that  $C_i = C_c$ ) to study the origin of edaphic drought impacts on GPP (SOL and/or NSOL) in forest ecosystems using EC flux measurements by replacing leaf-level variables by their ecosystems analogues using a big leaf framework (Knauer *et al.*, 2018b,a). The surface conductance ( $G_s$  analogous to  $g_s$ ) was estimated by inverting the Penman–Monteith equation (Monteith, 1965). We then inferred the bulk stomatal slope parameter ( $G_1$  analogous to  $g_1$ ) and the maximum apparent carboxylation rate of the ecosystem ( $V_{cmax,app}$ ) (Kosugi *et al.*, 2013) at a daily time step for each ecosystem. The study was restricted to the growing period excluding any autumn senescence or spring leaf emergence influence on the variation of  $V_{cmax,app}$ .

In addition, drought intensity was quantified using the REW as proposed by Granier *et al.* (1999a), which is a normalized index of soil water deficit varying from 0 to 1 that allows for edaphic status inter-site comparisons. This index was used in previous studies (Granier *et al.*, 1999a, 2007a) and, based on their results, we hypothesize that both E and GPP reductions will occur when REW falls below  $\sim 0.4$ .

The objective of this work is to examine the response of  $G_1$ , as a measure for SOL, and  $V_{cmax,app}$ , as a measure for NSOL, to soil water deficit using EC data collected in forests during the 2018 European drought. More specifically, we intend to answer the following questions: (1) how was REW impacted by the drought in forest sites in 2018? (2) Can we confirm the REW threshold of  $\sim 0.4$  for GPP reductions found in previous studies (Granier *et al.*, 1999a, 2007a)? (3) To what degree did SOL and NSOL impact GPP during the drought? (4) What were  $G_1$  and  $V_{cmax,app}$  responses to REW function shapes and how did these responses vary across sites?

## 2. Material and methods

### 2.1. Site and data description

Data have been processed by the Ecosystem Thematic Centre of the Integrated Carbon Observation System (ICOS) and form the 2018 drought ICOS/Fluxnet dataset (Drought 2018 Team & ICOS Ecosystem Thematic Centre, 2019), which is a compilation of EC fluxes, meteorological and edaphic data during the 2018 European drought at half-hourly resolution. Only sites with a sufficiently resolved vertical profile of soil water content sensors were selected. The main site characteristics are summarized in table 3.1. Flux data followed the standard FLUXNET processing (Aubinet *et al.*,

1999), including friction velocity ( $u^*$ ) filtering (Papale *et al.*, 2006) and GPP determination by night-time flux partitioning (Reichstein *et al.*, 2005). Only data marked with highest quality flags were used for this study. Latent heat fluxes were not corrected for energy balance closure.

**Table 3.1:** Main characteristics of the flux tower sites included in this study. The LAI corresponds to the maximum LAI typically observed at the sites. Please refer to the online version of the article for references.

site ID	country	latitude	longitude	dominating species	LAI m <sup>2</sup> m <sup>-2</sup>	soil texture	rooting depth m	ref
BE-Bra	Belgium	51.308	4.52	<i>Pinus sylvestris</i>	3	sand	1	[36]
BE-Vie	Belgium	50.305	5.998	<i>Fagus sylvatica</i> / <i>Pseudotsuga menziesii</i>	5	silty clay loam	1.4	[37]
CZ-Lnz	Czechia	48.682	16.946	<i>Quercus robur</i> / <i>Fraxinus angustifolia</i> / <i>Carpinus betulus</i> / <i>Tilia cordata</i>	6.5	sandy loam	1.2	[38]
CZ-Raj	Czechia	49.444	16.697	<i>Picea abies</i>	5	sandy loam	0.7	[39]
CZ-Stn	Czechia	49.036	17.97	<i>Fagus sylvatica</i>	5.5	sandy loam	0.7	[40]
DE-Hai	Germany	51.0792	10.453	<i>Fagus sylvatica</i>	6	clay loam	0.7	[41,42]
DK-Sor	Denmark	55.486	11.645	<i>Fagus sylvatica</i>	5	sandy clay loam	1	[43]
FR-Bil	France	44.494	-0.956	<i>Pinus pinaster</i>	2.5	sand	1.1	[44]
FR-Hes	France	48.674	7.065	<i>Fagus sylvatica</i>	6.5	silty clay loam	1.6	[45]
IT-Sr2	Italy	43.732	10.291	<i>Pinus pinea</i>	2.5	sand	1.2	

## 2.2. Quantification of drought

The intensity of edaphic drought was quantified by computing the REW content (Granier *et al.*, 2007a) at each time step and for the entire root depth using:

$$REW = \frac{\sum_i (SWC_i - SWC_{WP,i}) / (SWC_{FC,i} - SWC_{WP,i})}{h_{max}} \quad (3.1)$$

where  $i$  is the index of each soil layer over the rooting depth,  $SWC_i$  is the actual SWC,  $SWC_{WP}$  is the SWC at the wilting point,  $SWC_{FC}$  is the SWC at field capacity,  $\Delta h$  is the thickness of each layer and  $h_{max}$  is the maximum rooting depth. Each soil horizon was divided into soil layers corresponding to the number of sensors installed in the horizon. The layer boundaries were the horizon limits or the midway point between two sensors. Soils related data are summarized in the supplements (electronic supplementary material, table S1). For each layer,  $SWC_{WP}$  and  $SWC_{FC}$  were estimated using soil retention curves based on either measurements (by research teams) or modelling (based on soil textures) and checked for consistency with SWC 48 h after a rain event for  $SWC_{FC}$  and with minimum SWC values observed at the site for  $SWC_{WP}$  to avoid negative REW values. When not available, the maximum depth was defined as the bedrock depth (Granier *et al.*, 2007a). When data were available

(BE-Vie and BE-Bra), REW was corrected for the coarse fraction by applying a correction factor for each layer. According to Granier *et al.* (2007a, 1999a) it is expected that both GPP and  $G_s$  start to decrease when REW drops below 0.4. The evolution of REW at each site in 2018 is presented in the electronic supplementary material, figure S1.

### 2.3. Canopy surface variables

Detailed computation procedures for canopy surface variables are fully described in Knauer *et al.* (2018a). First the aerodynamic conductance to water transfer ( $G_{aw}$ ) was computed as a combination of an aerodynamic conductance to momentum (first term) and a boundary layer conductance (second term) as follows:

$$G_{aw} = \left( \frac{u_*^2}{u(z)} + (6.2u_*^{-0.667})^{-1} \right)^{-1} \quad (3.2)$$

where  $u_*$  is the friction velocity ( $\text{m s}^{-1}$ ), and  $u(z)$  the wind speed at measurement height ( $z$ ). Canopy surface conductance for water ( $G_s$ ,  $\text{m s}^{-1}$ ) was computed by inverting the Penman–Monteith equation (Monteith, 1965):

$$G_s = \frac{G_{aw}\gamma LE}{\Delta(R_n - G - S) + \rho_a c_p G_{ah} VPD_s - (\Delta + \gamma)LE} \quad (3.3)$$

where  $LE$  is the latent heat flux ( $\text{W m}^{-2}$ ),  $\gamma$  is the psychrometric constant ( $\text{Pa K}^{-1}$ ),  $s$  is the slope of the saturation vapour pressure curve at air temperature ( $\text{Pa K}^{-1}$ ),  $R_n$  is the net radiation ( $\text{W m}^{-2}$ ),  $G$  is the ground heat flux ( $\text{W m}^{-2}$ ),  $S$  is the sum of all storage terms ( $\text{W m}^{-2}$ ),  $C_p$  is the heat capacity of dry air ( $1005 \text{ J kg}^{-1} \text{ K}^{-1}$ ) and  $VPD_a$  is the VPD of ambient air (Pa).  $G$  was considered negligible when not available while  $S$  was not available and was set to 0 at all sites. The  $\text{CO}_2$  concentration at the canopy surface ( $C_s$ ), needed in the USO and diffusion equations, was computed as:

$$C_s = C_a + \frac{NEE}{G_{aw}/1.32} \quad (3.4)$$

where  $C_a$  ( $\mu\text{mol CO}_2 \text{ mol}^{-1}$ ) is the  $\text{CO}_2$  air concentration at the measurement height,  $NEE$  ( $\mu\text{mol CO}_2 \text{ m}^{-2}$ ) is the net  $\text{CO}_2$  ecosystem exchange and the factor 1.32 is the ratio of diffusivities of  $\text{CO}_2$  and water vapour in the boundary layer. The VPD at the canopy surface ( $VPD_s$ , Pa) was also computed (see Knauer *et al.* (2018a) for more details).  $G_s$  is a good predictor of bulk stomatal conductance only when evaporation is small compared to transpiration; data collected during a period of 48 h following a rain event were discarded. Secondly, the analysis was restricted to the growing season, avoiding senescence and leaf emergence periods. We defined this period as the days when the daily GPP (averaged over all the available years) smoothed with a 15 days

moving average window was higher than 70% of the 95th percentile of the daily GPP distribution.  $G_s$  data were also filtered excluding half hour with  $LE < 0$  or  $R_n < 0$ . Negative  $G_s$  values were filtered and  $G_s$  outliers were also discarded by removing data when  $G_s$  were higher than the 98<sup>th</sup> percentile of the  $G_s$  distribution.

## 2.4. Stomatal origin limitation

Similarly to previous work (Zhou *et al.*, 2013; H eroult *et al.*, 2013), reductions of GPP originating from SOL were assessed by analysing dependence on REW of the  $G_1$  parameter used in the USO model developed by Medlyn *et al.* (2011a) but adapted to the ecosystem scale using bulk ecosystem parameters (Knauer *et al.*, 2018a):

$$G_s = G_0 + 1.6 \left( 1 + \frac{G_1}{\sqrt{VPD_s}} \right) \frac{GPP_{high}}{C_s} \quad (3.5)$$

where,  $GPP_{high}$  is the GPP at high radiation ( $R_g > 500 \text{ W m}^{-2}$ ) and replaces net assimilation in the original leaf scale expression of the model,  $G_s$  replaces stomatal conductance and leaf surface variables were replaced by their corresponding canopy surface values (Wang *et al.*, 2009) ( $C_s$  the air  $\text{CO}_2$ ,  $VPD_s$ ). In this expression, the nocturnal stomatal conductance  $G_0$  was set to 0 as its magnitude can be considered negligible when compared to the other terms at saturating daylight conditions [27].  $G_1$ , the slope parameter, is a physiologically meaningful parameter as it was shown to be inversely related to  $\lambda$  (Medlyn *et al.*, 2011a) and to  $i\text{WUE}$  (Knauer *et al.*, 2018a).

$G_1$  was obtained by inverting equation (3.5) using half-hourly measurements. Because leaf respiration was neglected in equation (3.5), the equation was inverted only for high radiation data so that  $GPP_{high}$  should be much higher than leaf respiration. Negative  $G_1$  values were filtered and outliers were also discarded by removing data when absolute  $G_1$  was two times higher than the average absolute deviation from the median. Finally, daily  $G_1$  averages were then computed for days with at least five valid half-hourly values.

The response of  $G_1$  to REW was fitted with a segmented linear response curve (two segments) in order to test the presence of an REW threshold (break point) above which  $G_1$  is constant (no effect range) and under which stomatal regulation occurs ( $G_1$  decreases or increases) (Oosterbaan *et al.*, 1990). After a first fit, outliers were removed by exclusion of the  $G_1$  value having absolute residuals more than 2.2 times the standard deviation of the residuals distribution. A second fit was then done. Parameters obtained from a second fit were  $G_1^*$ , i.e. the average  $G_1$  value within the ‘no effect range’, the break point REW value ( $REW_{B,G_1}$ ) and the slope/intercept of the  $G_1$  decrease/increase. The presence of the break point was further tested by comparing the residuals of the model to those of a simple linear regression model using an F-test.

## 2.5. Non-stomatal origin limitation

Reductions of GPP originating from NSOL were studied by assessing the effect of water stress on apparent bulk  $V_{cmax}$  hereafter called  $V_{cmax,app}$ . It was obtained by inverting the expression of Rubisco-limited photosynthesis during high radiation conditions (Farquhar *et al.*, 1980):

$$V_{cmax,app} = \frac{GPP_{high} (C_i + K_m)}{(C_i - \Gamma^*)} \quad (3.6)$$

where  $V_{cmax,app}$  is expressed per  $m^2$  of soil and not of leaf as usual.  $K_m$  is the effective Michaelis Menten coefficient kinetics and  $\Gamma^*$  is the  $CO_2$  compensation point, which were both computed using temperature responses following (Bernacchi *et al.*, 2001),  $GPP_{high}$  is the GPP at  $R_g > 500 \text{ W m}^{-2}$  while  $C_i$  was computed using Fick's diffusion law (Farquhar & Sharkey, 1982):

$$C_i = C_s - \frac{GPP_{high}}{G_s/1.6} \quad (3.7)$$

Note that in equations (3.3)-(3.7), leaf respiration was neglected which should have a small effect on the results as, at high radiation, leaf respiration should be much smaller than GPP. Half-hourly values of  $V_{cmax,app}$  were then normalized for temperature to  $25^\circ\text{C}$  using an Arrhenius equation (Knauer *et al.*, 2018a) fitted for each decile of REW as  $V_{cmax,app}$  response to temperature was found to decrease under drought conditions. Finally,  $V_{cmax,app}$  was averaged on a daily basis and days with less than 5 half-hourly values were discarded. Considering the way we estimated  $V_{cmax,app}$ , a decrease in  $V_{cmax,app}$  indicates NSOL of GPP including either changes in mesophyll conductance or in actual  $V_{cmax}$  or other processes limiting GPP (apart from stomatal closure). The response of  $V_{cmax,app}$  to REW was assessed using the same segmented linear regression model as explained in the previous section.  $V_{cmax,app}^*$  and  $REW_{B,V_{cmax}}$  were defined as the average of  $V_{cmax,app}$  values (normalized at  $25^\circ\text{C}$ ) within the no effect range and the break point REW value, respectively.

## 2.6. Degree of stomatal and non stomatal origin limitation

To illustrate the degree of SOL and NSOL due to edaphic drought, following Zhou *et al.* (2013), a sensitivity analysis was performed to explore the impact of drought-induced changes in  $G_1$  and  $V_{cmax,app}$  on GPP. The impact of NSOL was assessed by comparing measured GPP to a theoretical non-affected value corresponding to a 'modelled' GPP computed using inverted equation (3.5) with  $V_{cmax,app}^*$  instead of  $V_{cmax,app}$  and a  $C_i$  obtained from equation (3.6) using observed  $G_s$  values (equation (3.3)) and measured GPP values. Similarly, the impact of SOL was assessed by comparing measured GPP to 'modelled' unaffected GPP computed using constant  $G_1^*$  values and ob-

served  $V_{cmax,app}$  values. The degree of limitation (DoL) was computed as the ratio of modelled ‘unaffected’ GPP against measured GPP and represents the factor by which GPP was divided because of SOL or NSOL.

### 3. Results

#### 3.1. $G_s$ and $GPP_{high}$

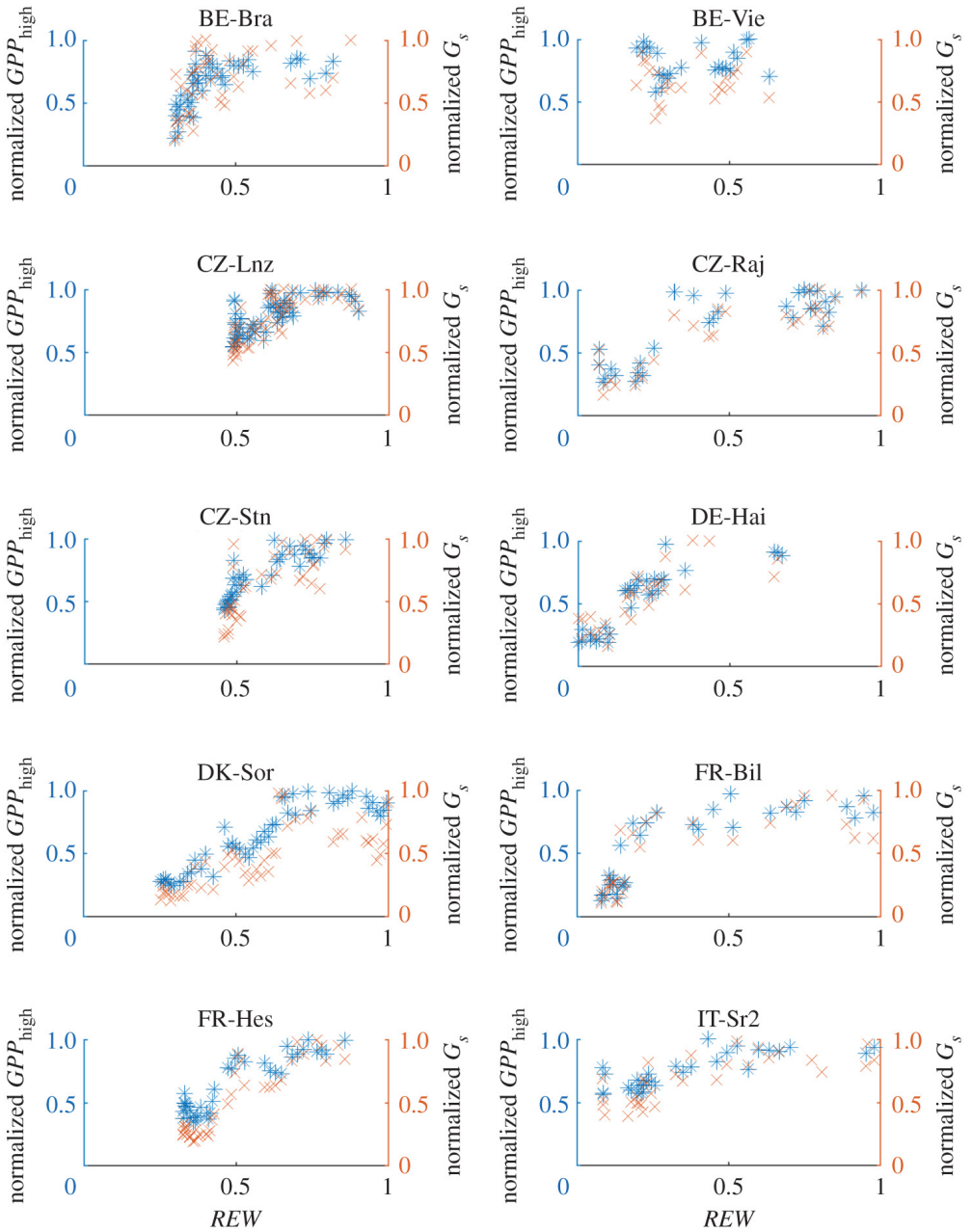
$GPP_{high}$  and  $G_s$  (for  $R_g > 500$ ) normalized by their respective maximum values in relation with REW are presented in figure 1. At all sites, we can observe that both  $GPP_{high}$  and  $G_s$  behave similarly. High values of both  $GPP_{high}$  and  $G_s$  were observed for high REW while both variables decreased simultaneously with REW. There was an exception at BE-Vie where such a pattern was not observed although both variables still behaved similarly. The lowest  $GPP_{high}$  and  $G_s$  values were observed at sites such as CZ-Raj, FR-Bil and DE-Hai where very low REW values (lower than 0.15) were reached. At IT-Sr2,  $G_s$  and  $GPP_{high}$  were still quite high (around half of maximum values) even for very low REW values most probably because, in this sandy soil, rooting depth was probably deeper than the deepest available SWC sensor (1.2 m, see electronic supplementary material, table S1) which caused an underestimation of REW.

#### 3.2. Response of $G_1$ to edaphic drought

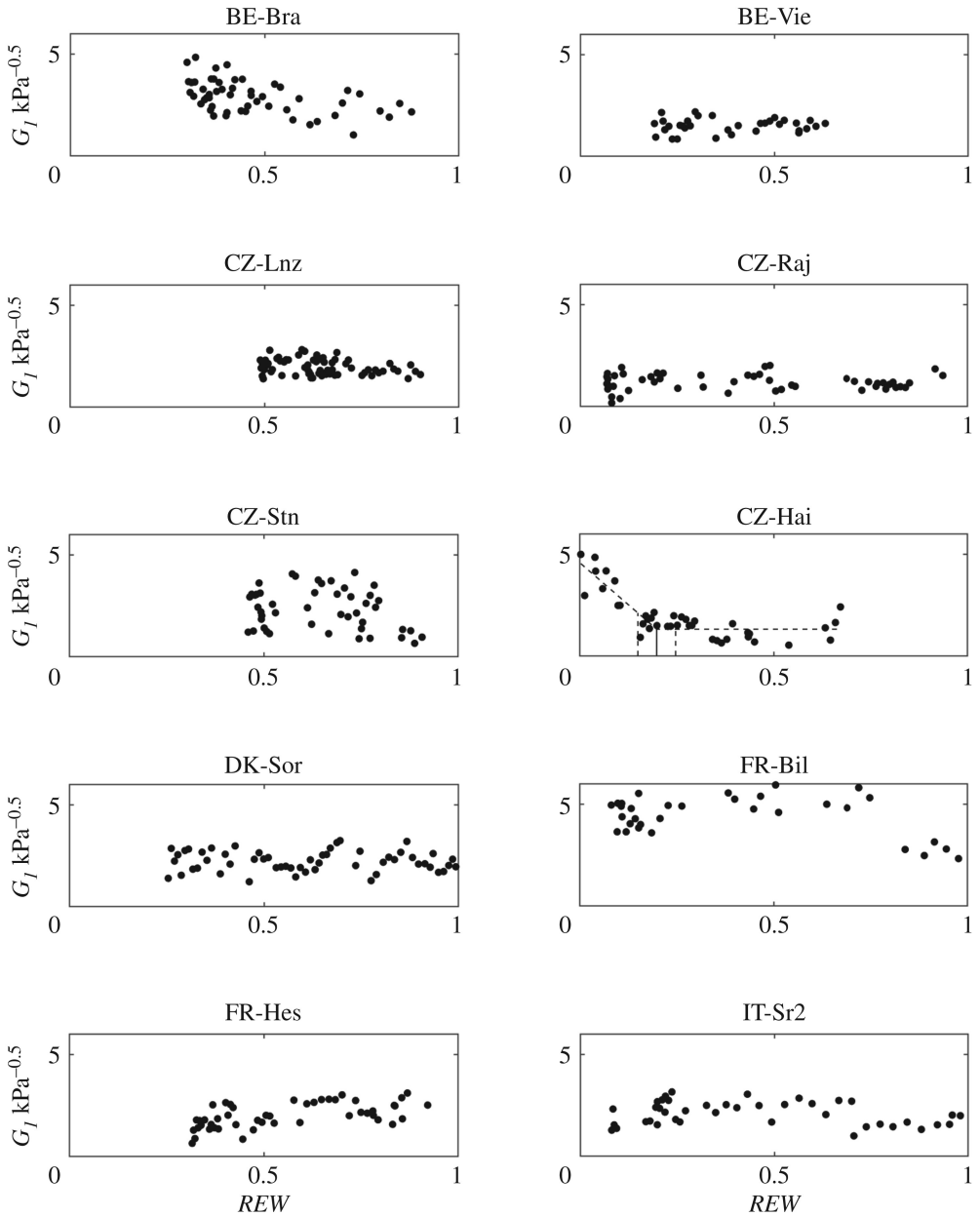
$G_1$  was found to be constant at all sites apart from DE-Hai even for sites where REW values lower than 0.4 were observed (figure 2). In BE-Bra  $G_1$  seems to be enhanced at low REW but the segmented model did not perform significantly better than the linear one. In DE-Hai,  $G_1$  was found to increase when REW dropped below a very low value of 0.2, which is quite close to the wilting point. Such low REW were also observed during the growing season at CZ-Raj and IT-Sr2 but no similar behaviour was observed. The lowest  $G_1^*$  (1.5 kPa<sup>0.5</sup> at CZ-Raj) was three times lower than the highest value (4.5 kPa<sup>0.5</sup> at FR-Bil, table 2).

#### 3.3. Response of $V_{cmax,app}$ to edaphic drought

The effect of NSOL caused by drought was studied by analysing the dependence of the temperature normalized  $V_{cmax,app}$  values on REW (figure 3). At all sites that experienced low REW conditions (below  $\sim 0.4$ ) apart from BE-Vie and IT-Sr2, constant  $V_{cmax,app}$  were observed for large REW values followed by a decrease when REW declined below an  $REW_{B,Vcmax}$  threshold. The  $REW_{B,Vcmax}$  were not significantly different (according to the confidence intervals) than the value of 0.4 which was found in previous studies, with an exception at DK-Sor where  $REW_{B,Vcmax}$  was higher ( $REW_{B,Vcmax} = 0.8 \pm 0.09$ , table 2). The high  $REW_{B,Vcmax}$  observed at DK-Sor



**Figure 3.1:** Dependence of  $GPP_{high}$  (blue, left axis) and  $G_s$  at high radiation ( $R_g > 500$ , red, right axis) normalized by their maximum value on REW for each site.



**Figure 3.2:** Dependence of  $G_1$  on REW for each site.  $REW_{B,G_1}$  are marked by a vertical solid line with 95% confidence intervals (dashed vertical lines). Regression lines are only shown when F-test p-values (comparison with linear regression) were smaller than 0.05.



**Table 3.2:** Maximum extractable water (EW), minimum observed REW in 2018 during the growing season,  $REW_{B, V_{cmax,app}}$  and  $REW_{B, G_1}$  (REW break points for VCMAX, APP and G1, respectively) given with 95% confidence intervals.  $V_{cmax,app}^*$  and  $G_1^*$  ( $V_{cmax,app}$  and  $G_1$  values in unstressed conditions) given with 95% confidence intervals. P-values are given for the F-test comparing the segmented mode (three parameters) to the linear model (two parameters).

site ID	EW mm	min REW —	$V_{cmax,app}^*$ $\mu\text{mol m}^{-2} \text{s}^{-1}$	$REW_{B, V_{cmax}}$ —	p-value —	$G_1^*$ $\text{kPa}^{-0.5}$	$REW_{B, G_1}$ —	p-value —
BE-Bra	133	0.30	$58 \pm 3$	$0.57 \pm 0.19$	<0.05	$2.9 \pm 0.2$	—	0.28
BE-Vie	215	0.19	$81 \pm 5$	—	1	$1.5 \pm 0.1$	—	1
CZ-Lnz	241	0.49	$116 \pm 10$	—	1	$2.0 \pm 0.2$	—	1
CZ-Raj	92	0.07	$80 \pm 7$	$0.36 \pm 0.11$	<0.001	$1.2 \pm 0.8$	—	1
CZ-Stn	236	0.46	$120 \pm 7$	—	1	$2.3 \pm 0.3$	—	1
DE-Hai	143	0.00	$86 \pm 17$	$0.36 \pm 0.07$	<0.001	$1.3 \pm 0.2$	$0.2 \pm 0.05$	<0.001
DK-Sor	176	0.25	$104 \pm 3$	$0.80 \pm 0.09$	<0.001	$2.2 \pm 0.2$	—	1
FR-Bil	159	0.08	$76 \pm 7$	$0.45 \pm 0.15$	<0.05	$4.5 \pm 0.4$	—	1
FR-Hes	338	0.33	$121 \pm 13$	$0.44 \pm 0.08$	<0.05	$2.0 \pm 0.5$	—	1
IT-Sr2	108	0.08	$89 \pm 6$	—	1	$2.1 \pm 0.2$	—	1

might result from an overestimation of REW as the shallowest available SWC probe was at 15 cm depth (see electronic supplementary material, table S1) and was not able to catch the beginning of the progressive drying of the upper layers that contain a large amount of roots. The most impacted site was DE-Hai where REW almost reached the wilting point ( $REW \sim 0$ ) with very low  $V_{cmax,app}$  values ( $\sim 15 \mu\text{mol m}^{-2} \text{s}^{-1}$ ) probably because of shallow soil and rooting depth (0.6 m).

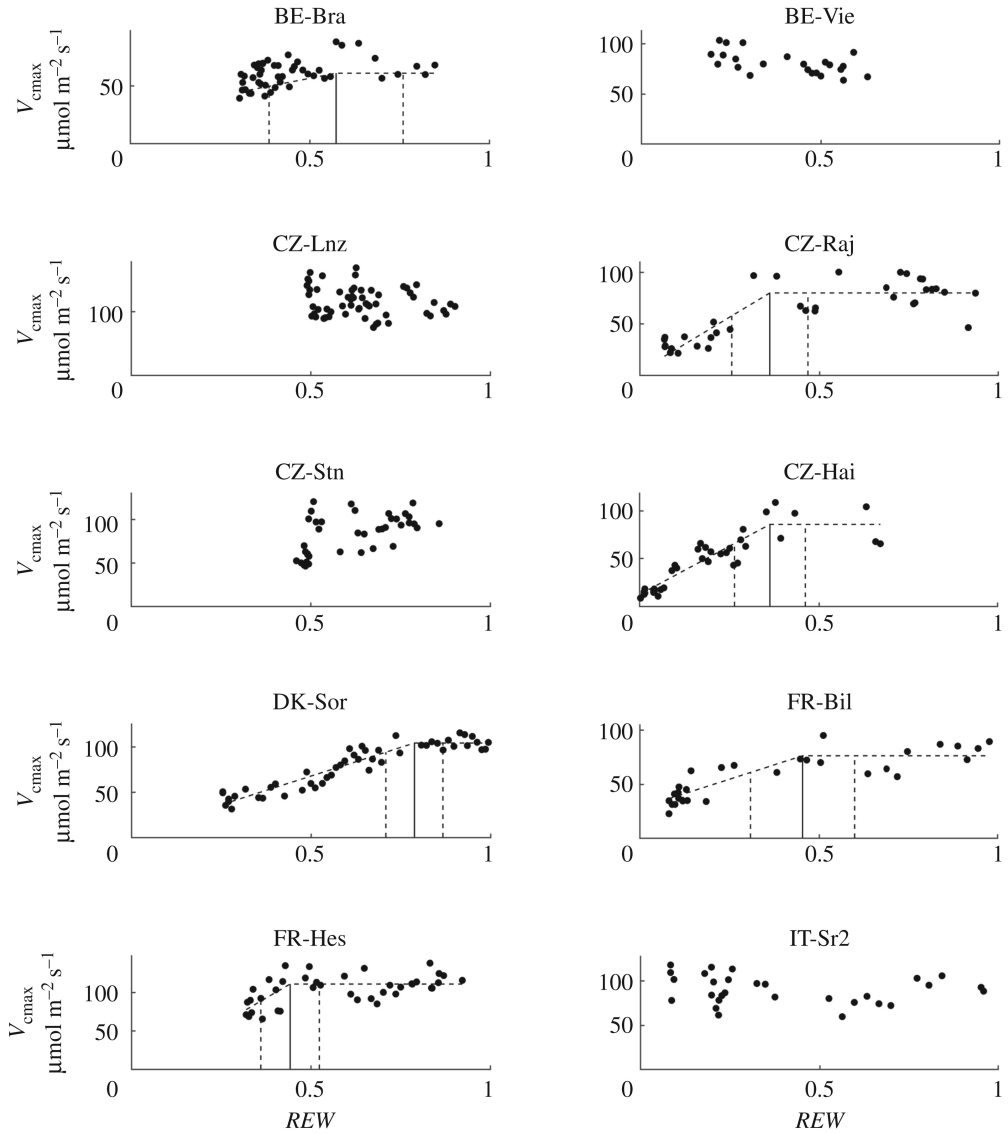
### 3.4. Degree of stomatal and non-stomatal limitation

DoL reached values of 5 for NSOL at DE-Hai while it remained close to 1 for SOL at all sites (figure 4). This analysis therefore confirms that, at all sites, NSOL was the dominant mechanism. As a result, reducing  $V_{cmax,app}$  while maintaining  $G_1$  constant could capture the variations of both  $GPP_{high}$  and  $G_s$  with drought (identical conclusions are obtained if we focus the analysis on  $G_s$  instead of GPP, data not shown). It is also worthwhile noticing that the increasing  $G_1$  observed at DE-Hai (and at BE-Bra to a lesser extent) did not lead to important changes in GPP.

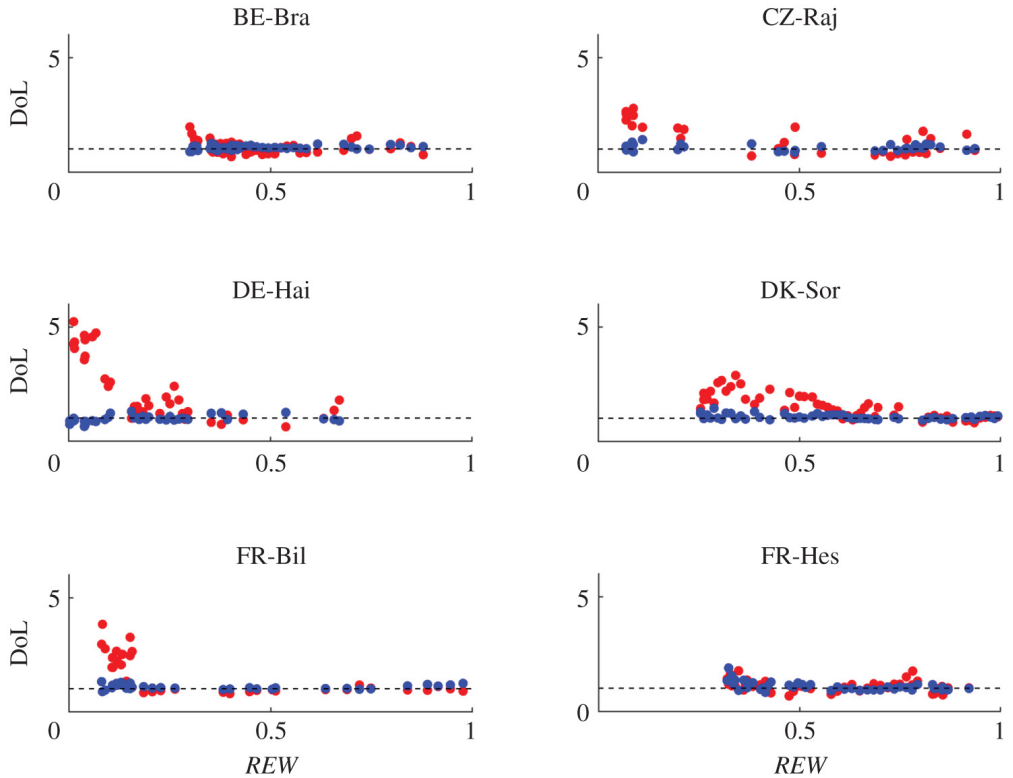
## 4. Discussion

### 4.1. Methodological limitations

Although the responses of  $V_{cmax,app}$  and  $G_1$  to REW were relatively consistent, some sites showed unexpected behaviours. For example,  $REW_{B, V_{cmax,app}}$  at DK-Sor



**Figure 3.3:** Dependence of  $V_{cmax,app}$  normalized at 25°C on REW for each site.  $REW_{B,V_{cmax}}$  are marked by a vertical solid line with 95% confidence intervals (dashed vertical lines). Regression lines are only shown when F-test p-values (comparison with linear regression) were smaller than 0.05.



**Figure 3.4:** Degree of limitation (DoL) by both SOL and NSOL. The degree of limitation was computed as the ratio of modelled GPP on measured GPP at high radiation. Modelled GPP was computed by using either fixed  $V_{cmax,app}^*$  (red points, NSOL) and varying  $G_1$  or fixed  $G_1^*$  and observed  $V_{cmax,app}$  (blue points, SOL).

was much higher than expected (0.85) because some SWC sensors experienced failures during the drought. At IT-Sr2, no limitation of GPP was found although very low REW values were estimated, probably because the SWC sensor profile was not deep enough to capture the whole rooting depth. Multiple and deeper sensor profiles (with matching wilting points and field capacities) would certainly help to reduce these uncertainties. Complementary measurements such as predawn leaf water potential and soil matric potential (Rogers *et al.*, 2017; Anderegg *et al.*, 2017), when REW approaches values close to 0.4 at the site, would also be useful.

The big leaf approach used in this study also has several limitations (Knauer *et al.*, 2018a) which could be critical when comparing leaf scale-derived parameters to big leaf canopy scale estimates (Knauer *et al.*, 2018a,b; Medlyn *et al.*, 2017b) or when attributing a behaviour to a specific species. First, the approach is only able to derive bulk parameters and is unable to distinguish the vertical and horizontal distribution of the properties. Horizontal heterogeneity is especially crucial at mixed forest sites where different species could show different responses to drought (and different root depth and therefore REW) which would blur their respective responses in the measured signal. This is especially critical at BE-Vie where the two most frequent wind directions (southwest and northeast) correspond to different stands (coniferous and beech stands) with possibly different root depth, REW and weather conditions (Aubinet *et al.*, 2018). Separating the data between each sector did not however improve the relation because of the lack of data (data not shown).

At sites with dense canopy and high LAI, vertical gradients of the parameters ( $G_s$ ,  $G_1$  and  $V_{cmax,app}$ ) could result from vertical gradients within plants of the same species or from physiological differences across species (Knauer *et al.*, 2018a). Sun leaves, developed under high irradiance, usually exhibit higher iWUE (lower  $G_1$ ) than those developed under shady conditions (Sellin *et al.*, 2010), primarily because of higher photosynthetic capacities. However, more critical for this study, little is known about to what degree these vertical gradients (within and across species) could affect the response of  $G_1$  and  $V_{cmax,app}$  to drought. To our knowledge, in most Earth system models, the same reduction functions of photosynthesis during edaphic drought (either NSOL or SOL) are used for sun and shade leaves (Rogers *et al.*, 2017). More complex multiple layers and/or sun-shade models as well as additional data gathered at multiple canopy layers would be needed to assess this question more closely.

Moreover, soil and vegetation components cannot be distinguished so that critical variables such as  $G_s$  (and variables depending on it, such as  $G_1$  and  $C_i$ ) will inevitably contain some signal from the soil. This signal can be reduced by filtering the data after rain events (Knauer *et al.*, 2018a) and should be small for dense canopies with LAI higher than 2–3 (Kelliher *et al.*, 1995) (which is the case in all sites) and even smaller when the upper soil layer dries.

Finally, systematic errors (energy balance non-closure Wilson *et al.* (2002)) in EC

fluxes are also major sources of uncertainties that affect  $G_1$  and  $V_{cmax,app}$  magnitudes. However, it was found that, at multiple flux tower sites, the surface energy balance was not modified during the 2018 drought (Graf *et al.*, 2020). This source of error is therefore unlikely to affect  $G_1$  and  $V_{cmax,app}$  responses to REW.

Nevertheless, despite all the limitations of the big leaf approach detailed above, this framework was very suitable for this multi-site study as it relied on very few ancillary data (Knauer *et al.*, 2018a). If the comparison of  $G_1$  and  $V_{cmax,app}$  (which inherently has a different meaning than leaf-level  $V_{cmax,app}$ ) is not straightforward, analysing the dynamics of these parameters inferred from in situ EC data during drought provides very useful information about how forest ecosystems reacted to these events.

## 4.2. Implications of non-stomatal origin limitation for the modelling of gross primary productivity and transpiration

In this study, similarly to Granier *et al.* (2007a), we found that GPP and  $G_s$  reductions can be expected when REW drops below  $\sim 0.4$ . To account for these reductions, using empirical reduction factors (ranging from 1 to 0) when soil water content falls below a given threshold is a widely used approach (Rogers *et al.*, 2017). However, it is questionable whether the reduction factors should be applied to SOL and/or to NSOL. This was previously investigated in Mediterranean ecosystems (reichstein *et al.*, 2002; Reichstein *et al.*, 2003; Keenan *et al.*, 2009) and it was found that, calibrating the model on either GPP or transpiration (E) (not both) by considering only SOL during edaphic drought conditions systematically led to overestimates of WUE which did not allow the correct simulation of both fluxes. Surprisingly, it was found that applying NSOL only was sufficient to correctly simulate both GPP and E.

In this study, we found that reducing  $V_{cmax,app}$  when REW dropped below  $\sim 0.4$  and using a constant  $G_1$  parameter (from the USO model (Medlyn *et al.*, 2011a)) allowed the capture of both GPP and  $G_s$  reductions at European forest sites. Similar conclusions were also found by Chen *et al.* (2019a) in four different ecosystems (temperate grassland, tropical savannah, boreal and one temperate forest). More specifically, relatively consistent behaviour was observed at the three beech (*Fagus sylvatica*) forest sites (FR-Hes, DK-Sor and DE-Hai) where NSOL were the main source of photosynthesis reductions with relatively constant  $G_1$ . Similarly, in a study carried out on adult beech using leaf-level measurements, the Ball–Berry slope was found to be almost insensitive to soil water potential (Op de Beeck *et al.*, 2010). Our results are in agreement with Granier *et al.* (2007a) who observed constant WUE even for very low REW during the 2003 drought and with Hentschel *et al.* (2016) who also found unchanged annual WUE derived from tree ring carbon isotopic composition.

Studies were also performed at the leaf scale to study the impact of drought on NSOL and SOL. In their meta-analysis, Zhou *et al.* (2013) found highly variable responses of  $g_1$  (leaf level) for woody species ranging from rather constant to severely decreasing

$g_1$  with drought. Decreasing  $V_{cmax,app}$  (NSOL) were however found for all species. It was also highlighted that NSOL was the main factor limiting photosynthesis under severe stress in 10 Mediterranean herbs and shrubs species (Galmés *et al.*, 2007; Egea *et al.*, 2011) and, more importantly for this study, for four tree species (Drake *et al.*, 2017). Unfortunately, a direct comparison with our results cannot be carried out as the studied species were different. Such direct comparisons have been carried out in a woodland dominated by Acacia trees by Tarin *et al.* (2020a) who found a close agreement between  $G_1$  and  $g_1$  estimated from ecosystem (EC big leaf) and leaf-level approaches, respectively.

According to Keenan *et al.* (2010b), NSOL could be caused by the variations of a finite mesophyll conductance with soil water availability (Flexas *et al.*, 2004) which, if not taken into account, leads to wrong estimates of actual  $V_{cmax}$  (Flexas *et al.*, 2012) (which was implicitly taken into account by using  $V_{cmax,app}$ ). In addition, the hypothesis that, under severe droughts, GPP can be directly impacted by biochemical limitations which cause the reduction of actual  $V_{cmax}$  should not be discarded (Lawlor & Tezara, 2009). Separating NSOL between these two mechanisms (mesophyll conductance and actual  $V_{cmax}$ ) was not done in this study. Currently, without additional leaf-level data to better understand the mechanisms underlying mesophyll conductance changes during droughts, we use an apparent  $V_{cmax,app}$  (Rogers *et al.*, 2017; Drake *et al.*, 2017; Keenan *et al.*, 2010a).

### 4.3. Optimal stomatal behaviour during drought and intrinsic water-use efficiency

We did not find a general pattern of systematically decreasing  $G_1$  during drought or, in other words, an increasing  $iWUE$  for stomatal closure (increasing  $\lambda$ ) across ecosystems as theoretically predicted (Manzoni *et al.*, 2011; Mäkelä, 1996). To the contrary, we found constant  $G_1$  (and therefore  $\lambda$ ) values at most sites, and even increasing values at DE-Hai (and BE-Bra to a lesser extent). This result (a constant  $G_1$ ) is rather surprising as it would suggest that changes in stomatal conductance responses were not needed to model  $G_s$  under long-term water stress events (Keenan *et al.*, 2009) and that  $\lambda$  does not increase with drought. We argue that this is caused by the fact that NSOL were not considered by Manzoni *et al.* (2011) and Mäkelä (1996) in their analyses as stomatal closure (reduced  $G_s$ ) is known to regulate leaf water flows in response to soil water availability (Buckley, 2019); without such mechanisms, leaves would be quickly dehydrated. However, one should consider that stomatal closure in response to drought does not necessarily lead to a decrease in  $G_1$  as, in USO, any reduction of  $V_{cmax,app}$  lead to a reduction in stomatal conductance (Zhou *et al.*, 2013). At very low REW values, previous studies showed that  $C_i$  could even increase because of NSOL (Flexas, 2002), which would explain the increase of  $G_1$  we observed at DE-Hai.

Another more complex approach to stomatal conductance modelling is to model

stomatal conductance in the function of leaf water potential which is expected to regulate  $G_s$  (TUZET *et al.*, 2003). This approach requires a complete model of water flow from the soil through the plant to the atmosphere (Tuzet *et al.*, 2017). This kind of model was tested by reducing the stomatal slope of the Ball–Berry–Leuning model (Leuning, 1995) with leaf water potential, but the model did not account for NSOL (Anderegg *et al.*, 2017). Recently, Dewar *et al.* (2018) proposed a new optimization model in which stomatal behaviour maximizes photosynthesis and where the costs of stomatal closure arise from NSOL (mesophyll conductance and/or carboxylation rate) and/or loss of hydraulic conductance (Sperry *et al.*, 2016). This results in a parameter, equivalent to  $G_1$ , which is expressed as a function of measurable variables such as hydraulic conductivity, leaf water potential and  $V_{cmax}$ . This model has been successfully tested on saplings for different plant functional types (Gimeno *et al.*, 2019) and fitted well sub-daily leaf scale observations; however, this still needs to be tested for longer term in situ ecosystem droughts. This could not be done in this study as leaf and soil water potentials were lacking. It does, however, highlight a promising research path for the future.

## 5. Conclusion

In this study, we used a big leaf framework to investigate the origin of edaphic drought impacts on GPP (stomatal origin limitation and non-stomatal origin limitation) in European forest ecosystems during the 2018 drought. In agreement with Granier *et al.* (2007a), we found that GPP and  $G_s$  were both greatly affected by soil moisture depletion at many sites. We went a step further by showing that these reductions could be faithfully modelled by decreasing  $V_{cmax,app}$  (NSOL) when the REW dropped below around 0.4 while keeping the  $G_1$  (SOL) parameter from the USO model (Medlyn *et al.*, 2011a) constant. These results were rather unexpected as it would suggest that stomatal closure was not responsible for GPP reductions with drought. We argue that this was caused by the fact that  $G_1$  was not representative of stomatal behaviour during drought because GPP was not only regulated by stomatal closure but also by NSOL. Nevertheless, these results strengthen the increasing evidence that NSOL should be included in stomatal conductance/photosynthesis models to faithfully simulate both GPP and water fluxes in forest ecosystems.

### **CRedit authorship contribution statement**

L.G.dlM. was the leader of the data analysis and writing of the paper. Q.B. and B.H. participated in the design of the study, gave intellectual inputs and feedbacks to the paper. B.L. designed the study and followed the data analyses and writing process closely. All the other co-authors were involved in the data acquisition as principal investigators for each site and actively contributed to the paper by commenting on intermediate versions.

### **Declaration of Competing Interest**

We declare we have no competing interests

### **Data availability**

Data will be made available on request.

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**Non-stomatal processes are responsible for the decrease in gross primary production of a potato crop during edaphic drought**



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## 1. Personal contributions

The maintenance of the EC site has been carried out by the technical team of the Loncée ICOS station, namely Henri Chopin and Gaëtan Bogaerts. The Ecosystem Thematic Centre (ETC) performed the data processing and quality control of the EC site of Loncée. The EC data used in this study is available on the ICOS Carbon Portal (see references in the paper below). I performed the data analysis and wrote the manuscript. Bernard Longdoz supervised the research, helped in interpreting the results and reviewed the different versions of the manuscript. Bernard Heinesch participated to the reviewing processes and helped in interpreting the results.

## 2. Abstract

Soil water stress is one of the main constraints on agrosystem functioning, causing a reduction in GPP. It is a key factor for the selection of drought-tolerant plant varieties and the adaptation of irrigation management strategies. Therefore, identifying the physiological factors limiting GPP is crucial. LSMs commonly implement a beta ( $\beta$ ) stress function to reproduce the effects of drought on GPP. It is still unclear whether GPP limitations originate from a direct stomatal response (SOL) or from other non-stomatal causes (NSOL). Moreover, the shape and thresholds of the  $\beta$  function are a major cause of uncertainty in LSMs. This study investigated the effects of edaphic drought on GPP limitations of a potato crop by using eddy covariance data from the Loncée ICOS station (BE-Lon) in Belgium for four years (2006, 2010, 2014 and 2018) and by calibrating the  $\beta$  function on NSOL and SOL. The decrease in the REW in the soil induced a decline in the apparent maximum carboxylation rate (NSOL) from similar REW thresholds for all growing seasons of potato crops. A sensitivity analysis showed that the non-inclusion of REW decrease on NSOL in the modeling of GPP led to important overestimations of carbon sequestration from a threshold corresponding to 46% of the maximum available soil water. The stomatal sensitivity to photosynthesis (SOL) remained constant or even increased in 2010 and 2018. The carbon and water fluxes were decoupled when REW decreased, exhibiting a strict control of NSOL on hourly dynamics of GPP. This study provides REW thresholds to identify drought stress episodes and to help for designing irrigation management strategies. Our results advocate for a better representation of the influence of drought on photosynthesis processes of potato crops to improve the accuracy of model predictions during drought.

### 3. Introduction

Many regions across the world are facing more intense and frequent episodes of water stress due to climate change (Chiang *et al.*, 2021). Edaphic drought has become a major limiting abiotic stress for agricultural production (Fahad *et al.*, 2017). It is also known to cause restrictions in carbon assimilation, impacting the photosynthetic capacities of terrestrial ecosystems, especially those located in the mid-latitude region of the Northern Hemisphere (Yu *et al.*, 2017). Consequently, severe water stress events reduce crop yield and sometimes offset carbon sinks or cause shifts from sink to source, inducing a positive feedback effect on global warming (Reichstein *et al.*, 2013). Predicting and anticipating the impacts of drought on ecosystems' abilities to store carbon requires a detailed understanding of the physiological processes involving CO<sub>2</sub> uptake through photosynthesis and their interconnections.

