

Université de Liège
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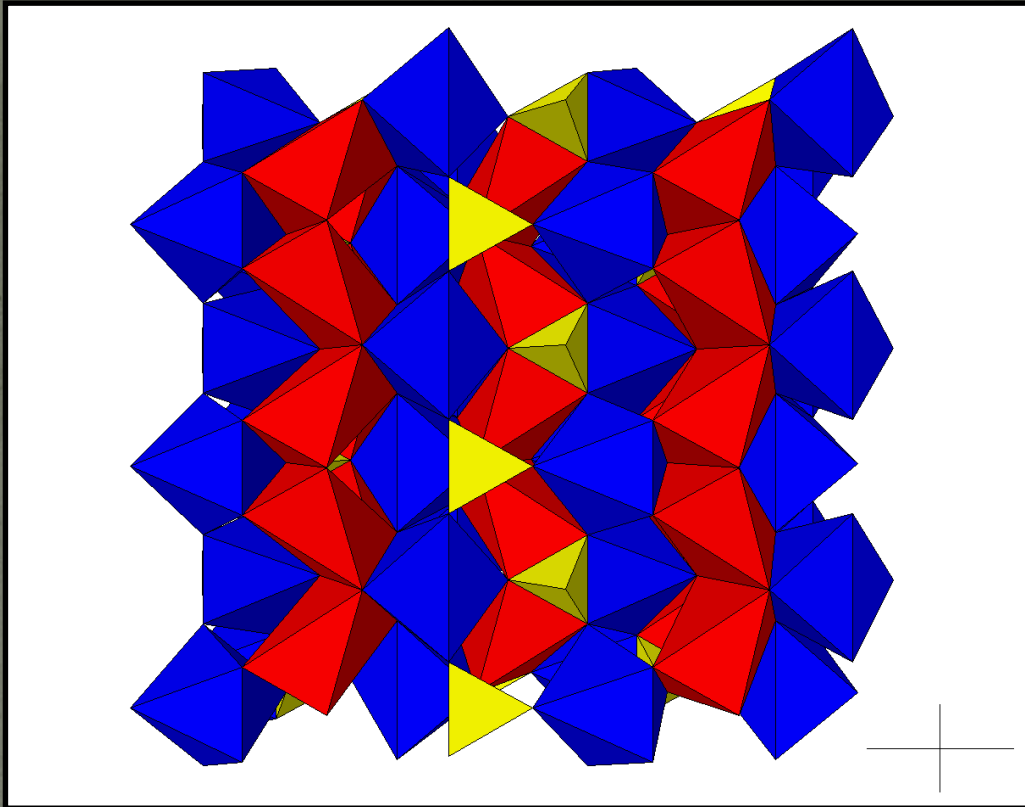


The stability of Si-rich triphylite in granitic pegmatites : an experimental investigation of the $\text{Li}(\text{Fe},\text{Mn})(\text{PO}_4)\text{-}(\text{Fe},\text{Mn})_2(\text{SiO}_4)$ olivine-type solid solution

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IMA-2018, Melbourne, August 15th, 2018

The olivine structure



Red octahedra: M1 (Li, Na, Fe, Mn, Mg)
Blue octahedra: M2 (Fe, Mn, Mg)
Yellow tetrahedra: (Si, P)

Phosphates

- Triphylite, $\text{LiFe}^{2+}(\text{PO}_4)$
- Lithiophilite, $\text{LiMn}(\text{PO}_4)$
- Natrophilite, $\text{NaMn}(\text{PO}_4)$
- Karenwebberite, $\text{NaFe}^{2+}(\text{PO}_4)$

Silicates

- Fayalite, $\text{Fe}^{2+}_2(\text{SiO}_4)$
- Forsterite, $\text{Mg}_2(\text{SiO}_4)$
- Tephroite, $\text{Mn}^{2+}_2(\text{SiO}_4)$

S.G. $Pmnb$

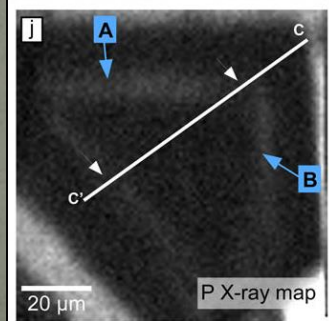
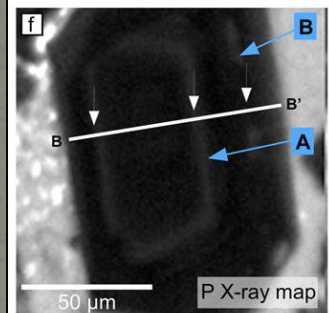
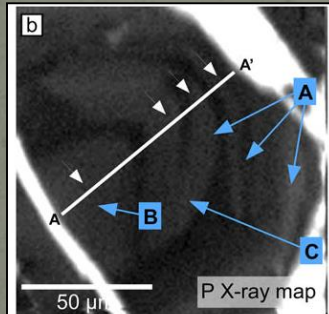
$a = 6.092 \text{ \AA}$

