1	Exchange between drainage systems and surface flows during urban flooding: Quasi-
2	steady and dynamic modelling in unsteady flow conditions
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19 Abstract

20 The accurate modelling of urban flooding constitutes an integral part of flood risk assessment and management in residential and industrial areas. Interactions between drainage networks and 21 22 surface runoff flows are commonly modelled based on weir/orifice equations; however, this 23 approach has not been satisfactorily validated in unsteady flow conditions due to uncertainties in estimating the discharge coefficients and associated head losses. This study utilises experimental 24 data of flow exchange between the sewer flow and the floodplain through a manhole without a lid 25 to develop two alternate approaches that simulate this interaction and describe the associated 26 27 exchange flow. A quasi-steady model links the exchange flow to the total head in the sewer pipe 28 and the head losses in the sewer and the manhole, whilst a dynamic model takes also into account 29 the evolution of the water level within the manhole at discrete time steps. The developed numerical models are subsequently validated against large-scale experimental data for unsteady sewer flow 30 31 conditions, featuring variable exchange to the surface. Results confirmed that both models can 32 accurately replicate experimental conditions, with improved performance when compared to existing methodologies based only on weir or orifice equations. 33

34 Keywords

35 Sewer/Surface flow interactions; Urban flood modelling; Drainage systems; Head losses;

36 Unsteady flow; Urban hydraulics

38 **1. Introduction**

39 Urban flooding events tend to become more frequent due to the increase of urbanization and changes in rainfall patterns linked with climate change. An accurate quantification of flood risk is 40 important for assessing relative vulnerability under given and predicted rainfall events in order to 41 42 develop cost effective asset investments and flood mitigation approaches. In urban areas, hydrodynamics associated with flood events is particularly complex because such events 43 commonly include interactions between surface flows/runoff and flows within urban drainage 44 networks (Schmitt et al. 2004, Rubinato et al. 2019). The risk of flooding is commonly evaluated 45 using hydraulic modelling tools, which utilize a number of empirical and semi-empirical 46 47 relationships (and associated parameters) to simulate processes such as runoff and 48 frictional/turbulent energy losses, including relationships to describe interactions between surface flows and drainage networks (Djordjevic et al. 2005, Leandro et al. 2009, Seyoum et al. 2012). 49 50 However, despite recent advances in remote sensing and open access data (Moy de Vitry et al. 2017, Moy de Vitry and Leitao 2020), there is a general paucity of high-resolution datasets for 51 flood model validation (Tscheikner-Gratl et al. 2016). Typical datasets consisting of point depth 52 of flow observations during flood events are insufficient to fully overcome parameter non-53 identifiability/equifinality issues in complex flood models, or provide a detailed evaluation of 54 modelling representations for individual model components such as above/below ground flow 55 56 exchange (Beven 2006, Dottori et al. 2013, Arrault et al. 2016).

57 Datasets collected during detailed, controlled laboratory studies can be used to validate 58 hydraulic models (Martins et al. 2017, 2018) and/or provide an improved understanding of 59 physical processes and flood model parameters such as energy loss coefficients (Hare 1983). 60 However, transferring findings from scaled laboratory studies into practice can be challenging. For

61 example, due to its significance in urban flood modelling, a number of experimental studies have investigated common representations of flow exchange between surface flows and urban drainage 62 networks through hydraulic structures such as manholes and gullies. Flow exchange through such 63 structures can be bi-directional (net exchange to the surface or the drainage network), and 64 hydraulic conditions associated with these situations are generally unsteady and highly three 65 66 dimensional (Lopes et al. 2015, Beg et al. 2018). Within network scale urban flood models such interactions are commonly represented by simple weir/orifice equations with net flows given as a 67 function of hydraulic/pressure head difference between surface flow and the drainage network, the 68 69 geometrical properties of the structure and an energy loss term (Nasello and Tucciarelli 2005, Chen et al. 2007, Seyoum et al. 2012). To evaluate this approach some studies have calibrated/validated 70 urban drainage/flood models based on physical models of urban catchments with multiple 71 exchange structures (Bazin et al. 2014, Fraga et al. 2015, Noh et al. 2016, Dong et al. 2021), whilst 72 other work has directly measured flow rates through individual interaction structures, allowing a 73 74 direct comparison between exchange equations and measurements. For example, Rubinato et al. (2017) conducted a detailed experimental study of the weir and orifice equations representation of 75 exchange flows through a scaled open manhole in both drainage and surcharging conditions. In 76 77 steady flow conditions with subcritical surface flows, discharge coefficients were calibrated based on flow, pressure and depth measurements in the pipe network and on the surface and found to be 78 constant over the range of tested flow conditions. However, calculated coefficients were sensitive 79 80 to flow depth/pressure values used within the calibration, which may in practice be calculated with different methods and vary over the longitudinal profile of the hydraulic structure. In addition, 81 82 when calibrated relationships were used to validate a numerical model against a range of unsteady

83 flow events, meaningful differences were observed between predicted and observed exchange84 volumes.