Plants respond to low soil water content through a complex chain of mechanisms constituting their adaptation strategies (see Harper *et al.* (2021) for a review). For example, plants usually close their stomata when the atmospheric demand for water (represented by the VPD) overcomes the supply capacity from the soil, which prevents cavitation in the plants' hydraulic transfer system (Hetherington & Woodward, 2003). A consequence of stomatal closure is the reduction of CO<sub>2</sub> diffusion from the leaf-surrounding atmosphere to sub-stomatal cavities, which leads to a decrease in carbon assimilation (Flexas *et al.*, 2004). Moreover, it is well known that stomatal conductance to CO<sub>2</sub> transfer ( $G_s$ ) is highly correlated to GPP under stable vapor pressure deficit and soil water status (Wong *et al.*, 1979). These findings have been the basis of the empirical parametrization of  $G_s$  using carbon assimilation as a key element in addition to environmental factors (see Damour *et al.* (2010) for a review). Within this framework, any decrease of  $G_s$  can be caused either by a modification of the environmental conditions or by an inhibition of photosynthetic capacities. The USO model (Medlyn *et al.*, 2011a) reconciled this empirical approach with the optimality theory of Cowan & Farquhar (1977), which states that a plant should adjust its stomatal opening to maximize carbon assimilation while minimizing water losses over a constant time interval. In this USO framework, the  $G_1$  parameter represents the slope of the linear relationship between  $G_s$  and carbon assimilation normalized by VPD and CO<sub>2</sub> concentration and is related to the marginal carbon gain per unit of water transpired (Medlyn *et al.*, 2011a). A reduction of carbon assimilation resulting from a decrease in  $G_1$  corresponds to a direct stomatal control of CO<sub>2</sub> diffusion and will be referred in this paper to as a SOL. Carbon assimilation can also be impacted through the inhibition of photosynthetic capacities in two possible ways: changes in mesophyll conductance restricting CO<sub>2</sub> diffusion from sub-stomatal cavities to the carboxylation sites in the chloroplast (Flexas *et al.*, 2008), and modification of the biochemical reaction activities in the chloroplast. Both are generally implemented in photosynthesis modeling by a reduction of the apparent maximum carboxylation rate  $V_{cmax,app}$  as in the FvCB

model of photosynthesis (Farquhar *et al.*, 1980; Flexas *et al.*, 2004). This origin of CO<sub>2</sub> assimilation limitation is further associated in this paper to a NSOL. It should be noted that NSOL also induces a diminution of  $G_s$  as carbon assimilation is recognized to be one of the variables influencing stomatal opening. The comparison between  $V_{cmax,app}$  and  $G_1$  dynamics during stressed periods gives insights on the importance of each cause of carbon photosynthesis perturbation (respectively NSOL and SOL). Using  $G_1$  instead of  $G_s$  as an indicator of SOL allows to partition the coupling between  $G_s$  and GPP into changes in photosynthetic capacities and marginal water cost of carbon gain (Zhou *et al.*, 2013, 2014).

LSMs commonly implement a beta ( $\beta$ ) stress function to account for soil water stress effects on carbon and water fluxes. This function uses the SWC normalized by the difference between the field capacity and the wilting point in the root layer (Egea *et al.*, 2011) to characterize the amount of water available for plant uptake. When SWC is below the field capacity, the  $\beta$  factor directly downregulates  $V_{cmax,app}$  and/or  $G_1$  (Egea *et al.*, 2011; Verhoef & Egea, 2014) to reproduce drought effects on carbon and water fluxes originating from respectively NSOL and SOL. However, Granier *et al.* (1999a) and Gourlez de la Motte *et al.* (2020) showed that plants can maintain a high level of carbon and water flux even when SWC is below the field capacity, leading to potential underestimations in carbon and water flux estimates from LSMs. Vidale *et al.* (2021) and Trugman *et al.* (2018) underlined this aspect by identifying that this stress function was one of the key uncertainties in photosynthesis and transpiration predictions in LSMs. Moreover, there is no consensus in the scientific community on whether soil water stress should impact stomatal functioning, photosynthetic capacities, or both, to accurately predict carbon and water fluxes during drought (Zhou *et al.*, 2013; Peters *et al.*, 2018). Therefore, the determination of the  $\beta$  stress factor (corresponding to a REW function) representing NSOL and SOL for different PFTs is a key factor for improving model predictions (Rogers *et al.*, 2017; Li *et al.*, 2022b).

Potato (*Solanum tuberosum*) is the third most important crop in the world in terms of food production, with more than 370 million metric tons produced in 2019 (FAOSTAT, 2021). In Europe, most potato crops are cultivated on loam soils with a predominance of clay, which allows the soil to hold an important amount of water for plant uptake. Potatoes cultivated on these soils represent more than 400 000 hectares of arable lands in Europe (Goffart *et al.*, 2022). Known to be a drought sensitive crop because of its shallow root system, potato yield and the quality of the tuber can be dramatically impacted by water stress (Obidiegwu, 2015). Therefore, any limitation on its production threatens food supply, impacting more than one billion people around the world (Lutaladio & Castaldi, 2009).

Up to now, no studies have investigated which processes are at the origin of the effects of drought on potato photosynthesis limitations estimated from eddy covariance data (EC). In this paper, we aimed at investigating the effects of drought on potato car-

bon assimilation by analyzing the data collected by the Loncée flux tower and meteorological station (BE-Lon) located in central Belgium, which is part of the Integrated Carbon Observation System (ICOS – Franz et al., 2018). More specifically, this paper aims at: (i) partitioning the impacts of REW on GPP between SOL ( $G_1$ ) and NSOL ( $V_{cmax,app}$ ) and (ii) defining a REW threshold below which these two limitations occurred.

## 4. Material and methods

### 4.1. Site description

The Loncée ICOS flux tower (Level 2 ICOS station) is installed in the middle of a cropland plot located in Loncée, about 50 km SE of Brussels in Belgium (50° 33' 5.71"N, 4° 44' 46.07"E, 167 m asl). It has been equipped with an eddy covariance (EC) system and a meteorological station since 2004 and is integrated into the CarboEurope-IP and FLUXNET networks. The cropland plot has an area of 12 hectares. The EC system has a fetch of 240 m and 200 m in the directions of the prevailing winds which are SW and NE (Buysse *et al.*, 2017). For any atmospheric conditions, the cropland area contribution to the EC footprint fluxes was large (91.1% in 2006, 90.5% in 2010, 90.7% in 2014, 93.3% in 2018, data not shown). The climate is temperate oceanic, with mean annual temperature and precipitation of respectively 10.2°C and 743 mm. The soil is a Luvisol (FAO classification) divided into two horizons (one plough layer from 0 to 35 cm and one layer enriched in clay particles from 35 to 100 cm – Table 4.1).

**Table 4.1:** Soil physical proprieties at BE-Lon: field capacity ( $\theta_{fc}$ ), wilting point ( $\theta_{wp}$ ), sand, silt, clay content and bulk density for the two soil horizons (0-35 cm and 35-100 cm). Field capacities and wilting points were measured by the WPC4 and HYPROP2 sensors (Meter environment., Hopkins Ct, NE, US) on soil samples collected in 2019.

Depth (cm)	$\theta_{fc}$ (cm <sup>-3</sup> cm <sup>-3</sup> )	$\theta_{wp}$ (cm <sup>-3</sup> cm <sup>-3</sup> )	Sand (%)	Silt (%)	Clay (%)	Bulk density (g cm <sup>-3</sup> )
0–35	45.3	11.7	8.4	80.5	11.2	1.3
35–85	40.4	13.0	6.1	75.7	18.2	1.5

### 4.2. Crop management

The Loncée station has been cultivated for more than 80 years, with cropping management based on a 4-year rotation. For at least the past 20 years (Aubinet *et al.*, 2009), the rotation has been: winter wheat (*Triticum aestivum*) / sugar beet (*Beta vulgaris*) / winter wheat (*Triticum aestivum*) / seed potato (*Solanum tuberosum*). Potatoes

were cultivated in 2006, 2010, 2014 and 2018. The farming operations during the four growing seasons are described in Table 4.4. The tubers were planted in ridges with a space of 60 to 70 cm between the ridges and of 30 to 40 cm between the tubers.

### 4.3. Meteorological and fluxes measurements

Micrometeorological measurements were collected at the half-hourly timescale, including air humidity and air temperature (RHT2, Delta-T Devices Ltd., Cambridge, UK) at 1.3 m height (and 2.8 m height in 2018), incident photosynthetic photon flux density (PAR Quantum sensor SKP 215, Skye Instruments Limited, Llandrindod Wells, UK) and SWC (EnviroSCAN Probe, Sentek Sensor Technologies, Stepney, SA, AU) at five different depths in 2018 (5, 15, 25, 55 and 85 cm). In 2006, 2010 and 2014, SWC was measured using time domain reflectometers (ML2 ThetaProbe, Delta-T Devices Ltd, Cambridge, UK) at three different depths (5, 20 and 40 cm). For all years, the soil water sensors were located at the middle of the plot near the EC station.

The net CO<sub>2</sub> ( $F_c$ ) and water vapor (LE) fluxes between the ecosystem and the atmosphere were determined on a half-hourly basis by the EC technique (Aubinet *et al.*, 2012) using high frequency data of vertical wind speed, CO<sub>2</sub> and water vapor concentrations at the ecosystem-atmosphere interface. These variables were measured at 20 Hz by respectively a sonic anemometer (Solent Research R3, Gill Instruments Lymington, UK) placed at a height of 2.93 m on a mast located at the center of the field, and an infrared gas analyzer (LI-7000 before 2014 and LI-7200 after, LI-COR, Lincoln, NE, US) placed close to the anemometer. Half-hourly fluxes of  $F_c$  and LE were determined by postprocessing raw 20 Hz data by the ICOS Ecosystem Thematic Center using the ONEFlux pipeline (Pastorello *et al.*, 2020) and are available in the ICOS Carbon Portal (Dumont *et al.*, 2023a). More precisely, the storage flux was neglected (i.e.,  $F_c$  corresponds to the net ecosystem exchange NEE) and GPP was obtained from the partitioning of NEE using the nighttime method (NT, Reichstein *et al.* (2002)). Moreover, data characterized by weak atmosphere turbulence level (low friction velocity  $u^*$ ) using the variable  $u^*$  threshold method (VUT) were discarded. The variables GPP\_NT\_VUT\_REF and LE\_F\_MDS with quality flags of 0 were selected from the dataset. Data was not gapfilled because only measurements can be used to study ecosystem functioning and no flux value cumulated over time were necessary.

### 4.4. Measurements of vegetation growth

Vegetation growth is commonly measured by the LAI which corresponds to the ratio of leaf area to unit ground surface area (Breda, 2003). However, the LAI does not allow to separate the green photosynthetic components (including chlorophyll cells) from the rest of the canopy structure. The Green Area Index (GAI), which corresponds to the green surface area of the vegetation, is a better proxy to assess the photosynthetically active component of the ecosystem. In this study, the GAI of the potato crop was mea-

sured six times during each growing season by a destructive sampling method. These measurements have been linked to the growing degree-days which is a main climatic driver of crop development (McMaster, 1997). The base temperature (temperature required for crop growth) was set to 7 °C (Sands *et al.*, 1979). A third-order polynomial function was fitted to GAI and growing-degree-days to model GAI dynamics and get a continuous function for each growing seasons.

In the results section, the timescale will be further expressed as Day after Emergence (DAE), which counts the number of days after the emergence of the first leaves of the crop, this latter corresponding to the date when GAI changed from a null to a positive value. The growing seasons were divided into two stages: the vegetative stage when GAI increased (leaf development and period before tuber development, roughly between DAE 0 and 40), and the reproductive stage when GAI decreased (tuber multiplication and elongation, roughly after DAE 40). This definition slightly differs from the reality since tuber growth initiation is expected to start before GAI reaches its maximum value (Obidiegwu, 2015). However, this assumption will not affect the results of this study since this denomination was only used for a descriptive purpose.

## 4.5. Quantification of soil water availability

The REW for plant uptake is a drought index which corresponds to the fraction of available water for plant uptake from the soil surface up to the maximum rooting depth. REW is expressed as a fraction of the maximum available water calculated by the difference between field capacity and wilting point (Granier *et al.*, 1999a). The maximum rooting depth was estimated by analyzing the dynamics of SWC within each soil horizon during periods of potato crop development without precipitation. The decrease in SWC was observed only within the first soil horizon (0–35 cm, Table 4.1), suggesting that root water uptake only occurred in surface soil layers. Therefore, we assumed that the roots can uptake soil water only up to 35 cm depth (Table 4.1). This assumption could lead to REW overestimation at the beginning of the growing season as most roots are located within shallow soil layers. However, SWC data showed a rapid extension of the water uptake into deeper soil layers and the beginning of the growing season is not a drought-prone period. REW was calculated as follows (Granier *et al.*, 1999a):

$$REW = \frac{\bar{\theta} - \theta_{fc}}{\theta_{fc} - \theta_{wp}} \quad (4.1)$$

where  $\theta_{wp}$ ,  $\theta_{fc}$  and  $\theta$  are respectively the wilting point, the field capacity and the average of SWC measurements in the first soil horizon (Table 1). The units and a description of the variables used in this study can be found in Table 4.5. REW ranges from 1 ( $\theta = \theta_{fc}$ ) to 0 ( $\theta = \theta_{wp}$ ), even if values higher than 1 can be observed just



after a rainfall event, when soil water content is above field capacity. REW is directly related to the water stress  $\beta$  factor used in many LSMs (Verhoef & Egea, 2014):

$$\beta = \begin{cases} 1 & \bar{\theta} > \theta_{fc} \\ REW^p & \theta_{wp} < \theta_{fc} \\ 0 & \bar{\theta} < \theta_{wp} \end{cases} \quad (4.2)$$

with  $p$  the exponent allowing to implement a non-linear dependence of model parameters on REW. However, many species can maintain constant photosynthetic parameters even when  $\theta < \theta_{fc}$  (Granier *et al.*, 1999a; Gourlez de la Motte *et al.*, 2020). Investigating the effects of drought on SOL and NSOL to determine specific REW thresholds from which model parameters linked to stomatal opening and apparent carboxylation rate are impacted by water stress is pivotal to avoid the propagation of uncertainties in models. This study focused on the characterization of these threshold by assuming that  $p = 1$ , as implemented in many LSMs (Vidale *et al.*, 2021; Oliver *et al.*, 2022).

## 4.6. "big leaf" approach

The "big leaf" is a concept where the canopy is simplified as a single leaf at the interface with the surrounding atmosphere. It is used to infer ecosystem physiological proprieties from flux tower measurements (Knauer *et al.*, 2018a). The air characteristics over the "big leaf" surface and the conductance related to the transfer of water vapor through respectively the stomata of the "big leaf" (named canopy conductance to water vapor  $G_{sw}$ ) and the boundary layer (named aerodynamic conductance to water vapor  $G_{aw}$ ) can be inferred from the micro-meteorological measurements performed above the canopy.

### 4.6.1. Determination of the "big leaf" surface conditions and aerodynamic conductance

Following the Thom model (Thom, 1972),  $G_{aw}$  is equivalent to the aerodynamic conductance for sensible heat ( $G_{ah}$ ) and can be determined by:

$$G_{aw} = G_{ah} = \left( \frac{u}{u_*^2} + 6u_*^{-0.667} \right) \quad (4.3)$$

where  $u_*$  is the friction velocity and  $u$  the wind speed above the canopy.  $G_{ah}$  was used to calculate the H<sub>2</sub>O concentration ( $e_s$ ) and the CO<sub>2</sub> concentration ( $C_s$ ) at the "big leaf" surface as follows (Knauer *et al.*, 2018a):

$$e_s = e_a + \frac{LE\gamma}{\rho_a G_{aw} c_p} \quad (4.4)$$

$$C_s = C_a + \frac{NEE}{G_{aw}} \frac{RT_a}{p} \quad (4.5)$$

with  $LE$  the latent heat flux (measured by the EC system),  $\rho_a$  the air density,  $c_p$  the heat capacity of dry air,  $\gamma$  the psychrometric constant and  $R$  is the perfect gas constant.  $e_a$ ,  $C_a$ ,  $T_a$  and  $p$  correspond respectively to the air vapor pressure, the  $\text{CO}_2$  concentration, the air temperature and the air pressure measured above the canopy.  $\frac{RT_a}{p}$  in Eq. (4.5) converts  $G_{aw}$  from  $\text{m s}^{-1}$  to  $\text{mol m}^{-2} \text{s}^{-1}$ .

The canopy temperature was determined from the longwave fluxes at the canopy surface (Knauer *et al.*, 2018a):

$$T_s = \sqrt[4]{\frac{LW_{out} - (1 - \epsilon)LW_{in}}{\epsilon\sigma}} \quad (4.6)$$

with  $LW_{in}$  and  $LW_{out}$  the longwave downward and outgoing fluxes measured by the meteorological station,  $\epsilon$  the emissivity of the canopy and  $\sigma$  the Stefan-Boltzmann constant. Missing  $T_s$  values due to lack of longwave fluxes measurements were gap filled with  $T_a$ . The vapor pressure deficit at the “big leaf” surface ( $VPD_s$ ) was calculated as the difference between the saturated vapor pressure and the air vapor pressure at the canopy surface ( $e_s$ ).

#### 4.6.2. Canopy conductance

$G_{sw}$  was computed by inverting the Penman-Monteith equation (Monteith, 1965):

$$G_{sw} = \frac{G_{ah}\gamma LE}{\Delta(R_n - G - S) + \rho_a c_p G_{ah} VPD_s - (\Delta + \gamma) LE} \quad (4.7)$$

where  $\Delta$  is the slope of the saturated vapor pressure curve at  $T_a$  calculated from Allen & of the United Nations (1998),  $R_n$  is the net radiative flux and  $G$  the ground heat flux (both measured). The flux corresponding to the energy stored in the canopy ( $S$ ) was neglected. Inferring  $G_{sw}$  from Eq. (4.7) is only possible when  $LE$  corresponds to the transpiration flux of the ecosystem (i.e., when soil and canopy evaporation are negligible comparing to transpiration). Therefore, the analysis focused only on the growing season and data during precipitation events and the subsequent 48 h (when evaporation is significant) were discarded. To ensure the selection of data with meaningful  $LE$  values, periods with low photosynthetic photon flux density ( $PPFD < 200 \mu\text{mol m}^{-2} \text{s}^{-1}$ ), low surface temperature ( $< 5 \text{ }^\circ\text{C}$ ), high relative humidity ( $RH > 95 \%$ ),  $LE < 0$  and  $R_n - G < 0$  were also discarded (Knauer *et al.*, 2018b). Moreover, the violation of the energy balance closure assumption in Eq. (4.7) may affect the interpretability of  $G_{sw}$  (Knauer *et al.*, 2018b). This issue was addressed by computing the slope of the regression between the sum of the turbulent fluxes and available energy

and by discarding half hourly data when the standardized residuals from the linear regression between were superior to 3.

#### 4.6.3. Maximum apparent carboxylation rate

Under high irradiance, photosynthesis is limited by the activity of the RuBisCO enzyme (Farquhar *et al.*, 1980). In this case, the FvCB model applied at the “big leaf” surface estimates GPP under high irradiance ( $GPP_{sat}$ ) from the apparent maximum carboxylation rate ( $V_{cmax,app}$ ), the intracellular CO<sub>2</sub> concentration ( $C_i$ ), the CO<sub>2</sub> compensation ( $\Gamma^*$ ) point and the effective Michaelis-Menten coefficient for RuBisCO kinetics ( $K_m$ ):

$$GPP_{sat} = \frac{V_{cmax,app}(C_i - \Gamma^*)}{C_i + K_m} \quad (4.8)$$

Both  $K_m$  and  $\Gamma^*$  were determined following an air temperature-based dependence (Bernacchi *et al.*, 2001). Eq. (4.8) assumes that leaf photorespiration is negligible under high irradiance.  $V_{cmax,app}$  was calculated by inverting Eq. (4.8):

$$V_{cmax,app} = \frac{GPP_{sat}(C_i + K_m)}{C_i - \Gamma^*} \quad (4.9)$$

where  $GPP_{sat}$  is the GPP when PPFD is higher that a threshold ( $PPFD_{sat}$ ) from which GPP saturated. This threshold was calculated from the light response curves of GPP for each growing season. The Mitscherlich model was used to mathematically describe the relationship between GPP and PPFD. Data was averaged within classes with a constant width of 200  $\mu\text{mol m}^{-2} \text{s}^{-1}$  (respectively  $\overline{GPP}$  and  $\overline{PPFD}$ ):

$$\overline{GPP} = a(1 - e^{-b(\overline{PPFD}-c)}) \quad (4.10)$$

with  $a$ ,  $b$  and  $c$  three parameters to adjust. The PPFD threshold was defined as the PPFD value when GPP deviated from the asymptote of the model ( $a$ ) as follows:

$$PPFD_{sat} = c - \frac{\ln(\frac{SE_a}{a})}{b} \quad (4.11)$$

with  $SE_a$  the standard deviation around  $a$ . In Eq. (4.9),  $C_i$  was determined following a Fick diffusive law (Farquhar & Sharkey, 1982), representing the CO<sub>2</sub> diffusion through stomata:

$$C_i = C_s - \frac{1.6GPP_{sat}}{G_{sw}} \quad (4.12)$$

where  $G_{sw}$  is determined by Eq. (4.7), 1.6 is the ratio between H<sub>2</sub>O and CO<sub>2</sub> conductances and  $C_s$  is the CO<sub>2</sub> concentration at the leaf surface, determined based on Eq.

(4.5).  $V_{cmax,app}$  values were further normalized at 25 °C ( $V_{cmax,app,25}$ ) by adjusting the Arrhenius model on  $V_{cmax,app}$  and  $T_s$  data averaged within  $T_s$  classes with a width of 2 °C (Medlyn *et al.*, 2002).

## 4.7. NSOL during drought

The effects of drought on NSOL can be characterized only if the influence of other factors on the dynamics of  $V_{cmax,app}$  is discarded from the analysis. The impact of  $u_*$  is limited by selecting fluxes during low turbulence periods, while  $V_{cmax,app}$  was normalized at 25 °C to account for temperature effects on NSOL. Moreover, only half-hourly data at high irradiance (PPFD higher than  $PPFD_{sat}$ ) were selected. The only remaining potential influence (besides REW) comes from the GAI variability which can be observed only for a study of the seasonal dynamics of  $GPP_{sat}$ . Therefore, the analysis should be focused on periods with relatively low variations of the phenological development of the crop. To overcome this potential influence, data with GAI above a threshold ( $GAI_{sat}$ ) were selected to ensure that the amount of green area is high enough so that any increase in the green canopy cover had a negligible effect on  $GPP_{sat}$  dynamics.  $GAI_{sat}$  was determined by the same method as  $PPFD_{sat}$  using Eqs. (4.10) and (4.11) by replacing PPFD by GAI. Days when  $GAI < GAI_{sat}$  were discarded since it may induce a bias in the analysis of the influence of REW on SOL and NSOL. The impact of edaphic drought on GPP originating from NSOL was assessed by studying the influence of REW on  $V_{cmax,app,25}$  using the data selected.

## 4.8. SOL during drought

SOL was defined in this paper as a decrease in  $G_1$ , the slope parameter of the USO model which was adapted from the leaf to the canopy level by using the “big leaf” framework (Medlyn *et al.*, 2017b; Knauer *et al.*, 2018a):

$$G_{sw} = G_0 + 1.6 \left( 1 + \frac{G_1}{\sqrt{VPD_s}} \right) \frac{GPP_{sat}}{C_s} \quad (4.13)$$

where  $G_0$  is the minimum stomatal conductance which was set to 0 (Medlyn *et al.*, 2017b) and  $G_1$  represents the ecosystem stomatal sensitivity to GPP at high irradiance normalized by  $VPD_s$  and  $C_s$ .  $G_1$  was calculated by combining Eqs. (4.12) and (4.13):

$$G_1 = \frac{\frac{C_i}{C_s} \sqrt{VPD_s}}{1 - \frac{C_i}{C_s}} \quad (4.14)$$

The impact of edaphic drought on  $GPP_{sat}$  originating from SOL was assessed by studying the influence of REW on  $G_1$ .

## 4.9. Detection of REW thresholds

A linear segmented model including one breakpoint and one segment at a constant value for high REW (asymptote) was adjusted between half-hourly  $V_{cmax,app,25}$  or  $G_1$  and REW. The asymptote of the model characterizes the unstressed  $V_{cmax,app,25}$  and  $G_1$  values, respectively named  $V_{cmax,app,25}^*$  and  $G_1^*$ . This procedure allows the detection of a breakpoint, below which SOL and NSOL increase or decrease linearly with REW. A Fisher F-test was conducted to verify if this segmented model was significantly different from a linear regression with no breakpoint. The quality of the segmented regressions was also assessed by the p-values of the adjusted coefficients and the  $R^2$  of the model.

## 4.10. Sensitivity analysis

A sensitivity analysis was performed by calculating the ratio between theoretical  $GPP_{sat}$  values that would be modeled without accounting for water stress effect on  $V_{cmax,app,25}$  and  $G_1$  ( $GPP_{V_{cmax,app,25}^*}$  and  $GPP_{G_1^*}$ ), with observed  $GPP_{sat}$  values.  $GPP_{V_{cmax,app,25}^*}$  was determined by using Eq. (4.8) where  $V_{cmax,app,25}$  was replaced by  $V_{cmax,app,25}^*$  and by assuming that  $G_0$  and ecosystem respiration are negligible.  $C_i$  in Eq. (4.8) was deduced from  $GPP_{V_{cmax,app,25}^*}$ ,  $C_a$ ,  $G_{aw}$ ,  $G_1$  and  $VPD_s$  by combining Eqs. (4.12), (4.5) and (4.13):

$$C_i = \left( C_a - \frac{GPP_{V_{cmax,app,25}^*}}{G_{aw}} \right) \left( 1 - \frac{1}{1 + \frac{G_1}{\sqrt{VPD_s}}} \right) \quad (4.15)$$

Inserting Eq. (4.15) in Eq. (4.8) gives the following equation:

$$GPP_{V_{cmax,app,25}^*} = V_{cmax,app,25}^* \frac{\left( C_a - \frac{GPP_{V_{cmax,app,25}^*}}{G_{aw}} \right) \left( 1 - \frac{1}{1 + \frac{G_1}{\sqrt{VPD_s}}} \right) - \Gamma^*}{\left( C_a - \frac{GPP_{V_{cmax,app,25}^*}}{G_{aw}} \right) \left( 1 - \frac{1}{1 + \frac{G_1}{\sqrt{VPD_s}}} \right) + K_m} \quad (4.16)$$

which was solved for  $GPP_{V_{cmax,app,25}^*}$ . In a similar way,  $GPP_{G_1^*}$  was determined by solving Eq. (4.16) where  $V_{cmax,app,25}^*$  was replaced by observed  $V_{cmax,app,25}$  and  $G_1$  by  $G_1^*$ . The comparison between  $GPP_{V_{cmax,app,25}^*}$  (or  $GPP_{G_1^*}$ ) and the measured  $GPP_{sat}$  estimates the bias induced by the non-inclusion of NSOL (or SOL) in the modeling of  $GPP_{sat}$ . More precisely, a ratio  $GPP_{G_1^*} / GPP_{sat}$  or  $GPP_{V_{cmax,app,25}^*} / GPP_{sat}$  higher than 1 indicates an overestimation of  $GPP_{sat}$  when REW effects on  $G_1$  or  $V_{cmax,app,25}$  are not considered.

## 4.11. Coupling between transpiration and photosynthesis

Stomata opening regulates both CO<sub>2</sub> and water vapor fluxes and is affected by the evaporative demand of the atmosphere. This leads to a high degree of coupling at the hourly timescale between  $GPP\sqrt{VPD_s}$  and LE (Nelson *et al.*, 2018). Changes in environmental conditions may affect this level of coupling. Temperature variations, or appearance of cloud cover (decrease of PPFD) may induce a decoupling between  $GPP\sqrt{VPD_s}$  and LE (Reinhardt & Smith, 2008; Urban *et al.*, 2017a; Krich *et al.*, 2022; Marchin *et al.*, 2022). Moreover, as drought intensifies and affects biochemical factors, carbon fixation may be inhibited independently of LE. Therefore, avoiding the potential effects of PPFD and  $T_s$  is required to investigate the effects of drought on the coupling between  $GPP\sqrt{VPD_s}$  and LE. This has been realized by regrouping many days with large ranges of  $T_s$  and PPFD daily variations.

The level of coupling was quantified by the Spearman's rank correlation coefficient (Spearman, 1904) between LE and normalized carbon assimilation ( $GPP\sqrt{VPD_s}$ ) (Nelson *et al.*, 2018):

$$DWCI = \rho(LE, GPP\sqrt{VPD_s}) \quad (4.17)$$

with *DWCI* the Diurnal Water-Carbon Index (Nelson *et al.*, 2018). The fluxes were normalized by their maximum value for each day before using Eq. (4.17), which allowed for using fluxes measurements across growing seasons. The data already discarded for the  $G_{sw}$  calculation process was removed, and the period of analysis was restricted to the growing seasons. A linear segmented regression was adjusted on the *DWCI* dependence on REW (all years regrouped) following the method presented in 2.9 to detect if the decrease in REW impacted the degree of coupling between carbon and water fluxes.

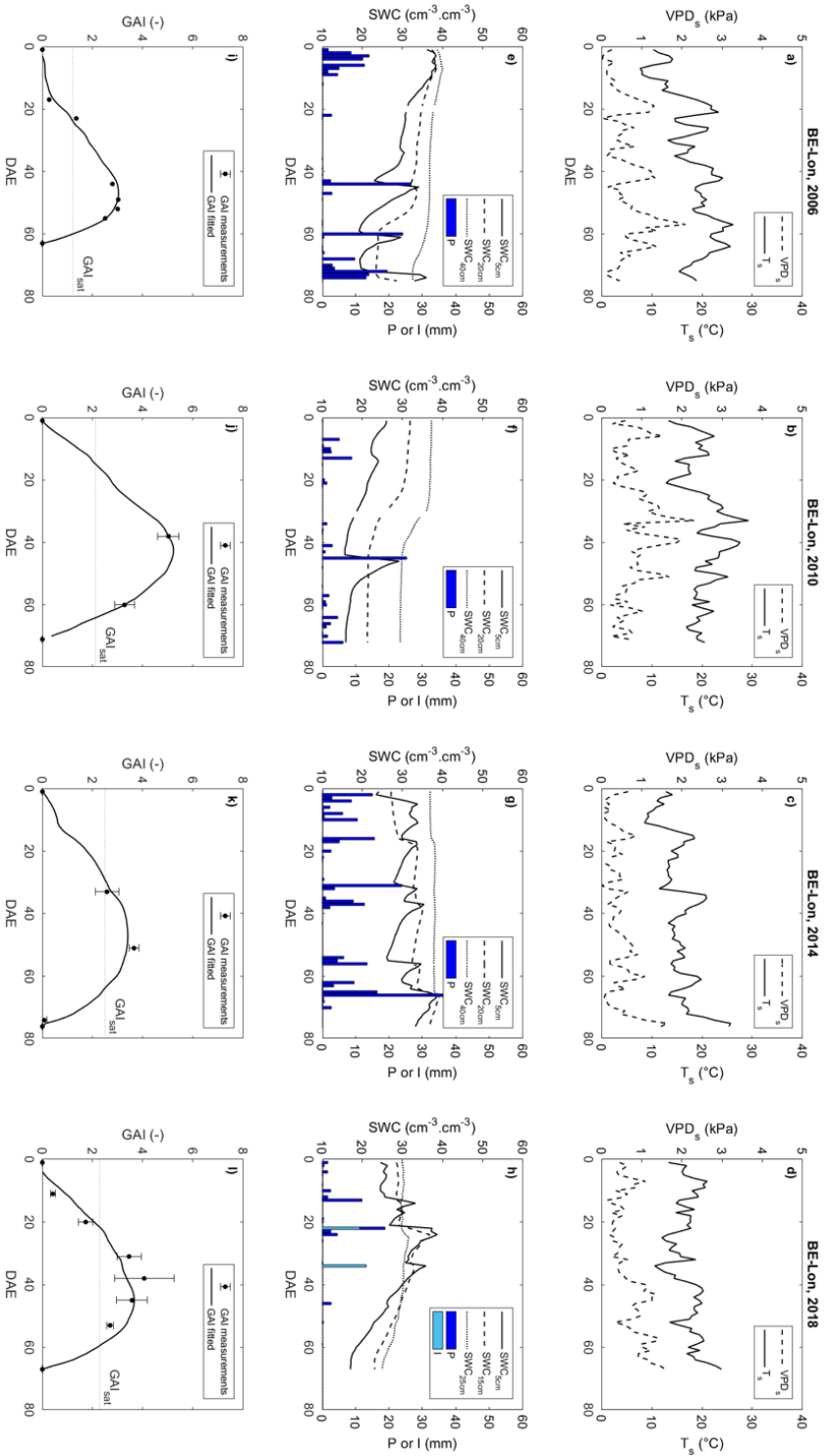
## 5. Results

### 5.1. Ancillary data and photosynthesis dynamics

#### 5.1.1. Crop characteristics and pedo-climatic conditions

The maximum GAI value was measured in 2010 (5.02) while other years showed relatively similar maximum values between 3 and 4 (Fig. 4.1, Table 4.2). The year 2006 had the longest growing season with 74 days, compared to the three other years, which had a growing season of 71 days in 2010, 64 days in 2014 and 66 days in 2018 (Table 2). These differences are explained by farming activities and haulm killing dates that change throughout the years. The 2018 growing season had the lowest rainfall, with cumulative precipitation of 68.4 mm, slightly lower than 2010 (76.6 mm), and considerably below 2006 (186.2 mm) and 2014 (222.4 mm) (Table 4.2). In 2018,

Chapter 4. Non-stomatal processes are responsible for the decrease in gross primary production of a potato crop during edaphic drought



**Figure 4.1:** Vapor pressure deficit at the canopy surface ( $VPD_s$ ), canopy temperature ( $T_s$ ) (subplots a to d) and soil water content (SWC) (subplots e to h) daily means. Daily sums of precipitation ( $P$ ) and irrigation ( $I$ ) are displayed on the second line (subplots e to h). Interpolated green area index  $GAI$  ( $GAI$  fitted) and measurements ( $GAI$  measurements) with their standard deviation are shown on the third line (subplots i to l). DAE corresponds to days after emergence.

the site was irrigated two times (24.5 mm) by the aspersion technique (Fig. 4.1), for a total amount of water received by the crop of about 92.9 mm (precipitation and irrigation). SWC at 5 cm depth decreased with the lack of precipitation and reached approximately 16 % in 2018 and 2010 (Fig. 4.1), which was slightly higher than the wilting point ( $\theta_{wp} = 11.7$  between 0 and 35 cm depth – Table 4.1). The minimum SWC values at deeper depths remained close to 25 %, suggesting that deeper soil layers in the first horizon were able to hold water above the wilting point during the growing seasons. The periods of lack of precipitation were associated to an increase in  $VPD_s$  and  $T_s$  (regularly in 2010, in late 2006 and 2018) when their daily averaged values both reached respectively about 1.5 kPa and 25 °C (Fig. 4.1). The trends in  $VPD_s$ ,  $T_s$  and SWC illustrated that climate conditions generated periods with potential conditions for both atmospheric and edaphic droughts during the vegetative stage in 2010 and during the reproductive stage in 2006 and 2018. The lowest REW values were quite similar in 2006, 2010 and 2018 growing seasons at respectively 0.40, 0.31 and 0.30 (Fig. 4.2). In 2014, REW was always higher than 0.60 due to favorable weather conditions.

**Table 4.2:** Crop characteristics: growing season (from emergence to haulm killing), cumulative precipitation ( $P_{cum}$ ), irrigation (I), maximum green area index ( $GAI_{max}$ ) and variety.

Growing season (days)	$P_{cum}$ (mm)	I (mm)	$GAI_{max}$ (-)	Variety (% of total area)	
2006	74	186.2	-	3.02	Spunta (62.8 %) / Primura (11.6 %) / Kondor (11.6 %) / Draga (14 %)
2010	71	76.6	-	$5.02 \pm 0.42$	Draga (100 %)
2014	64	222.4	-	$3.55 \pm 0.19$	Draga (72.8 %) / Bianchidea (20 %) / Kennebec (7.2 %)
2018	66	68.4	24.5	$4.05 \pm 1.18$	Agria (53.9 %) / Unidea (26.4 %) / Draga (19.7 %)

### 5.1.2. Photosynthesis and canopy conductance dynamics

A deviation from the expected bell-shape  $GPP_{sat}$  curve following GAI dynamics can be observed in 2018 and 2006 (during the reproductive stage) and in 2010 (during the vegetative stage) where  $GPP_{sat}$  dropped from 25 to 35  $\mu\text{mol m}^{-2} \text{s}^{-1}$  to 5–15  $\mu\text{mol m}^{-2} \text{s}^{-1}$  (Fig. 4.2). These sharp declines coincided with the decrease in REW (roughly from 0.65 to 0.45 in 2006, and from 0.60 to 0.30 in 2010 and 2018) and  $G_{sw}$  (roughly from 0.60 to 0.25  $\text{mol m}^{-2} \text{s}^{-1}$  in 2006 and from 0.25 to 0.10  $\text{mol m}^{-2} \text{s}^{-1}$  in 2010 and 2018, Fig. 4.2), which corresponded to the closure of the stomata when REW values were low. The relatively small amount of data recorded in 2014 (due to numerous precipitation events, and important gaps in the shortwave and longwave flux measurements) and 2006 (due to numerous days when cumulative precipitation was very low; e.g.,  $<0.3 \text{ mm day}^{-1}$  on DAE 39, 41, 56 or 64) explain that very few  $G_{sw}$  data points were available for the analysis. These observations highlighted two different



timings in drought appearance with varying amplitude and intensity, characterized by a decrease in  $GPP_{sat}$  and  $G_{sw}$ .

## 5.2. Drought impact on SOL and NSOL at the seasonal scale

### 5.2.1. Impact of drought on $V_{cmax,app,25}$

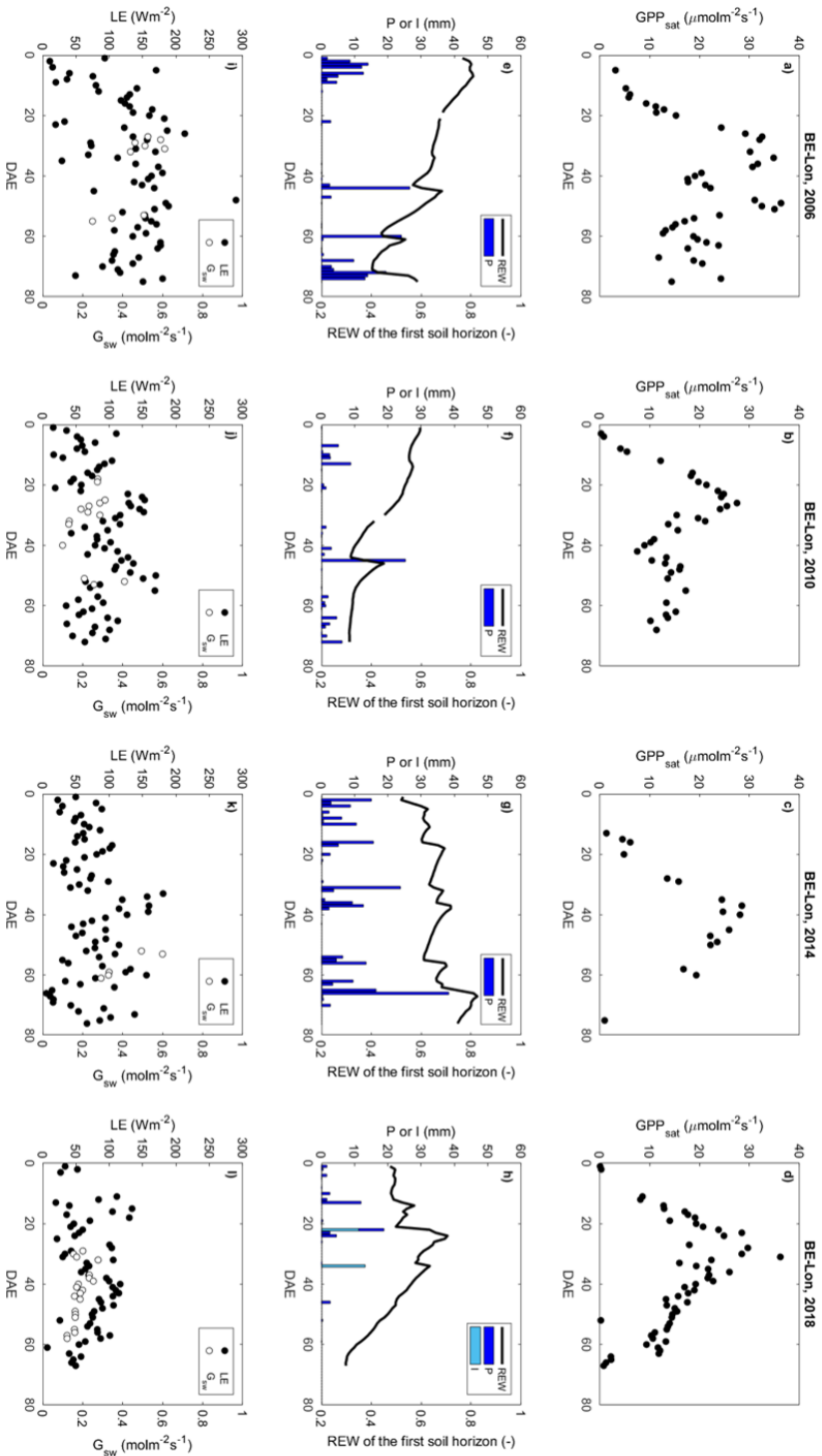
$V_{cmax,app,25}$  decreased with the decline in REW. The linear segmented model significantly described the dependence between  $V_{cmax,app,25}$  and REW in 2018, 2010 and 2006 (Table 4.3 – Fig. 4.2). REW breakpoints ranging between  $0.58 \pm 0.01$  and  $0.45 \pm 0.02$  were found whatever the development stages of the crop when drought occurred, indicating a relatively uniform tipping point below which the photosynthetic capacities of the canopy decrease due to edaphic drought (Table 4.3). The asymptote of the model ( $V_{cmax,app,25}^*$ ) or averaged value was in a narrow range (from  $98.7 \pm 1.69 \mu\text{mol m}^{-2} \text{s}^{-1}$  in 2018 to  $128 \pm 6.11 \mu\text{mol m}^{-2} \text{s}^{-1}$  in 2010 – Table 4.3). In 2014, no effect of REW on  $V_{cmax,app,25}$  was observed due to recurrent precipitation events and reduced amount of data (see explanation above). The slopes of the linear segmented model ranged from  $324 \pm 36.2 \mu\text{mol m}^{-2} \text{s}^{-1}$  in 2018 to  $821 \pm 137 \mu\text{mol m}^{-2} \text{s}^{-1}$  in 2010). The systematic pattern of decreasing  $V_{cmax,app,25}$  with REW with quite similar unstressed values and breakpoints whatever the years demonstrated the existence of a clear uniform dependence of photosynthetic capacities on soil water availability either during the vegetative (2010) or reproductive stage (2006 and 2018).

### 5.2.2. Impact of drought on $G_1$

$G_1$  was significantly impacted by REW in 2010 and 2018. More precisely,  $G_1$  increased during respectively the vegetative and reproductive stages (Fig. 4.3). The REW breakpoints were very similar for these two years ( $0.44 \pm 0.05$  in 2010 and  $0.45 \pm 0.03$  in 2018, Table 4.3, Fig. 4.3). The asymptote value was slightly higher in 2010 ( $1.69 \pm 0.24 \text{ kPa}^{0.5}$ ) compared to 2018 ( $2.37 \pm 0.05 \text{ kPa}^{0.5}$ ). No significant effect of REW on  $G_1$  was observed in 2006 and 2014. The averaged  $G_1$  values in 2006 and 2014 ( $4.35 \pm 0.80 \text{ kPa}^{0.5}$  and  $4.44 \pm 1.45 \text{ kPa}^{0.5}$  respectively, Table 4.3) were clearly higher than the asymptote of the segmented model in 2010 and 2018.

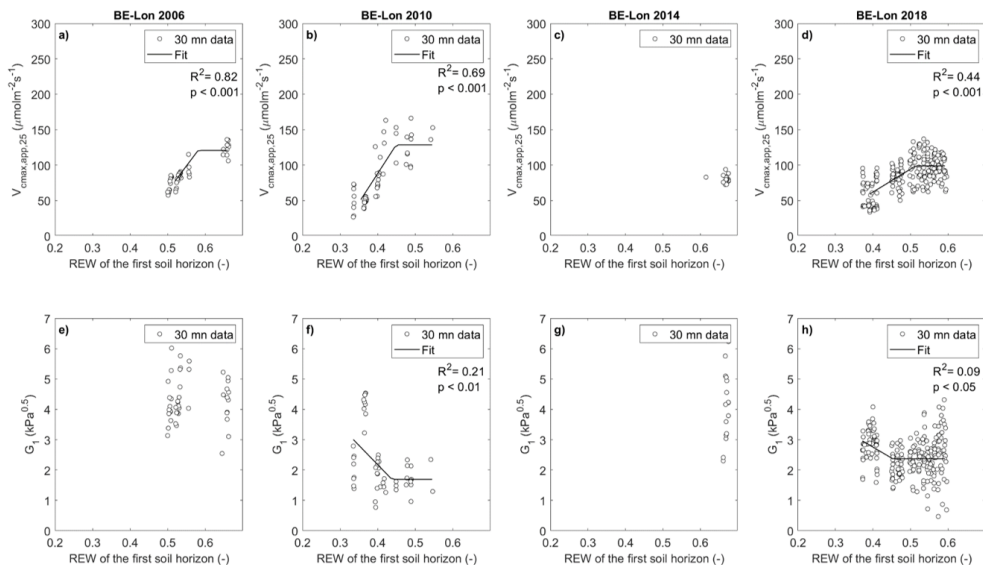
### 5.2.3. Impact of drought on GPP modeling

$GPP_{V_{cmax,app,25}^*} / GPP_{sat}$  increased when REW passed below a REW breakpoint of  $0.46 \pm 0.01$  for all years when edaphic drought induced NSOL, which is logically a similar breakpoint compared to those for  $V_{cmax,app,25}$  (Fig. 4.4- Table 4.3). This ratio even exceeded 2 in 2010 and 2018 when the REW reached its minimum values during water stress events (Fig. 4.4). In opposition, the ratio  $GPP_{G_1} / GPP_{sat}$  remained relatively constant for all the years and whatever the soil water availability. (Fig. 4.4). The REW breakpoint of the sensitivity analysis characterizes a tipping point illustrating the importance of including NSOL of GPP for potatoes during edaphic drought.

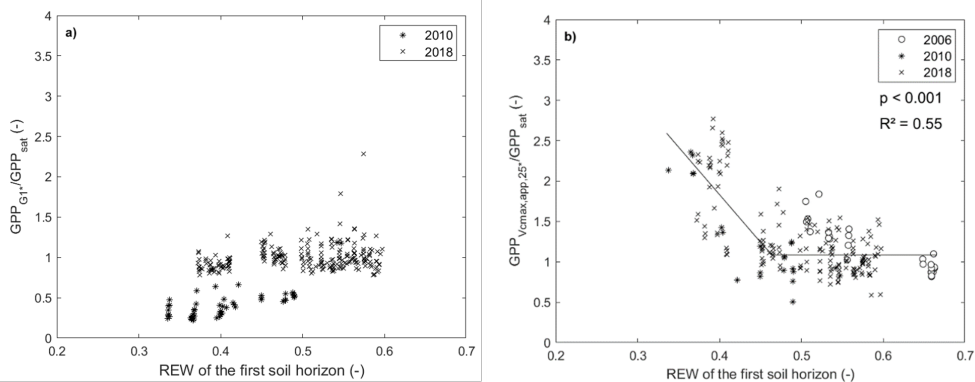


**Figure 4.2:** Gross primary production under high irradiance ( $GPP_{sat}$ ; subplots a to d), relative extractable water (REW; subplots e to h), latent heat flux (LE) and canopy conductance to water vapor ( $G_{sw}$ ; subplots i to f) daily means. Daily sums of precipitation (P) and irrigation (I) are displayed on the second line (subplots e to h). DAE corresponds to days after emergence.

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**Figure 4.3:** Apparent maximum carboxylation rate at 25°C ( $V_{cmax,app,25}$ ; subplots a to d) and slope parameter of the USO model ( $G_1$ ; subplots e to h) dynamics in function of the REW for the years studied. p refers to the p-value comparing the segmented model to a linear model (p-val<sub>mod</sub> in Table 4.3).



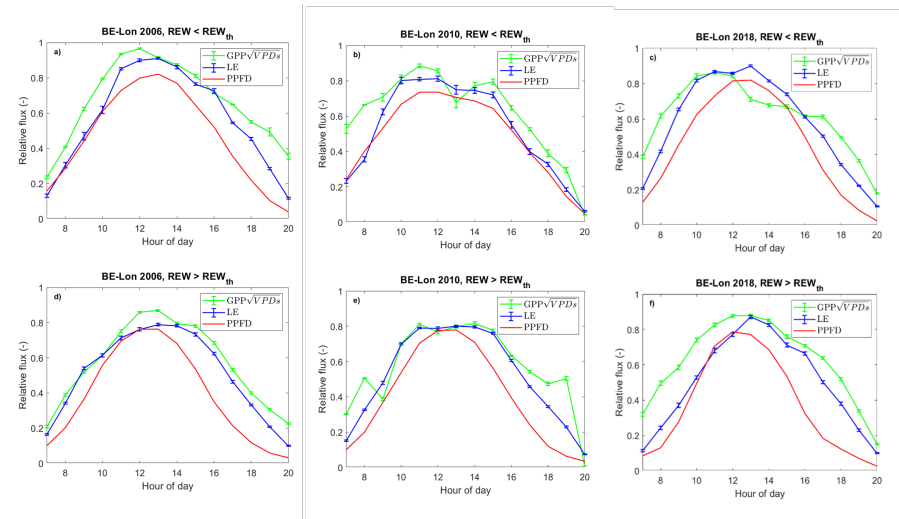
**Figure 4.4:** Sensitivity analysis for the modeling of GPP by setting  $V_{cmax,app,25}$  and  $G_1$  to a constant value when REW is non limiting (respectively ( $GPP_{V_{cmax,app,25}}*$  and  $GPP_{G_1}$ )). The ratios  $GPP_{V_{cmax,app,25}}*/GPP_{sat}$  (subplot a) and  $GPP_{G_1}*/GPP_{sat}$  (subplot b) were only calculated for the years when a decrease in  $V_{cmax,app,25}$  and/or an increase in  $G_1$  were observed. p refers to the p-value comparing the segmented model to a linear model (p-val<sub>mod</sub> in Table 4.3).

		$V_{cmax,app,25}$				$G_1$				
	REW breakpoint	Asymptote ( $V_{cmax,app,25}$ ) or averaged value	Slope	p <sub>val<sub>mod</sub></sub>	R <sup>2</sup>	REW breakpoint	Asymptote ( $G_1$ ) or averaged value	Slope	p <sub>val<sub>mod</sub></sub>	R <sup>2</sup>
	(-)	( $\mu\text{mol m}^{-2}\text{s}^{-1}$ )	( $\mu\text{mol m}^{-2}\text{s}^{-1}$ )	(-)	(-)	(-)	( $\text{kPa}^{0.5}$ )	( $\text{kPa}^{0.5}$ )	(-)	(-)
2006	0.58 ± 0.01 ***	121 ± 2.93 ***	704 ± 94.1 ***	< 0.001	0.82	/	4.35 ± 0.80	/	/	/
2010	0.45 ± 0.02 ***	128 ± 6.11 ***	821 ± 137 **	< 0.001	0.69	0.44 ± 0.05 ***	1.69 ± 0.24 ***	-12.7 ± 5.63 *	< 0.01	0.21
2014	/	80.2 ± 5.25	/	/	/	/	4.44 ± 1.45	/	/	/
2018	0.51 ± 0.01 ***	98.7 ± 1.69 ***	324 ± 36.2 ***	< 0.001	0.44	0.45 ± 0.03 ***	2.37 ± 0.05 ***	-7.27 ± 3.41 *	< 0.05	0.09
Sensitivity analysis on $V_{cmax,app,25}$										
	REW breakpoint	Asymptote	Slope	p <sub>val<sub>mod</sub></sub>	R <sup>2</sup>					
	(-)	(-)	(-)	(-)	(-)					
2006, 2010, 2018	0.46 ± 0.01 ***	1.08 ± 0.02 ***	-11.8 ± 1.31 ***	< 0.001	0.55					
DWCI										
	REW breakpoint	Asymptote	Slope	p <sub>val<sub>mod</sub></sub>	R <sup>2</sup>					
	(-)	(-)	(-)	(-)	(-)					
2006, 2010, 2018	0.45 ± 0.04 ***	0.91 ± 0.01 ***	2.04 ± 0.95 **	< 0.01	0.41					

**Table 4.3:** Statistical properties of the regressions: REW breakpoint, asymptote (or averaged value), slope, model p-value (p<sub>val<sub>mod</sub></sub>) and R<sup>2</sup> for  $V_{cmax,app,25}$ ,  $G_1$ , sensitivity analysis on  $V_{cmax,app,25}$  and for the DWCI. Parameters are given with their standard deviation. For the p-values, \* \*\* : <0.001, \*\* : <0.01, \* : <0.05, ns : >0.05.