$b = 10.429 \text{ \AA}$

$c = 4.738 \text{ \AA}$

Phosphorous in olivine-type silicates

Semarkona LL chondrite

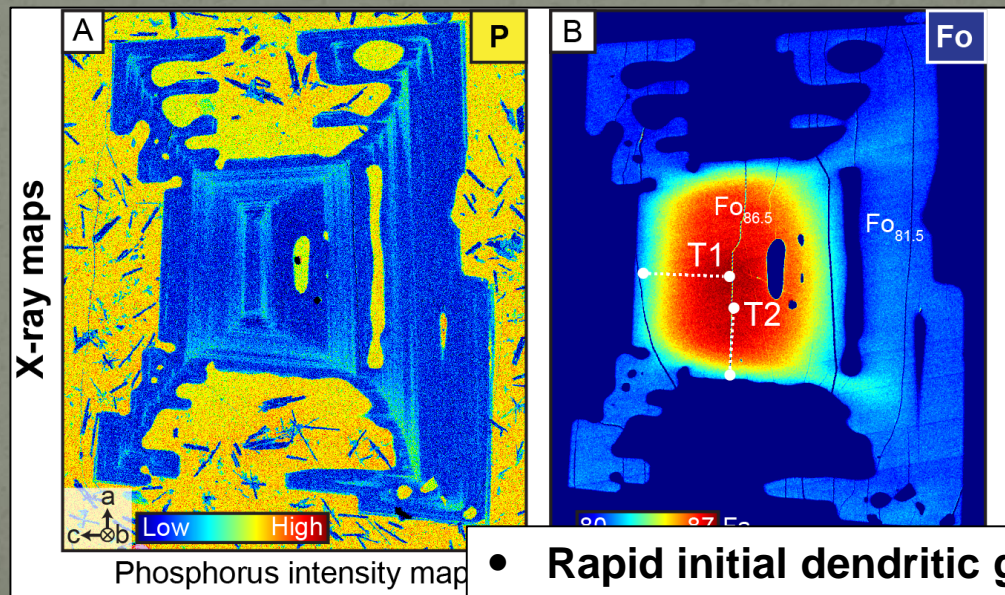


McCanta *et al.* (2016)

P₂O₅ in olivines

- Generally below 1 wt. %
- In meteorites: up to 5 wt. %

Kilauea volcano, Hawaii



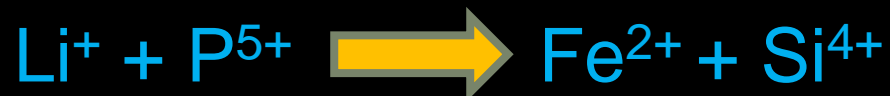
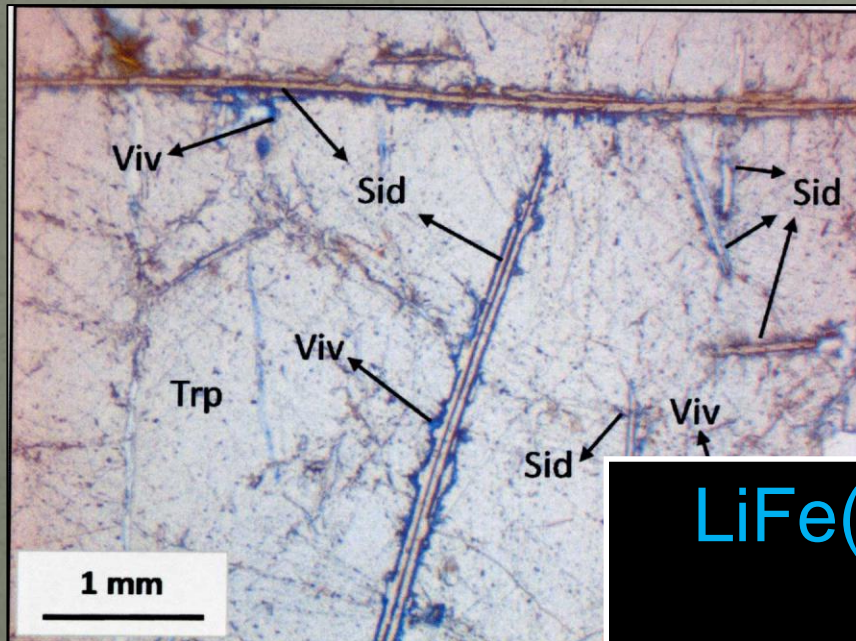
Shea *et al.* (2015)

- Rapid initial dendritic growth
- Slower P diffusion

Silicon in olivine-type phosphates

SiO₂ in triphylite

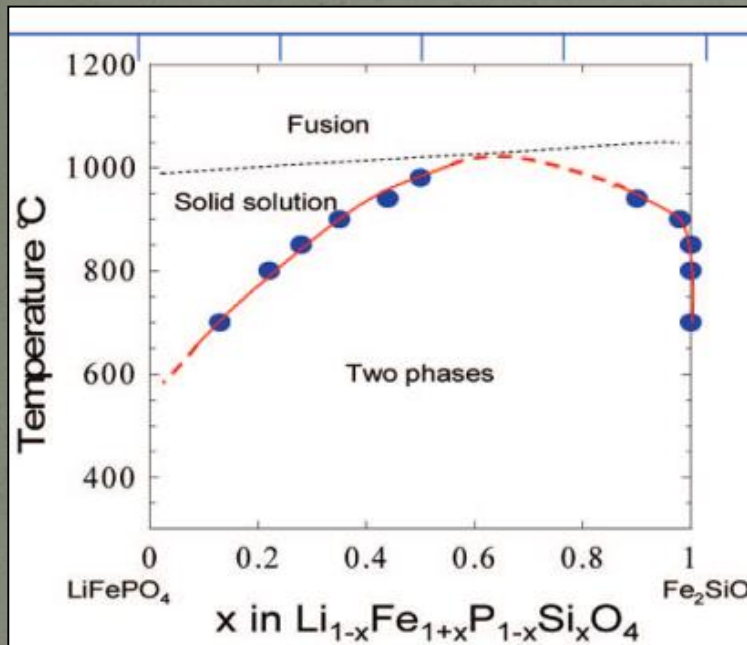
- Generally below 1 wt. %
- Pøibyslavice granite: 1 to 17 wt. % (Škoda *et al.*, 2013)