Several other laboratory studies have been conducted to calibrate weir/orifice equations for a 85 range of grate types and steady flow conditions. Martins et al. (2014) focused on drainage flows 86 87 into a gully pot, while Gomez et al. (2019) and Rubinato et al. (2018a) investigated drainage flows though different grate types. Kemper and Schlenkhoff (2019) specifically analyzed supercritical 88 flows over road drainage grates. All these studies have provided a wide range of recommended 89 discharge coefficients, likely due to the sensitivity of energy loss processes to the geometry of 90 91 different structures, but also potentially due to methodological differences in the definition of surface and drainage system hydraulic head (e.g. measurement location), how geometrical 92 93 properties are defined (e.g. calculation of void spaces) and how partially/fully submerged openings of valves influence the flow conditions. Hence the accurate representation of flood inundation 94 95 processes in urban areas may require site specific calibration of discharge coefficients (Dong et al. 2021), which is unfeasible in most practical applications. 96

Based on the studies described above, a number of issues concerning the suitability of 97 weir/orifice type methodologies to describe above/below ground flow interaction can be identified 98 as follows: 1. Outstanding uncertainty regarding the discharge coefficient, which past work has 99 100 shown to differ significantly from common standard values for classical weirs/orifices and thus requires site specific calibration or experimental/numerical studies for accurate identification 101 (Martins et al. 2014, Gomez et al. 2019). 2. Lack of understanding of the hydraulically effective 102 area of a drainage inlet during shallow flows and how this changes with flow depth and/or velocity 103 104 (Martins et al. 2018). 3. The sensitivity of the calculated exchange discharge (and/or discharge coefficient) to hydraulic head (Bazin et al. 2014, Rubinato et al. 2018b), which can vary 105

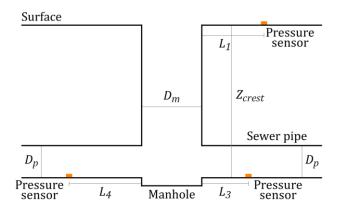
significantly within and across hydraulic structures (Marsalek 1985). 4. A lack of successful
validation of the approach in unsteady flow exchange conditions through individual structures
(Rubinato et al. 2017).

Flows which surcharge from or enter into drainage systems may also be considered as 109 110 diverging or converging junction flows respectively, both of which have been extensively studied on a fundamental level (McNown 1954, Graber 2010) although commonly in pipe diameters much 111 smaller than those found in drainage networks. Bazin et al. (2014) separated the path of 112 surcharging pipe flow into successive sections, corresponding each to a specific head loss: linear 113 114 head loss in the pipe, head loss at a division or a junction or head loss between the surface and the 115 sewer grate. It is also possible to conceptualize an interaction node such as a manhole as a storage 116 element with levels that rise and fall depending on net inflows and outflows through a time varying event. Hence, by utilizing more universal concepts associated with energy losses in diverging and 117 118 converging flows, more generally applicable methodologies may be determined.

119 This paper develops and presents two analytical modelling approaches to represent exchange 120 flows between piped drainage and surface flows via exchange structures such as manholes and 121 gullies. A number of experimental datasets described in Rubinato et al. (2017) are used to calibrate 122 the models in steady flow conditions and validate the models using a series of unsteady flow 123 events. The models performance is compared to both experimental data as well the widely used 124 weir/orifice based approaches that represent the current state of the art.