### 5.2.4. Coupling between carbon and water fluxes

In 2010 and 2018, NSOL induced a decoupling between  $GPP\sqrt{VPD_s}$  and LE, characterized by a shift of the  $GPP\sqrt{VPD_s}$  maximum toward morning hours (Fig. 4.5). During these years, a clear change in the  $GPP\sqrt{VPD_s}$  dynamics was observed, with a maximum value being reached in the early morning (Fig. 4.6). This decoupling between carbon and water fluxes was not observed in 2006 and 2014, which can be explained by favorable weather conditions in 2014 (absence of water stress) and probably the relatively low intensity of the drought effect on  $V_{cmax,app,25}$  in 2006 ( $V_{cmax,app,25}$  remained higher than  $50 \mu\text{mol m}^{-2} \text{s}^{-1}$ ). DWCI decreased when REW passed below a threshold of  $0.45 \pm 0.04$  (Fig. 6) which was similar to the thresholds for the  $GPP_{V_{cmax,app,25}} * /GPP_{sat}$  ratio and  $V_{cmax,app,25}$  dependence on REW (Table 4.3).



**Figure 4.5:** Latent heat flux (LE), photosynthetic photon flux density (PPFD) and normalized GPP ( $GPP\sqrt{VPD_s}$ ) hourly means and standard deviation when REW is lower than REW breakpoints of segmented linear regressions for  $V_{cmax,app,25}$  (subplots a to c) or when REW is not limiting (subplots d to f) for the three years when non-stomatal limitations were observed (Table 4.3).

## 6. Discussion

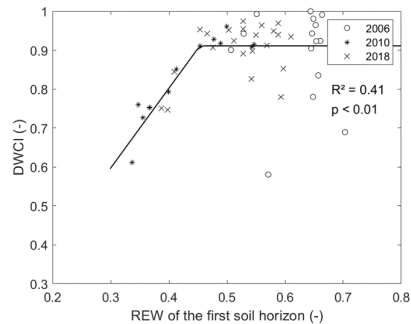
### 6.1. Photosynthesis limitations and methodological considerations

Whatever the year analyzed, we found a global pattern of reduction of photosynthesis capacities due to the influence of soil water availability. Only  $V_{cmax,app,25}$  decreased

with REW, which suggests that the  $\beta$  stress function should only be applied when REW passes below the breakpoint of  $0.46 \pm 0.01$  (for the varieties of potatoes cultivated at BE-Lon and the soil type present on this site). Not implementing water stress effect on  $V_{cmax,app,25}$  would significantly overestimate GPP. This provides strong evidence that NSOL played a major role (or even unique because SOL has not been observed) in carbon assimilation dynamics of potato under high irradiance and during edaphic drought during the four years studied. Moreover, our study was performed at the canopy level which is representative of the physiological reaction of the entire ecosystem to drought since the “big leaf” approach includes all the leaves within the footprint of the EC tower. These conclusions are in accordance with previous studies carried out on forests Gourlez de la Motte *et al.* (2020) and on various non-cropland ecosystems (Chen *et al.*, 2019a), tending to show that NSOL should not be neglected for many plant functional types. At the leaf-level, numerous studies have evaluated the importance of NSOL on photosynthesis using coupled gas-exchange and fluorescence measurements for various crop (Flexas *et al.*, 2008; Wang *et al.*, 2018a) and tree species (Grassi & Magnani, 2005; Perez-Martin *et al.*, 2014; Zait & Schwartz, 2018; Zhu *et al.*, 2021).

While most of the studies conducted on potato aimed at identifying the physiological response to imposed drought through the dynamics of photosynthetic capacities or stomatal closure (Ramírez *et al.*, 2016; Li *et al.*, 2017; Boguszewska-Mañkowska *et al.*, 2018; Silva-Díaz *et al.*, 2020; Aliche *et al.*, 2020), the calibration of the  $\beta$  stress function during edaphic water stress for this crop is lacking. Our study provides a function of REW that can be directly translated to models of photosynthesis in LSMs.

The decrease in  $V_{cmax,app,25}$  illustrates an alteration of the photosynthesis capacities of the crop, either by a restriction of CO<sub>2</sub> diffusion through the mesophyll, real biochemical alteration of the photosynthetic apparatus, or both. Excessive diffusive limitation of CO<sub>2</sub> by the mesophyll may lead to an accumulation of ATP and NADPH which can no longer be used by the Calvin cycle, causing energy imbalance, oxidative stress and damage to cell components (Pinheiro & Chaves, 2011). Moreover, biochemical alteration of photosynthetic capacities can be caused by photo-inhibition and/or irreversible inactivation of PSII, as has been identified during prolonged water stress (Obidiegwu, 2015). Our study did



**Figure 4.6:** Diurnal water carbon index (DWCI) daily dynamics in function of the relative extractable water (REW).  $p$  refers to the  $p$ -value comparing the segmented model to a linear model ( $p$ -val<sub>mod</sub> in Table 4.3).

not allow the identification of the NSOL causes as techniques used to quantify mesophyll conductance (involving carbon isotopes or fluorescence) were not tested. No global application of a model would require a heavy calibration for all plant functional types, a procedure that has not yet been achieved. Future studies should focus on continuous measurements of water extraction, photosynthetic capacities, mesophyll and stomatal conductance during dry periods and subsequent recovery episodes, so drought impacts on ecosystems can be fully understood and mechanistically modeled (Flexas *et al.*, 2018b). The relatively similar REW breakpoints for all growing seasons suggests that the same amount of water in the root zone was required to ensure that saturated photosynthesis remained unaffected by drought, whatever the timing of drought appearance or the potato varieties. This also indicates that the water needs for potato crops were similar during the vegetative (2006) or reproductive (2010, 2018) stages. This observation coincides with early studies stating that “soil moisture should never be allowed to drop below 50 % of the available range of moisture” for potato crop (Singh, 1969), confirming that potato is sensitive to water stress. Models such as Aquacrop usually calibrate similar functions of REW to simulate drought stress on crop physiology with upper thresholds of 0.6 for canopy expansion, 0.7 for canopy senescence, or 0.55 for stomatal control (Raes *et al.*, 2009). The degree of calibration of these thresholds is defined as minimum and requires further investigations for potatoes (Montoya *et al.*, 2016; Raes *et al.*, 2009). The REW breakpoint of 0.46 proposed by this study could be used to design irrigation management strategies from SWC measurements in surface soil layers to avoid the risk of crop abiotic damage. This information is complementary to high-throughput phenotyping techniques (Musse *et al.*, 2021) for selecting drought-tolerant varieties from the criteria of the lowest REW threshold for maintaining an optimum photosynthetic rate. The generalization of our results could be enhanced by applying the procedure and analysis proposed in this paper on other crop sites of the ICOS network where potato was cultivated, which could provide complementary REW thresholds for other soil types. The method used for calculating REW considers that all the roots have the same ability to extract water from the first soil horizon, whatever their age, diameter, architecture, location, or function. This assumption can be challenged, as it has been shown that the root water uptake can be dependent on root density, morphology, and physiology (Kumar *et al.*, 2015). For example, Stalham & Allen (2004) showed that 5 % of the deepest roots can account for a half of the total water uptake for the potato cultivar Cara. Some models do consider these dependences but are difficult to implement at a plot scale due to the high number of inputs and parameters required. Nonetheless, REW remains a useful indicator to determine the amount of available water in the soil from soil hydraulic or granulometric properties that are available for all the ICOS sites.

Some bias in the results can originate from errors in the estimation of  $C_i$  (Eq. (4.12)), and in the calculation of  $G_{sw}$  by Eq. (4.7). The Penman-Monteith equation assumes that the minimal gas diffusion through stomatal apertures ( $G_0$ ) is set to zero and that

no transfer through the cuticle (in parallel with diffusion through stomata) is considered. However, as the cuticle may represent the main pathway for transpiration during drought (Boyer, 2015a; Duursma *et al.*, 2019), these assumptions could possibly lead to an underestimation of  $G_{sw}$  and then  $C_i$  by Eq. (4.12) (Boyer, 2015a). To counter this potential drawback,  $C_i$  could be directly measured at the leaf level by a gas (Wang *et al.*, 2018b). However, estimating  $G_0$  by data adjustment on the USO model remains challenging due to statistical considerations (Duursma *et al.*, 2019), which explains that  $G_0$  is often set to zero or to a constant value in coupled photosynthesis-stomatal conductance models (Medlyn *et al.*, 2017b; Duursma *et al.*, 2019).

## 6.2. Stomatal conductance modeling during drought

While the empirical Ball-Berry (Ball *et al.*, 1987) stomatal conductance model has been used as a baseline in climate models since the mid-1990s (Bonan *et al.*, 2014), many LSMs currently use the USO model (Eq. (4.13)) to estimate  $G_{sw}$  (Kala *et al.*, 2015a; Lawrence *et al.*, 2019). However, the calibration of the response of  $G_1$  to edaphic water stress remains a key element for improving model predictions in a climate change context. Although most models downregulate  $G_1$  with soil moisture (following the stomatal optimality theory; Mäkelä (1996)), numerous studies have observed a different  $G_1$  dependence to soil moisture using leaf-level or ecosystem fluxes data (Zhou *et al.*, 2013; Chen *et al.*, 2019a; Gourlez de la Motte *et al.*, 2020). In line with these studies, our results showed that  $G_1$  remained constant (or even increased) with the decrease in REW for potato. This observation suggests the absence of regulation of stomata opening during drought and indicates a deviation from the optimality theory as the increase of  $G_1$  for low REW corresponds to a rise of the water cost of carbon gain. This behavior can be explained by a decrease in  $V_{cmax,app,25}$  (inducing a decline of the carbon gain through the alteration of efficiency of RuBisCO to fix CO<sub>2</sub> in the Calvin cycle) more important than the corresponding increase in the water transport capacity. This hypothesis is supported by Aliche *et al.* (2020), who observed an adaptation of the hydraulic architecture of four potato cultivars in response to dehydration through an increase in xylem flux density and an increase in the ratio between small and large xylem vessels, preventing from cavitation and facilitating the transport of water and assimilates. When  $V_{cmax,app,25}$  was relatively low, NSOL induced a reduction of  $GPP_{sat}$  leading to a decrease in  $G_{sw}$  computed by Eq. (4.13), but partly compensated by an increase in  $G_1$ . This may explain the decoupling between  $GPP\sqrt{VPD_s}$  and LE intra-day dynamics with decreasing REW and emergence of NSOL which induced an inhibition of photosynthetic capacities independently from transpiration. Similar results were presented by Nelson *et al.* (2018), who reported a decrease of DWCI during the 2003 drought in Puéchabon and for dry savanna/grassland plant functional types during dry episodes.



## 7. Conclusion

In this study, a “big leaf” approach including the FvCB model for photosynthesis and the USO model for stomatal conductance was used to infer the physiological properties of a potato crop for four consecutive growing seasons at the ICOS site of Loncée in Belgium. Drought induced a decrease in  $V_{cmax,app,25}$  with similar REW breakpoints and unstressed values for all years, which highlighted a uniform pattern of drought effects on GPP. A global value of 46 % of available soil water was obtained from a sensitivity analysis, which highlights the importance of implementing REW effects on  $V_{cmax,app,25}$  to reproduce drought effects on GPP. The coupling of this representation with a constant value or a linear increase with REW of the slope parameter ( $G_1$ ) in USO model was sufficient to simultaneously reproduce both carbon and water flux dynamics at seasonal and diurnal scales during drying-up episodes. The inclusion of the effects of drought on NSOL in LSMs for potatoes is recommended to properly implement the effects of drought on water and carbon cycles.

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### **CRedit authorship contribution statement**

**Quentin Beauclaire:** Conceptualization, Formal analysis, Software, Writing – original draft, Writing – review and editing, Validation. **Bernard Heinesch:** Resources, Writing – review and editing. **Bernard Longdoz:** Resources, Validation, Supervision, Writing – review and editing.

### **Declaration of Competing Interest**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

### **Data availability**

Data will be made available on request.

### **Acknowledgments and funding**

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## 8. Supplementary materials

Date	Farming activities	
11-Apr-06	Application	Asparagite (0N/0P/25K +10% calcium oxide + 30% sulphuric oven-dry)
1-May-06	Planting	Potato: Spunta/Primura/Kondor/Draga
2-May-06	Application	Liquid Nitrogen (117 Units ha <sup>-1</sup> )
	Weeding + Pre emergence treatment	Defi (4l ha <sup>-1</sup> ) + Linuron (1l ha <sup>-1</sup> )
13-May-06	Earthing up	
14-May-06	Weeding + Pre emergence treatment	Defi (4l ha <sup>-1</sup> ) + Linuron (1l ha <sup>-1</sup> )
1-Jun-06	Application	Insecticide and fungicide treatments
9-Jun-06		
15-Jun-06		
20-Jun-06		
27-Jun-06		
3-Jul-06		
8-Jul-06		
19-Jul-06		
6-Aug-06	Topkilling	Purivel (2.5 kg ha <sup>-1</sup> ) + Spotlight (0.25 l ha <sup>-1</sup> )
13-Aug-06	Topkilling	Purivel (2.5 kg ha <sup>-1</sup> ) + Reglone (2 l ha <sup>-1</sup> )
15-Sep-06	Harvest	Spunta/Primura/Kondor/Draga
12-Apr-10	Application	Asparagite (0N/0P/25K +10% calcium oxide + 30% sulphuric oven-dry)
24-25-Apr-10	Planting	Potatoes : Draga
5-May-10	Application	Liquid Nitrogen (97.5 Units ha <sup>-1</sup> )
8-May-10	Definitive ridging	
11-May-10	Weeding + Pre emergence treatment	Defi (4l ha <sup>-1</sup> ) + Linuron (1l ha <sup>-1</sup> )
3-Jun-10	Weeding	Sencor (0.15 kg)
3/7/13/22/29-Jun10	Insecticide and Fungicide treatments	Sumi Alfa (Insect.) / Revus (Fung.) + Sumi Alfa (Insect.) / Mancozeb (Fung.) + Tepeki

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		(Insect) / Dequiman (Fung.) + Sumi Alfa (Insect.) / Tepeki (Insect.)
6/13/23-Jul-10	Insecticide and Fungicide treatments	Tepeki (Insect.) / Karaté (Insect.) + Acrobat (Fung) / Ninja (Insect.)
3-Aug-10	Insecticide treatment	Fastac
10-Aug-10	Chemical haulm destruction	Reglone + Spotlight
13-Aug-10	Chemical haulm destruction	Reglone + Spotlight
3-4-5-Sept-10	Harvest	
2-Apr-14	Application	Liquid Nitrogen (183.3 l ha <sup>-1</sup> , 71.5 Units N ha <sup>-1</sup> )
5-7- Apr -14	Soil preparation, planting and earthing up	Potato : Draga/Bianchidea/Kennebec
5-7- Apr -14	Disinfection	Monarch (0.2 l t <sup>-1</sup> ) + Mirador (2.22l ha <sup>-1</sup> )
8- Apr -14	Weeding	Sencor (0.4 l ha <sup>-1</sup> ) + Centium (0.25 l ha <sup>-1</sup> ) + Reglone (on <i>Fumaria officinalis</i> ) (1.5 l ha <sup>-1</sup> )
12-May-14	Insecticide	Lambda 50 EC (0.2 l ha <sup>-1</sup> ) + Mineral Oil (5l ha <sup>-1</sup> )
12-May-14	Weeding	Titus (40 g ha <sup>-1</sup> )
17-May-14	Insecticide	Lambda 50 EC (0.2 l ha <sup>-1</sup> ) + Mineral Oil (5l ha <sup>-1</sup> )
23-May-14	Insecticide	Lambda 50 EC (0.2 l ha <sup>-1</sup> ) + Mineral Oil (5l ha <sup>-1</sup> )
29-May-14	Insecticide	Lambda 50 EC (0.2 l ha <sup>-1</sup> ) + Mineral Oil (5l ha <sup>-1</sup> )
29-May-14	Fungicide	Mancozèbe (2 kg ha <sup>-1</sup> )
6-Jun-14	Fungicide	Acrobat (2kg ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )
6- Jun -14	Insecticide	Sumi Alpha (0.2 l ha <sup>-1</sup> )
11- Jun -14	Fungicide	Acrobat (2.5 kg ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )
11- Jun -14	Insecticide	Sumi Alpha (0.2 l ha <sup>-1</sup> )
11- Jun -14	Fungicide	Manfil (2.5 kg ha <sup>-1</sup> )
17- Jun -14	Fungicide	Acrobat (2 kg ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )
17- Jun -14	Insecticide	Sumi Alpha (0.2 l ha <sup>-1</sup> )
23- Jun -14	Fungicide	Acrobat (2 kg ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )

## Stomatal and non-stomatal responses of typical temperate C3 crops to soil water stress

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23- Jun -14	Insecticide	Okapi (1 l ha <sup>-1</sup> )
30- Jun -14	Fungicide	Ranman (Component A: 0.2 l ha <sup>-1</sup> , Component B: 0.15 l ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )
30- Jun -14	Insecticide	Okapi (1 l ha <sup>-1</sup> )
7-Jul-14	Fungicide	Ranman (Component A: 0.2 l ha <sup>-1</sup> , Component B: 0.15 l ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )
15-Jul-14	Topkilling	Reglone (2 l ha <sup>-1</sup> )
15-Jul-14	Fungicide	Shirlan (0.5 l ha <sup>-1</sup> )
19-Jul-14	Topkilling	Spotlight (1 l ha <sup>-1</sup> )
19-Jul-14	Fungicide	Ranman (Component A: 0.2 l ha <sup>-1</sup> , Component B: 0.15 l ha <sup>-1</sup> ) + Mineral Oil (5 l ha <sup>-1</sup> )
21-22-Aug-14	Harvest	Draga/Bianchidea/Kennebec
12-Apr-18	Application	Solid (N 0, P 7, K 36 à 350 kg/ha + 5 MGO)
22-23-Apr-18	Planting	Potato: Agria/Unidea/Draga
9-10-May-18	Earthing up	
9-10-May-18	Application	Liquid Nitrogen (200 l ha <sup>-1</sup> )
13-May-18	Weeding	Promane (2 l ha <sup>-1</sup> ) + Défi (4 l ha <sup>-1</sup> ) + Centium (0.135 l ha <sup>-1</sup> )
23-May-18	Treatment	Mineral oil (3 l ha <sup>-1</sup> )
26-May-18	Treatment	Mineral oil (4 l ha <sup>-1</sup> )
31-May-18	Treatment	Mineral oil (8 l ha <sup>-1</sup> )
31-May-18	Insecticide	Teppeki (0.15 l ha <sup>-1</sup> )
1-2-Jun-18	Irrigation	Cv. Draga and Unidea
4-Jun-18	Treatment	Mineral oil (8 l ha <sup>-1</sup> )
4-Jun-18	Fungicide	Mancozeb (2 kg ha <sup>-1</sup> )
8-9-10-Jun-18	Irrigation	Draga/Unidea/Agria
13-Jun-18	Treatment	Mineral oil (8 l ha <sup>-1</sup> )
13-Jun-18	Fungicide	Prozeb (1.7 kg ha <sup>-1</sup> ) + Revus (0.4 l ha <sup>-1</sup> )
13-Jun-18	Insecticide	Sumi-alpha (0.2 l ha <sup>-1</sup> )
19-Jun-18	Treatment	Mineral oil (8 l ha <sup>-1</sup> )

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19-Jun-18	Fungicide	Cymopur (0.25 l ha <sup>-1</sup> ) + Ranman (0.4 l ha <sup>-1</sup> )
19-Jun-18	Insecticide	Karaté (0.07 l ha <sup>-1</sup> )
20-21-Jun-18	Irrigation	Cv. Unidea
26-Jun-18	Treatment	Mineral oil (8 l ha <sup>-1</sup> )
26-Jun-18	Fungicide	Cymopur (0.2 l ha <sup>-1</sup> ) + Prozeb (2 kg ha <sup>-1</sup> )
26-Jun-18	Insecticide	Tepeki (0.15 l ha <sup>-1</sup> )
7-Jul-18	Treatment	Mineral oil (10 l ha <sup>-1</sup> )
7-Jul-18	Insecticide	Tepeki (0.15 l ha <sup>-1</sup> )
19-Jul-18	Treatment	Mineral oil (10 l ha <sup>-1</sup> )
19-Jul-18	Fungicide	Prozeb (2 kg ha <sup>-1</sup> )
21-Jul-18	Topkilling	Reglone (1.5 l ha <sup>-1</sup> ) + Spotlight (0.8 l ha <sup>-1</sup> )
26-Jul-18	Topkilling	Reglone (1.5 l ha <sup>-1</sup> ) + Spotlight (0.8 l ha <sup>-1</sup> )
3-Aug-18	Crushing	Cv. Agria
4-Aug-18	Topkilling	Reglone (1 l ha <sup>-1</sup> ) + Spotlight (0.8 l ha <sup>-1</sup> ) : Agria and Unidea
8-Aug-18	Topkilling	Reglone (1 l ha <sup>-1</sup> ) + Spotlight (0.8 l ha <sup>-1</sup> ) : Agria, Unidea and Draga
22-Aug-18	Harvest	Draga/Unidea
11-Sep-18	Harvest	Agria

**Table 4.4:** Farming activities during the 2006, 2010, 2014 and 2018 growing seasons of potato crop at BE-Lon.

Symbol	Description	Value [Units]
<b>REW CALCULATION</b>		
$\bar{\theta}$	Averaged soil water content	[-]
$\theta_{wp}$	Soil water content at the wilting point	[-]
$\theta_{fc}$	Soil water content at the field capacity	[-]
$REW$	Relative extractable water	[-]
<b>EDDY COVARIANCE FLUXES AND "BIG-LEAF" MODEL</b>		
$GPP$	Gross Primary Production	$[\mu\text{mol m}^{-2}\text{s}^{-1}]$
$LE$	Latent heat flux	$[\text{W m}^{-2}]$
$H$	Sensible heat flux	$[\text{W m}^{-2}]$
$G$	Ground heat flux	$[\text{W m}^{-2}]$
$NEE$	Net Ecosystem Exchange	$[\mu\text{mol m}^{-2}\text{s}^{-1}]$
$G_{ah}$	Aerodynamic conductance for sensible heat transfer	$[\text{m s}^{-1}]$
$G_{aw}$	Aerodynamic conductance for water vapor	$[\text{m s}^{-1}]$
$C_s$	CO <sub>2</sub> concentration in the air at the canopy surface	$[\mu\text{mol mol}^{-1}]$
$C_i$	CO <sub>2</sub> concentration in the substomatal cavities	$[\mu\text{mol mol}^{-1}]$
$G_{sw}$	Canopy conductance to water vapor transfer	$[\text{mol m}^{-2}\text{s}^{-1}]$
$K_m$	Michaelis-Menten coefficient	$[\mu\text{mol mol}^{-1}]$
$\Gamma^*$	CO <sub>2</sub> compensation point	$[\mu\text{mol mol}^{-1}]$
$V_{cmax,app,25}$	Apparent carboxylation rate of the Rubisco at 25°C	$[\mu\text{mol m}^{-2}\text{s}^{-1}]$
$V_{cmax,app,25}^*$	Apparent carboxylation rate of the Rubisco at 25°C in absence of water stress	$[\mu\text{mol m}^{-2}\text{s}^{-1}]$
$G_1$	Slope parameter in the USO model	$[\text{kPa}^{0.5}]$
$G_0$	Minimum canopy conductance in the USO model (neglected)	$[\text{mol m}^{-2}\text{s}^{-1}]$
$G_1^*$	Slope parameter in the USO model in absence of water stress	$[\text{kPa}^{0.5}]$
$GAI_{sat}$	Threshold for the non-influence of GAI on GPP under high irradiance	[-]

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$PPFD_{sat}$	Threshold for the non-influence of PPFD on GPP	$[\mu\text{mol m}^{-2}\text{s}^{-1}]$
$DWCI$	Diurnal water-carbon index	$[-]$
<b>ENVIRONMENTAL CONDITIONS</b>		
$R_n$	Net radiation	$[\text{W m}^{-2}]$
$C_a$	$\text{CO}_2$ concentration in the air	$[\mu\text{mol mol}^{-1}]$
$e_a$	Air vapor pressure	$[\text{kPa}]$
$e_s$	Canopy vapor pressure	$[\text{kPa}]$
$T_a$	Air temperature	$[\text{K}]$
$T_s$	Canopy temperature	$[\text{K}]$
$u^*$	Air friction velocity	$[\text{m s}^{-1}]$
$u$	Wind speed	$[\text{m s}^{-1}]$
$VPD_s$	Vapor pressure deficit at the canopy surface	$[\text{kPa}]$
$\Delta$	Slope of saturated vapor pressure curve	$[\text{kPa K}^{-1}]$
$p$	Air pressure	$[\text{kPa}]$
<b>CONSTANTS</b>		
$c_p$	Heat capacity of dry air	$c_p = 1005$ $[\text{J kg}^{-1} \text{K}^{-1}]$
$\rho_a$	Air density	$\rho_a = 1.23$ $[\text{kg m}^{-3}]$
$\gamma$	Psychrometric constant	$\gamma = 0.067$ $[\text{kPa K}^{-1}]$
$\varepsilon$	Emissivity of the canopy in the infrared	$\varepsilon = 0.97$ $[-]$
$\sigma$	Stefan-Boltzmann constant	$\sigma = 5.67 \cdot 10^{-8}$ $[\text{W m}^{-2} \text{K}^{-1}]$
$R$	Perfect gas constant	$R = 8.314$ $[\text{J mol}^{-1} \text{K}^{-1}]$
<b>ACRONYMS</b>		

DAE	Day after emergence	$[-]$
GAI	Green leaf area index	$[-]$
SWC	Soil water content	$[-]$

**Table 4.5:** List of variables used in this study





## **New method to partition the origin of photosynthesis limitations of potato under soil water availability-limiting conditions**



## 1. Personal contributions

The setup of the drought experiment was realized by myself, with the help of Alain Debacq (former ICOS station technical team) and the AgricultureIsLife (AIL) research platform of the University of Liege. Leaf-level measurements were collected by myself with the help of Florian Vanden Brande during his master thesis. Soil samples for building soil water retention curves were collected by the EnvironmentIsLife (EIL) research platform of the University of Liege. The manuscript was reviewed by Bernard Longdoz, who also actively helped in setting the experiment and interpreting the results. Florian Vanden Brand also reviewed the manuscript. This manuscript is in preparation.

## 2. Abstract

The identification of the physiological processes limiting carbon assimilation under water stress is crucial for selecting drought-tolerant varieties and ensuring crop productivity. However, the influence of soil water availability on the limitation of photosynthesis is still not fully understood. This study aimed to investigate the origins of photosynthesis limitations for potato (*Solanum Tuberosum*) during a field drought experiment. Gas exchange and chlorophyll fluorescence measurements were performed at the leaf level on irrigated and non-irrigated plants to determine the response of CO<sub>2</sub> assimilation processes to the decrease in the REW in the soil. For the first time, a limitation analysis was performed to quantitatively assess the importance of each factor in limiting photosynthesis and stomatal conductance of potato. The decrease in REW induced a closure of stomata and a restriction of CO<sub>2</sub> diffusion by the mesophyll, which explained the early reduction of light-saturated photosynthesis. Biochemical capacities decreased only from severe soil water restrictions, which led to an increase in the CO<sub>2</sub> concentration in substomatal cavities and an increase in the stomatal conductance sensitivity to photosynthesis. We revisited the limitation analysis equations and showed that mesophyll conductance was the most important constraint on carbon and water exchanges regardless of soil water conditions. We emphasize the need to modify the partitioning methods to fully unravel the physiological controls on photosynthesis and stomatal conductance under water stress.

## 3. Introduction

European ecosystems are facing more intense and frequent water stress events due to altered rainfall patterns and rising temperature induced by anthropogenic climate change (Samaniego *et al.*, 2018). Precipitation shortage episodes perturbate plant water status and induce disruptions of the water and carbon cycles through the inhibition of carbon assimilation and transpiration (Bertolino *et al.*, 2019; Fahad *et al.*, 2017;

Trenberth *et al.*, 2014). As a result, ecosystems services such as food production and carbon storage are strongly impacted by lack of soil water (Chang & Bonnette, 2016; Hendrawan *et al.*, 2022; Kang *et al.*, 2021). Land-atmosphere feedbacks originating from the perturbation of such processes such as drought and heatwave may exacerbate climate change through water stress intensification (Anderegg *et al.*, 2019; Hartick *et al.*, 2022). An in-depth understanding of the effects of drought on ecosystems physiological functioning is required to predict future ecosystem service capacities and to improve climate models predictions (Ryu *et al.*, 2019).

Photosynthesis is the process by which plants convert CO<sub>2</sub> into carbohydrates. CO<sub>2</sub> supply to the photosynthetic system is a key factor influencing carbon assimilation and is strongly impacted by the physiological barriers on its diffusional pathway such as stomatal opening and diffusion within the mesophyll (Gago *et al.*, 2020; Nadal & Flexas, 2018). In addition, light-saturated photosynthesis is constrained by the efficiency of the Rubisco enzyme for fixing CO<sub>2</sub> in the Calvin cycle (Farquhar *et al.*, 1980). Uncertainties remain on the importance of each limiting factor under soil water-limiting conditions across PFTs (Rogers *et al.*, 2017).

Quantifying the limitations of drought on photosynthesis is also pivotal for assessing phenotype plasticity and for serving as a guide for the selection of drought-tolerant varieties (Lupo & Moshelion, 2024; Nguyen *et al.*, 2023). To that end, mechanistic modeling can be used for disentangling the complexity of the mechanisms regulating plant response to water stress (Stirbet *et al.*, 2020). In the FvCB model (Farquhar *et al.*, 1980), carbon assimilation under high irradiance ( $A_{sat}$ ) is mediated by the response of stomatal conductance ( $g_s$ ), mesophyll conductance ( $g_m$ ), and the maximum carboxylation rate of Rubisco ( $V_{cmax}$ ). The quantitative contribution of each of these factors in limiting photosynthesis under water stress can be estimated by first, developing the total derivative of  $A_{sat}$  as a sum of the total derivative of these factors and second, by estimating the dependence of these terms to soil water availability. This method, also known as limitation analysis (Grassi & Magnani, 2005; Jones, 1985), can be used to partition photosynthesis limitations between stomatal (i.e., a decrease in  $A_{sat}$  originating from  $g_s$ ) and non-stomatal factors (i.e., a decrease in  $A_{sat}$  originating from  $g_m$  and/or  $V_{cmax}$ ).

Stomata are the gates of CO<sub>2</sub> diffusion and water transpiration at the leaves surface, whose opening is regulated by a complex interplay of abiotic and biotic factors. For instance, it is well known that an increase in VPD drives the closure of stomata through the evaporation of water in the guard cells (McAdam & Brodribb, 2016). In addition, carbon assimilation regulates stomatal opening to balance the CO<sub>2</sub> diffusion with the efficiency of the Calvin cycle (Lemonnier & Lawson, 2023; Wong *et al.*, 1979). A mechanistic formulation of these dependencies has been proposed by Cowan & Farquhar (1977) who hypothesized that stomatal opening is regulated to maximize carbon gains and minimize water losses over a constant time interval under non-limiting soil

water conditions. This optimization theory is at the basis of the USO model where  $g_s$  is a function of VPD, CO<sub>2</sub> concentration at the leaf surface, carbon assimilation and the stomatal sensitivity to photosynthesis (Medlyn *et al.*, 2011a). This latter, which is the slope of USO model ( $g_1$ ), is linked to the water use strategy of the plant by being inversely proportional to the marginal carbon cost of water (Medlyn *et al.*, 2011a). During drying-up episodes, short timescales variations of  $g_1$  can be used as an indicator of plants adaptation strategy. In the framework of the optimality theory, plants can maximize carbon gains (increase in  $g_1$ ), minimize water losses (decrease in  $g_1$ ) or keep the same balance between carbon gains and water losses stable (constant  $g_1$ ). The sensitivity of  $g_1$  to soil water availability strongly depends on plant species and PFTs (H eroult *et al.*, 2013; Beauclaire *et al.*, 2023b; Gourlez de la Motte *et al.*, 2020; Zhou *et al.*, 2013). Although the formulation of the relationship between  $g_s$  and  $A_{sat}$ , and the water cost associated with the opening of stomata are still active research topics in the scientific community (Lamour *et al.*, 2022; Mrad *et al.*, 2019), the USO model has become a reference for representing stomatal behavior in LSMs (Kala *et al.*, 2015b; Lawrence *et al.*, 2019; Sabot *et al.*, 2022).

As  $g_s$  is mediated by carbon assimilation, a decrease in  $g_s$  can also be induced by biochemical or mesophyll limitations which regulate stomatal opening with the mesophyll demand for CO<sub>2</sub> (Lemonnier & Lawson, 2023; Medlyn *et al.*, 2011a; Zhou *et al.*, 2013). As a result,  $g_s$  and  $A_{sat}$  are strongly coupled and stomatal closure can either originate from an optimal stomatal adaptation, or from a disguised effect of mesophyll conductance and/or carboxylation rate of Rubisco (Zhou *et al.*, 2013; Medlyn *et al.*, 2011a). This complicates the identification of the origins of stomatal closure and photosynthesis limitations under water stress. Using  $g_1$  as the key variable for highlighting a stomatal control on photosynthesis allows to identify the feedback effect of non-stomatal factors on stomatal closure by linking photosynthesis limitations to the stomatal optimality theory (Zhou *et al.*, 2013). In addition, coupling the USO and FvCB models in the limitation analysis enables for a quantitative assessment of the effects of  $g_1$ , VPD,  $g_m$  and  $V_{cmax}$  on  $g_s$  and  $A_{sat}$ . To our knowledge, this study is the first to develop this approach. Both limitation schemes (i.e., either using  $g_1$  or  $g_s$  as the key variable for stomatal limitation) will be compared and discussed. In this paper, the limitations of photosynthesis originating from stomatal closure induced by a decrease in  $g_1$  or  $g_s$  are referred to as a SOL while an effect of  $g_m$  and/or  $V_{cmax}$  is referred to as a NSOL (Beauclaire *et al.*, 2023b; Gourlez de la Motte *et al.*, 2020).

SWC is a key eco-hydrological parameter impacting plant metabolism and more globally carbon and water fluxes (Zhou *et al.*, 2021). In particular, lack of soil water triggers complex mechanisms which regulate the water flow in the plant to avoid hydraulic failure (Mart inez-Vilalta *et al.*, 2014). When the soil edaphic proprieties are known, SWC can be used to determine the REW for plant uptake (Granier *et al.*, 2007a), which is often used in LSMs as a drought index to implement water stress effects originating either from SOL and/or NSOL on photosynthesis (Vidale *et al.*,

2021). The response of FvCB and USO model parameters to decreasing soil water availability is likely species and PFTs-specific, which makes REW a critical variable for modeling the response of terrestrial ecosystems to drought (Peters *et al.*, 2018; Rogers *et al.*, 2017; Vidale *et al.*, 2021; Zhou *et al.*, 2013).

Potato is one of the most important crops, providing food for more than one billion people around the world (Lutaladio & Castaldi, 2009). In Europe, more than 400 000 hectares of arable lands are used for potato cultivation (Goffart *et al.*, 2022). This crop is highly sensitive to water stress because of its shallow root system and its inability to extract water from deeper soil layers (Obidiegwu, 2015). In particular, tuber bulking is a critical stage of potato growth as it determines the yield and quality of the harvest (Gervais *et al.*, 2021). Partitioning the limitation components of the photosynthetic flux is crucial for selecting drought-tolerant varieties and ensuring food security. We have implemented this approach during a drought experiment on field-grown potatoes. The goals of this study were: (i) to describe the response of  $A_{sat}$ ,  $g_s$ ,  $g_m$ ,  $g_1$  and  $V_{cmax}$  to the decrease in REW, (ii) to perform a limitation analysis on  $A_{sat}$  by using  $g_s$  or  $g_1$ ,  $g_m$ ,  $V_{cmax}$  and VPD as explanatory variables and finally (iii) to define REW thresholds from which each of these limitations occurred.

## 4. Material and methods

### 4.1. Plant materials and experimental setup

Potato plants were grown on a 4-ha experimental plot located in Belgium, about 50 km southeast of Brussels (50° 33' 47.77"N, 4°42' 46.403"E). This cropland has been used for cultivating chicory, sugar beet and winter wheat. In total, 88 tubers of potato (*Solanum tuberosum*, cv Agria) were planted under a plastic polytunnel greenhouse of 12 m long and 5 m large.

SWC and soil temperature were measured by time domain reflectometers (ML3 ThetaProbe, Delta-T Devices Ltd, Cambridge, UK) placed at 10 cm and 30 cm depth. Under the plastic tunnel, air humidity and air temperature were measured by a resistive platinum thermometer and electrical capacitive hygrometer (HMP155, Vaisala Oyj, Helsinki, FI) at 1.5 m height. The tubers were planted on May 15, 2020, and the first leaves appeared on June 4, 2020, which was considered as the emergence (i.e. DAE of 0). Soil water availability was quantified by calculating the REW of the first soil horizon, where most of the root water uptake of potato is expected to occur (Beauclaire *et al.*, 2023b):

$$REW = \frac{\theta_{H1} - \theta_{wp,H1}}{\theta_{fc,H1} - \theta_{wp,H1}} \quad (5.1)$$

where  $\theta_{wp,H1} = 15.6$  and  $\theta_{fc,H1} = 35.01$  are respectively the wilting point and the field capacity of the first horizon ( $H_1$ : 0-30 cm) and  $\theta_{H1}$  is the SWC measured in  $H_1$ ,

which was considered as the weighted mean of SWC measurements measured at 10 cm and 30 cm depth (with a weight of 2/3 and 1/3 respectively).  $\theta_{wp,H1}$  and  $\theta_{fc,H1}$  were estimated from soil water retention curves by using the VG model (van Genuchten, 1980). Soil samples were collected before the experiment at 15 cm depth (3 replicates) and were saturated for at least 24h in distilled water. The pressure plates method (Richards (1948) – by the ISO 11274 standard) was applied and the measurements of suction head and SWC were recorded.  $\theta_{wp,H1}$  and  $\theta_{fc,H1}$  were estimated as the SWC at a pF (log of the suction head) respectively of 4.2 and 2.0. VG model parameters and retention curves of the three soil samples are given in the supplementary material (Fig. 5.6).

Over a first period of 35 days, all the plants were hand-watered to ensure that  $\theta_{H1}$  remained near field capacity. The drought treatment consisted in withholding irrigation to reproduce a long-term precipitation deficit on half of the plants. The other half was hand-watered during the experiment. The drought treatment started on DAE 40 (corresponding to the beginning of the tuber bulking stage) and stopped on DAE 74 (corresponding to the appearance of the first signs of senescence on the irrigated potato plants). All the plants have experienced the same PAR, temperature and VPD conditions under the plastic tunnel.

## 4.2. Leaf-level measurements

Gas exchange and chlorophyll fluorescence measurements were conducted during the tuber bulking stage at 14 different dates (between DAE 35 to DAE 74, Fig. 5.1) from 10 am to 4 pm. Only the youngest leaves in the upper part of the plant were selected by randomly sampling irrigated and non-irrigated plants. Measurements were performed by using a LI-COR LI-6400 equipped with a LI-6400-40 fluorescence chamber (LI-COR Inc., Lincoln, NE, USA). The following procedure was applied to each leaf sample. The CO<sub>2</sub> concentration in the chamber ( $C_s$ ) was set to 400  $\mu\text{molmol}^{-1}$ , the PPF in the PAR at 1200  $\mu\text{molm}^{-2}\text{s}^{-1}$  and the air humidity and temperature were maintained at ambient levels. After stabilization of the steady state fluorescence signal ( $F_s$ ), a multiphase flash with a saturation light of 9 000  $\mu\text{molm}^{-2}\text{s}^{-1}$  was applied, and the maximum fluorescence intensity under the light ( $F'_m$ ) was measured. In addition,  $A_{sat}$ , leaf temperature, stomatal conductance to water vapor ( $g_{sw}$ ), CO<sub>2</sub> concentration in sub-stomatal cavities ( $C_i$ ) and the vapor pressure deficit at the leaf surface ( $VPD_{leaf}$ ) were recorded. Stomatal conductance to CO<sub>2</sub> ( $g_s$ ) was determined by dividing  $g_{sw}$  by 1.6.

### 4.2.1. Non-stomatal limitations: $V_{cmax}$ and $g_m$

$V_{cmax}$  was determined by using a single measurement of gas exchanges at light saturation (Wilson *et al.*, 2000; De Kauwe *et al.*, 2016):

$$V_{cmax} = A_{sat} \frac{C_c + K_m}{C_c - \Gamma^*} \quad (5.2)$$

with  $K_m$  the Michaelis-Menten coefficient,  $\Gamma^*$  the  $\text{CO}_2$  compensation point and  $C_c$  the  $\text{CO}_2$  concentration in the chloroplast. Eq. 5.2 is based on a single measurement of  $\text{CO}_2$  assimilation at light saturation instead of using  $\text{CO}_2$ -response curves where  $V_{cmax}$  retrieval is impacted by the sensitivity of the fitting method (Miao *et al.*, 2009). Moreover, the leaf respiration ( $R_d$ ) was neglected, as it is much smaller compared to  $A_{sat}$  (Knauer *et al.*, 2018a; Von Caemmerer, 2013).  $K_m$  and  $\Gamma^*$  were estimated by using C3 plant-based temperature response curves (Bernacchi *et al.*, 2001).  $C_c$  was calculated by using the Fick law (Farquhar & Sharkey, 1982):

$$C_c = C_i - \frac{A_{sat}}{g_m} \quad (5.3)$$

where  $g_m$  is determined by using the “variable electron transport” method (Harley *et al.*, 1992):

$$g_m = \frac{A_{sat}}{C_i - \frac{\Gamma^*(J_F + 8A_{sat})}{J_F - 4A_{sat}}} \quad (5.4)$$

with  $J_F$  the electron transport rate estimated from PAR, leaf absorptance in the PAR ( $\alpha$ ) photochemical efficiency of PSII open centers  $\phi_{PSII}$ , and the fraction of the absorbed PAR allocated to PSII ( $\beta$ ) (Genty *et al.*, 1989; Valentini *et al.*, 1995):

$$J_F = \alpha\beta\phi_{PSII}PAR \quad (5.5)$$

In Eq. 5.5,  $\phi_{PSII}$  was determined from  $F'_m$  and  $F_s$  (Kramer *et al.*, 2004):

$$\phi_{PSII} = \frac{F'_m - F_s}{F'_m} \quad (5.6)$$

and  $\alpha\beta$  was determined from the linear relationship between  $\phi_{PSII}$  and the apparent quantum efficiency of the linear electron transport  $\phi_{e-}$  (Valentini *et al.*, 1995):

$$\alpha\beta = \frac{4}{k} \quad (5.7)$$

where 4 is the number of electrons needed per  $\text{CO}_2$  molecule fixed and  $k$  the slope of the linear relationship between  $\phi_{e-}$  and  $\phi_{PSII}$ . Under non-photorespiratory conditions,  $\phi_{e-}$  can be estimated by the apparent quantum efficiency of  $\text{CO}_2$  uptake  $\phi_{\text{CO}_2}$ , which is obtained by dividing the measured  $\text{CO}_2$  assimilation by the incident PPF (Genty *et al.*, 1989). Non-photorespiratory conditions were set by adding pure N2



(1% O<sub>2</sub>) to the LI-COR LI-6400 chamber. The meteorological conditions were maintained at ambient levels and the incoming PPFD was set to the following values: 2000, 1500, 1200, 1000, 800, 600, 400, 200, 100 and 0  $\mu\text{molm}^{-2}\text{s}^{-1}$ . Gas exchanges and fluorescence intensities were measured for each PPFD value and  $\phi_{CO_2}$  was calculated as the ratio of net carbon assimilation to *PAR*. The slope of fitted linear relationship between  $\phi_{e-}$  and  $\phi_{CO_2}$  (*k*) was used to determine  $\alpha\beta$  using Eq. 5.7. These measurements were conducted on three leaf samples for irrigated and non-irrigated potato plants and were repeated three times during the drought treatment (i.e., DAE 42, 64 and 73).

#### 4.2.2. Stomatal limitations: $g_1$

In the USO model,  $g_{sw}$  is a function of  $VPD_{leaf}$ ,  $C_s$  and  $A_{sat}$  (Medlyn *et al.*, 2011a):

$$g_{sw} = 1.6 \left( 1 + \frac{g_1}{\sqrt{VPD_{leaf}}} \right) \frac{A_{sat}}{C_s} \quad (5.8)$$

where the minimum stomatal conductance is neglected under high irradiance (Medlyn *et al.*, 2017b), and  $g_1$  is the stomatal sensitivity to photosynthesis, which is inversely related to the marginal WUE.  $g_1$  can be determined by combining the Fick law describing the CO<sub>2</sub> diffusion through stomata with Eq. 5.8, which gives (Medlyn *et al.*, 2017b):

$$g_1 = \frac{\frac{C_i}{C_s} \sqrt{VPD_{leaf}}}{1 - \frac{C_i}{C_s}} \quad (5.9)$$

### 4.3. Statistical analysis

$g_m$  was discarded when  $C_i$  was without the range 150-350  $\mu\text{molmol}^{-1}$ , which minimizes errors in  $R_d$  and  $\Gamma^*$ , and by extension in  $g_m$  (Harley *et al.*, 1992; Niinemets *et al.*, 2006; Veromann-Jürgenson *et al.*, 2017). Moreover,  $V_{cmax}$  and  $g_m$  were normalized at 25°C ( $V_{cmax,25}$ ,  $g_{m,25}$ ) by using an Arrhenius temperature response function parameterized on tobacco (Bernacchi *et al.*, 2001, 2002). Gas exchange and chlorophyll fluorescence-related variables (i.e.,  $A_{sat}$ ,  $g_s$ ,  $g_{m,25}$ ,  $g_1$  and  $V_{cmax,25}$ ) were averaged for each day of measurement and drought treatment (irrigated and non-irrigated), thus regrouping measurements performed under similar meteorological and edaphic conditions.

The response of  $A_{sat}$ ,  $g_s$ ,  $g_{m,25}$ ,  $g_1$  and  $V_{cmax,25}$  to the decrease in REW was assessed by using a linear-plateau model which consists in a constant value ( $y_{max}$ ) and a linear segment (with slope  $a$  and intercept  $b$ ) on either sides of a threshold ( $REW_{th}$ ). Such model has already been used to describe the response of SOL and NSOL to soil

water availability of potato crops at the ecosystem scale (Beauclaire *et al.*, 2023b), and is used to implement the response of LSMs parameters to drought (Vidale *et al.*, 2021). The statistical significance of the linear-plateau model was achieved by comparing its Akaike Information Criterion (AIC) corrected for low sample size (AICc, Burnham *et al.* (2011)) to the one of a higher parsimonious model (i.e., a linear model with one slope and intercept). The model with the lowest AICc was selected (Burnham *et al.*, 2011; Scoffoni *et al.*, 2012). Differences between models were considered as meaningful when their AICc differed by at least 7 (Burnham *et al.*, 2011). If the difference was less than 7, the segmented model was selected as such pattern has already been observed for potato (Beauclaire *et al.*, 2023b). Model performance was assessed by the  $R^2$  and the standard deviation of fitted parameters. The linear-plateau regression was fitted using the “nlsm” function from the “nlraa” package in R Studio (Archontoulis & Miguez, 2015; Miguez, 2023; R Core Team, 2020). Statistical difference between  $REW_{th}$  parameters was tested by calculating the p-value of a t-test using the fitted values and their corresponding standard deviation (Clogg *et al.*, 1995; Paternoster *et al.*, 1998).

#### 4.4. Limitation analysis

The first limitation scheme used in this study has been proposed by Jones (1985) where SOL were associated to a decrease in  $g_s$  either caused by a decrease in  $V_{cmax}$  or  $g_m$ . The relative variation of  $A_{sat}$  compared to its maximum value ( $dA_{sat}/A_{sat}$ ) is written as the sum of the relative variations of  $g_s$ ,  $g_m$  and  $V_{cmax}$  following (Grassi & Magnani, 2005; Jones, 1985):

$$\frac{dA_{sat}}{A_{sat}} = \frac{dg_s}{g_s} l_{gs} + \frac{dg_m}{g_m} l_{gm} + \frac{dV_{cmax}}{V_{cmax}} l_{V_{cmax}} = L_{gs} + L_{gm} + L_{V_{cmax}} \quad (5.10)$$

$$l_{gs} = \frac{\frac{g_t}{g_s} \frac{\delta A_{sat}}{\delta C_c}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.11)$$

$$l_{gm} = \frac{\frac{g_t}{g_m} \frac{\delta A_{sat}}{\delta C_c}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.12)$$

$$l_{V_{cmax}} = \frac{g_t}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.13)$$

where  $l_{gs}$ ,  $l_{gm}$  and  $l_{V_{cmax}}$  are respectively the relative stomatal, mesophyll and biochemical limitations (corresponding to dimensionless quantity between 0 and 1 that gives the proportion of the total limitation) while  $L_{gs}$ ,  $L_{gm}$  and  $L_{V_{cmax}}$  are the contributions of respectively the stomatal, mesophyll and biochemical limitations to the relative variation of  $A_{sat}$ .  $g_t$  is the total conductance to  $CO_2$  ( $(g_t)^{-1} = (g_s)^{-1} + (g_m)^{-1}$ )

and  $\delta A_{sat}/\delta C_c$  is the partial derivative of  $A_{sat}$  with respect to  $C_c$  using Eq. 5.2. In this study, Eq. 5.10 has been further normalized by  $dA_{sat}/A_{sat}$  to improve the interpretation of the data. The temporal dynamics of these relative variations can be explained solely by REW and VPD as the relationship to temperature has already been considered by normalizing  $V_{cmax}$  and  $g_m$  at 25°C, as well as the one to solar radiation by collecting the data at light saturation.