Baijot (2015)

Previous experiments

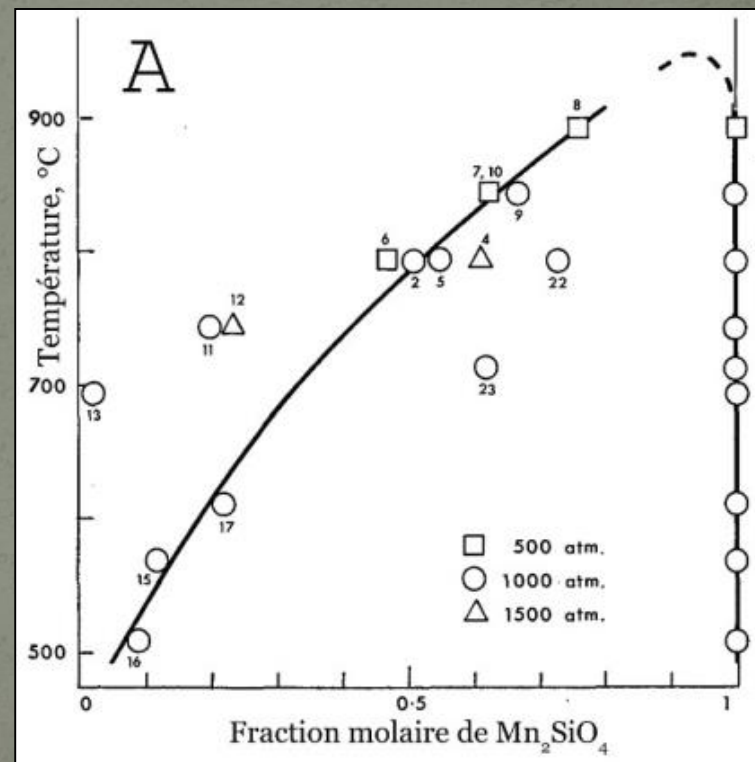
LiFe(PO₄) – Fe₂(SiO₄) system



Recham *et al.* (2008)

- Solid-state experiments
- Asymmetric miscibility gap
- No P in olivine below 900°C
- No Si in triphylite below 500°C

LiMn(PO₄) – Mn₂(SiO₄) system



Bradley *et al.* (1966)

New hydrothermal experiments

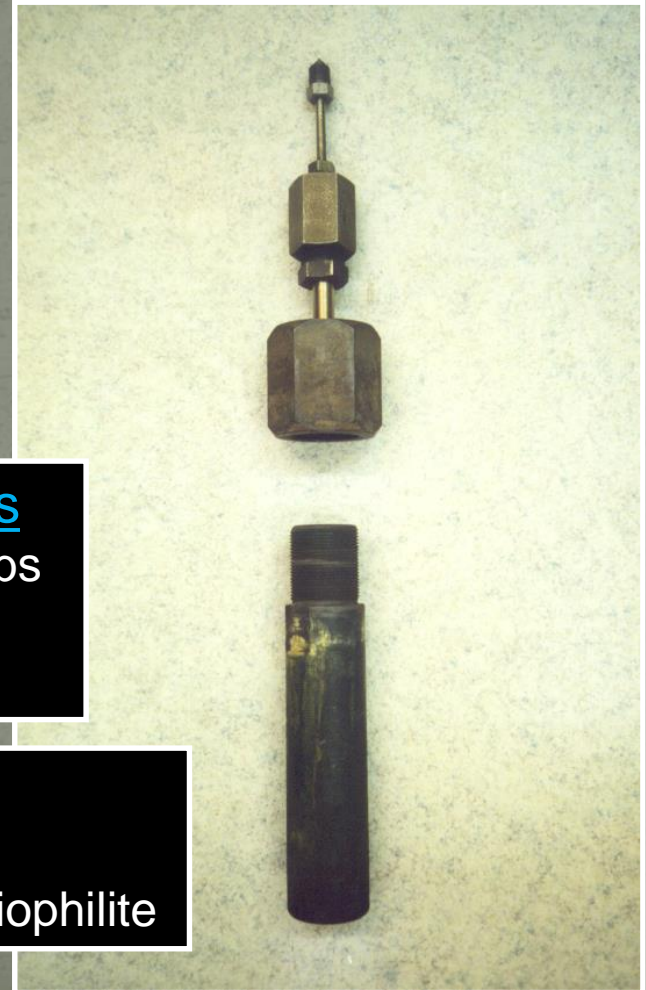


Hydrothermal experiments

- Tuttle-type cold-seal bombs
- $T = 400-700^{\circ}\text{C}$
- $P = 1 \text{ kbar}$