125 2. Data from large-scale laboratory experiments

126 The experiments were conducted within a facility constructed from PVC in the water 127 laboratory of the Civil and Structural Engineering Department at the University of Sheffield 128 (Rubinato 2015). This experimental facility consists of a sewer pipe system with no slope that is linked via a manhole to a hypothetical urban floodplain characterized by a slope of 0.001. Figure 129 1 provides a schematic diagram of the experimental setup. The sewer system was constructed 130 based on a 1/6 geometrical scale of a typical UK urban drainage system, while the urban floodplain 131 is 4 m wide and 8 m long and it is 0.478 m above the invert level of the pipe. This height is denoted 132 133 as Z_{crest} . The internal diameter of the manhole, D_m , is equal to 240 mm, while the internal diameter of the sewer pipe, D_p , is equal to 75 mm both upstream and downstream of the manhole. In the 134 following, the cross-section of the manhole and of the sewer pipe are denoted as A_m and A_p , 135 respectively. Flow discharges were measured with electromagnetic flowmeters in the floodplain 136 upstream, Q_1 , and downstream, Q_2 , of the manhole, and in the sewer pipes also upstream, Q_3 , and 137 downstream, Q_4 , of the manhole. Q_e denotes the exchange flow between the top of the manhole 138 and the surface flow. In the laboratory experiments, Q_e was not measured directly but, for steady-139 state flow conditions, it can be estimated from the difference between Q_3 and Q_4 which were both 140 measured. Pressure sensors provided the pressure head in the floodplain upstream of the manhole, 141 h_{p1} , and in the sewer pipes upstream, h_{p3} , and downstream, h_{p4} , of the manhole. The horizontal 142 distances of the pressure sensors from the nearest edge of the manhole were $L_1 = 340$ mm for the 143 144 pressure sensor in the floodplain and $L_3 = 230$ mm and $L_4 = 400$ mm for the pressure sensors in the sewer pipes upstream and downstream of the manhole, respectively (Figure 1). More details 145 146 about the experimental facility, the instrumentation and the test program are provided in Rubinato 147 (2015) and Rubinato et al. (2017).



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Figure 1. Schematic diagram of the experimental setup of Rubinato et al. (2017) (figure not to scale).

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Rubinato et al. (2017) analyzed both steady and unsteady flow cases and their extensive dataset 152 is reanalyzed in the present study. This dataset is the only currently openly available dataset on 153 unsteady flow through an individual exchange structure. The only additional data that are analyzed 154 herein and were not presented in Rubinato et al. (2017) are the pressure head data, h_{p4} , in the 155 156 downstream sewer pipe (Rubinato 2015). From the steady state tests, 15 cases with nonsurcharging sewer (Scenario 1 in Figure 2) and eight cases with surcharging sewer (Scenario 3 in 157 Figure 2) are considered here. For the nine unsteady flow tests, the flow on the floodplain was 158 159 maintained constant at 8.15 l/s while a flood hydrograph was run through the sewer pipe, replicating surface to sewer and sewer to surface flow exchange conditions during each unsteady 160 161 test. During the experiments, discharge and pressure measurements were recorded every dt = 0.05seconds. For steady flow experiments, the flow depth at the surface was measured at the location 162 163 of the pressure sensor (Figure 1), while for the unsteady flow experiments, the flow depth was estimated from the equation of Manning, with a Manning roughness coefficient equal to 0.009, 164 165 similar to an approach adopted in Rubinato et al. (2017).

166 **3.** Methods

167 The modelling methodologies analysed here aim at computing the exchange discharge, Q_e , 168 between a manhole and the surface during an unsteady event in a dual drainage system. The 169 exchange discharge Q_e is defined positive when the exchange flow goes towards the floodplain 170 (surcharging sewer). The two models that are presented in this section are referred to as "quasi-171 steady model" and "dynamic model", respectively.

In our quasi-steady model, we follow a similar approach as Rubinato et al. (2017), which regards 172 the pipe-manhole system as a flow junction ($Q_e \leq 0$) or bifurcation ($Q_e > 0$). Nonetheless, 173 compared to the weir/orifice flow exchange approach evaluated in Rubinato et al. (2017), we claim 174 that our quasi-steady model introduces more physics in the calculation of Q_e . Firstly, we link the 175 176 exchange flow discharge to the total head in the upstream pipe, while Rubinato et al. (2017) considered only the pressure head. Secondly, for the case of surcharging flow from the manhole 177 to the floodplain, we account explicitly for four different contributions to the overall head losses 178 179 between the pipe and the surface, similar to Bazin et al. (2014) for flow exchange between a street and an underground drainage pipe. Subsequently, we utilise the Bernoulli equation to describe the 180 flow exchange, whereas the approach tested in Rubinato et al. (2017) lumped the influence of all 181 head losses into a single calibrated orifice equation discharge coefficient. This is detailed in 182 183 Section 3.5. In our new dynamic model, we additionally take into account the evolution with time of the water level, h_m , in the manhole during an unsteady flow event. This evolution cannot be 184 expressed in quasi-steady models and it provides a better representation of transient effects in the 185 computation of the exchange discharge Q_e because the storage capacity in the manhole is 186 187 accounted for explicitly. Calibration of the models was performed in steady flow conditions, due to the ability to directly measure the flow rate and hence identify the energy loss parameters, 188