This approach has two drawbacks. First, the decrease in  $A_{sat}$  originating from stomatal closure through a decrease in  $g_s$  can be induced by  $g_m$  and  $V_{cmax}$ , which may result in the underestimation of the contribution of non-stomatal factors in limiting photosynthesis. Second, identifying the contribution of REW to the variation in  $L_{gs}$  is complex as VPD has varied during the experiment. To tackle these issues, we used  $g_1$  instead of  $g_s$  as SOL which allows, first to separate the feedback effect of NSOL on stomatal conductance, and second to consider the effect of VPD on stomatal closure (Zhou *et al.*, 2013). As a result, Eq. 5.10 has been modified by calculating the total derivative of  $g_s$  using the USO model, which gives (derived in Text S1):

$$\frac{dA_{sat}}{A_{sat}} = \frac{dg_1}{g_1} \left( \frac{l_{gs} C_i}{1 - g_s C_s} \right) + \frac{dg_m}{g_m} \left( \frac{l_{gm}}{1 - g_s} \right) + \frac{dV_{cmax}}{V_{cmax}} \left( \frac{l_{Vcmax}}{1 - g_s} \right) - \frac{dVPD}{VPD} \left( \frac{1}{2} \frac{l_{gs} C_i}{1 - l_{gs} C_s} \right) \quad (5.14)$$

$$\frac{dA_{sat}}{A_{sat}} = \frac{dg_1}{g_1} l_{g1,USO} + \frac{dg_m}{g_m} l_{gm,USO} + \frac{dV_{cmax}}{V_{cmax}} l_{Vcmax,USO} + \frac{dVPD}{VPD} l_{VPD,USO} \quad (5.15)$$

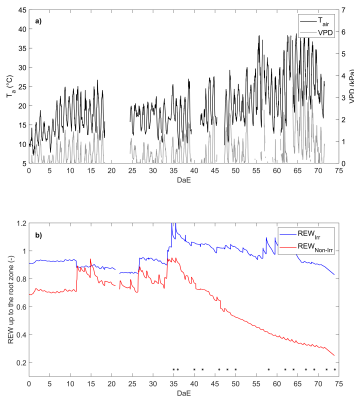
$$\frac{dA_{sat}}{A_{sat}} = L_{g1,USO} + L_{gm,USO} + L_{Vcmax,USO} + L_{VPD,USO} \quad (5.16)$$

where  $L_{g1,USO}$ ,  $L_{gm,USO}$ ,  $L_{Vcmax,USO}$  and  $L_{VPD,USO}$  are the contributions of respectively the optimal stomatal, mesophyll, biochemical and VPD limitations to the relative variation of  $A_{sat}$  using the USO model of stomatal conductance. Eq. 5.14 shows that  $dA_{sat}/A_{sat}$  can be written as the sum of the relative variations of  $g_1$ ,  $g_m$ ,  $V_{cmax}$  and VPD. Combining Eq. 5.10 and Eq. 5.16 allows to identify the effect of  $g_m$ ,  $V_{cmax}$ ,  $g_1$  and VPD on the contribution of stomatal closure to photosynthesis as follows:

$$L_{gs} = L_{gm,USO} - L_{gm} + L_{Vcmax,USO} - L_{Vcmax} + L_{VPD,USO} + L_{g1,USO} \quad (5.17)$$

$L_{gm,USO} - L_{gm} + L_{Vcmax,USO} - L_{Vcmax}$  is the effect of NSOL on stomatal closure while  $L_{VPD,USO} + L_{g1,USO}$  is the effect of VPD and  $g_1$  on stomatal closure according to the USO model. The relative variations in Eq. 5.15 and Eq. 5.10 are calculated from the difference between the value of the variable at a specific REW and the asymptote of the linear-plateau using  $\frac{dy}{y} = \frac{y_{max} - y}{y_{max} - y_{min}}$  with  $y$  being the ordinate at a specific

REW value and  $y_{max}$  the plateau of the segmented regression. In a similar fashion,  $dVPD/VPD$  is determined from the VPD-REW relationship. During precipitation shortage episodes, this relationship is decreasing (i.e., increase in VPD when REW decrease), which has been observed during the experiment for non-irrigated plants (Fig. 5.8). This relationship was confirmed by the data of the nearby eddy covariance station of Lonzé for similar edaphic proprieties (data not shown). Therefore, this linear relationship has been used to determine  $dVPD/VPD$  at each REW value.



**Figure 5.1:** Temporal evolution of air temperature ( $T_a$ ) and VPD under the plastic polytunnel greenhouse (panel a), and REW of the irrigated plot ( $REW_{irr}$ ) and non-irrigated plot ( $REW_{non-irr}$ ) (panel b). The asterisk indicates the days when leaf level measurements were conducted.

$A_{sat}$ . For all the other terms of Eq. 16, a decrease in  $V_{cmax}$ ,  $g_m$  or  $g_1$  induces a decrease in  $A_{sat}$ .

## 5. Results

### 5.1. Meteorological and edaphic conditions

The decline in soil water availability was synchronized with a period of progressive increase in VPD and air temperature under the plastic polytunnel greenhouse (Fig. 5.1a) up to a maximum value of 4.10 kPa and 39.02 °C respectively (Fig. 5.1a). Both irrigated and non-irrigated plants faced an increase in atmospheric dryness and air temperature. The REW of the non-irrigated plants decreased after stopping the

In Eq. 14, the ratio  $C_i/C_s$  also plays an important role in the limitation analysis as it directly influences  $L_{g1,USO}$  and  $L_{VPD,USO}$ . Using  $g_1$  as SOL implies that any stomatal constraint on  $A_{sat}$  should be associated to an increase of the ratio  $A_{sat}/g_s$ . Indeed, following the USO model framework, this constraint corresponds to a maximization of photosynthesis while minimizing water losses. As  $g_s$  and  $A_{sat}$  both regulate  $CO_2$  diffusion through stomatal apertures and  $CO_2$  fixation in the chloroplasts, the increase in  $A_{sat}/g_s$  is linked to a decrease in  $C_i/C_s$  illustrating an optimal stomatal control ( $C_i/C_s \sim 1 - A_{sat}/g_s$ ). The relationship between  $C_i/C_s$  and REW was also evaluated by fitting a linear-plateau model as described in 2.4. Note that  $L_{VPD,USO}$  is per essence negative because of the partial derivative of VPD with respect to  $A_{sat}$  as they are inversely related (i.e., VPD at the denominator in the USO model, Eq. 5.8). As a result, any increase in VPD induces a closure of stomata and a decrease in

irrigation and reached 0.24 at the end of the experiment while the REW of the irrigated plants remained higher than 0.83 due to continuous hand watering (Fig. 5.1-b).

## 5.2. Effect of decreasing REW on the dynamics of gas-exchange and fluorescence traits

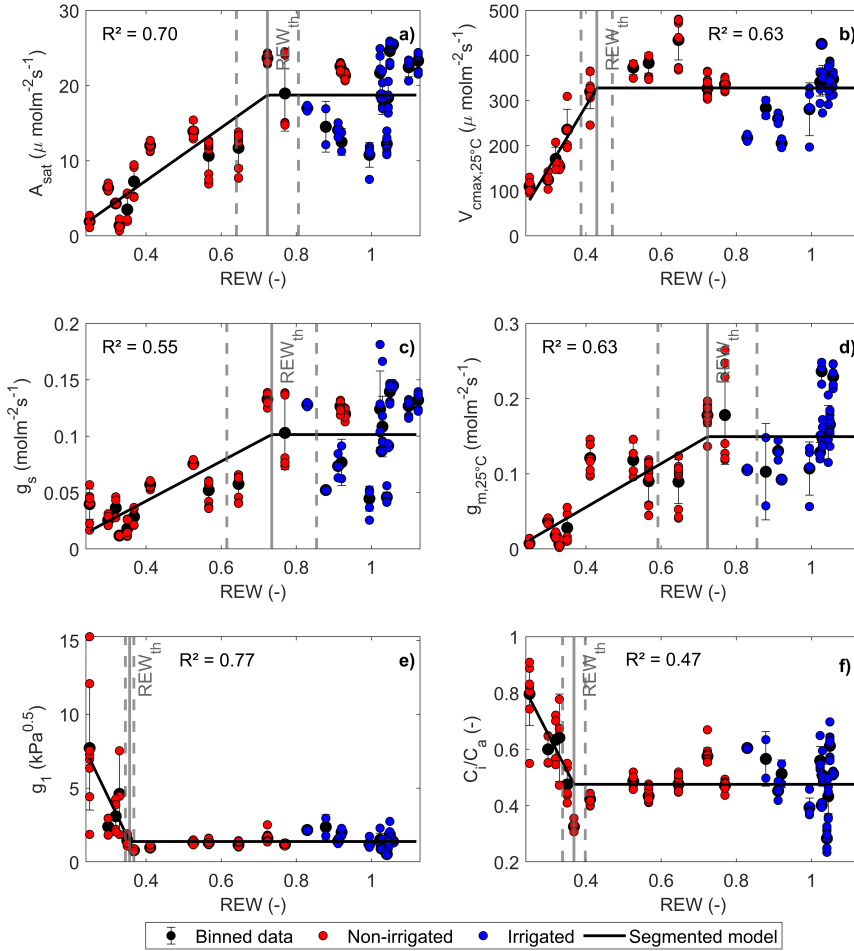
$\alpha\beta$  was not significantly different between irrigated and non-irrigated leaf samples at each DAE, and during the experiment (Fig. 5.7). Therefore, the mean of all  $\alpha\beta_{PSII}$  values was used in Eq. 5.5 (i.e.  $\alpha\beta_{PSII}=0.73 \pm 0.08$ ). The linear-plateau model had the lowest AICc representing the dependence of  $V_{cmax,25}$ ,  $g_1$  and  $C_i/C_s$  on REW. For  $g_s$ ,  $g_{m,25}$  and  $A_{sat}$ , the difference between the AICc of the linear-plateau and the linear model was less than 5.67 (Table 5.3). These differences were not considered as evidence for selecting a model over another (Burnham *et al.*, 2011). Therefore, the segmented model was chosen for reproducing the response of  $A_{sat}$ ,  $C_i/C_s$ ,  $g_s$ ,  $g_{m,25}$  to REW (Fig. 5.2, Table 5.1). The REW threshold at which  $A_{sat}$ ,  $g_s$ ,  $g_{m,25}$  started to decrease were higher than those of  $V_{cmax,25}$ ,  $g_1$  and  $C_i/C_s$  (Fig. 5.2, Table 5.1), which is confirmed by the p-values comparing these parameters (Table 5.2). Overall, CO<sub>2</sub> diffusion factors (i.e.,  $g_{m,25}$  and  $g_s$ ) were the first variables to decrease with REW. Because of a non-significance difference, the REW thresholds for  $g_s$  and  $g_{m,25}$  were averaged, corresponding to  $REW_{th,g_s,g_m} = 0.72 \pm 0.12$ . Biochemical limitation ( $V_{cmax,25}$ ) was only negatively impacted by severe soil water restrictions ( $REW_{th,V_{cmax}} = 0.43 \pm 0.04$ ).  $g_1$  and  $C_i/C_s$  increased from a smaller REW threshold compared to  $V_{cmax,25}$  ( $REW_{th,g_1,C_iC_s} = 0.37 \pm 0.02$ ; Fig. 5.2, Table 5.1, Table 5.2).

	$A_{sat}$ ( $\mu\text{mol m}^{-2} \text{s}^{-1}$ )	$V_{cmax,25}$ ( $\mu\text{mol m}^{-2} \text{s}^{-1}$ )	$g_s$ ( $\text{mol m}^{-2} \text{s}^{-1}$ )	$g_{m,25}$ ( $\text{mol m}^{-2} \text{s}^{-1}$ )	$g_1$ ( $\text{kPa}^{0.5}$ )	$C_i/C_s$ (-)
	$y = \begin{cases} y_{max} & x > REW_{th} \\ ax + b, & x \leq REW_{th} \end{cases}$					
$y_{max}(\pm SD)$	18.74 $\pm$ 1.00	328.04 $\pm$ 14.85	0.10 $\pm$ 0.01	0.15 $\pm$ 0.01	1.39 $\pm$ 0.17	0.48 $\pm$ 0.02
$a (\pm SD)$	35.22 $\pm$ 8.38	1391.7 $\pm$ 494.5	0.18 $\pm$ 0.06	0.29 $\pm$ 0.11	-52.67 $\pm$ 9.83	-2.59 $\pm$ 1.12
$b (\pm SD)$	-6.73 $\pm$ 3.86	-268.4 $\pm$ 163.4	-0.03 $\pm$ 0.03	-0.06 $\pm$ 0.05	20.15 $\pm$ 3.06	1.43 $\pm$ 0.35
$REW_{th} (\pm SD)$	0.72 $\pm$ 0.08	0.43 $\pm$ 0.04	0.73 $\pm$ 0.12	0.72 $\pm$ 0.13	0.36 $\pm$ 0.01	0.37 $\pm$ 0.03
R <sup>2</sup>	0.70	0.63	0.55	0.63	0.77	0.47

**Table 5.1:** Statistics of the linear-plateau regressions for the response of  $A_{sat}$ ,  $V_{cmax,25}$ ,  $g_s$ ,  $g_{m,25}$ ,  $g_1$  and  $C_i/C_s$  to REW. Parameters are given with their standard deviation (SD).

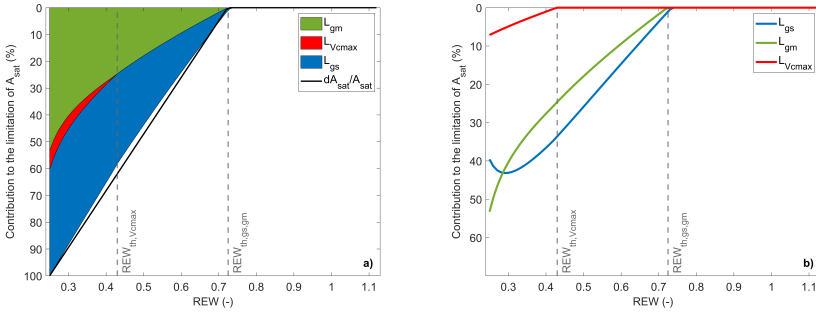
## 5.3. Limitation analysis

The first limitation analysis scheme used in this study consists in partitioning photosynthesis limitations under high irradiance between  $L_{gm}$ ,  $L_{V_{cmax}}$ , and  $L_{gs}$  (Jones, 1985).  $L_{gm}$  was always higher than  $L_{gs}$  above  $REW \sim 0.28$  where  $L_{gm}$  became pre-



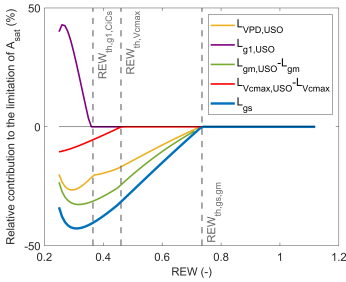
**Figure 5.2:** Response of  $A_{sat}$  (panel a),  $V_{cmax,25}$  (panel b),  $g_s$  (panel c),  $g_{m,25}$  (panel d),  $g_1$  (panel e) and  $C_i/C_s$  (panel f) to relative extractable water (REW). Red and blue dots indicate respectively non-irrigated and irrigated potato plants. The fitted curve represents the linear-plateau regression  $y = \begin{cases} y_{max}, & x > REW_{th}, \\ ax + b, & x \leq REW_{th} \end{cases}$ . Binned data is shown with the corresponding standard deviation (SD). The grey vertical lines indicate  $REW_{th} \pm SD$ .

dominant over  $L_{gs}$  (i.e., intersection of  $L_{gs}$  and  $L_{gm}$ , Fig. 5.3-b). When REW was minimum, 34% of the decrease in  $A_{sat}$  was explained by  $L_{gs}$ , 20% by  $L_{Vcmax}$  and 56% by  $L_{gm}$  (Fig. 5.3-a,b). This limitation scheme indicated that  $CO_2$  diffusion factors (i.e.,  $L_{gm}$  and  $L_{gs}$ ) explained most of the decrease in  $A_{sat}$  with a similar contribution.



**Figure 5.3:** Partitioning of  $A_{sat}$  limitations between  $L_{gs}$ ,  $L_{gm}$  and  $L_{Vcmax}$  in response to REW. Panel a) shows stacked limitation curves and panel b) shows unstacked limitation curves. The black line is the relative variation of  $A_{sat}$  compared to its maximum value (i.e.,  $y_{max}$  of the segmented regression) and the grey vertical lines indicates  $REW_{th} \pm SD$ .

However, using  $g_s$ ,  $g_m$ , and  $V_{cmax}$  in the partitioning analysis does not allow to fully identify the origin of the early stomatal closure, as  $g_s$  itself can be influenced by  $V_{cmax}$  and  $g_m$  through  $A_{sat}$  (Eq. 5.8). This hypothesis is supported by the similar REW threshold for  $g_s$  and  $g_{m,25}$ , which suggests that the two variables are closely related.

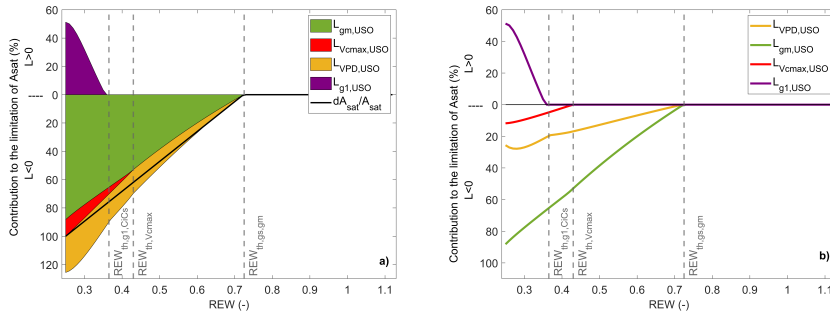


**Figure 5.4:** Partitioning of  $L_{gs}$  into  $L_{g1,USO}$ ,  $L_{gm,USO} - L_{gm}$ ,  $L_{Vcmax,USO} - L_{Vcmax}$  and  $L_{VPD,USO}$  in response to REW.

The grey vertical lines indicates  $REW_{th} \pm SD$ .

The positive contribution of  $g_1$  can be explained by the increase in

$dg_1/g_1$  (Fig. 2-e), which resulted in an increase in  $L_{g1,USO}$  (Eq. 15). Such increase in  $L_{g1,USO}$  was observed from  $REW_{th,g1,CiCs}=0.37 \pm 0.02$  (Table 5.1, Table 5.2) which corresponded to low  $A_{sat}$  ( $6.8 \mu\text{mol}^{-2}\text{s}^{-1}$ ) and  $g_s$  ( $0.04 \text{ mol}^{-2}\text{s}^{-1}$ ). Note that the sum of all curves in Fig. 5.5-b may not necessary equal 1 as the sum of limiting components when using the USO partitioning scheme did not exactly correspond to  $dA_{sat}/A_{sat}$  because of the uncertainties associated with the fitting of the linear-plateau segmented model on measurements (Fig. 5.2, Fig. 5.5-a).



**Figure 5.5:** Partitioning of  $A_{sat}$  limitations between  $L_{g1,USO}$ ,  $L_{gm,SUO}$ ,  $L_{Vcmax,USO}$  and  $L_{VPD,USO}$  in response to REW. Panel a) shows stacked limitation curves and panel b) shows unstacked limitation curves. The black line is the relative variation of  $A_{sat}$  compared to its maximum value (i.e.,  $y_{max}$  of the segmented regression) and the grey vertical lines indicates  $REW_{th} \pm SD$ .

## 6. Discussion

The determination of thresholds of soil water availability impacting  $\text{CO}_2$  assimilation is pivotal for calibrating the response of photosynthesis model parameters during drying-up episodes (Vidale *et al.*, 2021). The results of this study showed that soil water-limiting conditions induced a two-stage response of potato to water stress, with  $g_s$  and  $g_m$  being the first variables impacted by the decrease in REW followed by biochemical limitations through the decrease in  $V_{cmax}$  from a lower REW threshold. In addition, we used a new partitioning scheme where the total derivative of  $g_s$  was written as a function of its explanatory variables in the USO model (i.e.,  $g_1$ ,  $V_{cmax}$ , VPD and  $A_{sat}$ ). This method allowed to quantify the origins of the decrease in  $A_{sat}$  in response to changes in  $g_m$ ,  $V_{cmax}$ ,  $g_1$  and VPD. This partitioning was compared to the original formulation of photosynthesis limitations of Jones (1985) which attributed the origins of the reduction of  $A_{sat}$  to the relative variations of  $g_m$ ,  $V_{cmax}$  and  $g_s$ . The comparison between the two schemes provides an estimation of the importance of the factors influencing  $g_s$  and  $A_{sat}$ .



## 6.1. Limitation on CO<sub>2</sub> diffusion as a major constraint of photosynthesis

Stomatal closure is a well-known mechanism of potato to reduce transpiration under water stress (Gerhards *et al.*, 2016; Gordon *et al.*, 1997; Obidiegwu, 2015; Romero *et al.*, 2017; Vos & Oyarz n, 1987). Stomatal closure dynamics are complex and can be directly caused by the evaporation of the water held by guard cells, or by the loss of turgor pressure induced by sensing of signaling molecules (Bharath *et al.*, 2021; Ding & Chaumont, 2020; Obidiegwu, 2015; Pirasteh-Anosheh *et al.*, 2016; Zhang *et al.*, 2022a).

These mechanisms seem to be synchronized with those influencing the mesophyll conductance as evidenced by a similar REW threshold for  $g_s$  and  $g_m$  (Fig. 5.2). In particular, mesophyll and stomatal conductance share similar responses to abscisic acid (Flexas *et al.*, 2012; Li *et al.*, 2021a; Sorrentino *et al.*, 2016), internal CO<sub>2</sub> concentration (Engineer *et al.*, 2016; Tan *et al.*, 2017) or starch-derived molecules (Lawson *et al.*, 2014), which leads to similar responses during water stress (Flexas *et al.*, 2004; Wang *et al.*, 2018a; Xiong *et al.*, 2018).  $L_{gm}$  became predominant over  $L_{gs}$  under severe water stress, which was associated to a very low  $g_m$  and a strong restriction of CO<sub>2</sub> diffusion to chloroplasts

(Fig. 5.2, Fig 5.3). While this partitioning scheme indicated that photosynthesis limitations originated from  $g_s$  across most of the range of REW, it did not highlight the influence of non-stomatal factors on stomatal conductance. The origins of the decrease in  $g_s$  and  $A_{sat}$  can be identified by using the USO model equation in the limitation analysis. In particular, the USO partitioning scheme showed that most of stomatal closure dynamics can be attributed to a combined effect of  $g_m$  and VPD (Fig. 5.4). More specifically,  $L_{gm,USO}$  was always higher than the other limiting components (Fig. 5.4-b) which highlights the strong control of mesophyll conductance on stomatal closure through its influence on  $A_{sat}$  regardless of REW values. These results confirm the importance of the mesophyll constraint is numerous species in many PFTs (Cano *et al.*, 2013; Flexas *et al.*, 2009; Galmés *et al.*, 2007; Grassi & Magnani, 2005; Limousin *et al.*, 2010; Perez-Martin *et al.*, 2014; Wang *et al.*, 2018b; Zait & Schwartz, 2018; Zhu *et al.*, 2021), and emphasize the importance of including the effect of REW on  $g_m$  in LSMs (Knauer *et al.*, 2020). This study also provides a calibration of the water stress factor for potato and contributes to reducing the uncertainties when estimating carbon assimilation and transpiration under water stress (Vidale *et al.*, 2021).

p-value $REW_{th}$	$A_{sat}$	$V_{cmax,25}$	$g_s$	$g_{m,25}$	$g_1$	$C_i/C_s$
$A_{sat}$	-					
$V_{cmax,25}$	0.000***	-				
$g_s$	0.89 <sup>ns</sup>	0.000***	-			
$g_{m,25}$	0.99 <sup>ns</sup>	0.000***	0.92 <sup>ns</sup>	-		
$g_1$	0.00***	0.000***	0.000***	0.000***	-	
$C_i/C_s$	0.00***	0.000***	0.000***	0.000***	0.5 <sup>ns</sup>	-

**Table 5.2:** p-value of the t-test comparing  $REW_{th}$  between  $A_{sat}$ ,  $V_{cmax,25}$ ,  $g_s$ ,  $g_{m,25}$ ,  $g_1$  and  $C_i/C_s$ .\*\*\* indicates when the p-value is <0.001, \*\*: <0.01, \* : <0.05 and ns : >0.05

Additional information on the description of the physiological effects of mesophyll on stomatal closure can be found in (Lemonnier & Lawson, 2023). Since disentangling the primary metabolisms which synchronously control photosynthesis, stomatal and mesophyll conductance remains challenging, future studies would benefit from additional molecular or anatomical measurements to unravel the interplays between stomatal and non-stomatal factors (Gago *et al.*, 2020).

## 6.2. Severe restrictions in REW change the relationship between photosynthesis and stomatal conductance

Severe restrictions in soil water availability induced a decrease in  $V_{cmax}$  as well as an increase in  $g_1$  and  $C_i/C_s$  (Fig. 5.2). An increase in  $C_i/C_s$  can be observed under strong limitations in  $CO_2$  diffusion and decreasing photosynthetic activity (Bermúdez-Cardona *et al.*, 2015; Brodribb, 1996; Huang, 2020; Tan *et al.*, 2017). In particular,  $C_i/C_s$  increased when  $g_s$  was lower than  $0.04 \text{ molm}^{-2} \text{ s}^{-1}$ , which was already reported as a stomatal conductance threshold for such  $C_i$ -inflexion point in various species (Blankenagel *et al.*, 2018; Brodribb, 1996; Flexas & Medrano, 2002; Martin & Ruiz-Torres, 1992; Rouhi *et al.*, 2007) including potato (Ramírez *et al.*, 2016). In these conditions of photosynthesis inhibition, the excess of energy carried by sun irradiance must be metabolized by alternative processes such as the xanthophyll (Demmig-Adams *et al.*, 2012), lutein (García-Plazaola *et al.*, 2003) and photorespiratory cycles (Osmond *et al.*, 1980). This latter may contribute to the increase in  $C_i/C_s$  by emitting  $CO_2$  through the glycine decarboxylase enzyme which is not totally fixed by Rubisco (Busch *et al.*, 2017; Shi & Bloom, 2021).

The increase in  $g_1$  induced an increase in  $dg_1/g_1$  and  $L_{g1}$  when  $REW < REW_{th,g1,CiCs}$  (Fig. 5.2, Fig. 5.4-a,b).  $g_1$  is inversely related to the marginal carbon cost of water, which corresponds to the change in carbon gained per unit of water transpired, also known as marginal WUE (Medlyn *et al.*, 2011a). The increase in  $g_1$  can be explained either by: (1) an increase in transpiration per unit of carbon gained by photosynthesis, or (2) a decrease in photosynthesis per unit of water transpired (Medlyn *et al.*, 2011a). For example, increasing stomatal conductance to promote transpiration may help in cooling down leaves surface during heatwaves at the expense of increasing mortality risks through hydraulic vulnerability and cavitation (Marchin *et al.*, 2022; Urban *et al.*, 2017b). Numerous studies have highlighted such cooling effect on potato (Sprenger *et al.*, 2016; Zhang *et al.*, 2022b), which can ultimately lead to an increase in  $g_1$  (Marchin *et al.*, 2023). A decoupling between stomatal conductance and photosynthesis may be the consequence of an adaptive strategy (i.e., sacrificing water for leaf survival and future carbon gains), or the increasing viscosity of water at high temperatures which facilitates the transport of water in the vascular system (Marchin *et al.*, 2023). In our experiment, the lowest measurement of  $g_s$  was  $0.011 \text{ molm}^{-2}\text{s}^{-1}$ , which is higher than the reported value of minimum stomatal conductance for  $CO_2$  transfer

across plant species (i.e.,  $g_s = 0.008 \text{ molm}^{-2}\text{s}^{-1}$ , Duursma *et al.* (2019)) and suggests that stomata might not be fully closed. It is however unlikely that potato plants had access to water to sustain transpiration through stomata or cuticle because of the low REW values that were observed in these conditions (Fig. 5.2). Alternatively, the increase in  $g_1$  may be caused by a decrease in photosynthesis through the additional effect of NSOL on  $A_{sat}$  (Beauclaire *et al.*, 2023b; Gourlez de la Motte *et al.*, 2020), which intensifies the decoupling between carbon assimilation and stomatal conductance by decreasing WUE (Manzoni *et al.*, 2011). This hypothesis is supported by previous studies which have shown that irrigation enhances WUE for potato (Akkamis & Caliskan, 2023; Ati *et al.*, 2012).

The increase in  $g_1$  induced a positive contribution to  $dA_{sat}/A_{sat}$  (Fig. 5.5), suggesting that potato plants promoted the loss of water to the benefit of  $\text{CO}_2$  diffusion despite the risks for the hydraulic system and carbon apparatus when photosynthesis reached critical levels (Deva *et al.*, 2020; Reynolds-Henne *et al.*, 2010). However, this behavior contradicts with the optimality theory, which assumes that the curvature of photosynthesis versus transpiration is negative (Buckley *et al.*, 2017). This hypothesis is challenged when mesophyll conductance strongly constrains  $\text{CO}_2$  diffusion, resulting in an increasing sensitivity of photosynthesis to changes in  $g_m$  compared to  $g_s$  (Cowan & Farquhar, 1977). Therefore, low REW may induce a violation of the optimality theory which can be evidenced by an increase in  $g_1$  (Buckley *et al.*, 2017; Zenes *et al.*, 2020).

### 6.3. Methodological considerations

$g_m$  was determined by the ‘variable J’ method (Harley *et al.*, 1992) which is sensitive to variation in  $R_d$  and  $\Gamma^*$ . These two variables can be impacted by drought and heat, which was not considered in the method part. First, it has been shown that  $R_d$  can increase under water stress due to the additional release of  $\text{CO}_2$  from mitochondria by the photorespiratory cycle (Busch *et al.*, 2017; Pinheiro & Chaves, 2011; Schmiede *et al.*, 2023). Second, the sensitivity of  $\Gamma^*$  to temperature can change under critical levels (usually above  $30^\circ\text{C}$ , Walker *et al.* (2017)), invalidating the Bernacchi *et al.* (2001) parameterization on leaf temperature. Measuring the  $\text{CO}_2$  compensation point (Walker & Ort, 2015) and leaf respiration (Yin & Amthor, 2024) under drought could help resolving these uncertainties.

Lastly, the diffusion of water vapor through cuticle and epidermis may become significant compared to stomatal diffusion under water stress (Boyer, 2015a,b; Boyer *et al.*, 1997). As the transpiration flux measured by gas exchange measurement systems correspond to the sum of the diffusion through stomatal and cuticle conductance, strong overestimations of  $C_i$  can occur as the Fick law considers an identical gas phase path for  $\text{CO}_2$  and  $\text{H}_2\text{O}$ . Direct measurements of  $C_i$  by a modified gas exchange (BOYER & KAWAMITSU, 2011) or a modification of the Fick law by quantifying the

cuticle conductance (Wang *et al.*, 2018b) could increase the accuracy of  $C_i$  estimates under water stress.

## 7. Acknowledgements and fundings

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## 8. Supplementary materials

### 8.1. Mathematical development of the partitioning scheme using the USO and FvCB models

The mathematical development of the limitation analysis using the USO model (Medlyn *et al.*, 2011a) and the FvCB model (Farquhar *et al.*, 1980) can be demonstrated by starting from the equation describing the limitation of carbon assimilation under high irradiance ( $A_{sat}$ ) by Rubisco as:

$$A_{sat} = V_{cmax} \frac{C_c - \Gamma^*}{C_c + K_m} \quad (5.18)$$

where  $V_{cmax}$  is the maximum carboxylation rate of Rubisco,  $C_c$  the CO<sub>2</sub> concentration in the chloroplast,  $\Gamma^*$  the CO<sub>2</sub> compensation point,  $K_m$  the Michaelis-Menten coefficient of Rubisco kinetics and  $R_d$  is the leaf respiration. At 25°C (i.e., no effects of temperature on model parameters) and neglecting leaf respiration variations, the total derivative of  $A_{sat}$  is:

$$dA_{sat} = \frac{\delta A_{sat}}{\delta V_{cmax}} dV_{cmax} + \frac{\delta A_{sat}}{\delta C_c} dC_c \quad (5.19)$$

with  $\frac{\delta A_{sat}}{\delta V_{cmax}} = \frac{A_{sat}}{V_{cmax}}$  and  $\frac{\delta A_{sat}}{\delta C_c} = V_{cmax} \frac{K_m + \Gamma^*}{(C_c + K_m)^2}$ . The total derivative of  $C_c$  can be determined using the Fick's law (Farquhar & Sharkey, 1982) as:

$$C_c = c_s - \frac{A_{sat}}{g_s} - \frac{A_{sat}}{g_m} \quad (5.20)$$

with  $C_s$  the CO<sub>2</sub> concentration at the leaf surface,  $g_s$  and  $g_m$  the stomatal and mesophyll conductance to CO<sub>2</sub> transfer. When the dependance of  $g_s$  to VPD, CO<sub>2</sub> concentration and carbon assimilation is not considered (i.e., the USO model not implemented in the limitation analysis), the total derivative of  $C_c$  is:

$$dC_c = \frac{\delta C_c}{\delta A_{sat}} dA_{sat} + \frac{\delta C_c}{\delta g_s} dg_s + \frac{\delta C_c}{\delta g_m} dg_m \quad (5.21)$$

$$dC_c = \left( -\frac{1}{g_m} - \frac{1}{g_s} \right) dA_{sat} + \frac{A_{sat}}{(g_s)^2} dg_s + \frac{A_{sat}}{(g_m)^2} dg_m \quad (5.22)$$

Combining Eq. 5.22 and Eq. 5.19 gives:

$$dA_{sat} = \frac{A_{sat}}{V_{cmax}} dV_{cmax} + \frac{\delta A_{sat}}{\delta C_c} \left[ \left( -\frac{1}{g_m} - \frac{1}{g_s} \right) dA_{sat} + \frac{A_{sat}}{(g_s)^2} dg_s + \frac{A_{sat}}{(g_m)^2} dg_m \right] \quad (5.23)$$

Simplifying Eq. 5.23 and writing  $dA_{sat}$  on the left-hand side gives:

$$dA_{sat} + \frac{\delta A_{sat}}{\delta C_c} \left( \frac{1}{g_m} + \frac{1}{g_s} \right) dA_{sat} = \frac{A_{sat}}{V_{cmax}} dV_{cmax} + \frac{\delta A_{sat}}{\delta C_c} \frac{A_{sat}}{(g_s)^2} dg_s + \frac{\delta A_{sat}}{\delta C_c} \frac{A_{sat}}{(g_m)^2} dg_m \quad (5.24)$$

$$dA_{sat} = \frac{\frac{A_{sat}}{V_{cmax}} dV_{cmax}}{1 + \frac{\delta A_{sat}}{\delta C_c} \left( \frac{1}{g_s} + \frac{1}{g_m} \right)} + \frac{\frac{\delta A_{sat}}{\delta C_c} \frac{A_{sat}}{g_s} dg_s}{1 + \frac{\delta A_{sat}}{\delta C_c} \left( \frac{1}{g_s} + \frac{1}{g_m} \right)} + \frac{\frac{\delta A_{sat}}{\delta C_c} \frac{A_{sat}}{g_m} dg_m}{1 + \frac{\delta A_{sat}}{\delta C_c} \left( \frac{1}{g_s} + \frac{1}{g_m} \right)} \quad (5.25)$$

Finally, simplifying Eq. 5.25 in terms of relative changes in  $A_{sat}$  and implementing the total conductance to CO<sub>2</sub> transfer  $g_t = ((g_s)^{-1} + (g_m)^{-1})^{-1}$  gives:

$$\frac{dA_{sat}}{A_{sat}} = \frac{g_t \frac{dV_{cmax}}{V_{cmax}}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} + \frac{\frac{g_t dg_s}{g_s} \frac{\delta A_{sat}}{\delta C_c}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} + \frac{\frac{g_t dg_m}{g_m} \frac{\delta A_{sat}}{\delta C_c}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.26)$$

$$\frac{dA_{sat}}{A_{sat}} = \frac{dg_s}{g_s} l_{gs} + \frac{dg_m}{g_m} l_{gm} + \frac{dV_{cmax}}{V_{cmax}} l_{V_{cmax}} = L_{gs} + L_{gm} + L_{V_{cmax}} \quad (5.27)$$

with:

$$l_{gs} = \frac{\frac{g_t \delta A_{sat}}{g_s \delta C_c}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.28)$$

$$l_{gm} = \frac{\frac{g_t \delta A_{sat}}{g_m \delta C_c}}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.29)$$

$$l_{V_{cmax}} = \frac{g_t}{g_t + \frac{\delta A_{sat}}{\delta C_c}} \quad (5.30)$$

where  $L$  is the limitation component and  $l$  the relative limitation component. Eq. 5.27 is the limitation analysis scheme described in Grassi & Magnani (2005).

The second partitioning scheme considered in this study is obtained by implementing the USO model (Eq. 5.31) for calculating the total derivative of  $g_s$  ( $dg_s$ , Eq. 5.27) as:

$$g_s = \left(1 + \frac{g_1}{\sqrt{VPD}}\right) \frac{A_{sat}}{C_s} \quad (5.31)$$

$$dg_s = \frac{\delta g_s}{\delta g_1} dg_1 + \frac{\delta g_s}{\delta A_{sat}} dA_{sat} + \frac{\delta g_s}{\delta VPD} dVPD + \frac{\delta g_s}{\delta C_s} dC_s \quad (5.32)$$

Calculating the partial derivatives of each term and considering  $C_s$  as constant gives:

$$dg_s = \frac{A_{sat}}{C_s \sqrt{VPD}} dg_1 + \frac{\left(1 + \frac{g_1}{\sqrt{VPD}}\right)}{C_s} dA_{sat} - \frac{1}{2} \frac{A_{sat} g_1}{C_s VPD^{3/2}} \frac{dVPD}{VPD} \quad (5.33)$$

Writing Eq. 5.33 in terms of relative changes in  $g_s$  gives:

$$\frac{dg_s}{g_s} = \frac{g_1}{g_1 + \sqrt{VPD}} \frac{dg_1}{g_1} + \frac{dA_{sat}}{A_{sat}} - \frac{1}{2} \left( \frac{g_1}{g_1 + \sqrt{VPD}} \right) \frac{dVPD}{VPD} \quad (5.34)$$

Eq. 5.34 gives the relative variation of  $g_s$  with respect to the relative variations of  $g_1$ ,  $A_{sat}$  and VPD. Moreover, it can be shown that  $\frac{g_1}{g_1 + \sqrt{VPD}} = \frac{C_i}{C_s}$ :

$$\frac{C_i}{C_s} = 1 - \frac{\sqrt{VPD}}{\sqrt{VPD} + g_1} = \frac{g_1}{g_1 + \sqrt{VPD}} \quad (5.35)$$

which gives, after combining Eq. 5.34 into Eq. 5.35:

$$\frac{dg_s}{g_s} = \frac{C_i}{C_s} \frac{dg_1}{g_1} + \frac{dA_{sat}}{A_{sat}} - \frac{1}{2} \frac{C_i}{C_s} \frac{dVPD}{VPD} \quad (5.36)$$

Finally, combining Eq. 5.36 and Eq. 5.27 gives after simplification:

$$\frac{dA_{sat}}{A_{sat}} = \frac{dg_1}{g_1} \left( \frac{l_{gs}}{1-l_{gs}} \frac{C_i}{C_s} \right) + \frac{dg_m}{g_m} \left( \frac{l_{gm}}{1-l_{gs}} \right) + \frac{dV_{cmax}}{V_{cmax}} \left( \frac{l_{V_{cmax}}}{1-l_{gs}} \right) - \left( \frac{1}{2} \frac{l_{gs}}{1-l_{gs}} \frac{C_i}{C_s} \right) \frac{dVPD}{VPD} \quad (5.37)$$

$$\frac{dA_{sat}}{A_{sat}} = \frac{dg_1}{g_1} l_{g1,USO} + \frac{dg_m}{g_m} l_{gm,USO} + \frac{dV_{cmax}}{V_{cmax}} l_{V_{cmax},USO} + \frac{dVPD}{VPD} l_{VPD,USO} \quad (5.38)$$

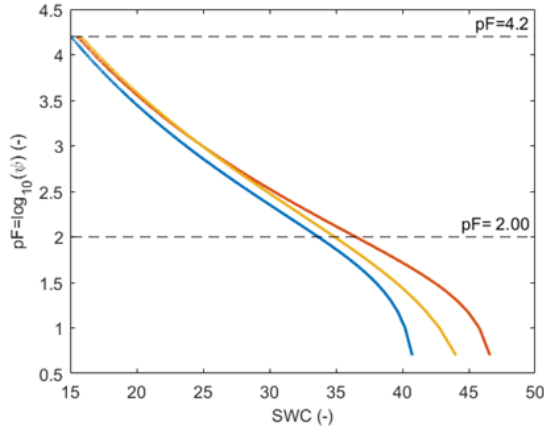
$$\frac{dA_{sat}}{A_{sat}} = L_{g1,USO} + L_{gm,USO} + L_{V_{cmax},USO} + L_{VPD,USO} \quad (5.39)$$

Eq. 5.38 and Eq. 5.38 link the relative variation of  $A_{sat}$  to the relative variations of VPD,  $g_1$ ,  $g_m$  and  $V_{cmax}$  (Farquhar *et al.*, 1980; Medlyn *et al.*, 2011a).

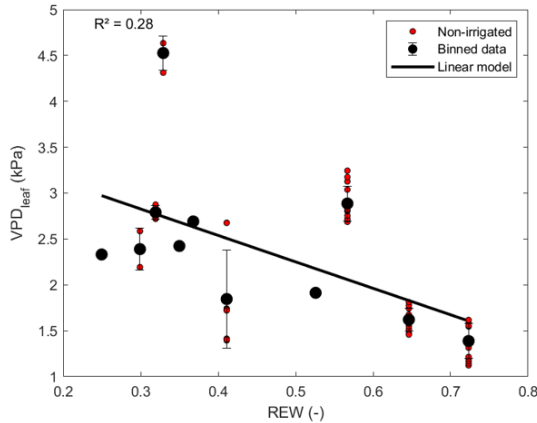
## 8.2. Figures and tables

$\Delta AICc =$ AICc-min(AICc)	$\Delta AICc$ Linear model	$\Delta AICc$ Segmented model
$A_{sat}$	0	-3.36
$V_{cmax,25}$	-7.45	0
$g_s$	0	4.72
$g_{m,25}$	0	5.67
$g_1$	23	0
$C_i/C_s$	6.22	0

**Table 5.3:** Akaike information criterion corrected for small sample size (AICc) of the linear and linear-plateau models describing the response of  $A_{sat}$ ,  $V_{cmax,25}$ ,  $g_s$ ,  $g_{m,25}$ ,  $g_1$  and  $C_i/C_s$  to REW.  $\Delta AICc$  is the difference between the lowest AICc and the linear or segmented model for each variable.

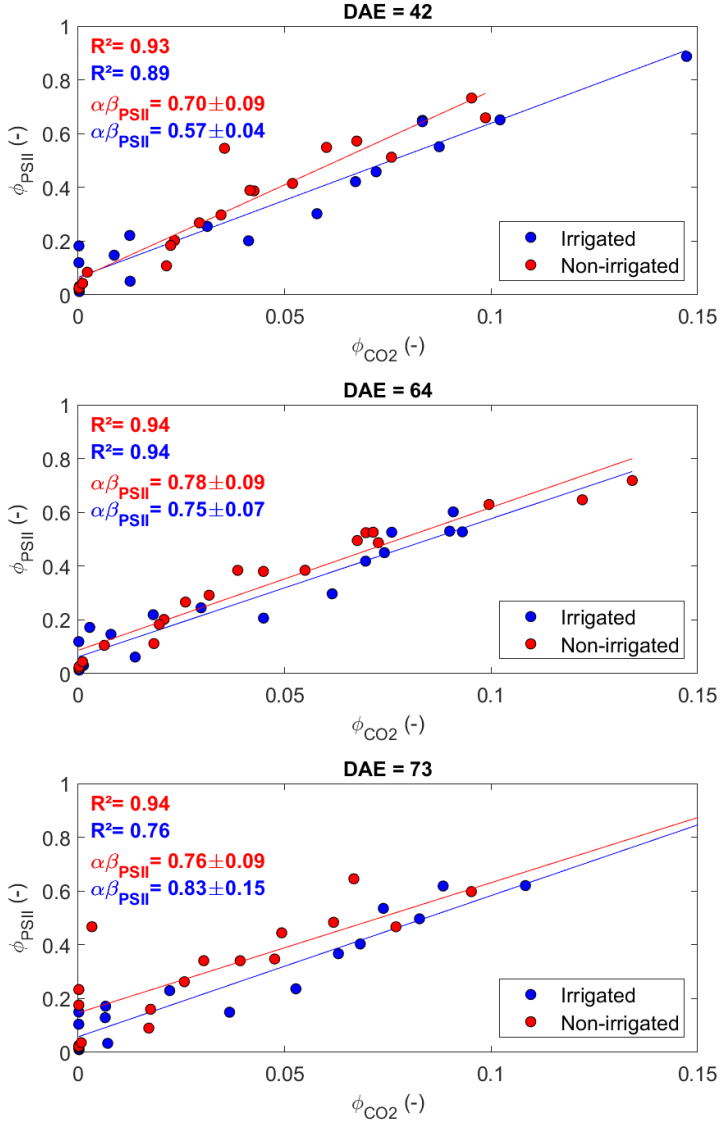


**Figure 5.6:** SWRC of the three soil samples collected at 15 cm depth. The field capacity and wilting point were calculated as the average values of soil water content SWC at a pF ( $\log(\psi)$ , with  $\psi$  the suction head) of respectively 2 and 4.2. SWRC were obtained from the van Genuchten model (van Genuchten, 1980) :  $SWC = SWC_r + \frac{SWC_s - SWC_r}{|1 + (\alpha|\psi|^n)|^m}$  with  $SWC_r$  the residual SWC,  $SWC_s$  the saturated SWC,  $\psi$  the suction head,  $\alpha$  related to the inverse of the air entry suction,  $n$  and  $m$  are empirical parameters. The parameters of the three SWRC are  $SWC_s=0.412$ ,  $SWC_r=0$ ,  $\alpha=0.028$ ,  $n=1.27$ ,  $m=0.13$  (blue curve),  $SWC_s=0.471$ ,  $SWC_r=0$ ,  $\alpha=0.04$ ,  $n=1.45$ ,  $m=0.118$  (orange curve), and  $SWC_s=0.460$ ,  $SWC_r=0$ ,  $\alpha=0.037$ ,  $n=0.84$ ,  $m=0.20$  (yellow curve).



**Figure 5.7:** Relationship between VPD at the leaf surface ( $VPD_{leaf}$ ) and REW for non-irrigated plants





**Figure 5.8:** Relationship between the fluorescence quantum yield of PSII ( $\phi_{PSII}$ ) and the apparent quantum yield of CO<sub>2</sub> assimilation ( $\phi_{CO_2}$ ) for three different DAE during the drought experiment. For each DAE, a linear regression ( $\phi_{PSII} = k\phi_{CO_2} + k'$ ) was fitted for irrigated (blue) and non-irrigated (red) data and the product of leaf absorptance to the fraction of the absorbed PAR allocated to PSII ( $\alpha\beta_{PSII}$ ) was calculated as  $4/k$ .



**Modeling gross primary production and  
transpiration from sun-induced chlorophyll  
fluorescence using a mechanistic  
light-response approach**



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## 1. Personal contributions

The installation of the spectrometer has been carried out by the technical team of the Loncée ICOS station, namely Henri Chopin and Gaëtan Bogaerts. The Ecosystem Thematic Centre (ETC) performed the data processing and quality control of the EC site of Loncée. The EC data is available on the ICOS Carbon Portal (see references in the paper below). Leaf-level measurements were collected by myself with the help of Natacha Pezzetti during her master thesis. Root biomass measurements were collected by Laura Delhez. The spectrometer was provided by the Forschungszentrum Jülich and Simon De Cannière, who also participated to the installation of the sensor. SIF data processing and quality control were performed by myself. The manuscript was reviewed by Bernard Longdoz, who also actively participated to the thesis by helping in setting the sensor installation and interpreting the results. Simon De Cannière and François Jonard helped in interpreting SIF data, discussing the results and reviewing the manuscript. Natacha Pezzetti and Laura Delhez reviewed the manuscript.

## 2. Abstract

SIF is a promising optical remote sensing signal which is directly linked to photosynthesis, allowing for the monitoring of GPP. Although empirical relationships between these variables have demonstrated the potential of SIF for site-specific GPP estimations, a better physiological understanding of the link between SIF and GPP would pave the way for a more robust model of photosynthesis. The MLR model is a novel approach which determines GPP from SIF by using only a small set of equations and parameters with physiological significance. This study combines the MLR model with the USO model to estimate both GPP and Tr at the ecosystem scale. Top-of-canopy SIF measurements were collected over a winter crop with a field spectrometer installed next to an eddy covariance station. MLR-USO model parameters were determined from gas exchange and active chlorophyll fluorescence measurements at the leaf level and interpolated on a half-hourly basis using solar irradiance and canopy temperature. GPP and Tr estimated by the MLR-USO model and eddy covariance measurements were highly correlated at half-hourly and daily timescales ( $R^2 \geq 0.91$ ,  $rRMSE \leq 13.7\%$ ) under a wide range of environmental conditions, including soil water stress. These results highlight the potential of the MLR-USO model as an important step towards

an improvement of our understanding of the coupling between the water and carbon cycles at the ecosystem scale and beyond.

### 3. Introduction

Each year, terrestrial ecosystems uptake approximately one-third of the carbon emitted by human activities (Friedlingstein *et al.*, 2022). GPP of terrestrial ecosystems represents the most important flux in the global carbon cycle (Beer *et al.*, 2010) and plays a central role in regulating atmospheric CO<sub>2</sub> concentration (Sha *et al.*, 2022).

Anthropogenic climate change increases the occurrence of climate extremes such as droughts or heatwaves, which causes a reduction in GPP (Williams *et al.*, 2014; Reichstein *et al.*, 2013). Whether terrestrial ecosystems will keep removing as much CO<sub>2</sub> from the atmosphere or become CO<sub>2</sub> sources is highly uncertain as most LSMs predict various future ecosystem uptake abilities in a climate change context (Ryu *et al.*, 2019). Quantifying and modeling GPP at global scales is therefore crucial to better understand how climate extremes constrain primary production in climate modeling.