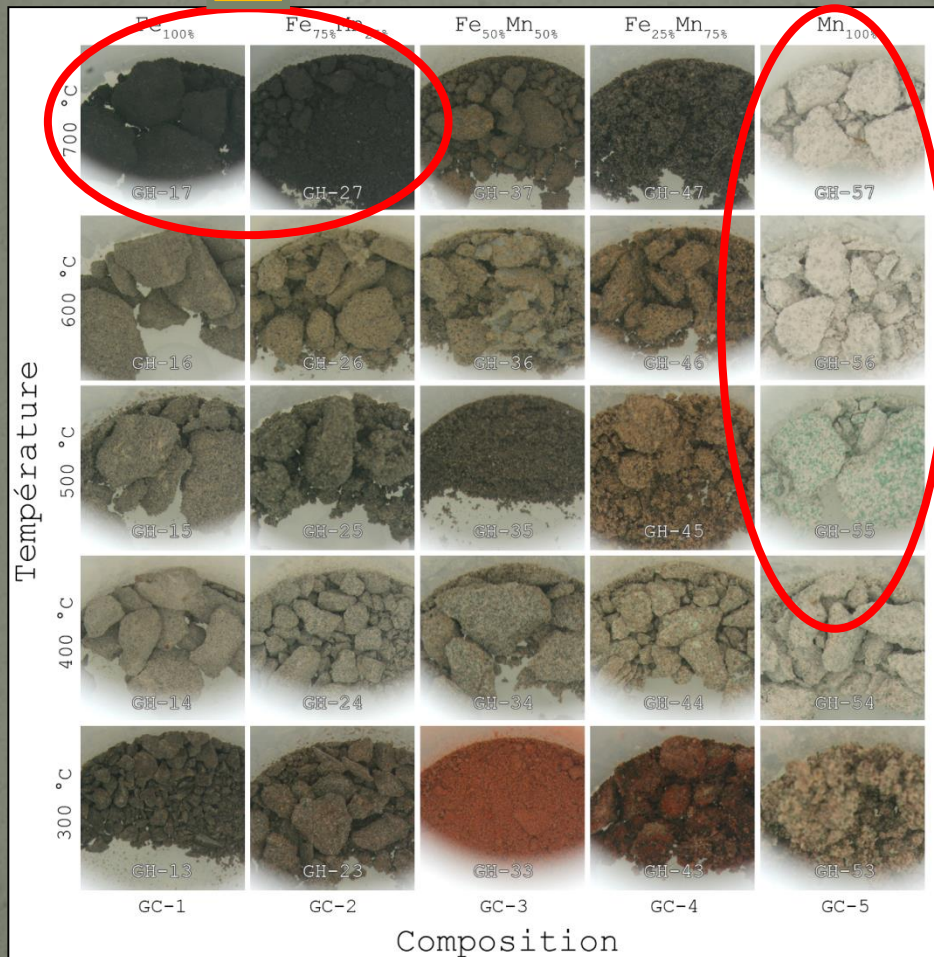
Starting material

- Mixture of chemicals
- 50 % fayalite/tephroite and 50 % triphylite/lithiophilite



Synthesized compounds

Rich in magnetite/jacobsite



Rich in olivine-type compounds

Synthesized compounds

- Fayalite/tephroite (1-71 %)
- Triphylite/lithiophilite (17-55 %)
- Quartz (< 16 %)
- Magnetite/jacobsite (< 36 %)
- FeO/MnO (< 17 %)
- Nambulite (< 22 %)