without the need for more complex model calibration methodologies (e.g. Noh et al. 2016). Model validation (or evaluation) is commonly undertaken after model calibration as a method of determining the ability of the model to replicate observed parameters without further modifications from the user (McMillan et al. 2016). To quantify any resulting uncertainties when simulating dynamic events, in this study validation was performed in unsteady flow conditions. The models are presented analytically in the following sections.

195 3.1. Flow scenarios

196 The model evaluated in Rubinato et al. (2017) as well as the two models presented here share the same typology of flow scenarios; but some refinements are introduced here. For this purpose, we 197 define notation H_m to refer to the flow head at the interaction node (manhole), which may be 198 199 approximated using different methods for the different models. The reference datum is the sewer pipe invert (Figure 1). Three different flow scenarios may occur (Figure 2), depending on the value 200 of the head H_m in the manhole with respect to the elevation of the floodplain, Z_{crest} , and the head 201 of the flow in the floodplain, H_s , which is equal to $(Q_1 / (W h_s))^2 / (2g) + Z_{crest} + h_s$, where W and 202 h_s are the width of and the flow depth in the floodplain, respectively. 203

- Scenario 1: free weir flow from the floodplain to the manhole ($Q_e < 0$), if the head in the manhole is lower than the level of the floodplain ($H_m \le Z_{crest}$);
- Scenario 2: submerged weir or orifice flow from the floodplain to the manhole ($Q_e < 0$), if 207 the head in the manhole is higher than the level of the floodplain but lower than the head 208 of the surface flow ($Z_{crest} < H_m \le H_s$);
- Scenario 3: overflow from surcharging sewer to the floodplain ($Q_e > 0$), if the head in the 210 manhole is greater than the head of the surface flow ($H_m > H_s$).

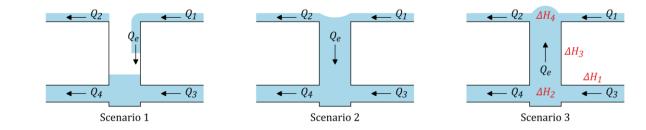




Figure 2. The three flow scenarios observed during a hydrograph (adapted from Rubinato et al. 2017). In Scenario 3, $\Delta H_1 - \Delta H_4$ denote the head losses that occur at different segments.

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3.2. Approximation of head within the manhole/pipe network based on the different models

The models presented herein and the models evaluated in Rubinato et al. (2017) differ in the level of detail in which they define and estimate for different flow conditions the total head of the flow in the manhole, and hence the head of the sewer network flow at the point it interacts with the surface flow:

Rubinato et al. (2017) calibrated the orifice flow exchange equation based on the pressure
 head *h*_{p3} at the location of the pressure sensor in the upstream pipe.

In the quasi-steady model, we approximate H_m by $H_3 - \Delta H_{tot}$. 222 • with $H_3 = (Q_3 / A_p)^2 / (2g) + h_{p3}$ the total head at the pressure sensor in the upstream pipe and 223 ΔH_{tot} the total head losses between the location of this pressure sensor and the point of flow 224 interaction (Figure 2) (see details in Section 3.5.1). For non-surcharging conditions (i.e., 225 Scenarios 1 and 2), head losses due to Q_e do not occur, as there is no upward flow in the 226 manhole, and the total head losses ΔH_{tot} are noted as ΔH_0 . 227

• In the dynamic model, H_m is simply taken equal to the water depth, h_m , in the manhole assuming the velocity head in the manhole is negligible. Note that the variable h_m is not considered in the other two models. 231

232 3.3. Mass balance in the manhole

Like in a pipe junction or bifurcation, the mass balance in the quasi-steady model can be writtenas:

$$Q_4 = Q_3 - Q_e \tag{1}$$

where the upstream discharge, Q_3 , in the pipe is a prescribed boundary condition and the exchange discharge Q_e is predicted by the model. Therefore, the discharge Q_4 in the downstream pipe can be determined directly from Eq. (1).