GPP is the total amount of carbon fixed by ecosystems and originates from complex photosynthetic processes that cannot be directly measured. The standard approach for estimating GPP at the ecosystem scale is the measurement of the NEE with the eddy covariance (EC) technique, and its partitioning between ecosystem respiration ( $R_{ECO}$ ) and GPP. More than 200 EC sites are currently part of the FLUXNET network, which regroups over 1500 site-years of data (Baldocchi, 2014; Pastorello *et al.*, 2020). However, flux towers are unevenly distributed across the globe, which explains that the temporal and spatial representativeness of flux measurements (the extrapolation of flux information to extended spatial and timescales) has become an increasingly important issue for the scientific community to provide flux information “everywhere and all the time” (Chu *et al.*, 2017, 2021). In addition, the size and shape of flux tower footprints (the source areas of EC fluxes) are directly affected by meteorological variables (such as wind direction or air turbulence) and surface roughness, which complicates the interpretation of GPP from local to global scales (Kong *et al.*, 2022). Most of the recent approaches for upscaling carbon assimilation to large scales use machine learning to build empirical models between remote sensing (RS) data and ground observations of GPP (Jung *et al.*, 2020)).

Photosynthesis is a complex chain of photochemical and enzymatical reactions, which are commonly categorized between the light reactions (which generate ATP and NADPH from light and water) and the dark reactions (which transform CO<sub>2</sub> into carbohydrates using ATP and NADPH). SIF originates from the emission of photons in the red and far-red regions, after the absorption of a fraction of PAR by chlorophyll a pigments during the light reactions of photosynthesis. Although SIF quantum yield is very low (1 to 2 %; Maxwell & Johnson (2000)) compared to PQ and thermal energy dissipa-

tion (i.e., NPQ), it is tightly coupled to GPP as these three processes are in competition for PAR. In particular, SIF is linked to the electron transport rate which fuels ATP and NADPH to the fixation sites in the chloroplasts (Gu *et al.*, 2019). While the relationship between SIF and PQ quantum yields at the leaf level is highly nonlinear on short timescales (Porcar-Castell *et al.*, 2014), large temporal and spatial scales data show a strong linearity between SIF and carbon assimilation (Lee *et al.*, 2015; Damm *et al.*, 2010; Li *et al.*, 2018b; Chen *et al.*, 2019a). Such relationship has been the basis of the calibration of empirical models between ground-based GPP and remotely sensed SIF for estimating GPP at global scale using the LUE approach (Porcar-Castell *et al.*, 2014; Jonard *et al.*, 2020; Li *et al.*, 2018b; Xiao *et al.*, 2019). However, this concept does not allow to identify the actual mechanistic link between SIF and GPP, and there is a lack of knowledge about the influence of physiological and environmental factors on these relationships. A decoupling between SIF and GPP can be observed during stress episodes such as heatwaves (Wohlfahrt *et al.*, 2018; Martini *et al.*, 2022), drought (Helm *et al.*, 2020; Marrs *et al.*, 2020; Chen *et al.*, 2021), light-limiting conditions (Chen *et al.*, 2019a), or for heterogeneous canopies (van der Tol *et al.*, 2009; Gu *et al.*, 2019). These examples illustrate the limits of using this approach across spatial scales and climate conditions.

A model must consider two aspects to estimate GPP from SIF with a mechanistic approach. First, the relationship between fluorescence and photochemistry must be characterized at the photosystem level. Second, the scattering and absorption effects of overlying leaves should be considered when the fluorescence photons travel through the canopy (Yang *et al.*, 2019). The Soil Canopy Observation, Photochemistry and Energy fluxes (SCOPE) model (van der Tol *et al.*, 2009) integrates both the physiological and the canopy structural aspects of the SIF emission in separate sub-models. These aspects are highly dynamic and change in function of the plant water status (De Cannière *et al.*, 2022; Wang *et al.*, 2022). Although the SCOPE model is often used in a reduced form when coupled to other models, it is computationally intensive and requires numerous parameters, forcing the user to make a series of assumptions (e.g., Lee *et al.* (2015); De Cannière *et al.* (2021)).

An alternative approach to link SIF and GPP is the MLR model (Gu *et al.*, 2019). While the SCOPE model uses sub-models to integrate the physiological and structural aspects of SIF emission, the MLR model only considers few parameters which describe the energy partitioning at the PSII level (e.g., the fraction of open PSII center  $q_L$  and the maximum photochemical quantum yield  $\phi_{PSII,max}$ ) and leaf-to-canopy radiative transfer processes (i.e., the escape probability factor  $f_{esc}$ ) to determine the actual electron transport rate between the two photosystems (PSII and PSI). This mechanistic representation of the light reactions contrasts with the FvCB model of photosynthesis (Farquhar *et al.*, 1980) where the potential electron transport rate is determined from PAR by an empirical relationship. The FvCB model requires numerous parameters that are highly variable across PFT and environmental conditions; such uncertainty

causes important disagreements in LSMs which use the FvCB model to estimate carbon assimilation (Walker *et al.*, 2021; Rogers *et al.*, 2017). By taking advantage of the physiological information carried by SIF, MLR model equations mechanistically describe photosynthetic processes while reducing the impact of parameter uncertainty (Han *et al.*, 2022b; Gu *et al.*, 2019). With this approach, GPP can be estimated from electron requirements for carboxylation and oxygenation regardless of carboxylation limiting stages (Gu *et al.*, 2019).

Stomatal opening controls both water and carbon exchanges at the leaf surface. Through stomatal regulation, plants balance water loss and carbon uptake (Cowan & Farquhar, 1977), leading to a close coupling between GPP and Tr. Recent models simulating stomatal behavior such as the USO model (Medlyn *et al.*, 2011a) are based on the optimization of the ratio between benefit (i.e., carbon gain) and cost (i.e., water loss) of stomatal opening and on environmental variables such as VPD, GPP and CO<sub>2</sub> concentration at the leaf surface. Coupling the MLR and USO models (MLR- USO) provides a new, process-based approach for the modeling of carbon and water fluxes from SIF.

SIF can be retrieved by high-resolution measurements of vegetation spectral properties from sensors on masts; aircrafts or satellites (Mohammed *et al.*, 2019; Jonard *et al.*, 2020). For instance, spaceborne spectrometers such as the TROPOMI or the NASA's Orbiting Carbon Observatory Orbiting Carbon Observatory-2 (OCO-2) measure SIF emission at a specific wavelength (e.g., 740 nm for TROPOMI) from kilometer-scale footprints. These data have been used for training machine learning algorithms linking SIF and GPP with empirical linear regressions (Jung *et al.*, 2020) or for investigating water and light limiting conditions on photosynthesis at large scales (Jonard *et al.*, 2022). The upcoming FLEX tandem mission with Sentinel-3 will enable the quantification of unique products related to SIF and surface temperature at a high spatial resolution (300 × 300 m; Drusch *et al.* (2017)). Such improvement will reduce the current mismatch between flux tower footprint and RS products. Moreover, many studies have been designed to validate future FLEX products by measuring SIF with top-of-canopy (TOC) sensors such as the FloX (JB Hyperspectral, Düsseldorf, Germany), whose development has been supported by the European Space Agency (ESA) (Schuttemeyer *et al.*, 2018). The MLR- USO model must be first evaluated at the ecosystem scale using TOC SIF measurements before future applications with RS SIF data at larger scales.

The MLR- USO model relies on a straightforward methodology for improving our understanding of the photosynthetic and transpiration processes by exploiting the physiological message carried by the SIF signal. To date, no studies have tested such a model. In this perspective, we used the MLR- USO model to estimate GPP and Tr from TOC SIF. This study focused on wheat, which is the most widely planted crop (Erenstein *et al.*, 2022) and an important crop functional type in climate models (Lu



*et al.*, 2017; Boas *et al.*, 2021). Our approach aimed at: (1) characterizing the relationship between SIF and GPP at the ecosystem scale for winter wheat, (2) measuring MLR- USO model parameters and assessing their dependence on meteorological variables and (3) validating MLR-USO model outputs with EC data.

## 4. Material and methods

### 4.1. Study site

All the measurements presented in this study were collected at the ICOS station of Loncée (BE-Lon, Level 2 ICOS station), located in the Wallonia region in Central Belgium (50° 33' N, 4°44' E, 167 m asl). The period studied is the 2022 growing season dedicated to winter wheat (*Triticum aestivum* cv. Skyscraper). The Loncée experimental site has been cultivated for more than 80 years based on a typical 4-year rotation: winter wheat / sugar beet (*Beta vulgaris*) / winter wheat / seed potato (*Solanum tuberosum*). The climate in Loncée is temperate oceanic, with mean annual air temperature and cumulative precipitation of about 10.2 °C and 743 mm respectively. The soil is silty clay, with two main horizons: a ploughed layer from 0 to 35 cm (FAO classification Ap) and a layer enriched in clay (FAO classification Bt) from 35 to 100 cm depth (Table 6.1). This study exclusively focused on the 2022 growing season dedicated to winter wheat (*Triticum aestivum* cv. Skyscraper), which was sown on October 28, 2021 and harvested on July, 24 2022. No irrigation was applied during the growing season. Farming activities are summarized in Table 6.2.

### 4.2. Eddy covariance fluxes and meteorological measurements

The contribution of the target cropland to the EC fluxes was large regardless of the atmospheric condition (more than 90% on average during the last 13 years, data not shown). The prevailing winds are SW and NE with a cropland fetch of respectively 240 m and 200 m. The sonic anemometer (Solent Research R3, Gill Instruments Lymington, UK) and gas analyzer (LI-7200, LI-COR, Lincoln, NE, US) of the EC station were selected following ICOS guidelines (Rebmann *et al.*, 2018). Half-hourly fluxes of net CO<sub>2</sub> and LE (from which Tr is thereafter determined) were computed from high frequency measurements (20 Hz) of vertical wind speed, CO<sub>2</sub> and H<sub>2</sub>O concentrations at 2.93 m above the ground. These raw data were processed by the ICOS Ecosystem Thematic Center (ETC) using the ONEFlux pipeline (Pastorello *et al.*, 2020) and are available in the ICOS Carbon Portal (Dumont *et al.*, 2023a).  $R_{ECO}$  and GPP were determined from NEE by using the nighttime partitioning method (Reichstein *et al.*, 2005). Storage CO<sub>2</sub> fluxes were neglected owing to the short measurement height. Flux data corresponding to low turbulence conditions (with low friction velocity  $u_*$ ) were discarded from the analysis using the variable  $u_*$  threshold method (VUT; Pastorello *et al.* (2020)), which is preferred when canopy cover changes throughout the

years such as crops. Moreover, no gap-filled fluxes were selected (i.e., quality flag of 0). The variables GPP\_NT\_VUT\_REF for GPP and LE\_F\_MDS for LE were selected from the dataset. These fluxes are referred to in this paper as  $GPP_{EC}$  and  $LE_{EC}$ .

Micrometeorological measurements were collected on a half-hourly basis, including incoming photosynthetic photon flux density in the PAR (photodiode-based sensor - PAR Quantum sensor SKP 215, Skye Instruments Limited, Llandrindod Wells, UK), precipitation (weighing rain gauge - TRwS415, MPS system sro, Bratislava, SK), air temperature and relative humidity (resistive platinum thermometer and electrical capacitive hygrometer - HMP155, Vaisala Oyj, Helsinki, FI), soil heat flux (self-calibrating soil heat flux plate- HFP01SC, Hukseflux Thermal Sensors B.V., Delft, NL), longwave downward and outgoing fluxes, net radiation (pyradiometer – CNR 4, Kipp and Zonen, Delft, NL) and SWC at three different locations and at five different depths (5, 15, 25, 55 and 85 cm) using silicon bandgap temperature and capacitance sensors (EnviroSCAN Probe, Sentek Sensor Technologies, Stepney, SA, AU).

### 4.3. Spectral measurements of SIF

Solar irradiance and reflected radiance from the canopy were measured by the FloX (JB Hyperspectral, Düsseldorf, Germany), which consists of two spectrometers: a QE Pro and a FLAME-S spectrometer (Ocean Optics, Inc., USA). QE Pro measures downwelling irradiance and upwelling radiance covering the range between 650 and 800 nm to observe the SIF signal in the O<sub>2</sub>-A (760 nm) and O<sub>2</sub>-B (687 nm) absorption bands while FLAME-S covers a wider range from 400 to 1000 nm including both visible and NIR wavelengths. The QE-Pro and FLAME-S spectrometers have a spectral sampling interval and spectral resolution of respectively 0.17 nm, 0.3 nm and 0.65 nm, 1.5 nm. The spectrometers are housed in a thermally regulated box which keeps temperature below 25°C to avoid spectral shifts and dark current drifts. The optical fibers were installed on a mast at 2 m height, at approximately 20 m from the center of the plot, and 5 m from the location where one of the three soil water content sensor profiles was installed. The downward looking bare fibers have an opening angle of 23° which allows to measure the canopy reflectance and SIF emission over a footprint of approximately 1.5 m<sup>2</sup>. The receptor of the upward looking fiber measuring the downwelling irradiance has a hemispherical field of view of 180°. The FloX was installed on February 24, 2022 and removed on July 18, 2022, few days before the harvest of winter wheat.

SIF was retrieved using the SFM from the canopy reflectance spectrum in the O<sub>2</sub>-A band which provides the most reliable SIF measurements (Cendrero-Mateo *et al.*, 2019; Chang *et al.*, 2020). The O<sub>2</sub>-A band was preferred over the O<sub>2</sub>-B band as it enables to track photosynthesis dynamics and limits the reabsorption effects within the canopy structure (Cendrero-Mateo *et al.*, 2019; Martini *et al.*, 2022). Moreover, dense cloud cover may complicate the retrieval of SIF. Therefore, days with a clearness index (i.e., the ratio between net radiation at the top of the canopy and extraterrestrial

net radiation, Chang *et al.* (2020)) lower than 0.3 were discarded from the analysis (Chen *et al.*, 2020; Yang *et al.*, 2021), as well as data when solar radiation is too low (before 9 am and after 3 pm; UTC+1). Nonlinearity and spectral shift corrections on raw FloX measurements were applied. A threshold of solar zenithal angle of 70° was set to avoid detector saturation (Chang *et al.*, 2021). TOC SIF retrieved with the SFM method within the O<sub>2</sub>-A band is further referred to as  $SIF_{TOC,760}$ .

#### 4.4. Downscaling of SIF from top-of-canopy to photosystem level

As SIF is affected by re-absorption and scattering effects of the canopy,  $SIF_{TOC,760}$  is indirectly related to the SIF emitted by leaves. Estimating GPP with the MLR model requires the determination of the SIF signal emitted only from PSII at the leaf level over the broadband spectrum emission of fluorescence (Gu *et al.*, 2019). Therefore,  $SIF_{TOC,760}$  values must be multiplied by three conversion factors representing (i) the fraction of PSII fluorescence emission contributing to the total SIF emission of the leaf ( $f_{PSII}$ ), (ii) the integration of the SIF signal at 760 nm over the broadband spectrum emission of fluorescence ( $f_{\lambda}$ ), and (iii) the effects of absorption and re-scattering by the canopy structure ( $f_{esc}^{-1}$ ). Total SIF ( $SIF_{TOT}$ ) was therefore calculated from  $SIF_{TOC,760}$  by (Liu *et al.*, 2022c):

$$SIF_{TOT} = \frac{SIF_{TOC,760} f_{PSII} f_{\lambda}}{f_{esc}} \quad (6.1)$$

The calculation steps of the three factors are detailed below.

##### 4.4.1. Separating the contribution of PSII to SIF measurement

Both PSI and PSII contribute to the emission of SIF in the NIR. Therefore, the influence of PSI on  $SIF_{TOC,760}$  must be considered before applying the MLR model which is valid only for PSII SIF emission. Therefore,  $f_{PSII}$  was calculated as the ratio of chlorophyll fluorescence quantum yields at the leaf level, while accounting for the contribution of PSI (Jia *et al.*, 2023). The detailed procedure of leaf-level measurements is given in section 4.9.

##### 4.4.2. From top-of-canopy to leaf surface

$f_{esc}$  is the escape probability that SIF photons emitted by PSII reach the top of the canopy.  $f_{esc}$  was determined as the ratio of NIR light originating from vegetation  $NIR_v$  to the fraction of incident PAR absorbed by the canopy  $f_{APAR}$  under the assumption of high leaf area index and low contribution of soil reflectance (Zeng *et al.*, 2019):

$$f_{esc} = \frac{NIR_v}{f_{APAR}} \quad (6.2)$$

$NIR_v$  was calculated from the Normalized Difference Vegetation Index (NDVI) and the canopy reflectance in the NIR ( $R_{NIR}$ ; Badgley *et al.* (2017)):

$$NIR_v = (NDVI - NDVI_0)R_{NIR} = \left( \frac{R_{NIR} - R_{RED}}{R_{NIR} + R_{RED}} - NDVI_0 \right) R_{NIR} \quad (6.3)$$

Where  $R_{RED}$  is the canopy reflectance in the red and  $NDVI_0$  is the minimum NDVI and accounts for the effects of bare soil on  $NIR_v$  (Badgley *et al.*, 2017).  $R_{NIR}$  and  $R_{RED}$  were measured by the FloX at respectively 760 nm and 680 nm.  $f_{APAR}$  was determined from NDVI by using a linear model which has already been used to estimate wheat crop yields (Moriondo *et al.*, 2007; Myneni & Williams, 1994):

$$f_{APAR} = 1.16 NDVI - 0.14 \quad (6.4)$$

$NDVI_0$  was set to 0.12, which is the NDVI value when  $f_{APAR}$  (Eq. 6.4) equals 0 (Myneni & Williams, 1994).

#### 4.4.3. Integrating the SIF signal over the broadband fluorescence emission spectrum

The final step of the conversion of  $SIF_{TOC,760}$  into  $SIF_{TOT}$  consists in integrating the SIF signal emitted by PSII at 760 nm over the broadband chlorophyll fluorescence emission spectrum, which requires broadband theoretical fluorescence emission spectrum of PSII. The SCOPE model (version 1.73; van der Tol *et al.* (2009)) was used to simulate a total of 6720 chlorophyll fluorescence spectra between 650 and 840 nm from a canopy with different leaf biochemical and structural proprieties, sun-canopy sensor geometry and structure. The fluorescence spectrum that provided the most meaningful representation of the dataset was retrieved by the Singular Vector Decomposition (SVD) technique (Liu *et al.*, 2022b; Zhao *et al.*, 2014). The first principal component explained more than 99% of the variance of the dataset (data not shown), which allowed to estimate the broadband SIF emitted from PSII at all wavelengths ( $SIF_{TOT,\lambda}$ ) by:

$$SIF_{TOT,\lambda} = \frac{SIF_{TOC,760} f_{PSII}}{f_{esc}} \frac{v_1(\lambda)}{v_1(760)} = \frac{SIF_{TOC,760} f_{PSII} f_\lambda}{f_{esc}}, \lambda \in [650, 840] \quad (6.5)$$

with  $\lambda$  the wavelength,  $v_1(\lambda)$  is the first right singular vector from the SVD and  $f_\lambda = \frac{v_1(\lambda)}{v_1(360)}$ . Eq. 6.5 becomes, after integration over the broadband range of SIF emission and unit conversion:

$$SIF_{TOT} = \pi \sum_{640}^{850} SIF_{TOT,\lambda} \frac{10^6 \cdot 10^{-9} \cdot 10^{-3}}{h \cdot c \cdot N_a} \quad (6.6)$$

with  $h$  the Planck constant,  $c$  the light velocity,  $N_a$  the number of Avogadro,  $SIF_{TOT,\lambda}$  in  $\text{mWm}^{-2}\text{nm}^{-1}\text{sr}^{-1}$ , and  $SIF_{TOT}$  in  $\mu\text{molm}^{-2}\text{s}^{-1}$ . The factor  $10^{-3}$  converts from milliwatt to watt,  $10^{-9}$  from nanometers to meters,  $10^6$  from moles to  $\mu\text{moles}$  and  $\pi$  integrates the signal over the hemispherical space.

## 4.5. Relationship between SIF and GPP

The relationship between measured GPP and total SIF (i.e.,  $GPP_{EC}$  and  $SIF_{TOT}$ ) was characterized by fitting a linear and a linear-plateau segmented model (Jonard *et al.*, 2022). This latter allows to identify a breakpoint in the relationship between two variables. Above the threshold, GPP is constant while SIF increases. Below the threshold, GPP and SIF are positively related. The most significant model was selected by using the lowest AIC, which avoids overfitting (Burnham *et al.*, 2002). Moreover, the SIF yield ( $SIF_y = \frac{SIF_{TOT}}{APAR}$  and the light-use efficiency of GPP  $LUE = \frac{GPP_{EC}}{APAR}$ , where  $APAR = f_{APAR}PAR$  is the absorbed irradiance), were calculated to reduce the effect of canopy structure and to provide a more physiological interpretation of the variability of  $GPP_{EC}$  and  $SIF_{TOT}$  throughout the growing season (Martini *et al.*, 2022).

## 4.6. GPP modeling

The MLR model calculates the actual electron transport rate ( $J_{SIF}$ ) from  $SIF_{TOT}$  by Gu *et al.* (2019):

$$J_{SIF} = q_L \frac{\phi_{PSII,max}(1 + k_{DF})SIF_{TOT}}{1 - \phi_{PSII,max}} \quad (6.7)$$

with  $q_L$  the fraction of PSII centers that are opened,  $\phi_{PSII,max}$  the maximum photochemical quantum yield, and  $k_{DF}$  equal to the ratio between  $k_D$  (representing the rate of constitutive heat thermal dissipation) and  $k_F$  (representing the rate of fluorescence emission), set to 19 as in Gu *et al.* (2019).  $q_L$  and  $\phi_{PSII,max}$  were calculated from leaf level measurements (see section 4.9).

Modeled GPP ( $GPP_{SIF}$ ) was determined from  $J_{SIF}$  based on the electron requirements for carboxylation and oxygenation processes (Gu *et al.*, 2019; Farquhar *et al.*, 1980):

$$GPP_{SIF} = J_{SIF} \frac{C_i - \Gamma^*}{C_i + K_m} = q_L \frac{\phi_{PSII,max}(1 + k_{DF})SIF_{TOT}}{1 - \phi_{PSII,max}} \frac{C_i - \Gamma^*}{4C_i + 8\Gamma^*} \quad (6.8)$$

with  $C_i$   $\text{CO}_2$  concentration in substomatal cavities directly measured at the leaf level (details in section 4.9) and  $\Gamma^*$  the  $\text{CO}_2$  compensation point.  $\Gamma^*$  was estimated from the canopy surface temperature ( $T_{can}$ ) (Bernacchi *et al.*, 2001):

$$\Gamma^* = 42.75 \exp \frac{37830(T_{can} - 278.15)}{298.15 R (T_{can} + 273.15)} \quad (6.9)$$

where  $T_{can}$  was determined from the longwave fluxes:

$$T_{can} = \sqrt[4]{\frac{LW_{out} - (1 - \epsilon)LW_{in}}{\epsilon \sigma}} \quad (6.10)$$

with  $LW_{in}$  and  $LW_{out}$  the longwave downward and outgoing fluxes,  $\epsilon$  the far-infrared emissivity equal to 0.97 and  $\sigma$  the Stefan-Boltzmann constant.

## 4.7. Transpiration modeling

The most widely used equation to estimate LE at the ecosystem scale is the PM equation (Monteith, 1965; Maes *et al.*, 2019). This equation models LE under the hypothesis that ecosystem fluxes originate from a single homogeneous plane (i.e., a single 'big-leaf', Knauer *et al.* (2018a)):

$$LE = \Delta (R_n - G) + \rho_a c_p VPD_{can} G_{ah} \Delta + \gamma \left(1 + \frac{G_{aw}}{G_{sw}}\right) \quad (6.11)$$

with  $\Delta$  the slope of the saturated pressure curve,  $R_n$  the net radiation,  $G$  the ground heat flux,  $\rho_a$  the air density,  $c_p$  the heat capacity of dry air,  $VPD_{can}$  the vapor pressure deficit at the canopy surface and  $\gamma$  the psychrometric constant.  $G_{aw}$  is the aerodynamic conductance for water vapor transport and  $G_{sw}$  refers to the canopy surface conductance to water vapor, which includes evaporation from the soil and transpiration from wet vegetation surfaces.  $VPD_{can}$  was determined from  $T_{can}$  and air relative humidity. The Thom equation (Thom, 1972) was used to determine  $G_{aw}$  was determined from wind speed ( $u$ ) and air friction velocity ( $u_*$ ):

$$G_{aw} = \left( \frac{u}{u_*^2} + 6u_*^{-0.667} \right)^{-1} \quad (6.12)$$

$G_{sw}$  (Eq. 6.11) was determined by using the USO model while neglecting the minimum stomatal conductance (Medlyn *et al.*, 2017b, 2011a):

$$G_{sw} \sim g_{sw} = 1.6 \left( 1 + \frac{g_1}{\sqrt{VPD_{leaf}}} \right) \frac{A_n}{C_s} \quad (6.13)$$

with  $g_1$  the slope parameter also related to the marginal water cost of carbon gain (Medlyn *et al.*, 2011a),  $g_{sw}$  the stomatal conductance to water vapor and  $A_n$  the net photosynthesis.  $VPD_{leaf}$  and  $C_s$  are respectively the VPD and  $\text{CO}_2$  concentration at the leaf surface (i.e., within the boundary layer). All variables in Eq. 6.13 were

measured at the leaf level.  $g_1$  was determined by adjusting a nonlinear regression on leaf-level measurements of  $A_n$ ,  $C_s$ ,  $VPD_{leaf}$  and  $g_{sw}$  (see section 4.9), giving  $g_1=2.28 \pm 0.09 \text{ kPa}^{0.5}$  (Fig. S1). The combination of Eq. 6.11 to Eq. 6.13 with  $A_n$  replaced by  $GPP_{SIF}$  from Eq. 6.8 and  $VPD_{leaf}$  replaced by  $VPD_{can}$ , gives  $G_{sw}$  and Tr ( $Tr_{SIF}$ ) from SIF measurements, under the assumption of dry canopy surface. Therefore, days when precipitation was measured and the subsequent 48h were discarded from the analysis (Knauer *et al.*, 2018b). In these conditions, it is assumed that the evaporation is negligible and that LE only corresponds to Tr.  $Tr_{SIF}$  was converted from  $\text{Wm}^{-2}$  to  $\text{mmolm}^{-2}\text{s}^{-1}$  using the latent heat of vaporization and the molar mass of  $\text{H}_2\text{O}$ .

## 4.8. Assessment of edaphic water stress

### 4.8.1. Relative extractable water

SWC was used to calculate the REW which represents the amount of water available for plant uptake in the root zone (Granier *et al.*, 1999a). The original equation of (Granier *et al.*, 1999a) has been modified to weight the amount of available water by the rooting depth within each soil layer, therefore providing a more realistic estimation of the soil water availability in the root zone. The detail of the calculation of REW can be found in the supplementary materials.

### 4.8.2. Relationship with ecosystem physiology

Physiological processes such as carbon assimilation, stomatal closure or SIF emission often show a two-steps response to the decrease in soil water availability with a clear transition from non-limiting to limiting soil water conditions from a threshold (Beauclaire *et al.*, 2023b; Gourlez de la Motte *et al.*, 2020; Jonard *et al.*, 2022). In this study, soil water-limiting conditions were defined by using the REW threshold from which a decrease in LUE was observed (Reitz *et al.*, 2023). As any other response ratio, LUE is strongly affected by changes in the denominator ( $APAR$ ) especially when it is small. Therefore, a logarithmic transformation was used to avoid skewness in the sampling distribution and to linearize the ratio so that LUE is affected equally by changes in  $APAR$  or  $GPP_{EC}$  (Hedges *et al.*, 1999). A linear and a linear-plateau segmented model were fitted to the  $\log_{10}$ -transformed daily means of LUE and daily means of REW, and statistical selection was made according to the lowest AIC similarly to the  $GPP_{EC} - SIF_{TOT}$  relationship (Stocker *et al.*, 2018; Jonard *et al.*, 2022; Reitz *et al.*, 2023). The linear-plateau segmented model had the lowest AIC and the REW threshold ( $REW_{th}$ ) was determined as the corresponding REW below which a decrease in REW induced a decrease in LUE. A threshold of  $0.57 \pm 0.06$  was found (Fig. 6.13). This lower limit of this threshold (i.e.,  $REW_{th}=0.51$ ) was used to discuss the impact of non-limiting (i.e.,  $REW > REW_{th}$ ) or limiting (i.e.,  $REW < REW_{th}$ ) soil water conditions on MLR- USO model robustness and on the  $GPP_{EC} - SIF_{TOT}$  relationship.

## 4.9. Gas exchange and fluorescence measurements at the leaf level

$q_L$ ,  $\phi_{PSII,max}$ ,  $C_i$  and  $f_{PSII}$  used in Eq. 6.1 and Eq. 6.8, as well as  $A_n$ ,  $C_s$ , and  $VPD_{leaf}$  in Eq. 6.13 were determined from active fluorescence and gas exchange measurements at the leaf level by using a LICOR LI-6400 device and a LI-6400-40 fluorescence chamber (LI-COR Inc., Lincoln, NE, USA). The following procedure was applied to each measurement. Only the youngest and most exposed leaves on the upper part of the canopy were selected, considering that they mostly contributed to the total emitted SIF signal due to their very large role in PAR interception, and their high content in chlorophyll and nitrogen (Li *et al.*, 2015). In the chamber, the  $CO_2$  concentration was set to 400 ppm, and the air humidity and temperature were maintained at ambient level. For the measurements on dark-adapted materials, the actinic light was turned off for at least 20 minutes. A multiphase fluorescence flash of  $9\,000\ \mu\text{mol m}^{-2}\ \text{s}^{-1}$  was then applied to measure the maximum and minimum fluorescence intensities in the dark ( $F_m$  and  $F_0$ ). After this, the actinic light was turned on and the PAR in the chamber ( $PAR_{leaf}$ ) was set to ambient level. After stabilization of the fluorescence and gas exchange signal, the steady state fluorescence intensity ( $F_s$ ) was measured before the application of the same multiphase fluorescence flash to measure the maximum fluorescence intensity under the light ( $F'_m$ ). The minimum fluorescence of a light-adapted sample ( $F'_0$ ) was measured after the application of a far-red pulse to excite PSI and draw the electrons from PSII to ensure that reaction centers remain fully oxidized. In addition,  $A_n$ , leaf transpiration, leaf temperature ( $T_{leaf}$ ) and  $C_s$  were measured. These variables were used to calculate  $g_{sw}$ ,  $C_i$  and  $VPD_{leaf}$  (Eq. 6.13). Leaf-level measurements were performed at different locations within the FloX footprint between 7 am and 5 pm. A total of 387 measurements were collected between April 14 and June 30, 2022 (Fig. 2). These fluorescence intensities should correspond only to PSII emission. This has been considered by hypothesizing that steady-state PSI fluorescence represents 24% of  $F_s$  (i.e.,  $F_{s,PSI} = 0.24F_s$  when performing active fluorescence measurements at the leaf level Pfündel *et al.* (2013)). This contribution reaches 45% in the NIR for SIF in the O<sub>2</sub>-A band (i.e.,  $F_{s,PSI,NIR} = 0.45F_s$ , Pfündel (2021)). Therefore,  $q_L$ ;  $\phi_{PSII,max}$  and  $f_{PSII}$  were calculated as following (Kramer *et al.*, 2004; Jia *et al.*, 2023)

$$q_L = \frac{F'_m - F_{s,PSI}}{F'_m - F'_0} \frac{F'_0}{F_{s,PSI}} \quad (6.14)$$

$$\phi_{PSII,max} = \frac{F_m - F_0}{F_m - F_{s,PSI}} \quad (6.15)$$

$$f_{PSII} = \frac{F_s - F_{s,PSI,NIR}}{F_s} = 1 - \frac{F_{s,PSI,NIR}}{F_s} \quad (6.16)$$



It is necessary to interpolate MLR model parameters determined by leaf-level measurements campaigns ( $q_L$ ,  $\phi_{PSII,max}$ ,  $C_i$  and  $f_{PSII}$ ) to calculate  $GPP_{SIF}$  (Eq. 6.8) and  $Tr_{SIF}$  (Eq. 6.11) at a half-hourly timescale before validation. This temporal interpolation has been carried out using the relationship between model parameters and potential environmental drivers ( $VPD_{leaf}$ ,  $PAR_{leaf}$ ,  $T_{leaf}$ , REW). An exponential model was used to represent the relationship between  $q_L$  and  $PAR_{leaf}$  (Chang *et al.*, 2021). As  $C_i$  is sensitive to leaf temperature with a peak between 20 °C and 30 °C for wheat (Huang *et al.*, 2021), a second-order polynomial model was used to interpolate  $C_i$  with  $T_{leaf}$ .  $f_{PSII}$  was linearly related to  $PAR_{leaf}$  by a first-order linear function. The equations are:

$$q_L = a_{qL} e^{-b_{qL} PAR_{leaf}} + c_{qL} \quad (6.17)$$

$$f_{PSII} = a_{fPSII} PAR_{leaf} + b_{fPSII} \quad (6.18)$$

$$C_i = a_{Ci} T_{leaf}^2 + b_{Ci} T_{leaf} + c_{Ci} \quad (6.19)$$

$q_L$ ,  $f_{PSII}$  and  $C_i$  were grouped in  $PAR_{leaf}$ ,  $T_{leaf}$  and  $VPD_{leaf}$  classes with an interval of respectively 50  $\mu\text{molm}^{-2}\text{s}^{-1}$ , 2°C and 0.1 kPa before fitting the exponential and polynomial models. Statistical significance of models was evaluated by comparing the AIC of the exponential (Eq. 6.17) and polynomial (Eq. 6.18, Eq. 6.19) models with the AIC of a linear (for Eq. 6.17, Eq. 6.19) and a zero-slope model (for Eq. 6.18). Only models with the lowest AIC are presented. The relationship between model parameters and other potential drivers (i.e., REW,  $PAR_{leaf}$ ,  $T_{leaf}$  and  $VPD_{leaf}$ ) was also tested (Fig. S3, Table 6.3).

#### 4.10. Energy partitioning at the leaf level

As mentioned in the introduction, NPQ, SIF and PQ are in competition for the incoming irradiance. Therefore, the sum of the three corresponding quantum yields ( $\phi_{NPQ}$ ,  $\phi_{SIF}$  and  $\phi_{PQ}$ ) equals 1 (Porcar-Castell *et al.*, 2014) and the relationship between carbon assimilation and fluorescence at the leaf level is tightly coupled to the dissipation of excessive energy by heat. It is then expected that the partitioning between  $\phi_{NPQ}$ ,  $\phi_{SIF}$  and  $\phi_{PQ}$  at the leaf level directly impacts the relationship between  $SIF_{TOT}$  and  $GPP_{SIF}$ . The quantum yields were calculated following (Gu *et al.*, 2019; Kramer *et al.*, 2004; Jia *et al.*, 2023):

$$\phi_{PQ} = \frac{F'_m - F_s}{F'_m - F_{s,PSI}} \quad (6.20)$$

$$\phi_{NPQ} = \frac{F_s - F_{s,PSI}}{F'_m - F_{s,PSI}} - \frac{F_s - F_{s,PSI}}{F_m - F_{s,PSI}} \quad (6.21)$$

$$\phi_{SIF} = \frac{1 - \phi_{PSII,max}}{(1 + k_{DF}) \left( \left( 1 + \frac{F_m - F'_m}{F'_m} \right) (1 - \phi_{PSII,max}) + q_L \phi_{PSII,max} \right)} \quad (6.22)$$

where  $NPQ = (F_m - F'_m)/(F'_m - F_{(s,PSI)})$  is the NPQ parameter. Note that  $\phi_{NPQ}$  and  $\phi_{PQ}$  were corrected for PSI emission and that  $\phi_{SIF}$  corresponds to the SIF quantum yield of PSII (Gu *et al.*, 2019). The relationship between  $\phi_{NPQ}$ ,  $\phi_{PQ}$ ,  $\phi_{SIF}$  and environmental drivers was evaluated using the correlation coefficients, and their corresponding p-values

### 4.11. Statistical analysis

The performance of the MLR- USO model was evaluated by calculating the  $R^2$  and the relative root mean square error (rRMSE) between  $GPP_{SIF}-GPP_{EC}$  and between  $Tr_{SIF}-Tr_{EC}$ , at both daily and half-hourly timescales. The rRMSE was calculated by dividing the root mean square error by the amplitude of EC data (i.e., the difference between the maximum and minimum value). The heteroskedasticity was assessed by calculating the p-value of the Breusch-Pagan test (Breusch & Pagan, 1979) to characterize a potential trend in residuals variability with  $GPP_{SIF}$  or  $Tr_{SIF}$ . Finally, the relationship between model residuals and potential drivers (REW,  $VPD_{can}$ ,  $T_{can}$ ,  $f_{esc}$ ) was evaluated using the correlation coefficients, and their corresponding p-values.

## 5. Results

### 5.1. Meteorological conditions and EC fluxes

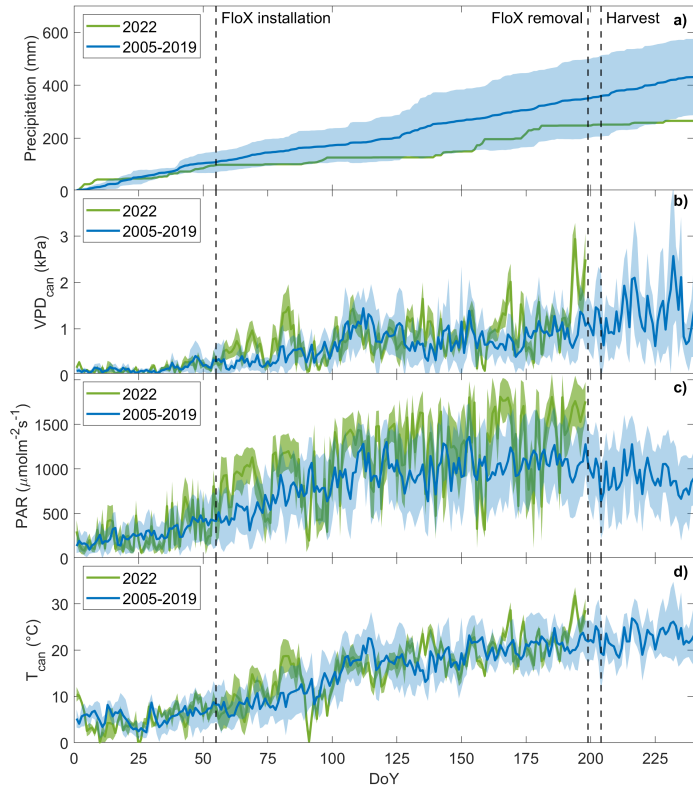
CO<sub>2</sub>, H<sub>2</sub>O fluxes and meteorological variables in 2022 were compared to the previous winter wheat growing seasons at BE-Lon by combining the FLUXNET2015 and the ICOS Carbon Portal datasets over the years 2005, 2007, 2009, 2011, 2013 and 2019 (Dumont *et al.*, 2023a; Pastorello *et al.*, 2020). The temporal evolution of cumulative precipitation,  $VPD_{can}$ , PAR and  $T_{can}$  in 2022 is presented in Fig. 6.1 and was compared to the average values over the reference period 2005-2019. In 2022, several episodes with very low precipitation can be observed, as highlighted by numerous horizontal lines on the cumulative precipitation curve (Fig. 6.1-a). At the harvest, the total cumulative precipitation in 2022 was about 251 mm, which is not significantly lower than the average cumulative precipitation for the reference period 2005-2019 ( $361 \pm 153$  mm).

Four peaks of  $VPD_{can}$  were observed in 2022 around DOY 83, 135, 169, 194, clearly above the average values of the reference period (Fig. 1-b). These sharp increases in  $VPD_{can}$  also corresponds to an increase in  $T_{can}$  and PAR (Fig. 6.1-c,d). The lack of precipitation during spring and summer 2022 caused a decrease in REW between DOY 55 to 85, DOY 98 to 139, and DOY 160 to 173 (Fig. 6.2-c). The most severe precipitation shortage episode was between DOY 98 to 139, which corresponds to more than 40 days without precipitation. During this period, REW decreased to a minimum value of 0.24. The lowest REW value was 0.14 at DOY 197 (Fig. 6.2-c).  $GPP_{EC}$  and  $LE_{EC}$  showed a strong seasonal pattern with a typical bell-shape curves in parallel with canopy cover development and NDVI dynamics (Fig. 6.2-a,b; Fig. 6.3-a). Maximum values of  $55.0 \mu\text{molm}^{-2}\text{s}^{-1}$  for  $GPP_{EC}$  and  $419 \text{ Wm}^{-2}$  for  $LE_{EC}$  were observed at DOY 133 (Fig. 6.2-a,b).

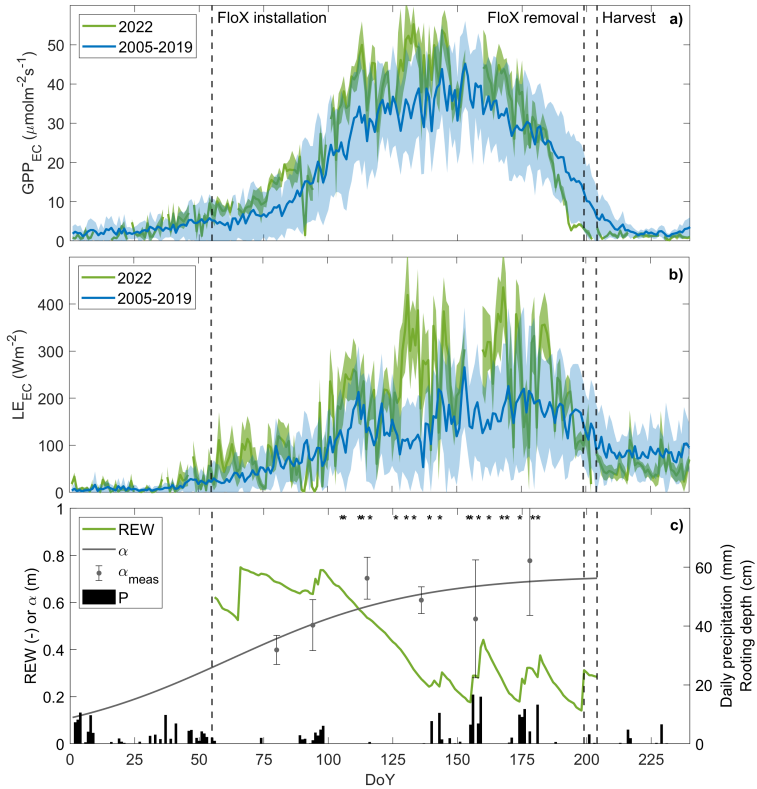
### 5.2. Temporal evolution of spectral data

Due to a malfunction of the QEPro spectrometer after DOY 175, only SIF data from DOY 55 to DOY 175 were analyzed. Moreover, data gaps between DOY 141-152 and 159-166 were caused by sensor maintenance (Fig. 6.3), and errant NDVI data between DOY 98 and 110 were replaced by data from the NDVI sensor installed on the EC station (laboratory-made sensor, see Soudani *et al.* (2012) for references). NDVI,  $NIR_v$  and  $f_{APAR}$  followed vegetation development with a gradual increase from DOY 55 to DOY 130 and a maximum value of respectively 0.89, 0.50 and 0.89 (Fig. 6.3-a,b), which characterizes the full development of the canopy. A progressive decrease due to senescence and yellowing can be observed from DOY 182 to DOY 198. The variability of  $f_{esc}$  was high before DOY 80, corresponding to the beginning of the growing season. Indeed,  $f_{esc}$  is very sensitive to short scale variations of NDVI (Eq. 6.2). This latter showed a strong variability in the early stage of vegetation growth, which can be

explained by the impact of soil reflectance and canopy scattering on the NDVI measurements. Therefore, NDVI measurements are very sensitive to changes in canopy structural properties (i.e., such as leaf angle inclination) when vegetation cover is low (Atherton *et al.*, 2022). This might explain the erratic pattern of  $f_{esc}$  before DOY 80. From DOY 80,  $f_{esc}$  followed crop phenology as NDVI and  $NIR_v$  increased, with values stabilizing between 0.50 and 0.39 (Fig. 6.3-b), which has already been observed in similar crops (Dechant *et al.*, 2020).  $SIF_{TOT}$  was impacted by both canopy cover development and half-hourly variations of meteorological conditions (Fig. 6.3-c). In particular,  $SIF_{TOT}$  gradually increased from DOY 55 to a maximum value of  $9.34 \mu\text{molm}^{-2}\text{s}^{-1}$  observed at DOY 123 (Fig. 6.3-c). The relationship between  $GPP_{EC}$  and  $SIF_{TOT}$  was characterized by an increase in  $GPP_{EC}$  for  $SIF_{TOT} \leq 6.90 \mu\text{molm}^{-2}\text{s}^{-1}$  before a progressive stabilization of  $GPP_{EC}$  at  $\sim 45.6 \mu\text{molm}^{-2}\text{s}^{-1}$  for  $SIF_{TOT} \geq 6.90 \mu\text{molm}^{-2}\text{s}^{-1}$  (Fig. 6.4-a). The nonlinearity was observed at very high  $APAR$  ( $\sim 1500 \mu\text{molm}^{-2}\text{s}^{-1}$ , corresponding to 3.6% of the  $APAR$  dataset) which may suggest a decoupling between  $GPP_{EC}$  and  $SIF_{TOT}$  induced by light absorption and/or canopy structure effects. The relationship between LUE and  $SIF_y$  was linear (Fig. 6.4-b), which highlights that the saturating shape was explained by light absorption rather than structural changes (Fig. 6.4-b). Finally, soil water-limiting conditions (i.e.,  $REW \leq REW_{th}$ ) did not induce a nonlinear pattern between  $SIF_y$  and LUE (Fig. 6.5).

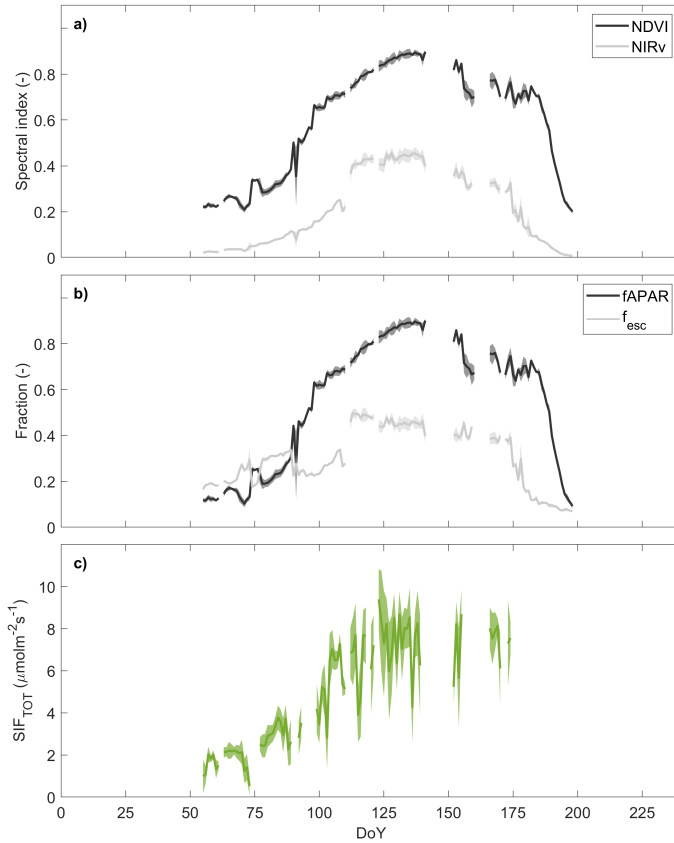


**Figure 6.1:** Temporal evolution in 2022 of daily cumulative precipitation (panel a), canopy vapor pressure deficit at the canopy surface ( $VPD_{can}$  - panel b), photosynthetic active radiation (PAR – panel c) and canopy temperature ( $T_{can}$  – panel d) compared to previous winter wheat growing seasons (2005-2019). Shaded areas correspond to the standard deviation around the daily means.

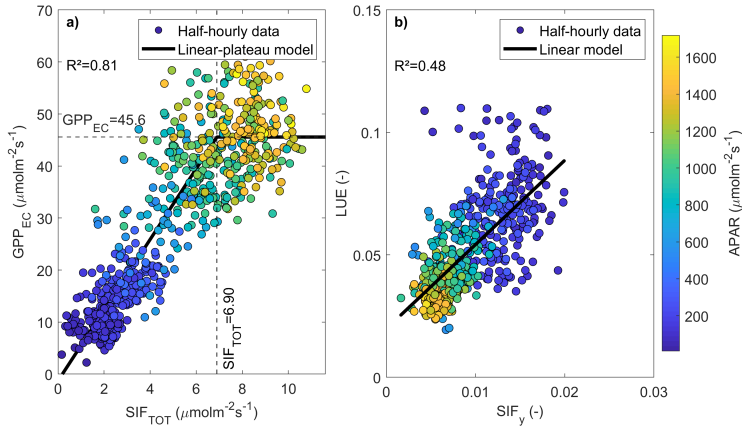


**Figure 6.2:** Temporal evolution in 2022 of gross primary production ( $GPP_{EC}$  – panel a), latent heat flux ( $LE_{EC}$  - panel b) and relative extractable water (REW- panel c). Subplot c) also displays maximum rooting depth measurements ( $\alpha_{meas}$ ) and extrapolated values ( $\alpha$ ), as well as daily cumulative precipitation (P) and days when leaf level measurements were performed (\*). In addition, panels a) and b) also display  $GPP_{EC}$  and  $LE_{EC}$  over the previous winter wheat growing seasons (2005-2019). Shaded areas correspond to the standard deviation around the daily means. The details of the calculation of REW are given in Text S1.

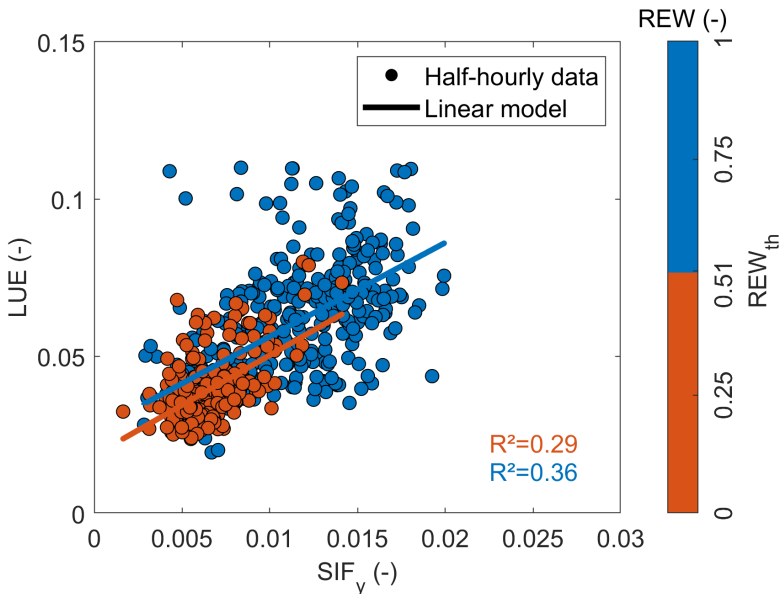
Chapter 6. Modeling gross primary production and transpiration from sun-induced chlorophyll fluorescence using a mechanistic light-response approach



**Figure 6.3:** Temporal evolution in 2022 of daily means of normalized difference vegetation index (NDVI) and near-infrared reflectance of vegetation ( $NIR_v$  – panel a), fraction of absorbed irradiance ( $f_{APAR}$ ), escape probability ( $f_{esc}$  – panel b), and total sun-induced chlorophyll fluorescence ( $SIF_{TOT}$  – panel c). Shaded areas correspond to the standard deviation around the daily means.