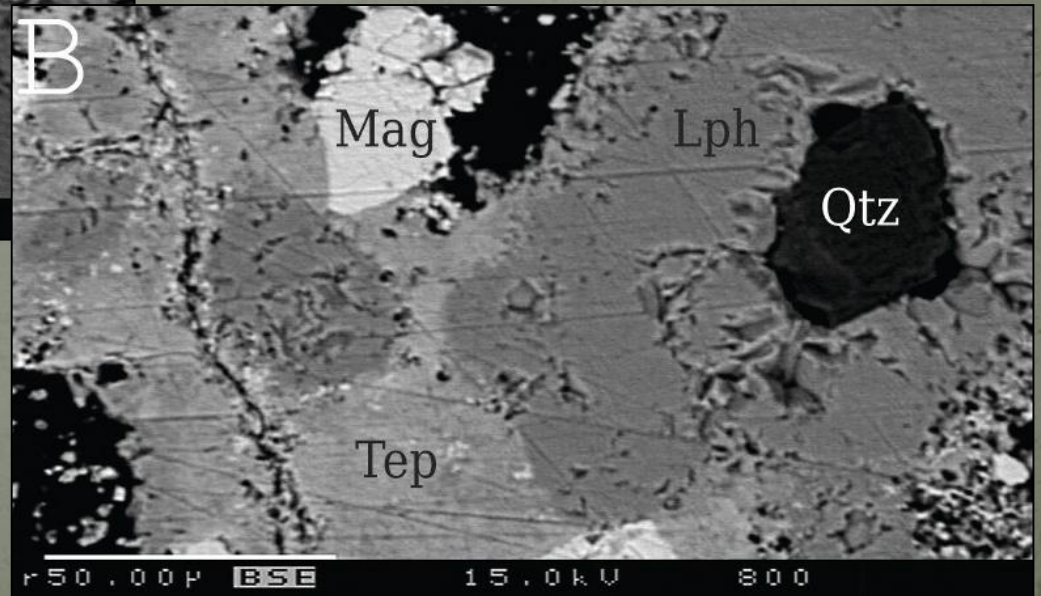
Synthesized phases

Idiomorphic lithiophilite crystals



Sample GH-37

Lithiophilite + tephroite assemblage



Sample GH-37

Nambulite, $\text{LiMn}_4\text{Si}_5\text{O}_{14}(\text{OH})$

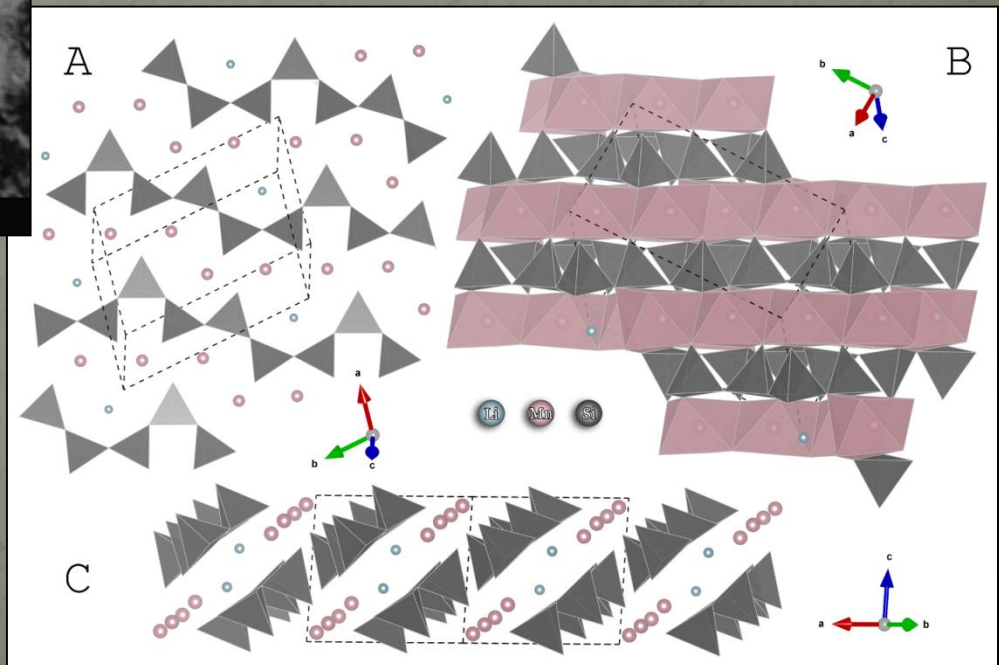
Nambulite crystals



Sample GH-54

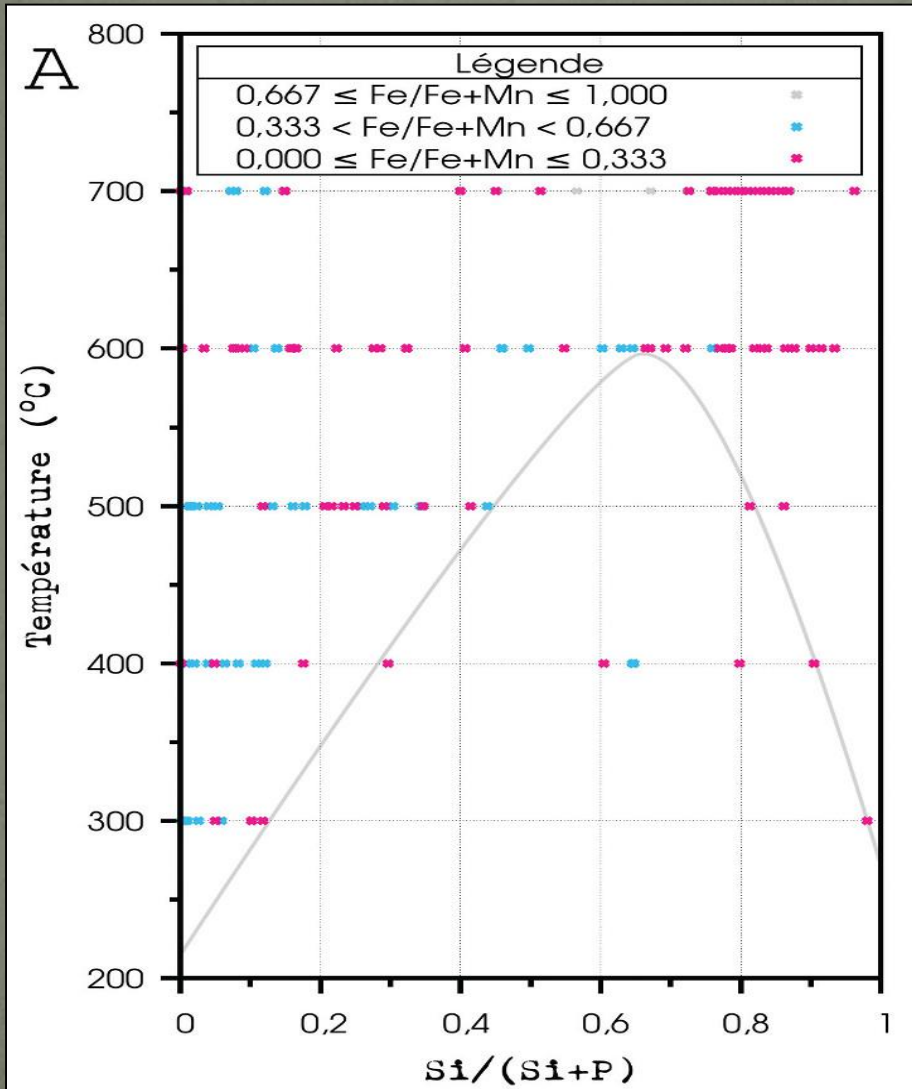
- Alteration of rhodonite
- Pyroxenoid structure
- MnO_6 layers and SiO_4 layers