Unlike the quasi-steady model, the dynamic model considers the manhole as a tank in whichthe volume of water varies in time with the contributing discharges, according to:

$$A_m \frac{dh_m}{dt} = Q_3 - Q_e - Q_4 \tag{2}$$

Like in the quasi-steady model, the upstream pipe discharge Q_3 is a prescribed boundary condition and the exchange discharge Q_e is estimated by the model. However, Eq. (2) is now necessary to update the value of h_m for the subsequent time step. Therefore, in the dynamic model, the discharge Q_4 in the downstream pipe needs to be computed separately. This step is further described in Section 3.5.2. In steady flow conditions, the mass balance in Eq. (2) reduces to Eq. (1) as in the quasi-steady model. 248 3.4. Non-surcharging sewer (surface to sewer exchange)

For non-surcharging flow conditions ($Q_e \le 0$), the quasi-steady and the dynamic model have similar exchange equations with the model tested in Rubinato et al. (2017), with the only difference being the utilization of the total head in the surface flow instead of just using the flow depth:

• in Scenario 1 ($H_m \le Z_{crest}$), a weir equation is used to describe free flow from the floodplain to the manhole:

254
$$Q_{e} = -\frac{2}{3}C_{1}\pi D_{m}\sqrt{2g(H_{s} - Z_{crest})^{3}}$$
(3)

with C_1 a discharge coefficient to be calibrated;

• in Scenario 2 ($Z_{crest} < H_m \le H_s$), a submerged weir equation is used when H_s - Z_{crest} 257 $\le A_m / (\pi D_m)$:

258
$$Q_e = -C_2 \pi D_m \left(H_s - Z_{crest}\right) \sqrt{2g \left(H_s - H_m\right)}$$
(4)

where C_2 is also a discharge coefficient whose value needs to be determined. When H_s - $Z_{crest} > A_m / (\pi D_m)$, a submerged orifice equation may be used (Rubinato et al. 2017); however, this threshold is not exceeded in our study.

Rubinato et al. (2017) used experimental observations in steady conditions to calibrate parameters C_1 and C_2 . Here, we recalibrated the value of C_1 with the same data as Rubinato et al. (2017) for Scenario 1 (see the expected values in Figure 6 in Rubinato et al. 2017), but with the surface flow head instead of the flow depth (Eq. (3)). The calibration performed by Rubinato et al. (2017) for parameter C_2 involved less data points and led to discontinuities in the computed exchange flow discharge at the transition between Scenarios 1 and 2 (see Figure 10 in Rubinato et al. 2017). 268 Therefore, in the models introduced here, we simply set $C_2 = 2/3 C_1$, which ensures continuity between the exchange discharges computed by Eqs. (3) and (4) when $H_m = Z_{crest}$. Compared to the 269 strategy followed by Rubinato et al. (2017), the continuity between Scenarios 1 and 2 is ensured 270 here at the expense of an accurate agreement with calibration data for Scenario 2; but the number 271 of available experimental data for Scenario 2 is limited and this scenario occurs in practice only 272 for a very short period of time during unsteady flow events at the onset and at the end of 273 surcharging flow (rising and falling limbs of the hydrograph). Therefore, the impact of this choice 274 on the overall accuracy of the computed exchange volume over a surcharging flow event is 275 276 expected to be very small. All calibration parameters are summarized in Table 1.

277 3.5. Surcharging sewer

278 3.5.1. Quasi-steady model

279 In the quasi-steady model, the overflow discharge from the surcharging manhole to the floodplain is computed from a Bernoulli equation written between the upstream sewer pipe (at the 280 281 pressure sensor location) and the surface flow, and by taking into account the total head losses, 282 ΔH_{tot} , along the flow path. Head losses occur at four different locations, as shown in Figure 2 in 283 Scenario 3, hence $\Delta H_{tot} = \Delta H_1 + \Delta H_2 + \Delta H_3 + \Delta H_4$. Linear head losses ΔH_1 are noted along the sewer pipe due to friction and head losses ΔH_2 occur due to flow division and expansion of the 284 cross-section at the junction where the sewer pipe meets the manhole. Additional energy is 285 dissipated as the water flows upward through the manhole with frictional linear head losses ΔH_3 , 286 and finally head losses ΔH_4 occur as the water exits the manhole to the street. Head losses between 287 the sewer pipe and the surface flow can therefore be described by (Idelchik 2007, Bazin et al. 288 2014): 289