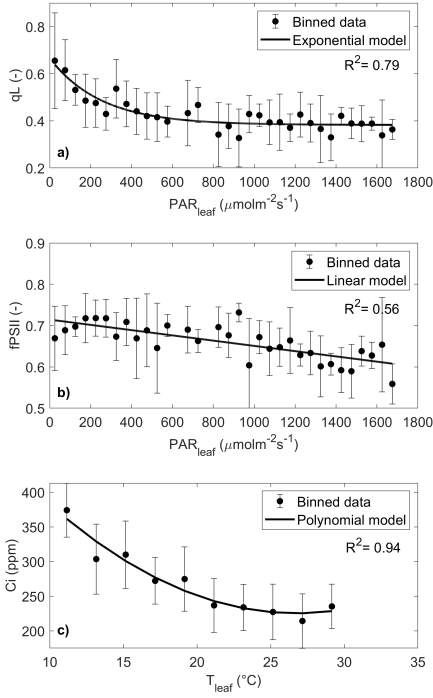
**Figure 6.4:** Relationship between gross primary production ( $GPP_{EC}$ ) and total sun-induced fluorescence ( $SIF_{TOT}$  - panel a), and between LUE and SIF yield ( $SIF_y$  - panel b). The color map indicates the range of absorbed PAR (APAR).  $R^2$  is the coefficient of determination of the linear or linear-plateau segmented model. The selected model was the one with the lowest AIC.



**Figure 6.5:** Relationship between LUE and SIF yield ( $SIF_y$ ) with relative extractable water (REW).  $R^2$  is the coefficient of determination of the linear model. The color map indicates the range of REW and is defined according to the REW threshold characterizing soil water-limiting conditions ( $REW_{th}$ )



### 5.3. Leaf-level measurements



**Figure 6.6:** Relationship between the fraction of open PSII centers ( $q_L$ ) and the fraction of PSII fluorescence emission ( $f_{PSII}$ ) with photosynthetic active radiation at the leaf level ( $PAR_{leaf}$  - panels a and b). Panel c) displays the relationship between CO<sub>2</sub> concentration in sub-stomatal cavities ( $C_i$ ) and leaf temperature ( $T_{leaf}$ ). Measurements are grouped into  $PAR_{leaf}$  or  $T_{leaf}$  classes with an interval of respectively  $50 \mu\text{molm}^{-2}\text{s}^{-1}$  and  $2^\circ\text{C}$ . The error bars show the standard deviation around the mean for each class.  $R^2$  is the coefficient of determination of the exponential or polynomial models.

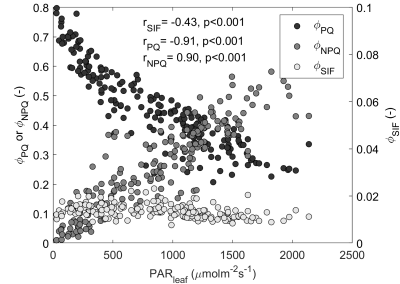
did not affect leaf-level quantum yields (Fig. 6.16).

$T_{leaf}$  explained a larger part of the variability of  $C_i$  compared to  $VPD_{leaf}$  and  $PAR_{leaf}$  (i.e., highest  $R^2$ , Fig. 6.6, Fig. 6.14, Table 6.2).  $\phi_{PSII,max}$  remained constant and was not affected by any environmental variables ( $\phi_{PSII,max}=0.81$ , Fig. S3).  $q_L$  exponentially decreased from 0.63 before stabilizing to 0.38 when PAR exceeded  $\sim 500 \mu\text{molm}^{-2}\text{s}^{-1}$  (Fig. 6.6-a), and  $f_{PSII}$  linearly decreased between 0.71 and 0.60 (Fig. 6.6-b).  $C_i$  decreased with increasing  $T_{leaf}$  up to  $\sim 20^\circ\text{C}$  before stabilizing above  $\sim 20^\circ\text{C}$  (Fig. 6.6-c). Fitted coefficients and  $R^2$  of the regressions (Eq. 6.17, 6.18 and 6.19) are given in Table S3. No significant relationship between  $q_L$ ,  $f_{PSII}$  and other environmental drivers was observed (Fig. 6.14). These equations were used to calculate  $q_L$ ,  $C_i$  and  $f_{PSII}$  at the half-hourly timescale using  $T_{can}$  (Eq. 6.10) and PAR for estimating  $GPP_{SIF}$  and  $SIF_{TOT}$ . Time series of leaf-level measurements is shown in Fig. 6.15.

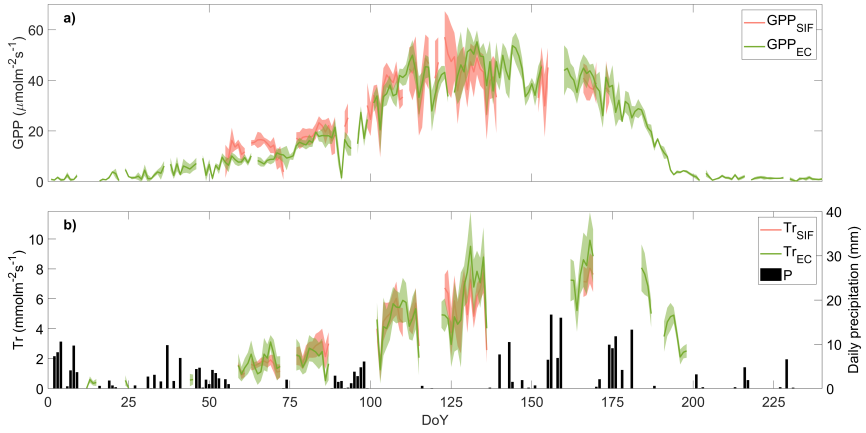
A positive correlation between PAR and  $\phi_{NPQ}$  was observed (Fig. 6.7), where  $\phi_{NPQ}$  increased with PAR from  $\sim 0$  to  $\sim 0.6$ . On the other hand,  $\phi_{PQ}$  decreased with increasing PAR from  $\sim 0.8$  to  $\sim 0.3$ , which resulted in a negative correlation (Fig. 6.7). A weaker negative correlation between  $\phi_{SIF}$  and PAR was also evidenced (Fig. 6.7). A significant correlation was also observed between leaf-level quantum yields and  $T_{leaf}$  and  $VPD_{leaf}$  (Fig. S5). In particular, the decrease in  $\phi_{SIF}$  and  $\phi_{PQ}$  with  $T_{leaf}$  and  $VPD_{leaf}$  was associated to the corresponding increase in  $\phi_{NPQ}$ . Moreover, the decrease in REW

## 5.4. MLR- USO model performances for estimating GPP and transpiration

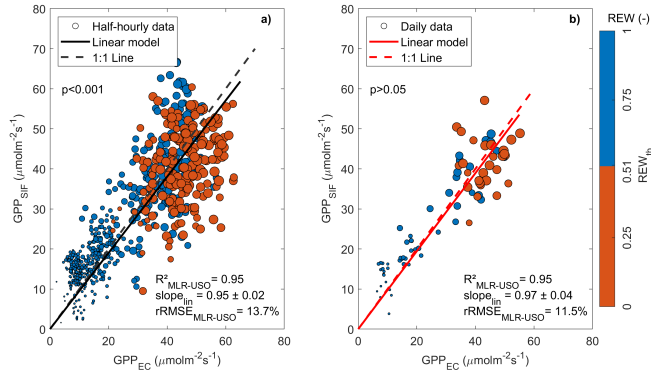
The MLR- USO model reproduced a large proportion of the temporal variability of  $GPP_{EC}$  and  $Tr_{EC}$  throughout the growing season and under various meteorological conditions (Fig. 6.8). These results are confirmed by the direct comparison between modeled GPP and Tr ( $GPP_{SIF}$  and  $Tr_{SIF}$ ) and EC measurements (Fig. 6.9 and Fig. 6.10). The MLR- USO model explained 95% of the variance of  $GPP_{EC}$  at the half-hourly timescale ( $R^2=0.95$ ) with a rRMSE of 13.7% and a slope of the linear regression between  $GPP_{EC}$  and  $GPP_{SIF}$  close to 1 (Fig. 6.9-a). At the daily timescale, the rRMSE was slightly lower, while the other statistics remained very similar ( $R^2=0.95$ , rRMSE=11.5%, slope= $0.97 \pm 0.04$ ) (Fig. 6.9-b). At the half-hourly timescale, 95% of the variance of  $Tr_{EC}$  was explained by the model ( $R^2=0.95$ ) with a rRMSE of 9.1% and a slope of  $0.89 \pm 0.02$  (Fig. 6.10-a). Similar statistics were observed at the daily timescale (Fig. 6.10-b). Low REW induced an underestimation of  $Tr_{SIF}$  compared to  $Tr_{EC}$ , which was more pronounced at the half-hourly timescale (Fig. 6.10). No deviation from the 1:1 line as  $APAR$  increased can be observed for both GPP and Tr (Fig. 6.9 and Fig. 6.10). The mean of residuals was very close to 0 for GPP ( $0.36 \mu\text{molm}^{-2}\text{s}^{-1}$  half-hourly /  $0.63 \mu\text{molm}^{-2}\text{s}^{-1}$  daily) and Tr ( $-0.11 \text{mmolm}^{-2}\text{s}^{-1}$  half-hourly /  $-0.15 \text{mmolm}^{-2}\text{s}^{-1}$  daily). An increase in residuals variability as  $GPP_{SIF}$  and  $Tr_{SIF}$  increased was observed, but only at the half-hourly timescale for GPP (i.e.,  $p<0.05$ , Fig. 6.9 and Fig. 6.10). GPP residuals were significantly correlated with REW,  $T_{can}$  and  $VPD_{can}$  but not with  $f_{esc}$  (Fig. 6.11-a,b,c,d). Similar correlations are observed for Tr residuals (Fig. 6.11-e,f,g,h). More specifically, the correlation was positive between model residuals and REW, and negative for  $T_{can}$  and  $VPD_{can}$ . A more important trend towards negative residuals for Tr was observed when REW was lower than  $REW_{th}$ , which can explain the slight underestimation of  $Tr_{SIF}$  compared to  $Tr_{EC}$  (Fig. 6.10). No correlation was observed at the daily timescale for any GPP or Tr residuals.



**Figure 6.7:** Relationship between quantum yields of fluorescence ( $\phi_{SIF}$ ), non-photochemical quenching ( $\phi_{NPQ}$ ) and photochemical quenching ( $\phi_{PQ}$ ) with photosynthetic active radiation at the leaf level ( $PAR_{leaf}$ ).  $r$  is the correlation coefficient and  $p$  is the corresponding p-value.



**Figure 6.8:** Temporal evolution in 2022 of gross primary production from eddy covariance ( $GPP_{EC}$ ) and estimated by the MLR- USO model ( $GPP_{SIF}$  - panel a). Panel b) shows the temporal evolution of daily transpiration from eddy covariance ( $Tr_{EC}$ ) and estimated by the MLR- USO model ( $Tr_{SIF}$ ), as well as the daily cumulative precipitation (P). Shaded areas correspond to the standard deviation around the daily means



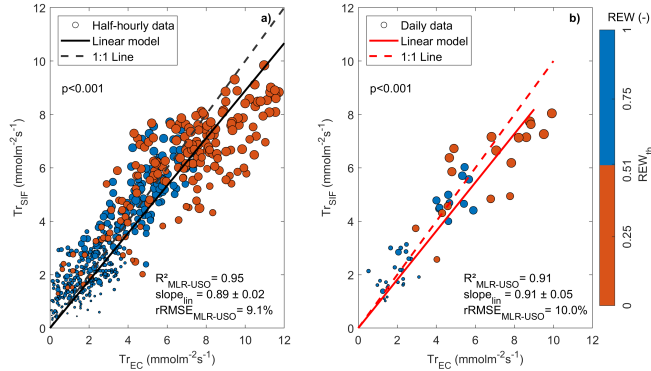
**Figure 6.9:** Scatterplot between gross primary production from eddy covariance ( $GPP_{EC}$ ) and gross primary production estimated by the MLR- USO model ( $GPP_{SIF}$ ) at the half-hourly (panel a) and daily (panel b) timescales. Dot size is proportional to the absorbed photosynthetic active radiation (APAR). The color map indicates the range of relative extractable water (REW) and is defined according to the REW threshold characterizing soil water-limiting conditions ( $REW_{th}$ ).  $R^2_{MLR-USO}$  and  $rRMSE_{MLR-USO}$  are respectively the coefficient of determination and the relative root mean squared error calculated from  $GPP_{SIF}$  and  $GPP_{EC}$ .  $slope_{lin}$  is the slope of the linear relationship between  $GPP_{SIF}$  and  $GPP_{EC}$ .  $p$  is the  $p$ -value of the Breusch-Pagan test for heteroskedasticity

## 6. Discussion

### 6.1. Relationship between MLR- USO model parameters and environmental drivers

A key step for applying the MLR model is the determination of the factors for estimating  $SIF_{TOT}$  from  $SIF_{TOC,760}$ . This step includes the conversion of photosystem-level SIF emission to canopy level, the conversion of narrowband SIF to broadband SIF considering a physiological and structural component, and finally the exclusion of PSI fluorescence.  $f_{\lambda}$  is constant and does not vary with meteorological variables while  $f_{esc}$  integrates the structural changes in the canopy and was retrieved from  $NIR_v$  and  $f_{APAR}$ , which are both reflectance-based measurements available at local and large scales (Zeng *et al.*, 2019; Dechant *et al.*, 2022). As the SIF signal includes the contribution of both photosystems in the O<sub>2</sub>-A band, the effect of PSI cannot be neglected.  $f_{PSII}$  was estimated by excluding the contribution of PSI to leaf-level fluorescence intensities which was set to 45% of  $F_s$  (Pfündel, 2021). Although this value is higher than previous observations (e.g., Genty *et al.* (1990); Franck *et al.* (2002); Pfündel *et al.* (2013)), it can be used to consider PSI contribution in the O<sub>2</sub>-A band as it was calculated from selective fluorescence measurements above 700 nm (Pfündel, 2021; Jia *et al.*, 2023).  $f_{PSII}$  moderately decreased with PAR, while being not affected by the increase in  $VPD_{leaf}$ ,  $T_{leaf}$  or REW (Fig. 6.6, Fig. S3). This dynamic is similar to the PAR- $\phi_{SIF}$  relationship which showed a weak negative correlation (Fig. 6.7) and can be explained by a strong coupling between  $F_s$  and PQ quantum yield under high irradiance when NPQ is limiting (Maguire *et al.*, 2020; Liu *et al.*, 2022c). This contrasts with other studies which modeled the fluorescence quantum yields ratio (Eq. 6.16) with temperature, PAR and CO<sub>2</sub> availability, making  $f_{PSII}$  highly sensitive to other environmental conditions (Liu *et al.*, 2022c; Bacour *et al.*, 2019; van der Tol *et al.*, 2014). Disentangling the importance of PSI contribution in the total SIF signal is one important issue to be considered in future studies (Porcar-Castell *et al.*, 2021), as it directly influences the magnitude of  $SIF_{TOT}$  especially in the O<sub>2</sub>-A band.

Given the wide diversity in physiological mechanisms across ecosystems, it is expected that the parameters of the MLR- USO model (i.e.,  $\phi_{PSII,max}$ ,  $q_L$ ,  $C_i$  and  $g_1$ ) are to be determined from a PFT-based calibration. We tackled this problem by taking in-situ measurements at the leaf level.  $T_{leaf}$  explained the most important part of  $C_i$  variability (highest R<sup>2</sup>, Fig. 6.6, Table 6.3) which is explained by the temperature response of stomata and Rubisco (Warren, 2006a). This combined effect leads to a temperature optimum for photosynthesis, typically around 25°C (Crous *et al.*, 2022), with some variability between species (Marchin *et al.*, 2022). Such relationship was also observed in other studies conducted on winter wheat (Urban *et al.*, 2018; Huang *et al.*, 2021).



**Figure 6.10:** Scatterplot between transpiration from eddy covariance ( $Tr_{EC}$ ) and gross primary production estimated by the MLR-USO model ( $Tr_{SIF}$ ) at the half-hourly (panel a) and daily (panel b) timescales. Dot size is proportional to the Absorbed Photosynthetically Active Radiation (APAR). The color map indicates the range of REW and is defined according to the REW threshold characterizing soil water-limiting conditions ( $REW_{th}$ ).  $R^2_{MLR-USO}$  and  $rRMSE_{MLR-USO}$  are respectively the coefficient of determination and the relative root mean squared error calculated from  $Tr_{SIF}$  and  $Tr_{EC}$ .  $slope_{lin}$  is the slope of the linear relationship between  $Tr_{SIF}$  and  $Tr_{EC}$ .  $p$  is the  $p$ -value of the Breusch-Pagan test for heteroskedasticity

The exponential decrease in  $q_L$  with PAR illustrates the closure of PSII centers when solar irradiance increases (Cendrero-Mateo *et al.*, 2015; Zivcak *et al.*, 2014; Flexas & Medrano, 2002; Brodrick *et al.*, 2022; Ouzounis *et al.*, 2015). We did not observe an effect of other environmental variables on  $q_L$  besides PAR (Fig. 6.14), in contrast with previous studies which showed a significant effect of heat and atmospheric dryness on this parameter (Han *et al.*, 2022b; Shin *et al.*, 2021; Zhou *et al.*, 2019a; Cano *et al.*, 2013). However, it might be possible that our experimental set-up didn't allow the identification of such effects as chlorophyll fluorescence measurements were carried out along with gas exchanges, which limited the number of measurements throughout the day and the growing season. Expanding the range of environmental variables under controlled conditions could uncover such additional effects (e.g., Han *et al.* (2022b)).

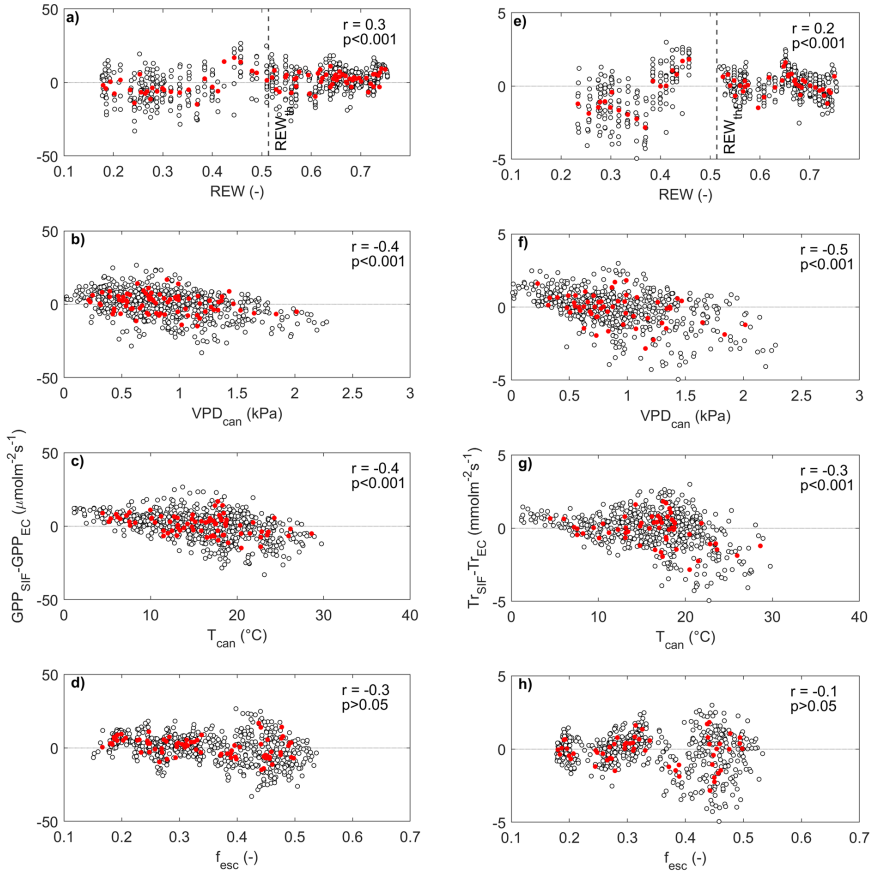
## 6.2. SIF-GPP relationship

The saturation of  $GPP_{EC}$  at high  $SIF_{TOT}$  (Fig. 6.4-a) and the linearity between LUE and  $SIF_y$  (Fig. 6.4-b) indicates that APAR is an important driver of the dynamics of  $GPP_{EC}$  and  $SIF_{TOT}$ . Such relationship has also been observed for croplands, savannas (Yao *et al.*, 2022) or needleleaf forests (Yao *et al.*, 2022; Kim *et al.*, 2021). This nonlinearity can be explained by the competitive mechanism between  $\phi_{NPQ}$ ,  $\phi_{PQ}$  and  $\phi_{SIF}$  at the leaf level for the incoming irradiance. We found an asymmetric pattern of  $\phi_{PQ}$  and  $\phi_{NPQ}$  dynamics with PAR (Fig. 6.7) as highlighted by numerous

studies (Baker, 2008; van der Tol *et al.*, 2014; Chen *et al.*, 2019b; Gu *et al.*, 2019). In light-limiting conditions, most of the incoming PAR is used for fueling the dark reactions of photosynthesis by carrying the energy through the electron transport chain to produce ATP and NADPH, resulting in a high PQ at low PAR. As more energy is received by chlorophyll molecules, PSII reaction centers close (decrease in  $q_L$ ) and the rate of absorbed photons decreases (decrease in  $\phi_{PQ}$ ). In these conditions, numerous regulatory mechanisms are used to dissipate the excess energy as heat to prevent the formation of reactive oxygen species (see Demmig-Adams & Adams III (2006) for a review), resulting in an increase in  $\phi_{NPQ}$  with increasing PAR. As  $\phi_{SIF}$  depends on  $\phi_{NPQ}$  and  $q_L$  (Eq. 6.22), the opposite behavior of these two parameters with PAR explains the weaker sensitivity of  $\phi_{SIF}$  to PAR (Fig. 6.7), and the nonlinearity between  $GPP_{EC}$  and  $SIF_{TOT}$  at high APAR (Fig. 6.4-a, Gu *et al.* (2019)).

The linearity between  $SIF_y$  and  $LUE$  was also observed under soil water-limiting conditions (i.e.,  $REW \leq REW_{th}$ , Fig. 6.5). Similar results were shown in deciduous forests (He *et al.*, 2019; Yang *et al.*, 2015), winter wheat crop (Shen *et al.*, 2022) or natural grassland (Verma *et al.*, 2017). A key factor impacting  $SIF_y$  and  $LUE$  is the maximum carboxylation rate of Rubisco ( $V_{cmax}$ ). A decrease in  $V_{cmax}$  is often observed in crops during drought (e.g., Beauclaire *et al.* (2023b); Zhou *et al.* (2013)), which could explain the decrease in LUE with REW as highlighted in this study (Fig. S2). Moreover, numerous studies using the SCOPE model (van der Tol *et al.*, 2009) have shown that SIF is sensitive to  $V_{cmax}$  (Zhang *et al.*, 2018, 2014; Koffi *et al.*, 2015; Wang & Xiao, 2021; Zhang *et al.*, 2016; Camino *et al.*, 2019), suggesting that  $SIF_y$  may decrease as well when soil water availability is reduced. Therefore, the linearity between  $SIF_y$  and LUE could be explained by a similar sensitivity of LUE and  $SIF_y$  to  $V_{cmax}$  during water stress.

Unlike the saturation of  $GPP_{EC}$  at high APAR and  $SIF_{TOT}$  (Fig. 6.4-a), a linear relationship between  $GPP_{SIF}$  and  $SIF_{TOT}$  was observed (Fig. 6.16-d). This suggests that the model did not capture the saturation of  $GPP_{EC}$  at high APAR. However, the corresponding impact on model performance was limited as only 3.6% of the APAR data was above the saturation point of  $\sim 1500 \mu\text{molm}^{-2}\text{s}^{-1}$  (Fig. 6.4-a). In the MLR model, the influence of PAR on  $GPP_{SIF}$  is implemented through the effects of  $q_L$  on  $J_{SIF}$  (Eq. 1.42, Eq. 6.14). This sensitivity is explained by the coupling between the dark and light reactions of photosynthesis. In particular, the production of energy-carrier molecules (i.e., ATP and NADPH) by the electron transport chain (which depends on  $q_L$ ) must balance carbon uptake, itself regulated by the stomatal dynamics (Gu *et al.*, 2019). As  $SIF_{TOT}$  is the product of  $\phi_{SIF}$ ,  $f_{esc}$ , and APAR (Gu *et al.*, 2019) it is expected that  $SIF_{TOT}$  increases with PAR. On the other hand, the decrease in  $q_L$  (Eq. 6.14, Fig. 6.6-a) compensates for the corresponding increase in  $SIF_{TOT}$  in the MLR model (i.e.,  $GPP_{SIF} \simeq q_L SIF_{TOT}$ , Eq. 1.43). As a result, the shape of the  $PAR - q_L$  relationship modulates the sensitivity of  $GPP_{SIF}$  to  $SIF_{TOT}$  (Gu *et al.*, 2019).



**Figure 6.11:** Scatterplot of MLR USO model residuals of gross primary production ( $GPP_{SIF} - GPP_{EC}$ ) and transpiration ( $Tr_{SIF} - Tr_{EC}$ ) compared to relative extractable water (REW- panels a, e), canopy vapor pressure deficit ( $VPD_{can}$  - panels b, f), canopy temperature ( $T_{can}$  - panels c, g) and escape probability ( $f_{esc}$  - panels d, h). Dark dots correspond to half-hourly data and red dots to daily data.  $r$  is the correlation coefficient and  $p$  is the corresponding p-value

The importance of the  $PAR - q_L$  relationship on  $GPP_{SIF}$  can be assessed by simulating  $q_L$  light-response curves based on the 95% confidence intervals of the fitted parameters (Eq. 6.17, Table 6.2, Fig. 6.17-a). These data are then fed into the MLR model (Eq. 6.8), which gives a modeled GPP ( $GPP_{SIF,mod}$ ) using modeled SIF determined from the light-use efficiency model (i.e.,  $SIF_{TOT,mod} = \phi_{SIF} \cdot f_{esc} \cdot f_{APAR} \cdot PAR$ , where  $f_{esc}$  and  $f_{APAR}$  are set to constant, and  $\phi_{SIF}$  is estimated from a linear relationship between  $\phi_{SIF}$  and PAR at the leaf level - Fig. 6.7, Fig. S6). This sensitivity analysis showed that the variability of the  $PAR - q_L$  model did not change the pattern of the light response curve of  $GPP_{SIF,mod}$ , which remained very similar to the  $PAR - SIF_{TOT,mod}$  relationship (Fig. 6.17-b). This similar sensitivity to PAR explained the strong linearity between the two variables regardless of the  $PAR - q_L$  response curves (Fig. 6.17-c). This suggests that the nonlinearity between  $SIF_{TOT}$  and  $GPP_{SIF}$  is not likely to be related to the measurements and parametrization of  $q_L$  as performed in this study.

One possible explanation of such discrepancy lies in the location of leaf-level measurements within the canopy. In particular, only sunlit leaves at the top of the canopy were selected for active fluorescence measurements, thus implicitly neglecting the contribution of the lower parts of the canopy to the SIF signal. However, wheat canopies can be characterized by leaves with various nitrogen content and structure (Furbank *et al.*, 2015; Chang *et al.*, 2022), and can have an heterogeneous contribution to the observed SIF signal. For instance, while sunlit leaves receive both diffuse and direct light, the proportion of absorbed diffuse light is higher for shaded leaves, which impacts their light absorption properties. In-situ measurements have shown that the sensitivity of  $q_L$  to PAR changes with the sunlit-shaded status of the leaf (Chang *et al.*, 2021; Zivcak *et al.*, 2014), which is also supported by model simulations on evergreen needleleaf forests (Chen *et al.*, 2024c). Such variability could explain the observed saturation of  $GPP_{SIF}$  at high PAR at the ecosystem scale. This hypothesis is supported by the sensitivity analysis on  $q_L$  conducted beyond the 95% confidence interval of  $PAR - q_L$  model parameters (Fig. 6.18), which shows that  $q_L$  is a key parameter for reproducing the SIF-GPP relationship. Neglecting such complexity in the response of a canopy-scale  $q_L$  to irradiance may explain the linearity between  $SIF_{TOT}$  and  $GPP_{SIF}$  at high APAR. This is even more relevant for dense canopies where the performances of the MLR USO model might be significantly affected under non-limiting light conditions. Future examination of the dependence of  $q_L$  on environmental conditions while differentiating light distribution through the canopy is needed (Chen *et al.*, 2024c; Chang *et al.*, 2021).

### 6.3. Performance of the model

The MLR model was used to determine GPP from TOC SIF measurement over a winter wheat crop. For the first time, this model has been coupled with the USO model



for stomatal conductance (Medlyn *et al.*, 2011a) to use both stomatal optimality theory and the MLR approach to estimate carbon and water fluxes with a single SIF measurement above the canopy. These findings should be considered in the light of previous studies, which already highlighted the promising capacities of the MLR model for predicting GPP at the canopy scale of an irrigated winter wheat crop (Liu *et al.*, 2022c), at the leaf scale for a grass and two tree species (Shi *et al.*, 2022), or for winter wheat grown under controlled conditions (L. Jia *et al.*, 2023). Another recent study of Han *et al.* (2022a) evaluated the robustness of the MLR model at the leaf level for 29 C3 and C4 plant species (not including winter wheat) representative of the major PFT across the globe and demonstrated that the MLR model was also capable of estimating net assimilation rate under diverse light and temperature conditions. The direct use of SIF to model Tr was also tested by (Feng *et al.*, 2021) who coupled the MLR model with the Ball-Woodrow-Berry (BWB) parametrization for stomatal conductance (Ball *et al.*, 1987). Although Tr estimated from the MLR BWB model was in a good agreement with EC data for one maize site and two forest sites, the results might have been impacted by neglecting the effects of changes in canopy structure (i.e., setting  $f_{esc}$  to constant, Feng *et al.* (2021)). Our study shows that accounting for canopy structural changes and model parameters dependence on environmental variables is critical for estimating GPP and Tr from SIF.

### 6.3.1. Robustness of the model during dry conditons

The decrease in LUE from  $REW_{th}$  (Fig. 6.13) highlights that winter wheat photosynthesis is sensitive to the decrease in soil water availability (Li *et al.*, 2023a; Yang *et al.*, 2023), which can lead to yield losses (Riedesel *et al.*, 2023). Statistical metrics indicated that the MLR USO model performed well even when soil water was limiting (i.e.,  $REW \leq REW_{th}$ ), or when  $T_{can}$  and  $VPD_{can}$  were high. The slopes of the linear regressions between MLR USO model estimates ( $GPP_{SIF}$ ,  $Tr_{SIF}$ ) and EC measurements ( $GPP_{EC}$ ,  $Tr_{EC}$ ) were very close to one (Fig. 6.9, Fig. 6.10), which indicates that the influence of meteorological conditions and changes in canopy structure were reproduced by the model through: (i) the dependence of model parameters on PAR and  $T_{can}$ , and (ii) the sensitivity of  $SIF_{TOT}$  to  $VPD_{can}$  and REW.

However, the correlation between model residuals and environmental drivers ( $T_{can}$ ,  $VPD_{can}$  and REW) suggests that soil water-limiting conditions and high  $T_{can}$  and  $VPD_{can}$  increased the error term of the MLR USO model. The cone-shaped scatterplot between model residuals and REW (Fig. 6.11) can be explained by uncertainties in the partitioning of NEE between  $GPP_{EC}$  and  $R_{ECO}$ . For instance, the partitioning of carbon fluxes from EC data relies on the accuracy of the relationship between  $T_{air}$  and  $R_{ECO}$  (i.e., nighttime method), which can be affected by soil water availability, diffuse radiation, soil temperature (Wohlfahrt & Galvagno, 2017) or the inhibition of daytime respiration (Keenan *et al.*, 2019). Although  $GPP_{EC}$  was retrieved using the OneFlux pipeline which uses relatively short moving windows to establish the relation-

ship between  $T_{air}$  and  $R_{ECO}$  (Pastorello *et al.*, 2020; Reichstein *et al.*, 2005), rapid changes in climate conditions such as rewetting after precipitation shortage episodes may induce a strong variability in this relationship, therefore impacting the accuracy of NEE partitioning (Tramontana *et al.*, 2020). This may explain the larger variability and the decreasing trend of the residuals for  $REW \leq REW_{th}$ , as numerous rainfall events were observed around DOY 150 after a constant reduction of soil water availability during the 40-days precipitation shortage period. In addition, MLR-USO parameters and  $VPD_{can}$  were determined using  $T_{can}$  from far-infrared radiation measurements at the canopy surface. While this approach has the advantage of increasing the versatility of the MLR USO model for a use at larger scales with RS of land surface temperature (Hulley *et al.*, 2019; Li *et al.*, 2013), the nighttime partitioning method uses  $T_{air}$  instead of  $T_{can}$  to determine ecosystem respiration and  $GPP_{EC}$  from NEE (Reichstein *et al.*, 2005). This difference may increase model residuals, especially when drought affects the coupling between  $T_{can}$  and  $T_{air}$  (Aprile *et al.*, 2013; Pradhan *et al.*, 2022). The choice of the driving temperature for  $R_{ECO}$  and the partitioning of NEE has been discussed in many studies and is still under debate in the scientific community (Lasslop *et al.*, 2012; Wohlfahrt & Galvagno, 2017). A recent paper of Kira *et al.* (2021) has explored the possibility to use SIF to empirically model GPP to improve the partitioning of NEE. In this perspective, the MLR model also represents a promising perspective to improve the partitioning of NEE by modeling GPP independently from NEE.

Finally, differences between  $Tr_{SIF}$  and  $Tr_{EC}$  may originate from the calibration of the USO model.  $g_1$  was estimated by using leaf-level measurements of gas exchanges, and the value of  $2.28 \text{ kPa}^{0.5}$  was in the range of the results of the study of (Medlyn *et al.*, 2017b) for crops, who calibrated the USO model from EC and leaf-level measurements. In our study, leaf-level measurements were conducted on sunlit leaves at the top of the canopy, which do not integrate the whole canopy variability of plant functional traits and meteorological conditions. In particular, canopy architecture induces variability in sun exposure, resulting in within-canopy gradients of meteorological conditions (Coble *et al.*, 2017) or plant traits (Niinemets *et al.*, 2004). Maximum stomatal conductance and stomatal sensitivity to PAR may be higher in the upper part of the canopy (Van Wilder *et al.*, 2022; Tarvainen *et al.*, 2013) whereas  $C_i$  is substantially lower in the deep canopy layers (Niinemets *et al.*, 2006). Therefore, the stomatal sensitivity to photosynthesis (i.e., the parameter  $g_1$ ) can be influenced by within-canopy variations in environmental conditions (Miner *et al.*, 2017), which ultimately results in uncertainties when scaling from leaves to the canopy. Limiting leaf-level measurements to sunlit leaves in the upper layers of the canopy may not capture such variability, and the potential effect of REW on  $g_1$ . In particular, it has been shown that  $g_1$  may increase under soil water-limiting conditions at the canopy scale (e.g., Beauclaire *et al.* (2023b); Gourlez de la Motte *et al.* (2020)), which could lead to an underestimation of canopy conductance during soil water stress. Considering this within-canopy variability might be important to model biosphere-atmosphere fluxes

for more dense and complex ecosystems such as forests (Bonan *et al.*, 2021). Future efforts should be focused on elucidating the relationship between model parameters and environmental conditions in dense canopies before testing the MLR USO model over a wider variety of ecosystems.

Despite this uncertainty in the calibration, the MLR USO model still reproduced a very high proportion of  $GPP_{EC}$  and  $Tr_{EC}$  variability while no parameters were downregulated by REW. It indicates that  $SIF_{TOT}$  was sensitive to the decrease in soil water availability either through changes in the physiological (i.e.,  $SIF_{TOC,760}$ ) or structural (i.e.,  $f_{esc}$ ) component of  $SIF_{TOT}$ . Although  $SIF_{TOC,760}$  may decrease due to an increase in NPQ (Porcar-Castell *et al.*, 2014; Xu *et al.*, 2021a), canopy structure through  $f_{esc}$  can also be affected by changes in leaf angle distribution usually caused by leaf rolling (Dechant *et al.*, 2020; Xu *et al.*, 2021a; Lu *et al.*, 2020) which is a common adaptative response of winter wheat for reducing transpiration and lowering leaf surface temperature during water stress (Ali *et al.*, 2022; Blum & Tuberosa, 2018). Disentangling drought effects between structural or physiological changes would require additional measurements of morphological traits such as leaf angle distribution (Kattenborn *et al.*, 2022; Zou *et al.*, 2014).

## 7. Conclusion

This study provides a method which combines existing mechanistic modeling frameworks for estimating GPP and Tr from SIF observations. The novelty of this approach lies in the joint representation of stomatal optimality and light use partitioning processes in a small set of mechanistically-based equations. Our results show that model parameters can be retrieved from canopy temperature and solar irradiance, which are available at large scales from RS data. A high correlation between MLR-USO model estimates and eddy covariance measurements was observed, with more than 90% of variance in observed GPP and Tr explained by the model ( $R^2 \geq 0.91$ ,  $rRMSE \leq 13.7\%$  at all timescales) across a wide range of environmental conditions, including light-limiting and soil water-limiting conditions. We also showed that the sensitivity of the fraction of open PSII centers to PAR plays a key role in regulating the response of GPP to SIF. This study confirms that the MLR-USO model can be used for estimating carbon and water fluxes from TOC SIF. More ecosystem-scale studies are needed to fully assess the potential of such model before an application at larger scales with satellite data. This study contributes to paving the way towards an improvement of the modeling of carbon and water fluxes across scales.

## 8. Acknowledgments and fundings

This research was conducted within the framework of the ICOS Wallonia project, which is supported by the Service Public de Wallonie, Belgium [1217769]. Simon

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## 9. Supplementary materials

### 9.1. Calculation of the REW

The REW (Granier *et al.*, 1999a) was calculated from SWC measurements and maximum rooting depth following:

$$REW = \begin{cases} \frac{\theta_{z_0} - \theta_{wp,z_0}}{\theta_{fc,z_0} - \theta_{wp,z_0}} & \text{if } \alpha \leq z_0 \\ \frac{z_{i-1} (\theta_{z_{i-1}} - \theta_{wp,z_{i-1}}) + |\alpha - z_i| ((\theta_{z_i} - \theta_{wp,z_i}))}{z_{i-1} (\theta_{fc,z_{i-1}} - \theta_{wp,z_{i-1}}) + |\alpha - z_i| ((\theta_{fc,z_i} - \theta_{wp,z_i}))} & \text{if } z_{i-1} \leq \alpha \leq z_i \\ & i \in [1 : 4] \end{cases} \quad (6.23)$$

where  $\theta_{z_i}$ ,  $\theta_{wp,z_i}$  and  $\theta_{fc,z_i}$  are respectively the measured soil water content, the wilting point and field capacity at the soil depth  $z_i$ , and  $\alpha$  is the maximum rooting depth.  $z_0$  corresponds to the depth of the first SWC sensor at 5 cm depth while  $i$  ranges from 1 to 4 (i.e., corresponding to the depths of SWC sensors at 15, 25, 55 and 85 cm). These depths were weighted by  $\pm 5$  cm to account for a zone of influence around the sensor. The wilting point and field capacity at each depth are shown in Table S1.  $\alpha$  was determined from root biomass, which was measured six times during the experiment ( $\alpha_{meas}$ , at day of year DOY 80, 94, 115, 136, 157 and 178) with samples taken within six soil layers (0-15 cm, 15-30 cm, 30-45 cm, 45-60 cm, 60-80 cm and 80-100 cm) across the field and soil horizons following a random sampling method. For each day of measurements, the cumulative RBD was calculated from biomass measurements and was extrapolated for all soil depth using a sigmoidal function  $y = \frac{a}{1 + \exp^{-b(z-c)}}$  with  $y$  the cumulative RBD,  $a$  the asymptote of the model,  $b$  the soil depth corresponding to  $y = \frac{a}{2}$ ,  $z$  the soil depth, and  $c$  a curvature parameter. For each day of measurement, a continuous RBD profile was calculated.  $\alpha$  was defined as the soil depth at which the root density does no longer significantly increase. Therefore,  $\alpha$  was calculated as the corresponding soil depth when  $y = a - \sigma_a$  with  $\sigma_a$  the standard deviation of  $a$ :

$$\alpha = c - \frac{\log \frac{\sigma_a}{a - \sigma_a}}{b} \quad (6.24)$$

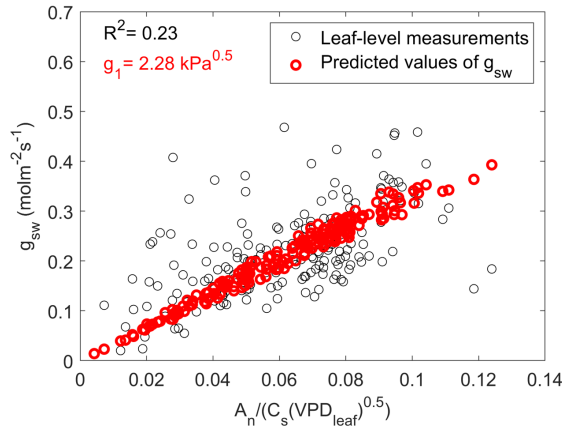
$\alpha$  was calculated for each day of the soil sampling campaign and was extrapolated from the date of sowing ( $\alpha = 0$  on October 21, 2021) to the date of harvest (July 24, 2022) with a sigmoidal function.

## 9.2. Calibration of the USO model on leaf-scale measurements

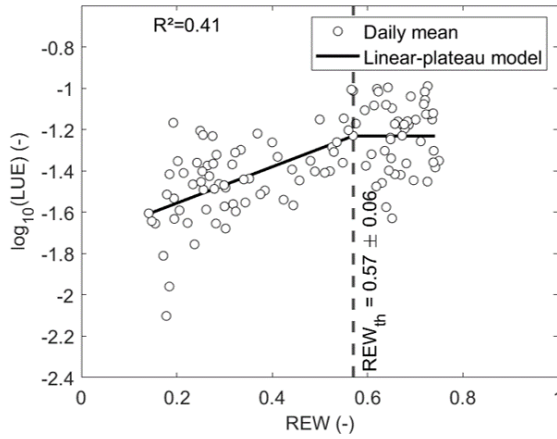
The USO model (Medlyn *et al.*, 2011a) was calibrated by adjusting a nonlinear regression between Eq. 6.25 and leaf-level measurements of stomatal conductance to water vapor ( $g_{sw}$ ), CO<sub>2</sub> concentration at the leaf surface ( $C_s$ ), leaf vapor pressure deficit ( $VPD_{leaf}$ ) and net photosynthesis ( $A_n$ ) following:

$$g_{sw} = 1.6 \left( 1 + \frac{g_1}{\sqrt{VPD_{leaf}}} \right) \frac{A_n}{C_s} \quad (6.25)$$

The relationship between  $g_{sw}$  and  $\frac{A_n}{C_s \sqrt{VPD_{leaf}}}$ , as well as predicted values of  $g_{sw}$  using Eq. 6.25 are shown in Fig. 6.12.

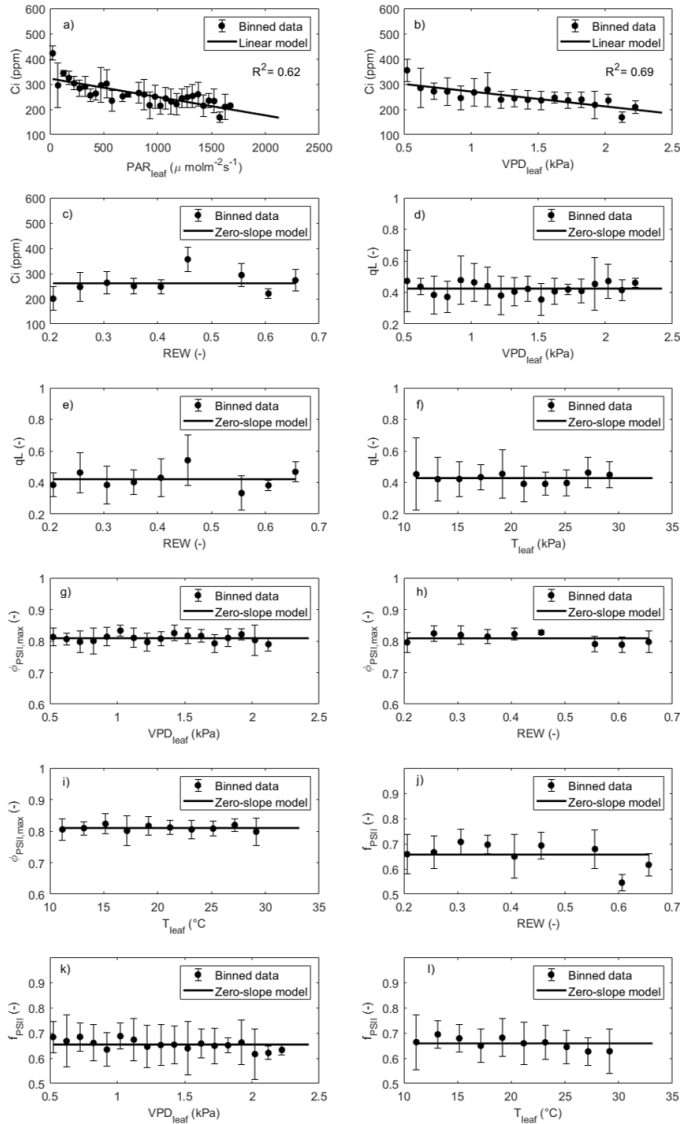


**Figure 6.12:** Relationship between stomatal conductance to water vapor ( $g_{sw}$ ), net photosynthesis ( $A_n$ ), CO<sub>2</sub> concentration at the leaf surface ( $C_s$ ) and leaf vapor pressure deficit ( $VPD_{leaf}$ ). The  $R^2$  is the coefficient of determination of the nonlinear regression of the USO model and  $g_1$  is the calibrated value. Red dots show the predicted values of  $g_{sw}$  using  $g_1=2.28 \text{ kPa}^{0.5}$

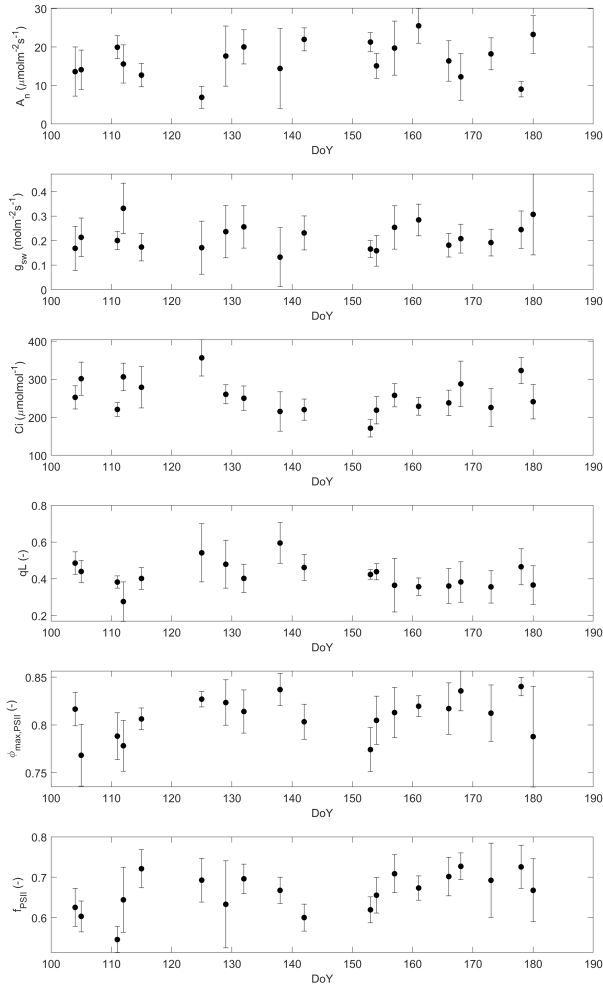


**Figure 6.13:** Relationship between the log10-transformed LUE ( $LUE = GPP_{EC}/APAR$ ) and the REW.  $R^2$  is the coefficient of determination of the two-regime model and  $\text{REW}_{\text{th}}$  is the REW threshold where a shift from soil water-limiting to non-limiting conditions is observed.

Chapter 6. Modeling gross primary production and transpiration from sun-induced chlorophyll fluorescence using a mechanistic light-response approach

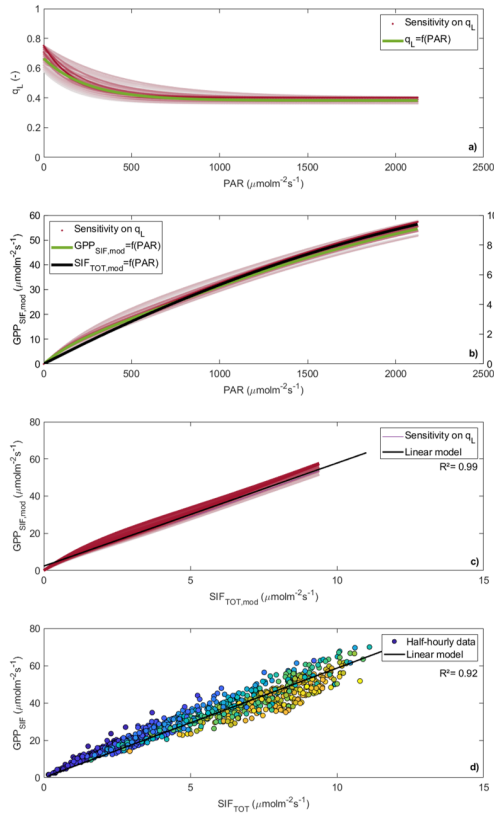


**Figure 6.14:** Relationship between MLR-USO model parameters ( $\text{CO}_2$  concentration in sub-stomatal cavities -  $C_i$ , fraction of open PSII centers -  $q_L$ , maximum photochemical quantum yield  $\phi_{PSII,max}$ , and fraction of fluorescence attributed to PSII -  $f_{PSII}$ ) and meteorological variables (leaf incoming irradiance -  $PAR_{leaf}$ , leaf vapor pressure deficit -  $VPD_{leaf}$  and relative extractable water - REW). Measurements are grouped into REW,  $PAR_{leaf}$ ,  $VPD_{leaf}$  or  $T_{leaf}$  classes with an interval of respectively 0.05, 50  $\mu\text{mol m}^{-2}\text{s}^{-2}$ , 0.1 kPa and  $2^\circ\text{C}$ . Only models with the lowest AIC are shown. The error bars show the standard deviation around the mean for each data class. Statistical metrics are given in Table 6.14.