Nambulite structure



Narita *et al.* (1975)

Phase diagrams

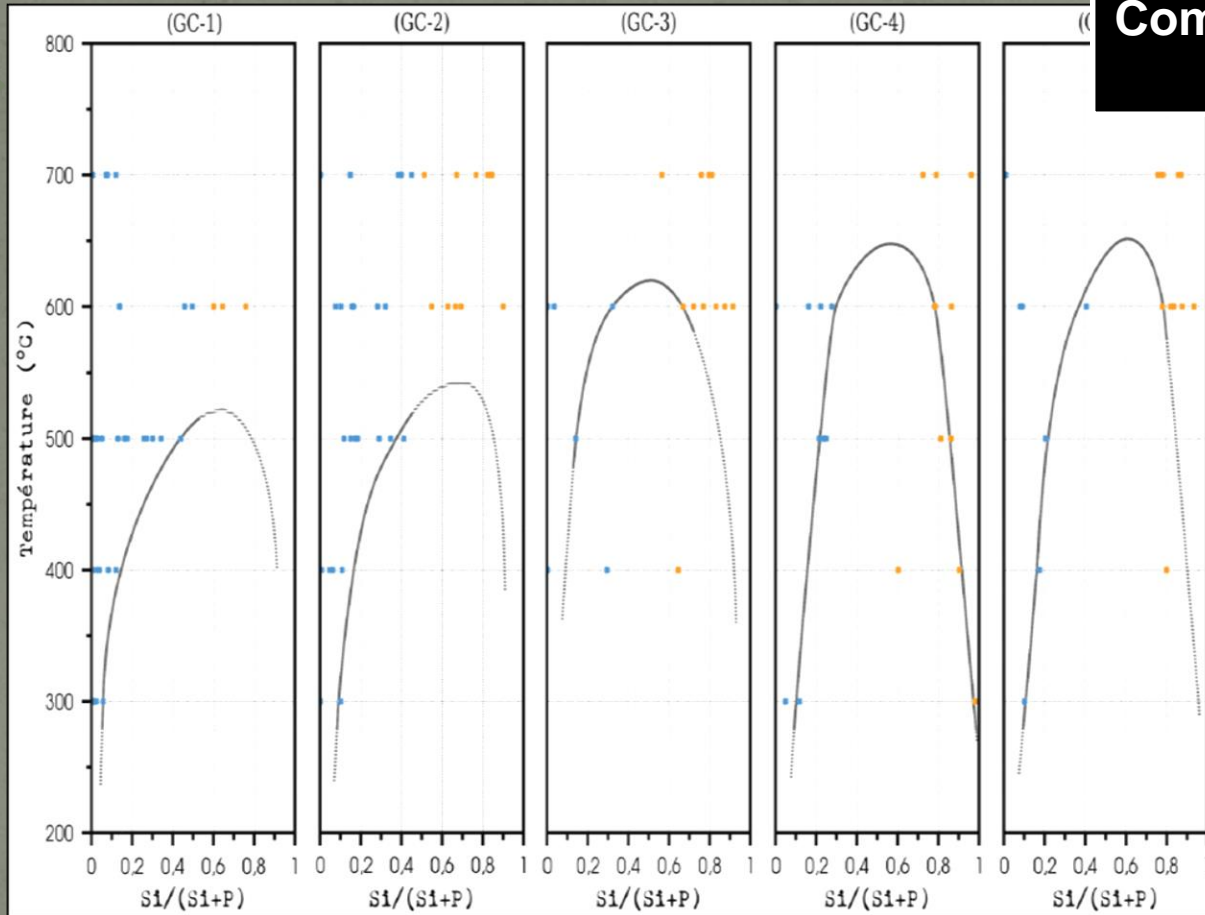


- Asymmetric miscibility gap
- No P in olivine below 300°C
- No Si in triphylite below 200°C
- Si-in-triphylite geothermometer