$$H_{3}-H_{s} = \underbrace{\frac{f_{p3}L_{3}}{D_{p}}\frac{Q_{3}^{2}}{2gA_{p}^{2}}}_{\Delta H_{1}} + \underbrace{\frac{Q_{3}^{2}}{2gA_{p}^{2}}}_{\Delta H_{2}} + \underbrace{\frac{f_{m}\left(Z_{crest}-D_{p}\right)}{D_{m}}\frac{Q_{e}^{2}}{2gA_{m}^{2}}}_{\Delta H_{3}} + \underbrace{\frac{k_{4}\alpha_{4}^{2}\frac{Q_{e}^{2}}{2gA_{m}^{2}}}_{\Delta H_{4}}}_{\Delta H_{4}}$$
(5)

where *f* is a friction coefficient denoted as f_{p3} and f_m for the upstream sewer pipe and the manhole, respectively. The friction coefficient, *f*, is estimated with the formula of Barr (Machiels et al. 2011) as a function of the roughness height k_s , which is considered equal to 0.0005 mm both for the sewer pipe and the manhole, based on a previous calibration performed by Beg et al. (2020).

295 In Eq. (5), the parameter α_4 is the ratio of flow velocity exiting the manhole to the flow velocity inside the manhole, with α_4^2 being equal to 0.95 (Idelchik 2007), while the coefficient k_4 is 296 associated with local head losses at the exit of the manhole and is considered equal to 1 (Idelchik 297 298 2007). Hence, the only remaining parameter to be determined in Eq. (5) is the coefficient k_2 299 associated with head losses due to expansion of the flow from the sewer pipe to the manhole. This 300 parameter is likely to require calibration because the available values in the literature either correspond to ratios A_m / A_p of less than one, e.g. in Idelchik (2007), or to pipes of equal cross-301 302 sectional area, e.g. in Hager (2010), while in this case the A_m / A_p ratio is greater than ten. As a 303 result, the available values may not be applicable.

304 *Calibration procedure*

By rearranging Eq. (5) with the standard values of parameters k_4 and α_4 , the numerical value of k_2 can be computed as:

307
$$k_{2} = \frac{2gA_{p}^{2}}{Q_{3}^{2}} \left[H_{3} - H_{s} - \frac{f_{p3}L_{3}}{D_{p}} \frac{Q_{3}^{2}}{2gA_{p}^{2}} - \frac{f_{m}(Z_{crest} - D_{p})}{D_{m}} \frac{Q_{e}^{2}}{2gA_{m}^{2}} - k_{4}\alpha_{4}^{2} \frac{Q_{e}^{2}}{2gA_{m}^{2}} \right]$$
(6)

where all quantities in the right-hand-side of Eq. (6) can be evaluated from experimental steady flow observations of Q_3 , h_s , h_{p3} and Q_e . Note that the total head H_s in the floodplain is measured upstream and not downstream of the manhole. As a result, potential head changes of the surface flow due to the interaction of floodplain flow with the surcharging jet are ignored; however, these head changes are negligible because the floodplain is very wide, and hence the overall difference in surface flow depth and velocity upstream and downstream of the manhole is negligible (Rubinato et al. 2018b).

When plotting the values of k_2 from Eq. (6) as a function of the portion of the pipe inflow discharge being exchanged with the surface, Q_e / Q_3 , it appears that the parameter k_2 varies almost linearly with Q_e / Q_3 (Section 4.1.1). Therefore, we consider the following linear function to parameterize k_2 :

$$k_2 = \alpha \frac{Q_e}{Q_3} + \beta \tag{7}$$

where parameters α and β need to be calibrated with experimental data, as shown in Section 4.1.1.

321 *Applying the quasi-steady model*

To classify the different flow scenarios (e.g. the transition point between non surcharging and surcharging flows) ΔH_0 (see Section 3.2) is first needed. In Scenarios 1 and 2 there is no upward flow in the manhole and hence no head losses due to Q_{e} . By substituting $Q_e = 0$ in Eqs. (5) and (7) , it follows that the head loss ΔH_0 , introduced in Section 3.2, may be evaluated by:

326
$$\Delta H_0 = \left(\beta + \frac{f_{p3}L_3}{D_p}\right) \frac{Q_3^2}{2gA_p^2} \tag{8}$$

To classify Scenarios 2 and 3 (transition to surcharging flows) H_m is compared to H_s , with H_m taken equal to $H_3 - \Delta H_0$. In this form, if H_m is lower than H_s , then there is no surcharge and no upward flow in the manhole, while if H_m is greater than H_s , then upward, surcharging flow occurs. In the latter case, the total head in the manhole H_m should then be approximated by $H_3 - \Delta H_{tot}$, which is smaller than $H_3 - \Delta H_0$, because of the inclusion of additional positive parameters in the head losses (Eq. (5)).