**Figure 6.15:** Time series of daily means of leaf-level measurements of net photosynthesis ( $A_n$ ), stomatal conductance to water vapor ( $g_{sw}$ ),  $\text{CO}_2$  concentration in the substomatal cavities ( $C_i$ ), fraction of open PSII centers ( $q_L$ ), maximum photochemical quantum yield ( $\phi_{PSII,max}$ ) and fraction of emitted fluorescence attributed to PSII ( $f_{PSII}$ ). The error bars show the standard deviation around the daily means.





**Figure 6.17:** Response of the relationship between fraction of the open PSII centers ( $q_L$ ) and solar irradiance ( $PAR$ ) to changes in regression parameters (Eq. 6.17 – Table 6.3) within their observed variability (i.e.,  $\pm 95\%$  confidence interval – panel a). Panel b shows the corresponding impact on the relationship between modeled GPP ( $GPP_{SIF,mod}$ ) and PAR.

$GPP_{SIF,mod}$  was estimated using the MLR model (Eq. 6.8) and total sun-induced chlorophyll fluorescence ( $SIF_{TOT,mod}$ ) simulated using  $SIF_{TOT,mod} = \phi_{SIF} f_{APAR} PAR f_{esc}$  (Gu *et al.*, 2019).  $\phi_{SIF}$  is the PSII fluorescence quantum yield which was estimated using a first-order polynomial model from the negative correlation between  $\phi_{SIF}$  measurements and PAR (Fig. 6.7). All the other parameters are kept constant, with values of  $C_i=265$  ppm,  $\Gamma^*=30$  ppm (temperature =  $20^\circ\text{C}$ ),  $f_{esc}=0.5$ ,  $f_{APAR}=0.7$ . PAR ranged from 0 to the maximum PAR measured with a step of  $5 \mu\text{molm}^{-2}\text{s}^{-1}$ . The green curve is the relationship between PAR and  $q_L$  measurements (panel a, Fig. 6-a), and between PAR and  $GPP_{SIF,mod}$  (panel b). The black curve is the relationship between  $SIF_{TOT,mod}$  and PAR (panel b). Panel c) shows the relationship between simulated  $GPP_{SIF,mod}$  and  $SIF_{TOT,mod}$ , and the corresponding linear regression. Each color from red to light-red is associated to a  $q_L$ -PAR curve with combination of regression parameters. Panel d) shows the relationship between GPP estimated by the MLR-USO model ( $GPP_{SIF}$ - Eq. 6.8) and measured total SIF ( $SIF_{TOT}$  - Eq. 6.1).

Horizon	Depth (cm)	$\theta_{fc}$ ( $\text{cm}^{-3} \text{cm}^{-3}$ )	$\theta_{wp}$ ( $\text{cm}^{-3} \text{cm}^{-3}$ )	Sand (%)	Silt (%)	Clay (%)	Bulk density ( $\text{g cm}^{-3}$ )
H <sub>1</sub>	0-35	45.3	11.7	8.35	80.5	11.2	1.27
H <sub>2</sub>	35-85	40.4	13.01	6.1	75.7	18.2	1.52

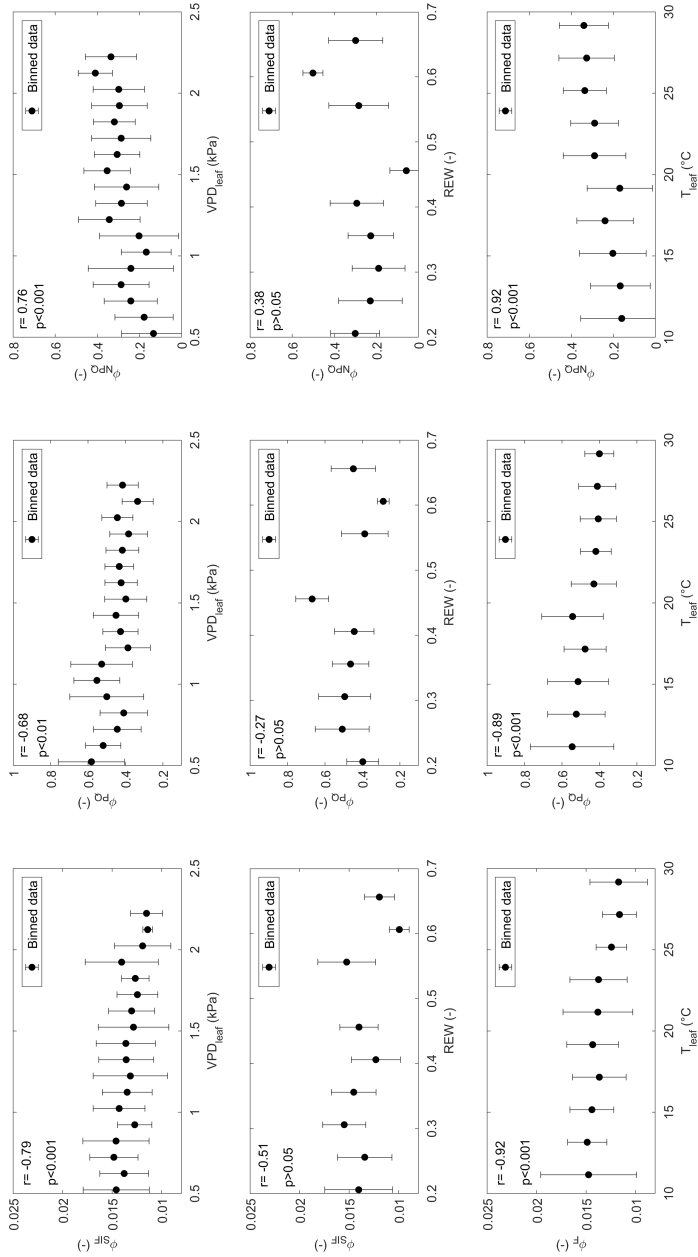
**Table 6.1:** Soil physical proprieties in the ICOS Lonzée site: field capacity ( $\theta_{fc}$ ), wilting point ( $\theta_{wp}$ ), sand, silt, clay content and bulk density for the two soil horizons (H1: 0-35 cm and H2: 35-100 cm). Field capacities and wilting points were measured by the WPC4 and HYPROP2 sensors (Meter environment., Hopkins Ct, NE, US) on soil samples collected in 2019.

Date	Activities	Product
28/10/21	Sowing	<i>Triticum aestivum</i> cv. Skyscraper
24/03/22	Fertilizer	Liquid Nitrogen 39/0/0 (150 l ha <sup>-1</sup> )
13/04/22	Weeding	Capri Duo (200 g ha <sup>-1</sup> ) + Accurate (10 g ha <sup>-1</sup> )
13/04/22	Growth regulator	Cycocel (1 l ha <sup>-1</sup> )
13/04/22	Biostimulating fertilizer	Amino Mag (2 l ha <sup>-1</sup> )
14/04/22	Fertilizer	Liquid Nitrogen 39/0/0 (100 l ha <sup>-1</sup> )
12/05/22	Fungicide	Aquino (1,2 l ha <sup>-1</sup> ) + Protendo (0,4 l ha <sup>-1</sup> )
17/05/22	Fertilizer	Nitrate 27 (200 kg)
31/05/22	Fungicide	Ascra (1,2 l ha <sup>-1</sup> ) + Protendo (0,5 l ha <sup>-1</sup> )
24/07/22	Harvest	<i>Triticum aestivum</i> cv. Skyscraper

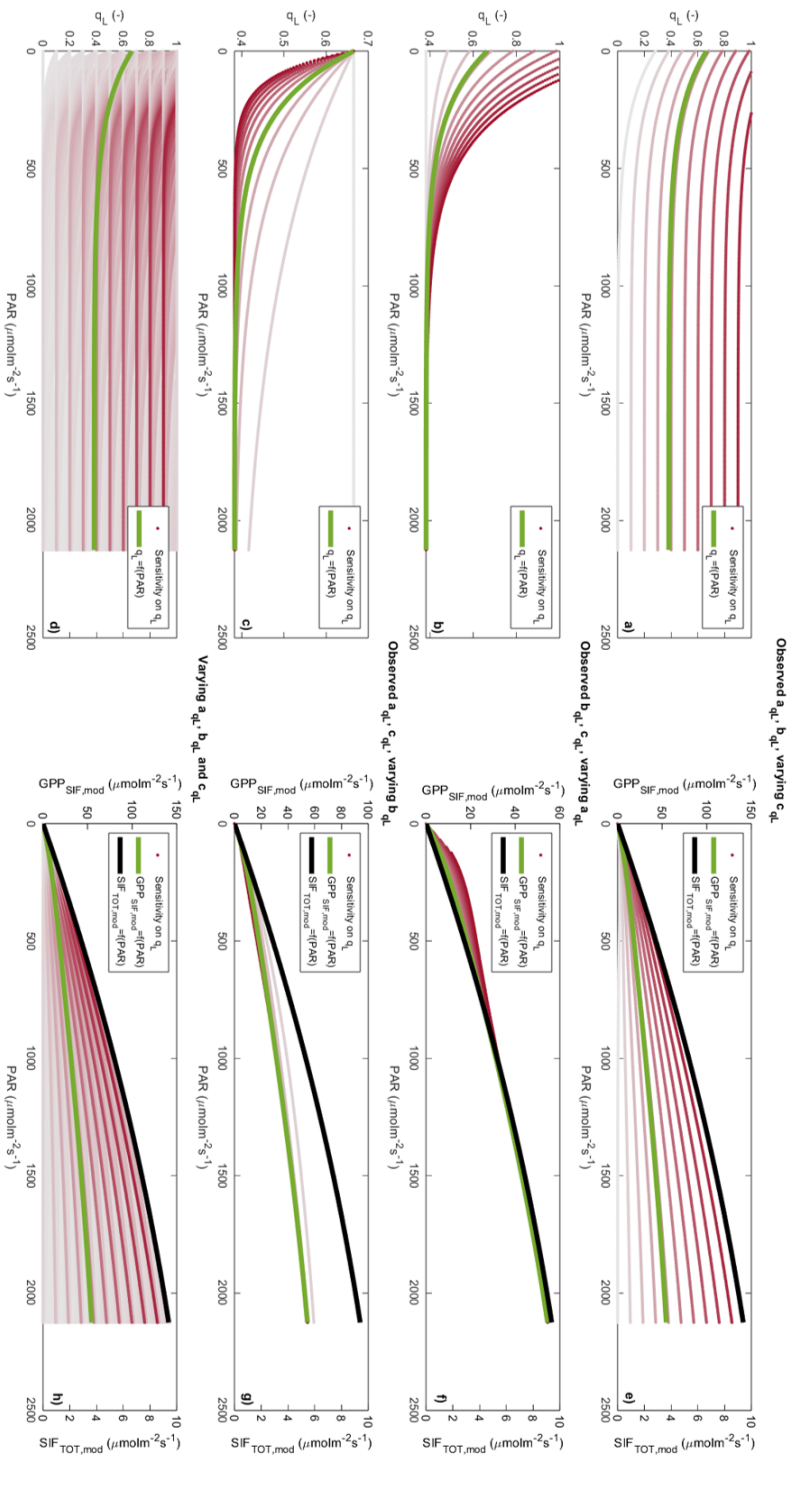
**Table 6.2:** Farming activities in the ICOS Lonzée site during the 2022 winter wheat growing season

MLR model parameters	a ± CI <sub>95%</sub>	b ± CI <sub>95%</sub>	c ± CI <sub>95%</sub>	R <sup>2</sup>
$C_i = a_{Ci} \cdot T_{leaf}^2 + b_{Ci} \cdot T_{leaf} + c_{Ci}$	0.55 ± 0.36	-26 ± 14.8	626 ± 141	0.94
$C_i = a_{Ci} \cdot PAR_{leaf} + b_{Ci}$	-0.07 ± 0.02	332 ± 21	-	0.62
$C_i = a_{Ci} \cdot VPD_{leaf} + b_{Ci}$	-58.2 ± 21.1	330 ± 31	-	0.68
$f_{PSII} = a_{fPSII} \cdot PAR_{leaf} + b_{fPSII}$	(-6.77 ± 2.25) · 10 <sup>-5</sup>	0.72 ± 0.02	-	0.56
$q_L = a_{qL} \cdot \exp^{-b_{qL} \cdot PAR_{leaf}} + c_{qL}$	0.28 ± 0.07	0.004 ± 0.002	0.38 ± 0.02	0.79

**Table 6.3:** Statistics of the regressions between MLR model parameters (CO<sub>2</sub> concentration in substomatal cavities -  $C_i$ , fraction of open PSII centers -  $q_L$  and fraction of emitted fluorescence attributed to PSII -  $f_{PSII}$ ) and meteorological variables (leaf temperature  $T_{leaf}$ , vapor pressure deficit  $VPD_{leaf}$ , and solar irradiance  $PAR_{leaf}$ ). Regression parameters are given with their 95% confidence intervals.



**Figure 6.16:** Relationship between leaf-level quantum yields (photochemistry -  $\phi_{PQ}$ , non-photochemical quenching -  $\phi_{NPQ}$  and chlorophyll fluorescence -  $\phi_{SIF}$ ) and meteorological variables (leaf vapor pressure deficit -  $VPD_{leaf}$ , REW and leaf temperature  $T_{leaf}$ ). Measurements are grouped into REW,  $VPD_{leaf}$  or  $T_{leaf}$  classes with an interval of respectively 0.05, 0.1 kPa and 2°C. The error bars show the standard deviation around the mean for each data class. r is the correlation coefficient and p is the corresponding p-value.



**Figure 6.18:** Response of the relationship between the fraction of the open PSII centers ( $q_L$ ) and solar irradiance (PAR) to changes in regression parameters (Eq. 17 – Table 6.3) beyond their 95% confidence interval (panels a to d). In panel (a),  $c_{qL}$  varies between 0 and 1. In panel (b),  $a_{qL}$  varies between 0 and 1. In panel (c),  $b_{qL}$  varies between 0 and 0.01. All the other parameters are taken from Table 6.3 if not specified otherwise. Panels (e) to (h) shows the corresponding relationship between  $GPP_{SIF_{mod}}$  and  $SIF_{TOT_{mod}}$ . Details about the sensitivity analysis calculation steps can be found in the legend of Fig. 6.17.

**Thesis achievements, methodological considerations and perspectives**



## 1. Thesis achievements

Improving the understanding of the effects of soil water stress on photosynthetic processes is one of the main challenges in photosynthesis research. By using a collection of gas exchange and ChlF measurement techniques at the leaf and ecosystem scales, this thesis contributed to advance representation of photosynthesis in LSMs (Rogers *et al.*, 2017) by : (i) the identification of the constraints on photosynthesis originating from non-stomatal factors and the calibration of their response to soil water availability, (ii) the development of a new partitioning method to emphasize the key role played by mesophyll conductance on limiting carbon assimilation and transpiration, and (iii) the use of RS data to improve the modeling of carbon assimilation and transpiration at the ecosystem scale.

### 1.1. The constraints on photosynthesis originating from non-stomatal factors under soil water stress

One of the objectives of this thesis was to determine which parameter of the USO and FvCB models should be modified to consider the effect of decreasing soil water availability on physiological processes. This approach was tested on potato and forest ecosystems (dominated by European beech (*Fagus sylvatica*), European spruce (*Picea abies*), European red pine (*Pinus sylvestris*) and maritime pine (*Pinus pinaster*)) either at the ecosystem scale (for forests and potato) or at the leaf scale (for potato solely) (Gourlez de la Motte *et al.*, 2020; Beauclaire *et al.*, 2023b). The limitations on photosynthesis originating from a stomatal control were associated to an effect of the parameter  $g_1$  of the USO model (or  $G_1$  at the ecosystem scale) while the limitations of photosynthesis originating from other non-stomatal factors were associated to  $V_{cmax}$  and  $g_m$  (or apparent  $V_{cmax,app}$  at the ecosystem scale) of the FvCB model. These models are used as sub-models to simulate leaf carbon and water exchanges in LSMs and provide insights on the effects of water stress on the whole plant functioning (Rogers *et al.*, 2017).

How to consider the effects of soil water stress on fluxes remains an important issue in modeling. Most LSMs use an empirical multiplier (i.e., the  $\beta$  factor determined from soil water availability, Eq. 1.49) which directly affects FvCB or USO model parameters when SWC drops below field capacity (i.e., REW<1). This induces abrupt transitions and neglects any plasticity of the physiological response of ecosystems to lack of soil water, ultimately leading to unrealistic drought responses (Zhou *et al.*, 2019b). These abrupt transitions and the uncertainty on which variable should be affected by  $\beta$  are key issues when modeling drought response of terrestrial ecosystems. As highlighted by Vidale *et al.* (2021), "The treatment of  $\beta$  (...) represents a key structural uncertainty in contemporary LSMs, in terms of predictions of gross primary productivity, energy fluxes and soil moisture evolution (...)".

Different approaches are implemented in LSMs for integrating  $\beta$ . For instance,  $\beta$  is

applied on  $g_1$  in CABLE (Wang *et al.*, 2011; De Kauwe *et al.*, 2015a), while in JULES and in JSBACH,  $\beta$  is used on  $A_n$  and  $g_s$  (Harper *et al.*, 2021; Knauer *et al.*, 2015). In ORCHIDEE  $V_{cmax}$  is decreased by  $\beta$  (Krinner *et al.*, 2005). Other approaches use  $\beta$  to affect biochemical signaling such as ABA synthesis, root growth or leaf senescence (see Harper *et al.* (2021) for references). These differences in the treatment of  $\beta$  largely originate from a poor description of the response of plant gas exchanges to drought due to a lack of model-oriented experiments or field studies (Powell *et al.*, 2013; Zhou *et al.*, 2019b).

In this thesis, we used gas exchange and leaf-level ChlF measurements to calibrate  $\beta$  on NSOL and SOL for forest and potato ecosystems, either at the canopy scale (chapter 4 and 3), or at the leaf scale (chapter 5). One important finding of this thesis is the consistent relationship between NSOL, SOL and REW. In particular, canopy-scale measurements of EC showed that the decrease in REW induced a decrease in  $V_{cmax,app}$  while  $G_1$  remained constant or slightly increased under severe water stress. It highlights that NSOL was the main constraint on GPP, a pattern which is likely to be consistent in C3 plant species (e.g., Zhou *et al.* (2013); Zait & Schwartz (2018); Wang *et al.* (2018b)). Neglecting the dependence of  $V_{cmax,app}$  to REW resulted in strong overestimation of GPP in temperate forest ecosystems (chapter 3), and potato (chapter 4). These results reflect that  $\beta$  should be applied on  $V_{cmax,app}$  to reproduce drought effects on GPP. The increase in  $G_1$  when REW  $\sim 0.40$  is more likely to be a consequence of high NSOL constraints, and can be interpreted in terms of water use strategy.

While it provides large datasets of gas exchanges data at the ecosystem scale, the EC technique does not allow the identification of the true origin of the decrease in  $V_{cmax,app}$ . Was it caused by an alteration of biochemical capacities (i.e., a decrease in  $V_{cmax}$ ) or by a strong restriction of CO<sub>2</sub> diffusion to the chloroplasts by the mesophyll (i.e., a decrease in mesophyll conductance  $g_m$ ) ? This issue is critical as numerous studies have shown that  $g_m$  is sensitive to water stress (see chapter 5). LSMs typically model  $g_m$  at the leaf level by considering the effects of temperature and water stress through empirical multipliers similar to  $\beta$  (Knauer *et al.*, 2019b). Therefore, realistic drought responses of photosynthesis and transpiration also rely on calibrating  $\beta$  on  $g_m$  (as it regulates CO<sub>2</sub> diffusion to chloroplasts (Eq. 1.27), and  $g_s$  through  $A_{sat}$  (Eq. 1.35)). We tackled this issue by setting-up a drought experiment on field-grown potatoes and by conducting joint measurements of gas exchange and actively-induced ChlF at the leaf level using a LI-COR LI-6400XT (Chapter 5). We showed that water stress induced a two-stage response of physiological processes with constraints on CO<sub>2</sub> diffusion (i.e., related to conductances) being the first reaction of potato plants to soil drought. Although this approach is essential to analyze the temporal dynamics of model parameters and their dependence to soil water conditions, it does not indicate which model parameters had the most important impact on photosynthesis. To do so, we first started by calculating the total derivative of  $A_{sat}$  as a function of  $g_s$ ,



$g_m$  and  $V_{cmax}$  to quantitatively assess the importance of each parameter in explaining the response of photosynthesis to water stress. This method is not new and has been already applied in numerous studies to decipher the importance of stomatal and non-stomatal constraints on photosynthesis (e.g., Grassi & Magnani (2005); Zait & Schwartz (2018); Wang *et al.* (2018b)). The novelty of the method proposed in chapter 5 is the coupling of the USO and FvCB models in the limitation analysis to identify the origins of the limitations on  $A_{sat}$  and  $g_s$ . The basis of this method is the partitioning of stomatal conductance into relative variations of VPD,  $g_1$  and  $A_{sat}$  as described in the USO model (Medlyn *et al.*, 2011b). The mathematical development is given in the supplementary material in chapter 5. The two partitioning schemes (with or without the implementation of the USO model) showed that mesophyll conductance was the most important constraint on  $A_{sat}$  and  $g_s$  (Fig. 5.5). These results go further than the studies presented in chapter 4 and chapter 3, and provide a quantitative analysis of the role played by mesophyll conductance on the decrease in photosynthesis and stomatal conductance under water stress. In addition, the importance of VPD in regulating stomatal conductance and photosynthesis can be quantitatively assessed.

Another interesting result is the consistent decrease in WUE (increase in  $g_1$ ) under severe drought, which contradicts with the expected increase in WUE in particular for crops (Yu *et al.*, 2020a), and suggests that extreme dry conditions altered the ability of potato to maintain a constant water use for a given carbon gained. This increase in  $g_1$  highlights the strong effect of NSOL on the decoupling between photosynthesis and transpiration (see chapter 4), almost solely driven by  $g_m$  (see chapter 5). In particular, it suggests that improving CO<sub>2</sub> diffusion to chloroplasts can increase WUE. This result is consistent with other studies which showed that  $g_m$  is a key limiting factor for WUE for crops (Leakey *et al.*, 2019) or for tree species (Zhu *et al.*, 2021). Engineering anatomical and biochemical factors associated to  $g_m$  such as cell wall thickness, porosity, the surface of chloroplasts exposed to the intercellular air spaces, aquaporins or carbonic anhydrase activity could increase CO<sub>2</sub> diffusion and by extension WUE (Petrík *et al.*, 2023). Additional measurements of cells anatomy, molecular signaling and water transport in plants would help in understanding the mechanisms driving the response of mesophyll and WUE to water stress.

Not only the role played by NSOL in limiting photosynthesis under water stress was highlighted, but a similar REW threshold of  $\sim 0.4$  was found for forests and potato at the ecosystem scale. This result for forest ecosystems is in accordance with previous studies which have shown that a decrease in Tr and GPP of temperate forest stands was observed from a REW of  $\sim 0.4$  (Bernier *et al.*, 2002; Rambal *et al.*, 2003; Granier *et al.*, 2007b, 1999b; Bréda *et al.*, 2006, 1995) and confirm that the threshold at which  $\beta$  starts impacting NSOL should be lower than field capacity (Verhoef & Egea, 2014). This threshold of  $\sim 0.4$  is often used to separate stressed and unstressed conditions when interpreting tree gas exchanges (e.g., Xu *et al.* (2017); Van Sundert *et al.* (2020); Aguilos *et al.* (2019)). However, there is little information regarding

the REW thresholds for NSOL or SOL on potato. We found two contrasting results with different thresholds at the leaf scale for diffusive ( $\text{REW}=0.72 \pm 0.12$ ) and biochemical constraints ( $\text{REW}=0.43 \pm 0.03$ ). These discrepancies are further discussed in section 2.2.1. These results are nonetheless particularly useful in crop models such as Aquacrop where the calibration level for these thresholds is defined as minimum for potato (Raes *et al.*, 2009). More studies should be conducted to provide further analysis on the similarities between REW thresholds for forests and potato. Importantly, this method could be applied to select potato varieties which provides the best drought-resistant characteristics (i.e., the lowest REW threshold), thus potentially improving crop yields and irrigation scheduling. Perspectives are given in section 3.2.

## 1.2. Modeling of GPP from SIF by a mechanistic framework

For more than four decades, ChlF has been used to physiological processes and more specifically stress responses of plants (Baker, 2008; Genty *et al.*, 1989). To date, most of ChlF measurements are performed by active methods at the leaf scale. Because these sensors use an artificial light to drive the reduction of PSII centers, the application of such technique for remote sensing platforms is limited (Cendrero-Mateo *et al.*, 2015; Magney *et al.*, 2017). Spectrometers with sub-nanometer resolutions have enabled the development of passive methods for retrieving SIF in the O<sub>2</sub>-A and O<sub>2</sub>-B absorption bands using the FLD or SFM method (see chapter 1, section 3.2.2). The MLR model now provides a mechanistic framework for interpreting SIF temporal variability in terms of changes in carbon assimilation (Gu *et al.*, 2019). Yet it remained to be tested under natural conditions and under soil-water limiting conditions. We also added a transpiration module to the model by using the USO equation for stomatal conductance, thereby allowing to estimate both GPP and Tr from SIF.

This thesis have demonstrated the potential of the MLR-USO model for estimating GPP and Tr from SIF measurements at the ecosystem scale. In particular, MLR-USO model estimates were highly correlated to eddy covariance measurements (i.e.,  $R^2>0.91$ ,  $r\text{RMSE}<13.7\%$ ) across a wide range of environmental conditions, including high solar irradiance, atmospheric and edaphic dryness (chapter 5). The model relies only on the measurement of SIF, PPFD and surface temperature. However, major uncertainties still remain, especially when considering an application at larger scales or on other ecosystems where within-canopy gradients of environmental conditions are observed and cannot be neglected. In particular, selecting only top-of-canopy leaves for model parameterization may result in strong uncertainties as it is well known that the understorey and midstorey also strongly contribute to the total SIF signal in forests (Morozumi *et al.*, 2023a; Liu *et al.*, 2019). In addition, the response of GPP to PPFD varies with the location of the leaf within the canopy where top-of-canopy leaves are mostly light-saturated while shaded leaves are light-limited. Therefore, light availability must be characterized by distributing top-of-canopy PPFD into the different layers

of the canopy (Zhang *et al.*, 2023b). These aspects are not captured by 'big leaf' models which simplify the canopy as a single layer (already discussed in chapter 1 and in chapter 6). In this perspective, several research questions remain. Does the MLR-USO model requires a species-specific or PFTs-specific calibration? Does the hypothesis of a predominant contribution of top-of-canopy leaves to the SIF signal still hold for other PFTs ? These questions and the recent advances published in the literature are further detailed in section 3.7.

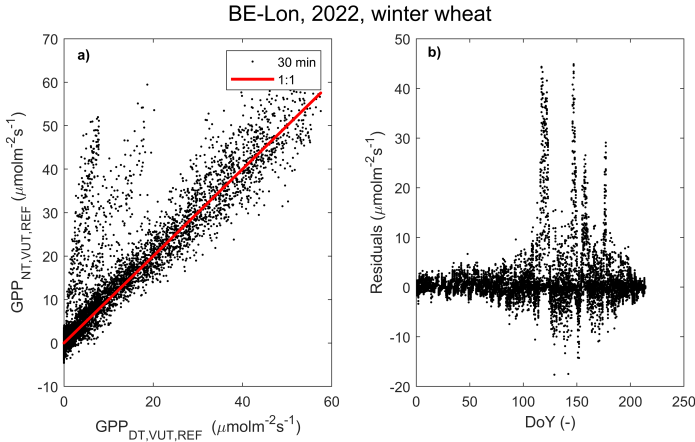
## 2. Methodological considerations

The results of this thesis were obtained by designing a framework which is based on several hypotheses and methodological choices. The main methodological considerations are : (i) the choice of the NT partitioning method for estimating GPP from NEE, the choice of REW as the explanatory variable for calibrating  $\beta$ , and (iii) the choice of joint measurements of gas exchanges and ChlF to calculate  $g_m$ .

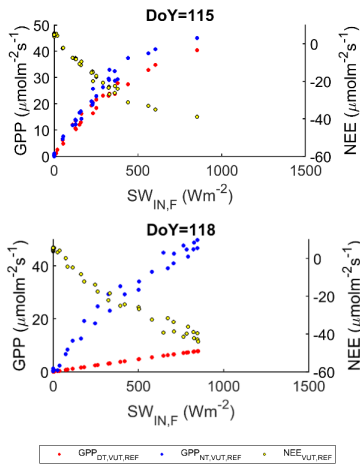
### 2.1. GPP partitioning methods

EC products are determined by a series of post-processing methods which are implemented into a common data treatment pipeline within the ICOS and FLUXNET networks (i.e., the OneFlux pipeline, Pastorello *et al.* (2020)) which allows a harmonization of data treatment and facilitates the comparison of EC products across sites. One of the key steps is the separation of the contribution of ecosystem respiration  $R_{ECO}$  from NEE by either the DT or NT partitioning methods (Reichstein *et al.*, 2005; Lasslop *et al.*, 2010). This last remains the most widely used in the flux community (Trifunov *et al.*, 2021). Partitioning methods strongly rely on several underlying assumptions regarding the relationship between NEE and few environmental drivers (see chapter 1 and chapter 6). While both methods are relatively similar for estimating  $R_{ECO}$  (they only differ by how  $R_b$  (Eq. 1.17) is estimated either from NT or DT data), strong differences may be observed for GPP.

A practical example of these discrepancies can be found in the 2022 dataset of BE-Lon. Despite a strong correlation between DT and NT GPP (Fig. 7.1-a), a decoupling is observed at three specific periods during the growing season: between DoY 116-123, DoY 146-149 and DoY 176-177 (Fig. 7.1-b), which suggests an underestimation of DT GPP. This is confirmed by the light-response curve of the two GPP, highlighting a clear unexpected decrease in DT GPP (Fig. 7.2).



**Figure 7.1:** Difference between DT and NT GPP at BE-Lon in 2022. The names of the variables in the post-processed dataset (Dumont *et al.*, 2023b) are as in the legend.

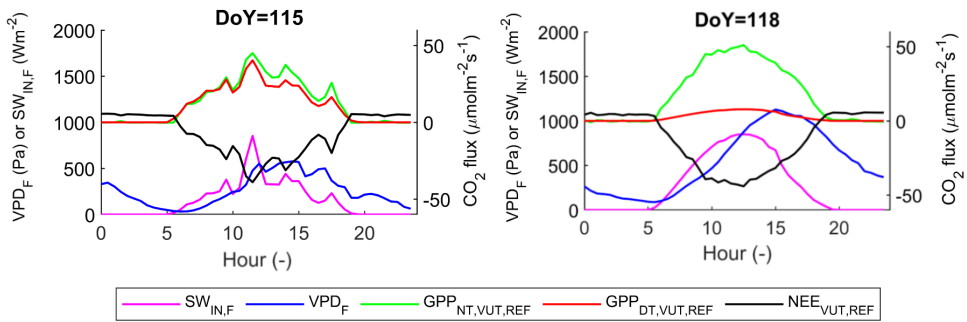


**Figure 7.2:** Light response curves of NT and DT GPP at BE-Lon in 2022 for two contrasting DoY. The names of the variables in the post-processed dataset (Dumont *et al.*, 2023b) are as in the legend.

These differences can only originate from the parameterization of the NEE light-response curve within a 4-days moving window leading to an unrealistic response to VPD (i.e., low  $\beta_{NEE}$ ) or to  $R_g$  (i.e., low  $\alpha_{NEE}$ ). In the DT partitioning method,  $T_{ref}$  is first fixed as the median temperature of the moving window and  $T_0$  is set to  $-46.02$  °C.  $E_0$  is then estimated from nighttime data for a shifting moving window of 4 days, before  $\alpha_{NEE}$ ,  $\beta_{NEE}$  and  $R_b$  (Eq. 1.18) are fitted using DT data. The 2022 dataset of BE-Lon did not show any unexpected low  $R_g$  or high VPD during the days when these anomalies were detected (as a recall the threshold VPD value for the DT approach is 1 kPa, (Wutzler *et al.*, 2018), Fig. 7.3). This issue still needs to be resolved. Therefore, the choice of the partitioning method for GPP may have strong implications in modeling or ecophysiological studies. For example, the use of DT GPP as validation dataset of MLR-USO would have weakened the performances of

the model compared to NT GPP (data not shown). This uncertainty logically propagates to carbon budgets where a difference of  $\sim 120$  gCm $^{-2}$ d $^{-1}$  between DT and NT cumulative GPP was observed at harvest (data not shown).

In the perspective of evaluating the differences between NT and DT GPP, Jung *et al.* (2024) have proposed an algorithm for flagging these inconsistencies. Complementary consistency flags (C2F) revealed that GPP,  $R_{\text{ECO}}$  and radiation variables were the most flagged data across the FLUXNET2015 dataset (Jung *et al.*, 2024). This suggests that, despite recent development in data post-processing and quality check controls, a fine analysis of GPP time series is a prerequisite before interpreting EC data. Partitioning NEE remains an key research topic in the scientific community, and many alternatives for estimating  $R_{\text{ECO}}$  and GPP from NEE have been developed. Most of the recent methods for partitioning NEE use proxies of GPP, based, for instance, on SIF (Zhan *et al.*, 2022; Kira *et al.*, 2021), carbonyl sulfide (Kohonen *et al.*, 2022), machine learning algorithms (Chen *et al.*, 2024a) or stable isotopes (Lee *et al.*, 2020).



**Figure 7.3:** Meteorological conditions at BE-Lon in 2022 during two contrasting DoY for DT and NT GPP. The names of the variables in the post-processed dataset (Dumont *et al.*, 2023b) correspond to the legend.  $SW_{\text{IN}}$  is incoming shortwave radiation (i.e.,  $\sim R_g$ ).

## 2.2. On the choice of REW for modeling NSOL and SOL under water stress

Characterizing soil water status in ecophysiology requires the measurement of a physical variable which considers the water availability for plant uptake. In this thesis, we chose to combine SWC with SWRC to determine REW, which characterizes the fraction of available water for plant uptake. REW provides a better estimation of soil water holding capacities compared to SWC by integrating soil physical properties (through  $\theta_{wp}$  and  $\theta_{fc}$ ) and roots water uptake capacities (through weighting  $\beta$  with root length or root biomass). In addition, field capacity and wilting point can be measured by a single soil sampling campaign and do not change over time (if no soil disturbances occur). As a result, this variable presents two key advantages: (i) it provides additional information regarding soil water availability to plant uptake compared to other approaches considering SWC solely, and (ii) it avoids the drawbacks of characterizing soil water status with SWP (see chapter 1). However, several uncertainties

and methodological considerations remain.

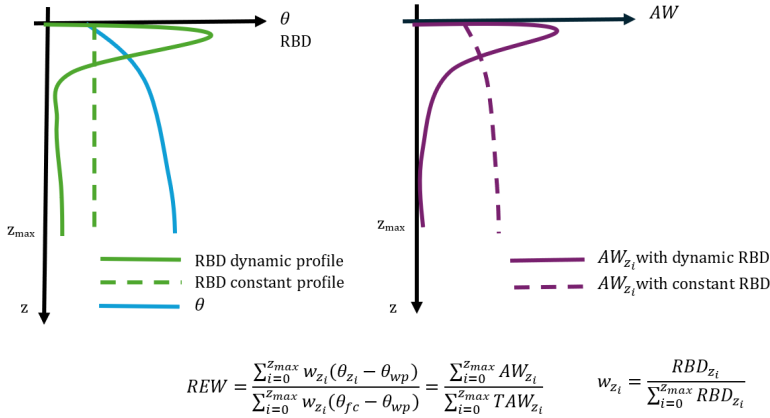
### 2.2.1. Diverging REW thresholds for similar ecosystems

The results of this thesis show that different REW thresholds were found for potato in chapter 4 and chapter 5 where leaf-level  $g_m$  decreased from a higher REW threshold than canopy-scale  $V_{cmax,app}$  (i.e., respectively  $REW = 0.72 \pm 0.12$  and  $REW = 0.43 \pm 0.03$ ). Similar uncertainties were also observed for forest ecosystems although the confidence intervals of the thresholds were overlapping and were not significantly different from 0.4. These differences can be explained by the design of the drought experiment setup (chapter 5), the measurement of SWC, or to the calculation of REW.

First, similar potato varieties were selected for the drought experiment (chapter 5) compared to the main variety planted in Loncée in 2018 (i.e., the variety Agria, chapter 4). This limits the differences in species-specific traits when comparing the results. However, no fertilizers, pesticides or fungicides were applied during the drought experiment while the EC site in Loncée is managed by conventional farming practises. This could have weakened the photosynthetic apparatus during the drought experiment (by e.g., a lack of nutrients), resulting in an increasing sensitivity to soil drying and a higher REW threshold.

Second, we calculated REW by considering either a constant weighting scheme (i.e., REW weighted only with the depth of soil layers inferred from SWC temporal dynamics, chapter 3, chapter 4 and chapter 5), or a dynamic weighting scheme (i.e., REW weighted by a dynamic RBD determined from root biomass measurements, chapter 6). Assuming a constant rooting depth in the topsoil might be challenged for potato, where root water uptake can be observed in deep soil layers (Stalham & Allen, 2004, 2001). In particular, differences in root anatomy can lead to very different Root Water Uptake (RWU) capacities. For instance, species with a high fine root hair fraction can access water strongly held by adsorptive forces by finer pores, leading to an increasing capacity to sustain the water flow under high atmospheric pressure (e.g., Marin *et al.* (2021) for barley, or Duddek *et al.* (2023) for maize). Another example is the lateral root formation which facilitates root water uptake (Agee *et al.*, 2021). When the contribution of soil water adsorptive forces compared to capillarity increases in a specific soil layer, plants may also stimulate the production of fine roots to promote water uptake from deeper soil layers where water is mostly held by capillarity. Although these patterns have been highlighted in many tree species (e.g., Peek *et al.* (2006); Leonova *et al.* (2022); Zwetsloot & Bauerle (2021)) and crops (Li *et al.*, 2022a), intra-species uncertainties remain (Gessler *et al.*, 2022; Walthert *et al.*, 2021). Acclimation to local climate conditions seem to play a key role in the drought tolerance of an ecosystem (Puchi *et al.*, 2024; Fan *et al.*, 2017). Overall, a wide variety of water uptake depth across PFTs and climate conditions can be expected (Bachofen *et al.*, 2024). In this perspective, measurements of rooting depth or RBD provides key information but relies on time and labor-consuming destructive biomass sampling campaigns (as detailed

in chapter 6). Instead, RBD models could be used to determine the water uptake capacities profile in the soil and REW weighting factors at a finer temporal resolution (Perona *et al.*, 2022; Kumar *et al.*, 2015). Non-destructive methods for measuring root biomass are also detailed in Cabal *et al.* (2021). This aspect is critical, as REW is very sensitive to the factors used to weight the available water for plant uptake in each soil layer (Fig. 7.4).



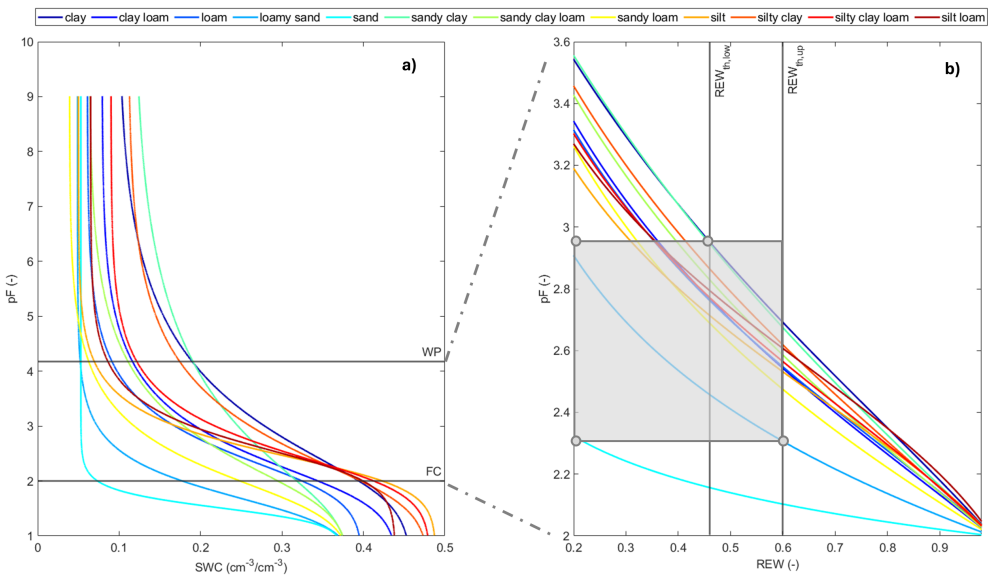
**Figure 7.4:** Conceptual illustration of a theoretical constant and dynamic RBD profile and the corresponding weighting factors ( $w_{z_i}$ ) of REW for each soil layer ( $i$ ) at a specific depth  $z_i$ . For simplification, the wilting point ( $\theta_{wp}$ ) and field capacity ( $\theta_{fc}$ ) are considered constant.

REW is calculated as the ratio of available water (AW) to total available water (TAW) integrated over the root zone and weighted by  $w_{z_i}$ , which are calculated based on a proportion of the RBD profile.

Finally, differences in SWC measurements could originate from sensors calibration equations, which are pivotal for linking the raw measurement of soil dielectric permittivity to SWC (Cosh *et al.*, 2005; J. D. Jabro *et al.*, 2005; RoTimi Ojo *et al.*, 2015). As the sensors used in these two studies were not the same (i.e., EnviroScan sensor in the 2018 dataset, ML2 and ML3 sensors for the other years in chapter 4; ML2 sensor in chapter 5), the corresponding uncertainty could have been exacerbated. In addition, uncertainties remain when estimating field capacity and wilting point from either laboratory or field observation-based methods (Parker & Patrigiani, 2021; Evett *et al.*, 2019; Kukul & Irmak, 2023). The measurement of SWC by TDR sensors can be also strongly overestimated in highly conductive soils such as loam and clay soils (Bittelli *et al.*, 2008). As a result, a series of uncertainties can be propagated throughout REW calculation. Comparing REW data across sites is therefore challenging (Ford & Quiring, 2019). These considerations advocate for a careful use of REW for characterizing water status.

### 2.2.2. From thresholds of REW to thresholds of SWP

Despite these methodological considerations which may explain the differences between REW thresholds, it is worth noting that these values remain in a narrow range when considering the lower limit of the leaf-scale and the upper limit of the ecosystem-scale thresholds (i.e., between 0.60 and 0.46). This range also includes most of thresholds found for forest ecosystems in chapter 3 (Fig. 3.2). This result can be interpreted in terms of water potentials. To this end, we built SWRCs by using VG model parameters of twelve typical soil types as defined by the United States Department of Agriculture (USDA) in the pedotransfer function Rosetta (Schaap *et al.*, 2001; Zhang & Schaap, 2017). These SWRCs were used to get the field capacity and wilting point of each soil textural class. REW was then reconstructed to interpret the thresholds in terms of SWP (Fig. 7.5). Importantly, it can be shown that the range of REW thresholds corresponds to a range of SWP between -19.6 kPa and -87.4 kPa for most of soil textural classes. Only sand soils had a smaller range of SWP thresholds between -12 kPa and -13.8 kPa (i.e., pF between 2.10 and 2.15, Fig. 7.5).



**Figure 7.5:** REW thresholds linked to SWP by SWRCs of twelve USDA soil textural classes. SWRCs were reconstruct by using VG model parameters from the Rosetta database (Schaap *et al.*, 2001; Zhang & Schaap, 2017).  $pF = \log_{10}(\psi)$  with  $\psi$  in cm. Panel a) shows the SWRCs of the twelve different soil types with field capacity (FC) and wilting point (WP). Panel b) shows the relationship between REW and pF. The shaded grey area illustrates the range of pF values corresponding to REW between 0.46 and 0.6 (i.e., the thresholds of NSOL found in this thesis for potato and forest ecosystems).



By comparison, other studies have found that the water transport in the roots is impeded from SWP thresholds ranging from -6 to -1000 kPa (Cai *et al.*, 2022). Although a SWP of -25 kPa has been reported as optimal growth conditions for potato (Wang *et al.*, 2007), data is still critically lacking. Much lower critical SWP values ranging from -470 kPa to -1200 kPa have been reported for tree species (Eckes-Shephard *et al.*, 2021; de Blécourt *et al.*, 2021; Schütt *et al.*, 2022). These values are however difficult to compare as they strongly differ by the choice of the response variable to changes in SWP. The results of this thesis underline two important outcomes: physiological processes through NSOL are impacted at a much smaller SWP than wilting point ( $pF=4.2$  or  $\psi_{soil}=-1550$  kPa) and calibrating  $\beta$  provides a relatively narrow range of SWP thresholds across soil types. Importantly, these values can also be used in models which uses  $\beta$  as a function of SWP or incorporates plant hydraulics (see section 3.4).

### 2.2.3. Downregulating NSOL with the $\beta$ scalar

In this thesis, we chose to calibrate the response of photosynthesis limitations to soil water stress by adjusting  $\beta$  on field data. The main feature of this approach is the determine the shape of  $\beta$  (itself a function of REW) and the parameter on which it should be applied to capture plant response to drought (either on SOL and/or on NSOL). This assumes that REW is a bulk variable encompassing many other processes in the SPAC continuum such as hydraulic or physiological limitations. In addition to the uncertainties when calculating REW (as detailed in section 2.2), the shape of  $\beta$ , and the thresholds of REW are expected to vary in time, or across PFT (Robinson *et al.*, 2019; Teuling *et al.*, 2006). The lack of an universal  $\beta$  formulation is one primary source of uncertainties in carbon cycling modeling, representing between 40-50 % of inter-model variability in GPP predictions (Trugman *et al.*, 2018; Paschalis *et al.*, 2020). Although modifying  $\beta$  to consider plant-specific resistances to drought through a site calibration is possible and recommended (section 3.3, Fig. 7.6), this approach globally lacks from theoretical foundation which may ultimately leads to unrealistic drought responses of gas exchanges in LSMs. Calibrating  $\beta$  for different physiological stages, species, plant species or climate zones represents a huge effort with its share of uncertainty not to be overlooked. An alternative to these empirically-based approaches for reproducing vegetation responses to drought considers the explicit representation of plant hydraulics in coupled photosynthesis-stomatal conductance models (Eller *et al.*, 2020; Sabot *et al.*, 2020). This is further detailed in section 3.4.

## 2.3. Mesophyll conductance and carbon isotopes

Mesophyll conductance was estimated from joint measurements of gas exchange and chlorophyll fluorescence under natural conditions of CO<sub>2</sub> concentration (Harley *et al.*, 1992). Alternatively,  $g_m$  can also be determined by exploiting the fractionation of carbon isotopes naturally occurring during photosynthesis.

Atmospheric CO<sub>2</sub> contains about 1.1 ‰ of <sup>13</sup>C (the heavier C isotope) and about 98.9 ‰ of <sup>12</sup>C (the lighter C isotope). During photosynthesis, <sup>13</sup>C is discriminated against <sup>12</sup>C because of (i) its slower diffusion to chloroplasts and (ii) the reduced capacity of Rubisco to fix heavy <sup>13</sup>C (O'Leary, 1981). As a result, photosynthetic products are typically enriched in <sup>12</sup>C compared to <sup>13</sup>C. The deviation of the sample isotope ratio compared to a known standard ( $\delta^{13}$ ) is related to the carbon isotope discrimination ( $\Delta_{iso}$ ) by:

$$\Delta_{iso} = \frac{\zeta(\delta_{in}^{13} - \delta_{out}^{13})}{1 + \delta_{out}^{13} - \zeta(\delta_{out}^{13} - \delta_{in}^{13})} \quad (7.1)$$

$$\zeta = \frac{C_{in}}{C_{in} - C_{out}} \quad (7.2)$$

with  $\delta_{in}^{13}$ ,  $\delta_{out}^{13}$ ,  $C_{in}$  and  $C_{out}$  are the measured carbon isotope composition and CO<sub>2</sub> concentration inside and outside the measurement chamber (Evans *et al.*, 1986). The isotopic composition of leaf samples can be measured by laser absorption spectroscopy (Bowling *et al.*, 2003; TAZOE *et al.*, 2011; Holloway-Phillips *et al.*, 2019) which can be coupled to the LI-6400 for joint measurements with CO<sub>2</sub> concentration (Sonawane & Cousins, 2019; Holloway-Phillips *et al.*, 2019). The main advantage of these methods is the ability to measure carbon isotope composition in real-time under controlled conditions set in the chamber, therefore allowing to track carbon fractionation at the sub-daily timescale (Diao *et al.*, 2024). Mesophyll conductance can be determined by the difference between carbon isotope discrimination assuming that  $g_m$  is infinite ( $\Delta_{iso,inf}$ ) and  $\Delta_{iso}$  as follows (Busch *et al.*, 2020):

$$g_m = \frac{1 + t}{1 - t} \frac{b - a_m - \frac{1+b}{(1+e')\alpha_R} e' \frac{R_d}{A}}{\Delta_{iso,inf} - \Delta_{iso}} \frac{A}{C_a} \quad (7.3)$$

where  $a_m$  represents fractionation associated with CO<sub>2</sub> diffusion in water (1.8‰),  $b$  represents the fractionation during carboxylation (30‰),  $e'$  the fractionation during day respiration,  $t$  is the ternary effect (Farquhar & Cernusak, 2012) and  $\alpha_R = 1 + \frac{R_d}{A} \frac{e'}{(1+e')}$ . It has been shown that both methods (gas exchange - fluorometry or gas exchange - laser absorption spectroscopy) yielded similar results, highlighting that they all converge towards a uniform  $g_m$  estimate (Marino *et al.*, 2020; Kromdijk *et al.*, 2019; Flexas *et al.*, 2008; Warren, 2006b). Joint measurements of carbon and oxygen isotopes also showed promising perspectives to partition  $g_m$  components and assess their response to environmental drives (Barbour *et al.*, 2016; Xiao & Zhu, 2017). Nonetheless, each of these methods has its own limitations, and  $g_m$  measurements should be interpreted with caution (Sun *et al.*, 2014; Flexas *et al.*, 2008; Pons *et al.*, 2009). For instance, isotope methods requires a precise knowledge of isotopic fractionation factors and an expansive instrumentation. In this perspective, joint measurements gas exchange and ChlF (actively-induced or with SIF - see section 3.5) still

remains a widely used approach for characterizing the impact of edaphic drought on mesophyll conductance. The recent review of Cousins *et al.* (2020) gives a series of recommendations and discuss the recent advances for measuring  $g_m$ .

### 3. Perspectives

In this section, we draw some perspectives to this work which are summarized in Fig. 7.13.