Phase diagrams

Complete miscibility at 550°C

Complete miscibility at 650°C

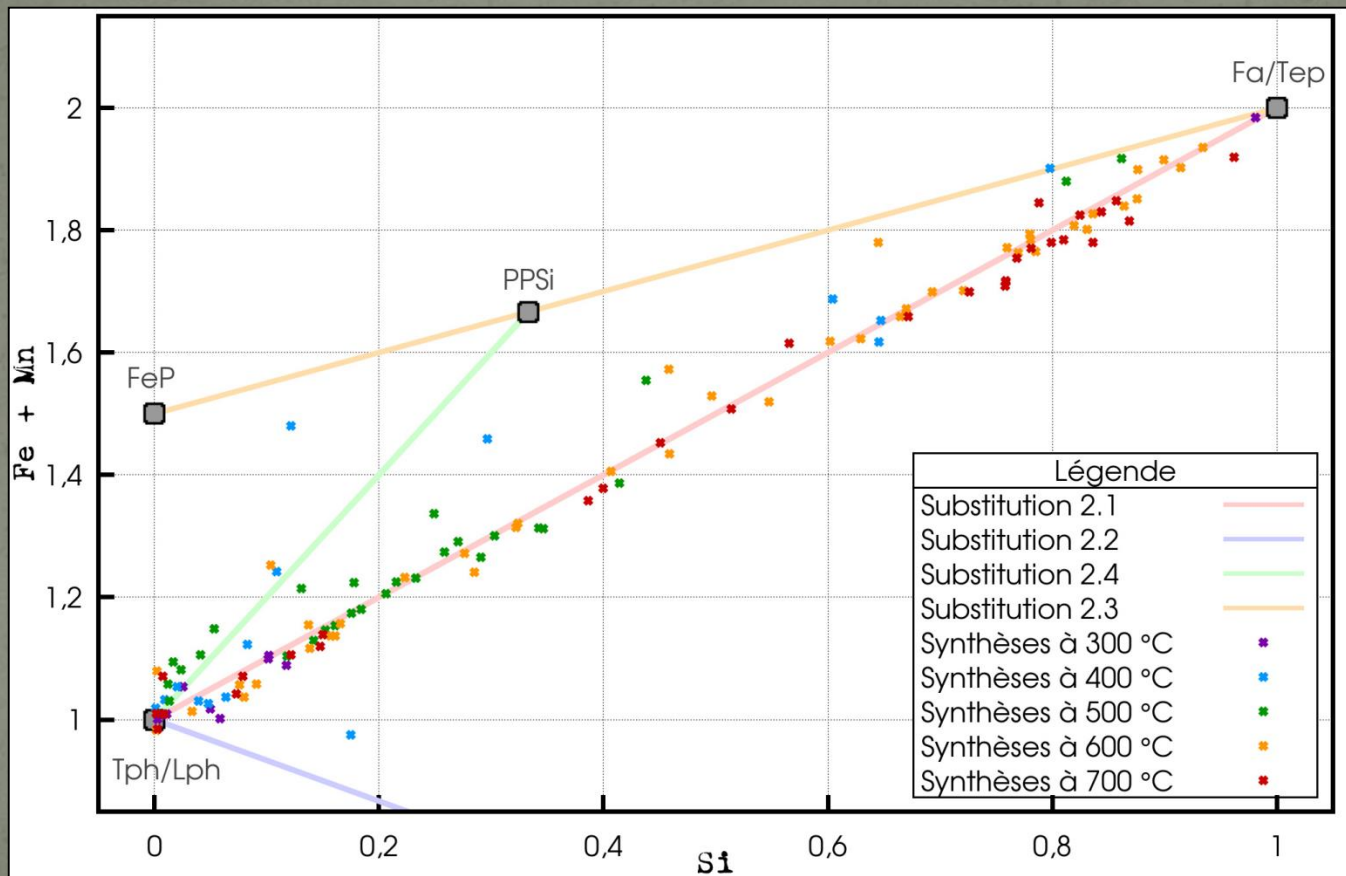
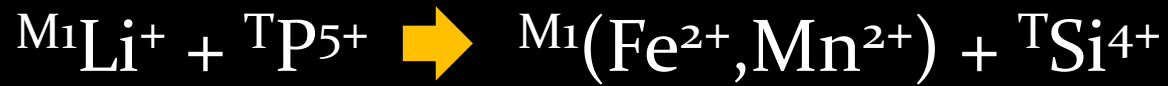