After the calibration of the parameter k_2 , the exchange discharge, Q_e , in surcharging flow conditions (Scenario 3) can be estimated with Eq. (5) through an iterative process by testing values of Q_e until the two sides of the equation converge. The needed input parameters are the flow discharge and pressure in the upstream sewer pipe, the flow discharge at the floodplain, and the geometric characteristics of the drainage system and the floodplain.

338 *3.5.2. Dynamic model*

Similarly to Rubinato et al. (2017), the dynamic model for surcharging sewer uses simply an orifice equation to estimate the surcharging discharge. Nevertheless, in this case the head in the equation is taken here equal to the water depth in the manhole, h_m .

342
$$Q_e = C_3 A_m \sqrt{2g \left(h_m - H_s\right)}$$
(9)

In this case, to determine Q_e , the discharge Q_4 is first needed to compute the water depth in the manhole with Eq. (2). Subsequently, the computed water depth can be used for the estimation of the exchange discharge with Eq. (9). In the dynamic model, the discharge Q_4 is estimated by applying Bernoulli equation between the top of the surcharging manhole jet and the position of the pressure sensor in the downstream sewer pipe (Figure 1) where a head, H_4 , boundary condition is 348 set. Local head losses between the manhole jet and H_4 are expressed similarly to the quasi-steady model. To simplify the computations, the head losses in the manhole, i.e., ΔH_3 and ΔH_4 in Eq. (5) 349 , are omitted because they are considered negligible as shown by the results of the quasi-steady 350 model (Section 4.1.1). Hence, Bernoulli equation between the water surface in the manhole and a 351 352 point in the downstream sewer pipe that is located at a distance L_4 from the downstream edge of the manhole takes into account local contraction losses as the flow exits the manhole, as well as 353 frictional losses in the pipe. Assuming the losses from the contraction form a similar relationship 354 355 with the flow partition as in Eq. (7), the Bernoulli equation with the aid of Eq. (1) can be written 356 as:

357
$$h_m - H_4 = \left(\alpha' \frac{Q_3 - Q_4}{Q_4} + \beta' + f_{p4} \frac{L_4}{D_p}\right) \frac{Q_4^2}{2gA_p^2}$$
(10)

358 where f_{p4} is the friction coefficient for flow in the downstream sewer pipe.

359 *Calibration procedure*

Based on observed data of steady surcharging flow, parameters α' and β' in Eq. (10) may be determined by a linear regression with $(Q_3 - Q_4)/Q_4$ being the independent variable. However, this requires the prior knowledge of h_m . This is attained by applying the Bernoulli equation between the location of the pressure transducer in the sewer pipe upstream of the manhole and the top of the surcharging manhole jet, as follows:

365
$$H_{3} - h_{m} = f_{p3} \frac{L_{3}}{D_{p}} \frac{Q_{3}^{2}}{2gA_{p}^{2}} + k_{2}^{*} \frac{Q_{3}^{2}}{2gA_{p}^{2}}$$
(11)

This equation has a similar structure to Eq. (5) but the head losses in the manhole and in the overflow are considered negligible, similar to Eq. (10), while the coefficient k_2 " differs from k_2 of Eq. (5) because of the utilization of h_m .

The combination of Eqs. (9) and (11), along with the division of both sides of the resulting equation with the velocity head in the upstream sewer pipe and the relationship $Q_3 - Q_4 = Q_e$ in steady surcharging flow conditions, lead to the following non-dimensional equation:

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$$(H_3 - H_s) \frac{2gA_p^2}{Q_3^2} - f_{p3} \frac{L_3}{D_p} = \frac{A_p^2}{A_m^2 C_3^2} \left(\frac{Q_3 - Q_4}{Q_3}\right)^2 + k_2^{"}$$
(12)

Based on observed data of steady surcharging flow, the discharge coefficient C_3 can now be estimated with polynomial regression analysis of Eq. (12), with $(Q_3 - Q_4)/Q_3$ being the independent variable. The water depth h_m in the manhole for steady flow conditions can be subsequently calculated with Eq. (9) or Eq. (11) and, finally, the parameters α' and β' in Eq. (10) can be estimated by linear regression.