#### 3.1. Drought resistance of croplands from EC data

This thesis opens up new perspectives for the quantification of stomatal and non-stomatal responses of other crops or PFTs. The method proposed in chapter 4 (and originally developed in chapter 3) could be applied on other agricultural ecosystems of the ICOS and FLUXNET networks to determine whether SOL or NSOL control GPP under water stress, and to calibrate  $\beta$  to reduce uncertainties in LSMs. In particular, most of ICOS sites provide SWC measurements and a description of the soil profile with soil textural proprieties in each soil layer. Using such approach with EC fluxes can be used to determine crops drought tolerance and resistance, with the crop with the lowest REW threshold being the most resistant to water stress (e.g., Gholipoor *et al.* (2013)). This method could be directly tested on nearly 30 EC cropland sites of ICOS and FLUXNET (Pastorello *et al.*, 2020). Drought-resistant tree populations could also be identified by applying this method on other forest ecosystems.

#### 3.2. REW thresholds as a tool for improving crop yields and irrigation scheduling

Reliable estimates of drought responses of physiological processes are also important for improving crop modeling. In particular, the Aquacrop model (Raes *et al.*, 2009) uses empirical factors to implement water stress effects on crop growth and simulate crop yield and irrigation scheduling. In a similar fashion to  $\beta$  in LSMs, these modifiers are either linear, concave or convex, and affect physiological processes when the soil water availability is lower than a percentage of the total amount of water available for plant uptake in the root zone (i.e., also known as Total Available Water (TAW), the denominator of REW). In Aquacrop, five water stress coefficients are used to affect canopy expansion, transpiration, stomatal closure, canopy senescence and pollination (Raes *et al.*, 2009). A set of upper and lower limits (when the stress function starts/stops operating) are defined for each crop species. Not only impacting crop yield estimates, these thresholds also influence the design of irrigation scheduling, which is provided by Aquacrop through an irrigation module (Alvar-Beltrán *et al.*, 2023; Mwiya *et al.*, 2020). Among the different scheduling approaches, one particularly relies on thresholds of SWC (i.e., the 'soil moisture-based' approach (Gu *et al.*, 2020)). The

key feature of this method is to maintain the SWC in the root zone within an acceptable range for crop growth. This threshold is typically set to the field capacity, which implies that the root zone is refilled as soon as REW is lower than 1 (Wang *et al.*, 2017; Haley & Dukes, 2012). Another approach consists in using a depletion factor, which is a threshold of REW used to calculate the 'readily available water' from the TAW (Allen *et al.*, 1998). The amount of water to be added directly depends on the depletion factor (i.e., very similar to  $REW_{th}$ ) which is often set to 0.5 (Allen *et al.*, 1998). For potato, existing irrigation scheduling strategies use a threshold of 0.70-0.85 which can vary at different growth stages (King *et al.*, 2020). We suggest that this threshold can be reduced to 0.46-0.6 for the potato plants studied in this thesis. Determining these thresholds is particularly important in the context of water scarcity, where every drop of water must be efficiently used for crop growth. These 'threshold-based' approaches can be used as a preventive tool to buffer the negative effects of drought on crop physiology. More studies based on the calibration of the response of physiological processes under water stress are needed.

### 3.3. Modifying $\beta$ for implementing more realistic responses to drought in modeling

Instead of setting the starting point of the downregulation of NSOL by  $\beta$  to the field capacity (Eq. 1.49), we suggest to use more realistic SWC values from which model parameters should be impacted by water stress. In particular, we calibrated a linear-plateau model to characterize the response of NSOL and SOL to REW. This model has a simple linear segment below the REW threshold ( $REW_{th}$ ), characterized by a slope ( $a$ ) and an intercept ( $b$ ). As  $\beta$  is an empirical scalar which affects model parameters when REW decrease,  $\beta$  should correspond to the relative variation of  $y$  compared to its maximum value  $y_{max}$ . In other words, if  $y$  has decreased by 50% compared to  $y_{max}$  because of REW,  $\beta$  should equal 0.5. This can be described mathematically as follows. The linear-plateau model is:

$$y = \begin{cases} y_{max}, & REW \geq REW_{th}, \\ a \cdot REW + b, & REW < REW_{th} \end{cases} \quad (7.4)$$

with  $a = \frac{y_{max} - b}{REW_{th}}$ .  $b$ ,  $y_{max}$  and  $REW_{th}$  are fitted parameters. By considering that (i)  $\beta$  must be expressed as a function of REW, and (ii)  $\beta$  must equal the relative variation of  $y$  compared to  $y_{max}$ , one can write:

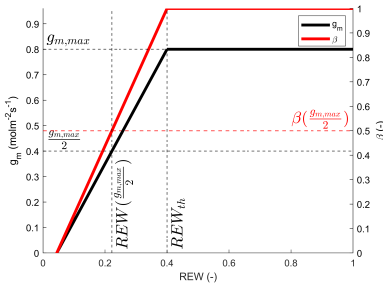
$$\beta = \frac{y}{y_{max}} = \frac{REW}{REW_{th}} \left( 1 - \frac{b}{y_{max}} \right) + \frac{b}{y_{max}} \quad (7.5)$$

This equation shows that  $\beta$  is dependent on three fitted parameters which are species (or PFT)-specific. Eq. 7.5 can be written in an integrated form using REW or SWC:

$$\beta = \begin{cases} 1 & REW \geq REW_{th} \\ \frac{REW}{REW_{th}} \left(1 - \frac{b}{y_{max}}\right) + \frac{b}{y_{max}} & REW < REW_{th} \end{cases} \quad (7.6)$$

$$\beta = \begin{cases} 1 & \theta \geq REW_{th}(\theta_{fc} - \theta_{wp}) + \theta_{wp} \\ \frac{(\theta - \theta_{wp})}{REW_{th}(\theta_{fc} - \theta_{wp})} \left(1 - \frac{b}{y_{max}}\right) + \frac{b}{y_{max}} & \theta < REW_{th}(\theta_{fc} - \theta_{wp}) + \theta_{wp} \end{cases} \quad (7.7)$$

In total, Eq. 7.7 requires two additional parameters compared to Eq. 7.5 : the field capacity and wilting point ( $\theta_{fc}$ ,  $\theta_{wp}$ ), which allow to convert REW into SWC. As an example, a theoretical relationship between  $g_m$  and REW, and the corresponding  $\beta$  are given in Fig. 7.6.



**Figure 7.6:** Example of a modified  $\beta$  factor considering a linear-plateau relationship between  $g_m$  and REW using  $g_{m,max} = 0.8 \text{ mol m}^{-2} \text{ s}^{-1}$ ,  $REW_{th} = 0.4$  and  $b = -0.1 \text{ mol m}^{-2} \text{ s}^{-1}$ .

Eq. 7.7 provides a more realistic calibration of the response of NSOL to the decrease in REW. Notably, Eq. 7.7 is very similar to the equation used in JULES (v4.6). In particular, a linear-plateau also describes the response of  $A_n$  to soil water availability, and  $\beta$  is operating only from an upper SWC threshold ( $\theta_{upp}$ ) which depends on a parameter ( $p_0$ ) allowing to delay the decrease in carbon assimilation (Harper *et al.*, 2021):

$$\theta_{upp} = (1 - p_0)(\theta_{fc} - \theta_{wp}) + \theta_{wp} \quad (7.8)$$

By comparing Eq. 7.8 and Eq. 7.7, it can be show that  $p_0$  is directly related to  $REW_{th}$  as:

$$p_0 = 1 - REW_{th} \quad (7.9)$$

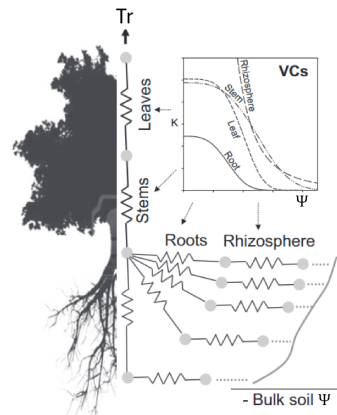
In JULES,  $p_0$  is set to 0.4, which means that  $\beta$  is operating from a  $REW_{th}$  of 0.6 (Harper *et al.*, 2021). Moreover,  $\beta$  equal zero when  $\theta = \theta_{wp}$  (i.e., when  $REW=0$ ), which also implies that the intercept  $b$  in Eq. 7.7 equal 0. It can be shown that the mathematical expression of  $\beta$  in JULES (Harper *et al.*, 2021) and Eq. 7.7 are exactly similar if  $b = 0$ . Therefore, Eq. 7.7 requires two more parameters compared to JULES ( $y_{max}$  and  $b$ ), but will inevitably yield significant improvement for representing the response of physiological processes to water stress. These additional parameters can be determined from EC and leaf-level measurements. Modifying the weighting scheme

and assessing the impacts on LSMs estimates is currently explored in numerous studies. Future studies should focus on testing the implementation of a modified  $\beta$  factor accounting for RBD and RWU plant capacities in LSMs. For example, better estimates of  $Tr$  under water stress have been found when using RWU instead of rooting fraction (similarly to rooting depth) for  $\beta$  in the LSM CLM5 (Sulis *et al.*, 2019). It is also urgent to unify soil water availability metrics to facilitate the interpretation of drought experiment results (Vicca *et al.*, 2012).

### 3.4. The profit-based optimization: an alternative to empirical solutions for implementing plant responses to drought in modeling

One of the alternatives of  $\beta$ -based empirical approaches for modeling plant responses to drought is plant hydraulic models, which represent the water transport in vegetation using flux-gradient relationships. Incorporating plant hydraulics in photosynthesis models provides a mechanistic bridge between physiological regulation of leaf gas exchanges and environmental drivers. These models aim to capture plant responses to drought through the representation of stomatal responses to LWP, which is the most direct variable for assessing leaf water status (Brodribb *et al.*, 2003; Klein, 2014). The LWP sensitivity to stomata is tightly coupled to xylem anatomy and vulnerability to cavitation. Here we describe the 'loss-gain' theory, where plant hydraulics modeling is based on the supply of transpiration while considering a 'cost' and a 'gain' function (Sperry *et al.*, 2017).

The water flow is supplied to the leaves as transpiration by the tension-cohesion mechanism along the SPAC where water flows from low to high potentials (Fig. 7.7). Importantly, the transpiration stream is driver by total plant resistance, along with soil and atmospheric water deficits (Wolfe *et al.*, 2023). The supply function therefore describes the relation between transpiration ( $Tr$ ) and xylem water potential at a specific soil water potential. It is calculated by integrating vulnerability curves within each component of the SPAC (i.e., the response of the hydraulic conductance  $k$  to the decrease in water potential  $\psi$ ) following (Sperry *et al.*, 2017):



**Figure 7.7:** Conductance network linking bulk soil water potential ( $\psi$ ) to sustain transpiration from the rhizosphere to the leaves. Each network component has a vulnerability curve (VC) describing the drop in hydraulic conductance  $k$  with increasing  $\psi$  (Sperry *et al.*, 2016).

$$Tr = \int_{\Psi_{up}}^{\Psi_{down}} k(\Psi)d(\Psi) \quad (7.10)$$

where  $k$  is described with a two-parameter vulnerability curve for xylem components (Sperry *et al.*, 2016, 2017):

$$k(\Psi) = k_{max}e^{[(\frac{-\Psi}{b})^c]} \quad (7.11)$$

with  $k_{max}$  the maximum hydraulic conductance,  $b$  corresponding to  $\Psi$  when  $\frac{k(\Psi)}{k_{max}} = 0.37$  and  $c$  controls the shape of the vulnerability curve. A vulnerability curve for rhizosphere is also proposed as a function of  $k_{max}$ , soil texture-specific parameters and SWP using a VG function. A common assumption of such model is the consideration of unsegmented xylem components (i.e., the same vulnerability curve for all xylem components) (Sperry *et al.*, 2017).

**On the demand-side**, the 'demand' function specifies the actual regulation of plant transpiration rate along the supply function, mostly achieved by stomatal movements. The role of the 'loss' function is to avoid a critical loss of hydraulic conductance leading to cavitation, and to utilize the full supply capacity of xylem for sustaining transpiration. As a result, the 'demand' and supply functions are mathematically dependant (Sperry & Love, 2015). The demand function characterizes the restrictive effect of increasing  $\psi$  on the transpiration flow as (Sperry *et al.*, 2017):

$$\Pi(\Psi) = \frac{k_{max} - k(\Psi)}{k_{max} - k_{crit}} \quad (7.12)$$

with  $k_{crit}$  the critical hydraulic conductance when  $\Psi = \Psi_{crit}$  at dessication. While this approach clearly identifies the fitness cost of moving water (i.e., the loss of conductivity), it ignores the role played by stomata for regulating photosynthesis and the corresponding carbon gain. To that end, a 'gain' function has been added in the model of Sperry *et al.* (2017).

**On the gain-side**, the 'gain' function considers the carbon gain through photosynthesis related to stomatal opening defined as Sperry *et al.* (2017):

$$\Omega(\Psi) = \frac{A(\Psi)}{A_{max}} \quad (7.13)$$

where  $A$  is the gross assimilation rate, determined by a photosynthesis model similar to the FvCB model (Collatz *et al.*, 1991). This latter is determined by coupling the supply function (giving the plant conductance to water) to the Fick law representing CO<sub>2</sub> diffusion through stomata.  $A$  is evaluated at each  $\Psi$  and  $A_{max}$  is the maximum  $A$  over the full range of SWP.

For each supply function of  $T_r$ , the leaf temperature is calculated from radiation, wind speed and air temperature following a flux-gradient equation. The leaf to air vapor pressure deficit is then determined, followed by the canopy conductance, ultimately used for the internal  $\text{CO}_2$  concentration in the gain function. Overall, the model requires environmental drivers, hydraulic cost and photosynthetic gain parameters (see Table 1 in Sperry *et al.* (2017)). Among the parameters specific to the hydraulic cost, only the shape of the vulnerability curve and  $k_{max}$  need to be set to a specific value ( $b$  is set by  $k_{max}$ , and  $k_{max}$  is set by  $V_{cmax}$ ). No other parameters is specie or PFT-specific.

At each fixed instant for a given set of environmental drivers, a family of water potential response curves is generated by the model, representing steady state values. The plant can only be on one single point among all the possibilities. The profit ( $Pr$ ) maximization criterion solve all these possibilities by maximizing the difference between the photosynthetic gain ( $\Omega(\Psi)$ ) and the hydraulic cost ( $\Pi(\Psi)$ ) (Wolf *et al.*, 2016; Sperry *et al.*, 2017):

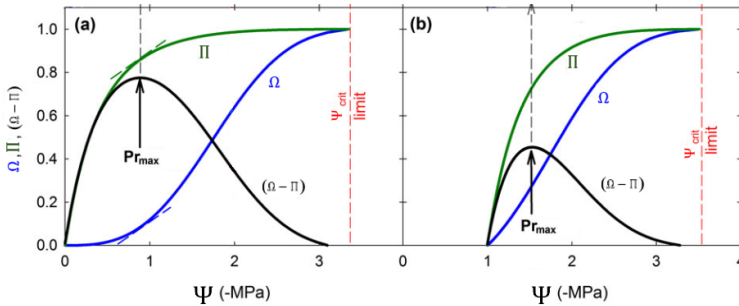
$$Pr = \max(\Omega(\Psi) - \Pi(\Psi)) \quad (7.14)$$

which occurs when the partial derivative of  $\Omega(\Psi)$  and  $\Pi(\Psi)$  are equal. This maximum difference occurs at a unique  $\Psi$  in the plant, which gives a set of model solutions at the corresponding instant and defines a hydraulic safety margin for maximizing carbon gains and minimizing hydraulic failure. The influence of soil water availability is considered in the model through changes in SWP which directly influences the rhizosphere water potential, resulting in a smaller optimal plant potential and a smaller stomatal conductance (Fig. 7.8).

The 'profit-maximization' approach applied to the 'cost-gain' theory in plant hydraulics does not require any calibration of empirical factors related to SWC on SOL or NSOL. Instead, only few plant traits (e.g.,  $k_{max}$ ) need to be calibrated on measurements of stomatal conductance, or determined by in situ measurements. SWC data can be converted into SWP by SWRC and fed in the model to predict plant response to drought. While the gain function is well identified (i.e., carbon gain through photosynthesis), many optimization models have been developed in the past years, where the 'cost' function has been associated to a parabola function of leaf water potential (Wolf *et al.*, 2016) or to a reduction of photosynthetic capacities (Dewar *et al.*, 2018). Finally, Joshi *et al.* (2022) set the profit-maximization hypothesis as the benefit from photosynthetic assimilation minus the costs of maintaining the photosynthetic capacities and the hydraulic pathway. This latter study yielded particularly good performances for predicting carbon assimilation, stomatal conductance and photosynthetic capacities under progressive soil drying (Joshi *et al.*, 2022).

These approaches provide undoubtedly a more realistic representation of plant responses to drought. However, do they outperform empirical approaches? In a broad review testing the performances of twelve empirical and optimization approaches, Sabot





**Figure 7.8:** The hydraulic cost function ( $\Omega$ , blue) is the fractional loss of hydraulic conductance that is calculated from the supply function  $Tr(\Psi)$ . The gain function ( $\Pi$ , green) is the fractional increase in  $A$  that is calculated from the  $A(\Psi)$  curve. The maximum profit is reached when the derivative of  $\Omega - \Pi$  function is null. At this point, stomata are assumed to maintain the plant at the optimum. The corresponding  $\Psi$  solution is used to calculate leaf temperature, leaf VPD,  $A$ ,  $Tr$ , internal  $\text{CO}_2$  concentration and stomatal conductance. The second panel (b) describes a situation when soil water potential shifts from 0 MPa (panel a) to -1 MPa (panel b). The optimal  $\Psi$  solution becomes more negative, and the six parameters mentioned previously decrease correspondingly. Adapted from Sperry *et al.* (2017).

*et al.* (2022) highlighted that five optimization schemes (including the profit maximization model described above) showed slightly better performances than the USO model for predicting leaf-level stomatal conductance. While the inclusion of plant hydraulics and the profit-maximization approach generally improves the modeling of vegetation gas exchanges (Bonan *et al.*, 2014; Eller *et al.*, 2020; Kennedy *et al.*, 2019; Sabot *et al.*, 2020), it comes at a high computational cost (Prentice *et al.*, 2015). In addition, the parameters of the vulnerability curves are difficult to measure because (i) they are typically measured at a single SPAC component and do not reflect the behavior of the whole plant, and (ii) they remain highly variable across species and PFTs (Couvreur *et al.*, 2018). Modifying  $g_1$  of  $V_{cmax}$  by  $\beta$ , either as a function of SWC or SWP (TUZET *et al.*, 2003; Zhou *et al.*, 2013; Anderegg *et al.*, 2017), still remains the most widely used method while yielding reasonably good results (Blyth *et al.*, 2021; Paschalis *et al.*, 2020). More studies are needed to test and implement these novel approaches in LSMs (Blyth *et al.*, 2021; Sabot *et al.*, 2022).

### 3.5. Canopy-scale mesophyll conductance

Characterizing the response of mesophyll conductance to meteorological drivers remains an important model deficiency in LSMs (Knauer *et al.*, 2020). As  $g_m$  is related to the electron transport rate, it could be estimated by using the mechanistic representation of the light reactions of photosynthesis provided by the MLR model (Eq. 1.42, chapter 6). Therefore, joint measurements of SIF and GPP at the ecosystem scale holds great promises for estimating  $g_m$  at a high temporal and large spatial resolution.

In particular,  $G_m$  (canopy-scale  $g_m$ ) could be estimated from SIF and GPP as follows:

$$G_m = \frac{GPP}{C_i - \frac{\Gamma^*(J_{SIF} + 8GPP)}{J_{SIF} - 8GPP}} \quad (7.15)$$

where  $C_i$  is canopy-scale  $\text{CO}_2$  concentration in substomatal cavities and is a function of canopy conductance to water vapor and GPP (see chapter 4). Therefore, using joint measurements of GPP and SIF at the ecosystem scale could also provide REW-response curves of NSOL based on automatic SIF measurements. This would certainly help in disentangling the role of mesophyll conductance in limiting  $\text{CO}_2$  supply, reducing LSMs uncertainties, and engineering mesophyll conductance to improve photosynthesis under water stress (Nguyen *et al.*, 2023; Lundgren & Fleming, 2020).

### 3.6. Estimating $R_{\text{ECO}}$ from NEE and SIF

By relying solely on meteorological drivers and SIF, the MLR model also represents an interesting approach for separating NEE into GPP and  $R_{\text{ECO}}$ . For now, EC data post-processing pipelines use either the NT or DT partitioning methods to estimate  $R_{\text{ECO}}$  and GPP from NEE (Reichstein *et al.*, 2005; Lasslop *et al.*, 2012; Pastorello *et al.*, 2020). As detailed in chapter 6 and section 2.1, these methods present some key uncertainties. Recent studies have used either machine learning algorithms (Zhan *et al.*, 2022) or empirical methods (Kira *et al.*, 2021) to determine  $R_{\text{ECO}}$  from NEE and SIF, which led to a significant improvement of  $R_{\text{ECO}}$  estimates under water stress and high temperatures. By using the MLR model for GPP and NEE measurements,  $R_{\text{ECO}}$  can be determined as:

$$R_{\text{ECO}} = \text{NEE} + \frac{C_c - \Gamma^*}{4C_c + 8\Gamma^*} \frac{q_L \text{SIF}_{\text{TOT}} \phi_{\text{PSII},\text{max}}(1 + k_{\text{DF}})}{1 - \phi_{\text{PSII},\text{max}}} \quad (7.16)$$

### 3.7. Versatility of the MLR-USO model: integrating the complexity of heterogeneous canopies

As listed in section 1.2, several key limitations remain for a use of the MLR-USO model on other ecosystems and at larger scales. In a recent study, Chen *et al.* (2024c) proposed an adaptation of the MLR model on three evergreen needleleaf forests in Canada (site code CA-Obs), United States (US-NR1) and South Korea (KR-TCK). SIF observations were collected by a PhotoSpec spectrometer at 26 m height (at CA-Obs and US-NR1) and by a QE-Pro spectrometer at 40 m height (at KR-ICK). The MLR model was implemented into a layered, a two-leaf and a layered-two-leaf model.

In the MLR model (Eq. 1.43), several parameters are expected to show a strong spatial variability within the canopy, notably  $q_L$  (which is related to leaf light absorption properties) and  $C_i$  (which is related to both stomatal and photosynthetic dynamics).

The most straightforward approach consists in assuming that the  $q_L$  light-response curve is the same for all the leaves of the canopy whatever their light exposure. The alternative is to measure  $q_L$  light response curves at different depth within the canopy, or on sunlit and shaded leaves. In addition, the origin of the SIF signal must be attributed to each of these different canopy components. Overall, the light distribution and the SIF emission within the canopy must be described.

### 3.7.1. The layered MLR model

In the layered model, the PAR within each layer  $i$  is calculated from top-of-canopy PAR following Chen *et al.* (2024c):

$$PAR_i = PAR e^{-\frac{0.5 \Omega_{Cl} LAI_{ai}}{\cos(SZA)}} \quad (7.17)$$

with  $SZA$  the solar zenith angle,  $LAI_{ai}$  the accumulated leaf area index from the top to layer  $i$ , and  $\Omega_{Cl}$  the clumping index (i.e., characterizing the clumping of leaves within the canopy). These two parameters are related to canopy structure and are expected to be PFT-specific and to vary throughout the growing season. This model therefore requires time series of  $\Omega_{Cl}$ , which can be retrieved from RS data (Chen *et al.*, 2024c) or measured at the field scale (Liang *et al.*, 2023). Eq. 7.17 can be used to determine  $q_L$  using either a single or multiple light-response curves. Instead of using a temperature-response curve, half-hourly  $C_i$  can be estimated from a PAR-based relationship as these two variables are expected to be strongly correlated (Beauclair *et al.*, 2023a). The decomposition of SIF emission within the canopy is described by dividing top-of-canopy  $SIF_{TOT}$  as:

$$SIF_{TOT,i} = SIF_{TOT} \frac{APAR_i \phi_{SIF,i}}{\sum_i^n APAR_i \phi_{SIF,i}} \quad (7.18)$$

where  $n$  is the number of layers,  $APAR_i$  is calculated as the difference between the  $PAR$  in layer  $i + 1$  and the  $PAR$  in the layer  $i$  (Eq. 7.17, and  $\phi_{SIF,i}$  is determined from a nonlinear function with  $PAR_i$  (Liu *et al.*, 2021). This layered version requires measurements of PAR and SIF at the top-of-canopy, as well as of the clumping index  $\Omega_{Cl}$  and the LAI. The total GPP is then calculated as the sum of all GPP in each  $i$ -th layer within the canopy.

### 3.7.2. The two-leaf MLR model

The two-leaf MLR model categorizes the leaves according to their light exposure (either shaded or sunlit) (Chen *et al.*, 2024c). The decomposition of the SIF signal is based on Eq. 7.18 with  $APAR$  divided into a sun and a shade component:

$$APAR_{sun} = (1 - a) PAR_{sun} LAI_{sun} \quad (7.19)$$

$$APAR_{shade} = (1 - a) PAR_{shade} LAI_{shade} \quad (7.20)$$

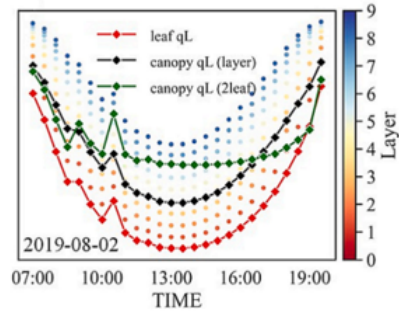
with  $a$  the leaf albedo. The calculation of  $PAR_{shade}$ ,  $LAI_{shade}$ ,  $PAR_{sun}$  and  $LAI_{sun}$  is detailed in Chen *et al.* (1999). The total GPP is then calculated as the sum of all GPP from sunlit and shaded leaves.

These models provide estimates of  $q_L$  that vary with leaf exposure or canopy depth (Fig. 7.9). Overall, these two models are expected to outperform the 'big leaf' model (Chen *et al.*, 2024c). Much work remains to be done, specifically for broadleaf forests typical of Western Europe. In addition, the understorey and midstorey are expected to strongly contribute to the measured SIF signal (Morozumi *et al.*, 2023b; Hornero *et al.*, 2021). This must be considered in the future versions of the model for a broader use on forest ecosystems.

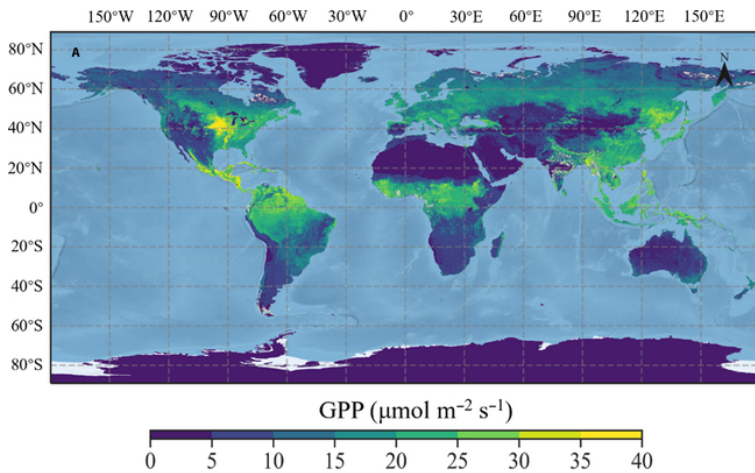
### 3.8. Versatility of the MLR-USO model: from the field to larger scales with RS SIF

#### 3.8.1. The first global GPP product based on the MLR model

In the perspective of using the MLR model at larger scales with satellite-based SIF, very interesting results are provided by a recent study of Chen *et al.* (2024b) who applied the MLR model with TROPISIF dataset (Guanter *et al.*, 2021) to estimate GPP (defined as the CMLR GPP product) from May 2018 to September 2021, and validated their estimates using ground-based GPP from 293 flux sites. The authors used the 'big leaf' version of the MLR model where  $q_L$  was parameterized by a machine learning techniques which consisted of (i) calculating  $q_L$  from Eq. 1.43 using measured ground based GPP and RS TROPOMI SIF, and (ii) taking 70% of the dataset to train a machine learning model using several explanatory variables such as VPD, NDVI, or PAR. The CMLR GPP dataset showed comparable performances when compared to other SIF-based GPP products. While relying on 'big leaf' assumptions and an empirical parameterization for  $q_L$ , their approach provides global maps of GPP with the MLR model at a  $0.05^\circ$  spatial resolution (Fig. 7.10). A significant improvement of the GPP product could be provided by sub-daily SIF data which would allow to track short-term responses of ecosystem functioning to climate drivers. In that case, the 'big leaf' underlying hypotheses are not likely to hold (Chen *et al.*, 2024b).



**Figure 7.9:** The diurnal patterns of 'big leaf'  $q_L$  (red lines), canopy-scale  $q_L$  from the layered model (black lines), canopy-scale  $q_L$  from the two-leaf model (green lines), and the  $q_L$  in different layers of the canopy (colored scatters). The colour of the dots represents the layer index, and the larger number represents the deeper location within the canopy. The LAI of each layer in this figure can be calculated as total LAI/10 with 10 being the total number of layers (Chen *et al.*, 2024c).



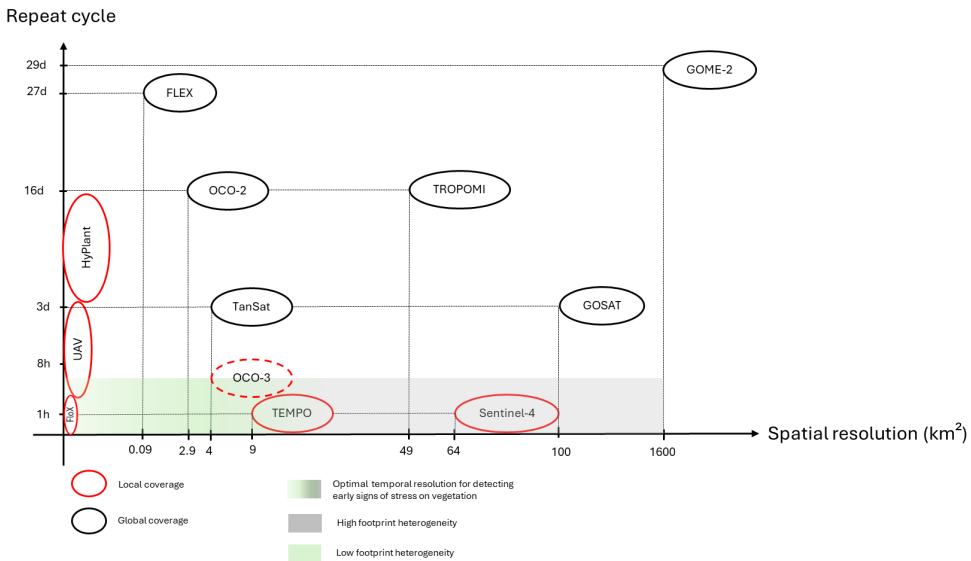
**Figure 7.10:** Global map of CMLR SIF-based GPP using the MLR model on 29 July 2019 at a  $0.05^\circ$  resolution (Chen *et al.*, 2024b).

### 3.8.2. Geostationary satellites for tracking sub-daily dynamics of plant gas exchanges

For now, the majority of RS SIF products are provided by satellites in sun-synchronous orbits. These data are either too coarse or with a long revisit time, which does not allow a detection of early stress responses of vegetation to drought. The most precise RS SIF measurement will be provided by FLEX ( $0.09 \text{ km}^2$  spatial resolution) but comes to an expense of a long revisit time (27 days) (Fig. 7.11). The optimal temporal scale for an early detection of stress effects on vegetation is at the sub-daily timescale where short-term responses of plants to meteorological drivers can be monitored (e.g., mid-day stomatal closure, or day-to-day downregulation of photosynthetic capacities).

Alternatively, an existing fleet of satellites have the capability in studying plant functioning on a sub-daily basis (Xiao *et al.*, 2021). To date, a few of them sample the Earth reflectance in the far-red region, which is suitable for measuring SIF. These are: Orbiting Carbon Observatory-3 (OCO-3) (low-Earth orbit), Tropospheric Emission Monitoring of Pollution (TEMPO) (geostationary) and Sentinel-4 (geostationary) (note that GeoCarb was supposed to be launched in 2024 but was cancelled by the NASA in November 2022). OCO-3 is currently onboard ISS and has been providing SIF measurements during the day above the tropics (i.e., between  $52^\circ\text{N}$ - $52^\circ\text{S}$ ) since August 2019 (Taylor *et al.*, 2020). OCO-3 SIF strongly correlates with OCO-2 SIF, and therefore with tower-based GPP and SIF across multiple PFT and climate zones (Li *et al.*, 2018a). However, OCO-3 observations are not continuous for a given location (Taylor *et al.*, 2020). TEMPO has been originally designed for air pollution monitoring over North America and was launched in early 2022 (Zoogman *et al.*, 2017). Finally, ESA Sentinel-4 (embarked upon Meteosat Third Generation-Sounder (MTG-S)) is a pay-

load with an ultra-violet visible near infrared spectrometer providing hourly sampling over Europe (Quesada-Ruiz *et al.*, 2020) to be launched in September 2024. As being very new geostationary satellite missions, most of exploring results are not yet published. They however hold great hopes for understanding the diurnal cycling of plant functioning and ecosystem processes (Li *et al.*, 2023b; Zoogman *et al.*, 2017). These data will open up new perspectives for unravelling the responses of vegetation under edaphic droughts or heatwaves from satellite measurements. The MLR model is going to be play a key role for interpreting these data in the near future.



**Figure 7.11:** Spatial and temporal resolution of SIF measurement systems. The colorbar ranges from dark grey to light green representing high to low footprint heterogeneity. The optimal resolution range for capturing short-term responses of vegetation to stress is in the green area. The size of circles is indicative. The scale of the axes is not proportional. OCO-3 sampling at a particular geolocation varies across all hours of the day at absolute latitudes less than  $\sim 52^\circ$ . Spatial resolutions and repeat cycles are taken from Mohammed *et al.* (2019) and Doughty *et al.* (2022).

### 3.8.3. Alternative RS products : monitoring vegetation temperature and heat dissipation

Within the infrared spectral regions, some other geostationary and low-Earth orbit satellites also have the capability for measuring the energy emitted by the Earth surface by sampling the 3-14  $\mu\text{m}$  range (also known as Thermal InfraRed (TIR)). All bodies at a temperature above 273 K emit radiation which is related to their temperature (Liang & Wang, 2020). As a result, RS of TIR can be used to retrieve the Land Surface Temperature (LST), which is an important variable within the Earth’s climate system

and an indicator of the energy and water exchange between the atmosphere and the land surface (Li *et al.*, 2023c). LST can be used to determine drought indexes such as the vegetation health index, which is very efficient for detecting vegetation droughts (Zeng *et al.*, 2023). Numerous applications in agricultural sciences also exist, such as monitoring of plant diseases and crop yields (Khanal *et al.*, 2017). The ECOSystem Spaceborne Thermal Radiometer Experiment on Space Station (ECOSTRESS) satellite provides LST with a pixel size 70 m several times a day. These measurements provide key information for farmers and forest managers such as a hourly monitoring of vegetation yellowing due to drought at an unprecedented precision, both in terms of spatial and temporal resolution (Xiao *et al.*, 2021). ECOSTRESS also provides estimates of evapotranspiration from energy budget-based models using LST (Kalma *et al.*, 2008). Recent studies have also applied a machine learning algorithm trained on meteorological variables and land cover to predict GPP from ECOSTRESS LST (Li *et al.*, 2021b), or to reconstruct global SIF datasets at a meter-range spatial resolution (Zhang *et al.*, 2023a). Nonetheless, SIF remains a better proxy of GPP as being directly related to carbon assimilation processes.

SIF is not the only hyperspectral variable to be strongly linked to short-term variations of GPP. In particular, the energy dissipated by heat through mostly the xanthophyll cycle at hourly scales is an important element in the energy dissipation pathways. When plants experience very high sun irradiance, photochemical quenching solely is not sufficient for dissipating the incoming energy. This excessive amount is dissipated mostly by the de-epoxidation of violaxanthin into zeaxanthin via antheraxanthin in the xanthophyll cycle (Demmig-Adams & Adams III, 2006), which is included into what is referred to as NPQ. The rapid response of the xanthophyll cycle allows photochemistry to keep operating at very high PAR flux densities. For nearly three decades, it is known that the activity of the xanthophyll cycle can be monitored by the Photochemical Reflectance Index (PRI), which is derived from reflectance at 531 nm (reduced by deepoxidation of xanthophyll pigments) and a reference wavelength (typically 570 nm) (Penuelas *et al.*, 1995). PRI is strongly correlated to vegetation LUE and GPP at the sub-daily timescale (Garbulsky *et al.*, 2011), and can capture a wide range of vegetation responses to environmental conditions (Yang *et al.*, 2020a; Mulero *et al.*, 2023; Zhang *et al.*, 2017; Magney *et al.*, 2016). However, interpreting PRI measurements at the seasonal scale might be complicated by its sensitivity to canopy structure or carotene content (Garbulsky *et al.*, 2011). Yet, PRI is very effective for capturing sub-daily temporal dynamics of NPQ which reveals the control of abiotic stresses (e.g., soil moisture, or high irradiance) on GPP. As the response of SIF to water stress or heatwaves still remains unclear (see chapter 6), this specific characteristic of PRI has been exploited to improve empirical SIF-GPP models specifically under water stress (Ma *et al.*, 2022; Wang *et al.*, 2020a; Schickling *et al.*, 2016). The strong correlation between PRI and  $V_{\text{cmax}}$  has also been highlighted (Chou *et al.*, 2020; Yu *et al.*, 2020b) which opens up new perspectives for mapping this parameter at large scales

or for assimilating PRI in the FvCB model. Current RS products provide daily PRI from Moderate Resolution Imaging Spectroradiometer (MODIS) (Gamon *et al.*, 2016; Vanikiotis *et al.*, 2021) or Environmental Mapping and Analysis Program (EnMAP) (Dotzler *et al.*, 2015) with a spatial resolution of 1 km and 30 m respectively. Importantly, what is currently lacking is a mechanistic modeling framework for using both PRI and SIF to estimate GPP. While we showed that SIF was likely responsive to dry soil conditions (chapter 6), future studies should explore the possibility to use PRI as a stress indicator, notably in the MLR model.

### **3.9. Assessing vegetation water status with RS of microwave observations**

Besides LWP, Vegetation Water Content (VWC) is also an important physiological indicators to assess plant water conditions. VWC can be measured either by destructive methods (e.g., by measuring the weight difference of the sample before and after drying in a heat chamber; Zhou *et al.* (2021)) or by non-destructive methods (e.g., terahertz radiation (Li *et al.*, 2020), electrical impedance spectroscopy (Serrano-Finetti *et al.*, 2023), hyperspectral spectroscopy (Junttila *et al.*, 2022), or microwaves (Calla *et al.*, 2008)). Destructive methods are time-consuming, lead to perturbations of the vegetation as the leaf must be removed from the plant before being analyzed, and cannot provide an estimation of VWC of large patches of vegetation. Such approaches do not allow to track short-term variations of VWC at diurnal scales.

The main advantage of VWC is its ability for being monitored by RS techniques. In particular, microwaves strongly interact with the dielectric properties of soil and vegetation (Jackson, 1993; Jackson & Schmugge, 1991). When considering remote sensing of soil and vegetation moisture, two types of sensors coexist. Radiometers measure the intensity of the natural thermal emission of a land surface (i.e., the emissivity - passive method) and have a coarse spatial resolution that increases with frequency. Typical radiometers spatial resolution is in the 10-100 km range. Radars measure the backscattering coefficient of a land surface from the power ratio between the emitted and returned electromagnetic wave (i.e., the reflectivity - active method). Synthetic aperture radars (SAR) use a radar technique that can provide finer spatial resolution (in the meter range). Such instruments only perform measurements during a very small fraction of their orbit around Earth because of the high energy demand of the SAR technique (Grasso *et al.*, 2021). Flying SAR instruments in tandem overcomes this issue (e.g., SAR onboard Sentinel-1). Dozens of passive and active microwave satellites have been operating over the past decades, with very different characteristics of band frequency, repeat cycle and spatial resolution (Mavrovic *et al.*, 2023). The most recent satellite microwave missions are Soil Moisture and Ocean Salinity (SMOS) and Soil Moisture Active Passive (SMAP) (passive only, the radar onboard SMAP experienced failure 208 days after launching), as well as Advanced SCATterometer (ASCAT) and



SAR-Sentinel (radar) (Mazzariello *et al.*, 2023).

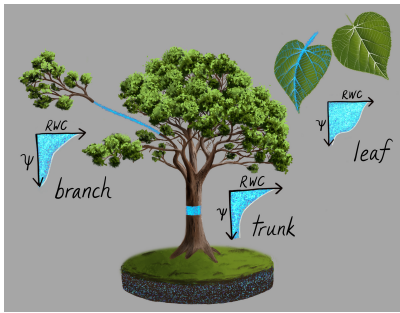
The penetration of microwaves within the land surface depends on their band frequency (the longer the microwave, the deeper the penetration). Therefore, microwave measurements include the signal of different components such as vegetation, soil, snow and atmosphere (Kerr *et al.*, 2012). Disentangling the complexity of this raw signal requires the use of radiative transfer models which calculate scattering, reflection and attenuation of the electromagnetic waves of the different components of the surface. The degree of attenuation of the microwave by the vegetation (i.e., the Vegetation Optical Depth (VOD)) is related to various attenuation and scattering factors such as vegetation structure, water content or sensor wavelength (Jagdhuber *et al.*, 2019).

VOD was originally considered as a noise in the microwave signal, but it recently gained interest as a promising ecological indicator linked to plant hydraulics (Konings *et al.*, 2019). Spaceborne passive and active microwave sensors have been used to provide VOD at large scales since the 1970s (Owe *et al.*, 2008; Frappart *et al.*, 2020). Recent studies have derived long-term VOD products from multiple sensors in the Ku-band (19 GHz), X-band (10.7 GHz) and C-band (6.9 GHz) (Moesinger *et al.*, 2020; Hu *et al.*, 2023; Liu *et al.*, 2023). L-band VOD from SMAP and SMOS is expected to be more correlated with dense ecosystem physiology than X-band or C-band VOD (Chaparro *et al.*, 2019).

Early *in situ* measurements on soybean have highlighted a strong linearity between VOD and VWC (Jackson & Schmugge, 1991), later confirmed on many species at the diurnal scale (e.g., Xu *et al.* (2021b); van Emmerik *et al.* (2015)). However, the co-variation of VOD with AboveGround Biomass (AGB) and Vegetation Relative Water Content (VRWC) complicates the interpretability of microwave measurements at large temporal and/or spatial scales. For instance, Tian *et al.* (2017) showed that seasonal variations in VOD were influenced by changes in forest phenology and vegetation water status. Therefore, the link between VOD and vegetation water status is strongly influenced by biomass changes (Konings *et al.*, 2019). However, the role of AGB in explaining VOD dynamics can be reduced at temporal scales when AGB variations are negligible. Many studies have used this approach to derive vegetation water status from VOD retrieved from tower-based or satellite microwave measurements (Momen *et al.*, 2017; Konings & Gentine, 2017; van Emmerik *et al.*, 2017). The reverse route (isolating the effect of vegetation water status for assessing biomass changes) is however more complicated (Konings *et al.*, 2019).

Notwithstanding the key recent advances in monitoring VOD, several limitations remain. First, the validation of RS VOD data has not yet been achieved as ground measurements are very limited. In this perspective, the use of affordable microwave sensors on the ground could be very helpful for validating satellite-based VOD products. For instance, Humphrey & Frankenberg (2023) and Yao *et al.* (2024) have shown that Global Navigation Satellite Systems (GNSS) can be used at the ecosystem scale

to capture sub-daily dynamics of VOD and canopy water content of forests. Such validation campaigns are required for assessing the accuracy of RS VOD products. In addition, the temporal and spatial frequency of microwave satellite measurements is one of main constraint on providing a VOD product that can be linked to plant physiology and to gas exchanges. For instance, satellites in sun-synchronous orbits such as SMOS provide a measurement at 6 AM and 6 PM (local time) every 3 days (Al-Yaari *et al.*, 2017). Assessing ecosystem responses to drought requires joint measurements of SIF and VOD at a sub-daily resolution to capture the dynamics of photosynthesis limiting factors such as stomatal closure through changes in LWP (Zhang *et al.*, 2019b). Monitoring such processes at a high frequency with radar geostationary satellites potentially has a direct implication for forecasting crop yields and forest fire risks (Nolan *et al.*, 2020; Konings *et al.*, 2021; Jackson *et al.*, 2004).



**Figure 7.12:** As VOD scales with VWC, retrieving plant water potential from VOD requires a vertical distributions of tissue-specific water retention properties through VRWC curves.  $\psi$  is the vegetation water potential within each tree structural units (Konings *et al.*, 2021).

Finally, linking VOD to trends in vegetation water status is not straightforward. In particular, the relationship between VOD and VWC actually varies with the time of day, the location of the emission source within the canopy (e.g., leaves, trunk or branches) and the water status of the vegetation. Most existing modeling approaches for retrieving VOD (e.g., zeroth-order radiative transfer model or water cloud models for radiometer and radar-based measurements, (Jagdhuber *et al.*, 2019)) represents the vegetation as a single water reservoir without disentangling the different structural units of the canopy. Future retrieval algorithms including a more mechanistic representation of microwave dynamics in vegetation with a description of ecosystem structure would help in estimating the VWC at different heights within the canopy (Konings *et al.*, 2021; Steele-Dunne *et al.*, 2017). This would be particularly helpful

for describing the hydraulic constraints along the SPAC and deciphering the driving forces of the water flow under water stress (Konings *et al.*, 2021). Characterizing Vegetation Water Retention Curves (VWRC) which link VWC to plant water potential throughout the canopy is also needed (Fig. 7.12).

Importantly, the combination of SIF and VOD at the sub-daily timescale provides a window on the health of terrestrial ecosystems by combining information on photosynthetic activity and VWC, which are two major drivers of the carbon and water cycles. The MLR-USO model is placed at the core of this thematic as a key tool for linking plant hydraulics and carbon assimilation. By being a proxy of VWC and LWP, VOD can capture the diel cycle of stomatal conductance which directly influences the CO<sub>2</sub>

availability to chloroplasts and the transpiration flux. Future studies should focus on adapting the MLR-USO model for assimilating SIF and VOD products at a sub-daily timescale.

## 4. General conclusion

The outcomes of this thesis show that, more than ever, delving into the complexity of the intertwined processes ruling ecosystem responses to a changing climate requires the coordination of different research teams involved in monitoring the soil and plant water status, performing flux analysis, characterizing root system, and collecting proximal/remote sensing measurements of hyperspectral variables. The development of these areas of research should be structured around EC stations, which provide the gold benchmark for greenhouse gases and energy fluxes measurements at the canopy level. We encourage the creation of a data portal where proximal sensing measurements are made available for researchers similarly to the ICOS network. This perspective is currently explored in the United States (Pierrat *et al.*, 2023) while China has moved forward with the creation in 2021 of the ChinaSpec network regrouping 26 sites (Zhang *et al.*, 2021). We also advocate for the constitution of complementary research teams and an improved cooperation between research centers. With several EC sites spanning across different PFTs, the BIODYNE department has a key role to play. Future research projects or equipment funds should be focused on measuring root biomass density, acquiring field spectrometers, microwave sensors or microtensiometers for durable, permanent installation at measurement sites. Future PhDs could be orientated towards these research topics. Considerable, but no less exciting challenges lie ahead.

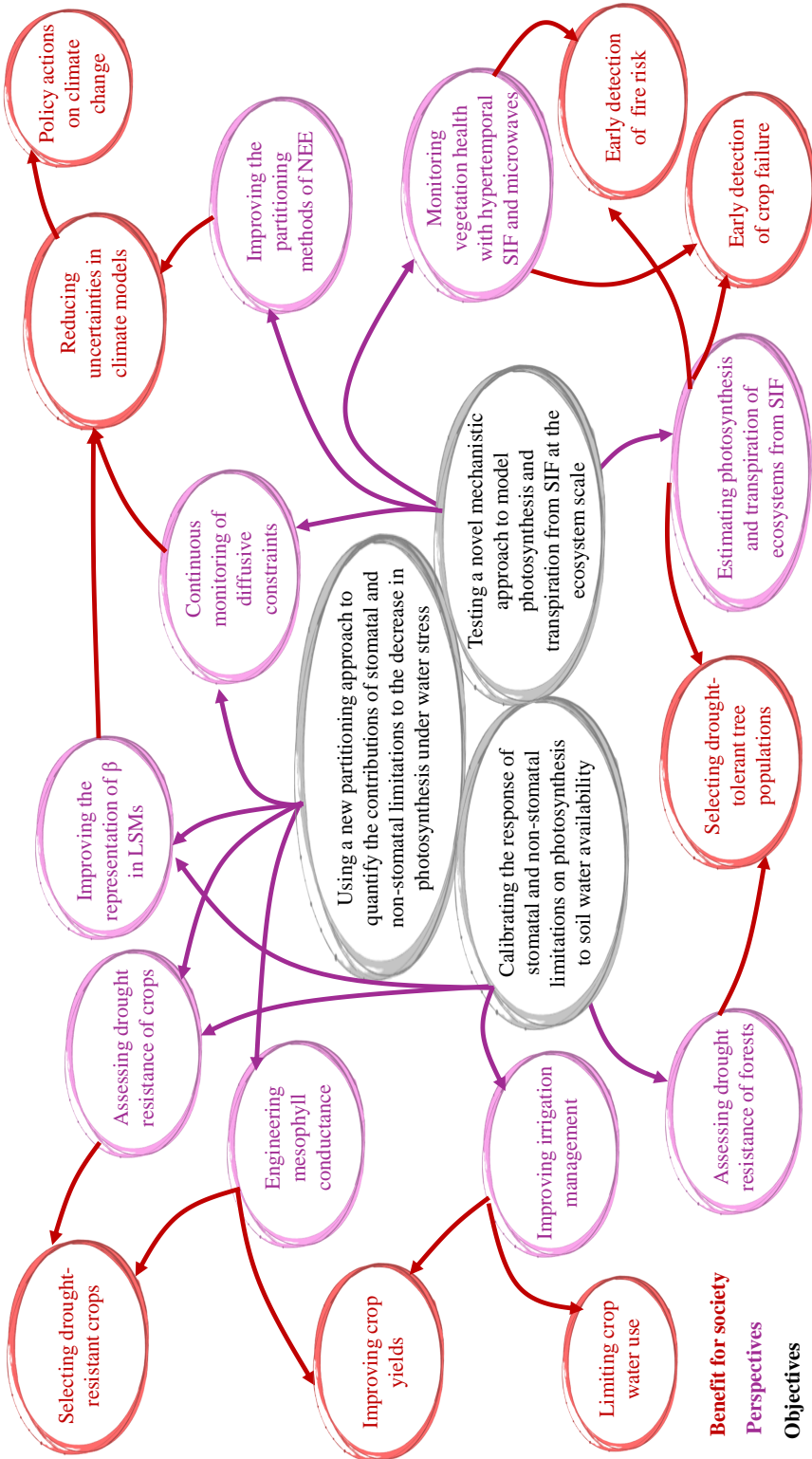


Figure 7.13: Perspectives of the thesis

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