Mn-rich



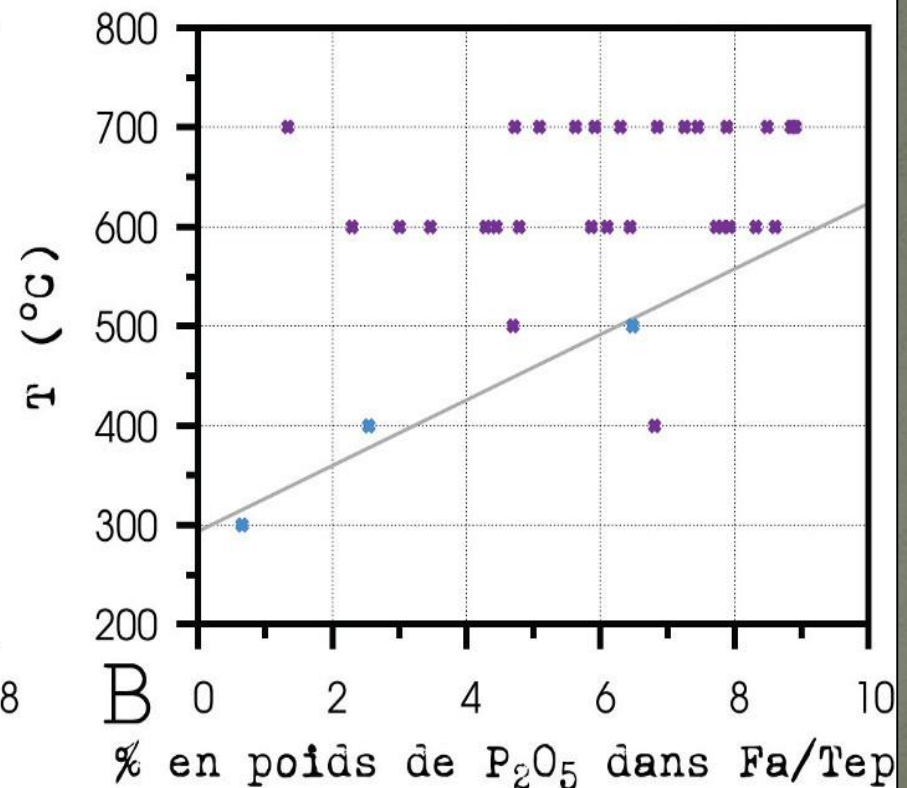
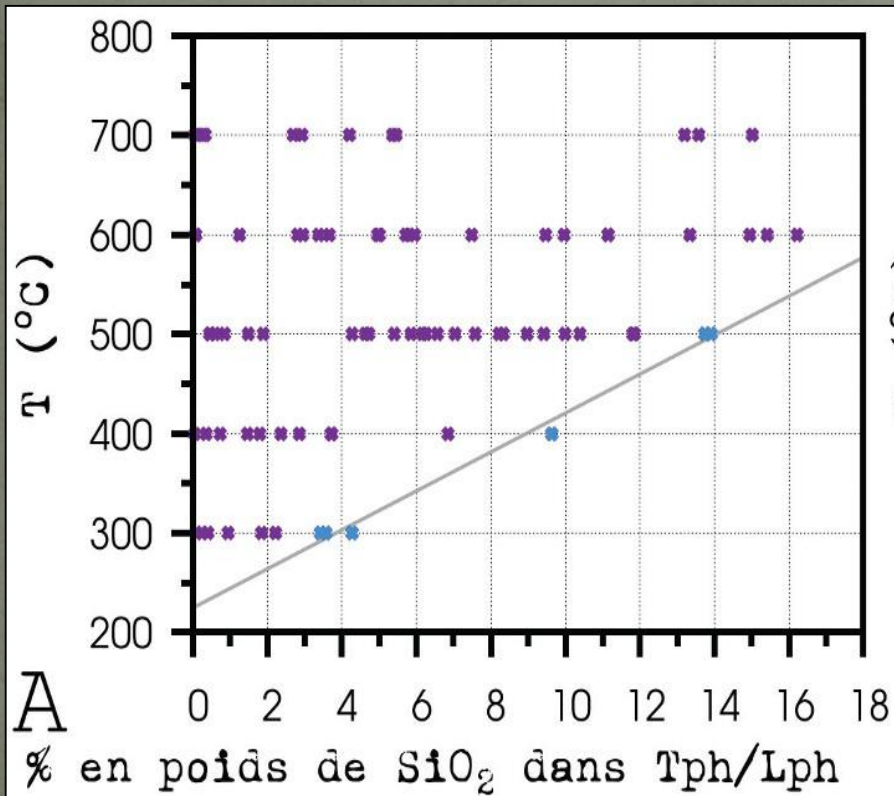
Fe-rich

Substitution mechanism



Maximal solubilities

- $T_{\text{phosphate}} = 19.6 \cdot [\text{wt. \% SiO}_2] + 224.8$
- $T_{\text{silicate}} = 33.0 \cdot [\text{wt. \% P}_2\text{O}_5] + 293.7$



Geothermometric applications



Si-bearing pegmatitic triphylites

• Brissago pegmatite, Switzerland:	0.05 to 0.13 wt. % SiO_2	218 – 236 °C
• Newport pegmatite, NH, USA:	0.33 wt. % SiO_2	223 - 240 °C
• Sapucaia pegmatite, Brazil:	0.02 to 0.77 wt. % SiO_2	217 – 249 °C
• Los Aleros pegmatite, Argentina:	0.78 wt. % SiO_2	231 – 249 °C
• Kolmozero pegmatite, Kola, Russia:	0.90 wt. % SiO_2	233 – 251 °C
• Tsaobismund pegmatite, Namibia:	0.12 to 1.0 wt. % SiO_2	219 – 254 °C

Si-bearing granitic triphylite

• Pøibyslavice, Czech Republic:	1.1 to 17 wt. % SiO_2	237 – 582 °C
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Very low temperatures → Maximal solubilities are not reached

Phosphates = « closed system », not in equilibrium with silicates

↓
Immiscibility phosphates/silicates?

Conclusions

- The triphylite + fayalite assemblage has been obtained experimentally
- Miscibility gap below 550 (Mn-rich compositions) or 650 °C (Fe-rich compositions)
- Si-in-triphylite geothermometer gives too low temperatures, since equilibrium is not reached with enclosing silicates.

Thank you for your attention!