378 *Applying the dynamic model*

Given knowledge of the discharge coefficient C_3 and the parameters α' and β' , the exchange discharge, Q_e , and the discharge in the downstream sewer pipe, Q_4 , for surcharging conditions can be estimated for each time step with Eqs. (9) and (10), respectively. The water depth in the manhole is updated at each time step in unsteady flow conditions with Eq. (2). The data requirements of the dynamic model are the flow discharge and pressure in the upstream sewer pipe, the pressure in the downstream sewer pipe, the flow discharge at the floodplain, and the geometric characteristics of the drainage system and the floodplain. 386

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Table 1. Parameterization of the head of the flow in the manhole, H_m , and calibration parameters for the examined models.

	Rubinato et al. (2017)	Quasi-steady model	Dynamic model		
H_m	h_{p3}	$H_3 - \Delta H_{\rm tot}$	h_m		
C_1	0.54	0.38	0.38		
C_2	0.056	2 / 3 × 0.38	2 / 3 × 0.38		
<i>C</i> ₃	0.167	-	0.168		

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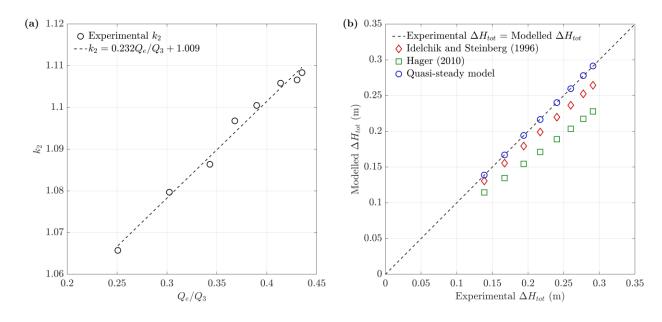
392 4. Results and discussion

4.1. Calibration of models with steady flow data for surcharging flow conditions

394 4.1.1. Quasi-steady model

For each of the eight steady-state surcharging flow tests conducted by Rubinato et al. (2017), 395 the numerical value of k_2 was computed from the experimental observations using Eq. (6). When 396 the values of k_2 are plotted against the observed ratios Q_e / Q_3 , the data points follow a linear trend, 397 398 as demonstrated in Figure 3a. This confirms the relevance of the parametrization proposed in 399 Eq. (7). By applying linear regression, the coefficients in Eq. (7) are evaluated as $\alpha = 0.232$ and $\beta = 1.009$ (Figure 3a). Subsequently, the computed total head losses from the sewer pipe to the 400 401 surface, as modelled based on the right-hand-side of Eq. (5) and with the aid of Eq. (7), are compared to the observed head difference $H_3 - H_s$. As shown in Figure 3b, the computed values 402 agree well with the measurements. 403

For the sake of comparison, the total head losses were also computed by estimating the parameter k_2 with the formulae of Idelchik (2007) and Hager (2010), as described in the Appendix. Both of these models underestimate the total head losses (Figure 3b). The calibrated quasi-steady model performs slightly better than the formula of Idelchik (2007) and significantly better than the formula of Hager (2010). Although less accurate than the model calibrated here, the formula of Idelchik (2007) still provides a useful value of k_2 to estimate the total head loss in the absence of calibration data.



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Figure 3. (a) Linear regression between parameter k_2 and Q_e / Q_3 for the quasi-steady model based on the experimental data of Rubinato et al. (2017) for steady surcharging flow and (b) Comparison of observed (left-hand-side of Eq. (5)) and computed (right-hand-side of Eq. (5)) total head losses from the sewer pipe to the surface for steady surcharging flow.

Figure 4 shows the head losses that occur at each segment of the system for Scenario 3 (Figure 2). The total head losses, ΔH_{tot} , depend mostly on the head losses in the second section of the system, ΔH_2 , where the sewer pipe meets the manhole. Specifically, for the eight steady flow experiments of Rubinato et al. (2017) with surcharging flow, ΔH_2 constitutes more than 95% of the total head losses, whereas ΔH_1 , is less than 5%, and ΔH_3 and ΔH_4 are approximately 0.01% and 0.1%, respectively. Due to the small contribution of ΔH_4 to the total head loss, the latter is not particularly sensitive to parameters k_4 and α_4 which justifies the use of standard values.

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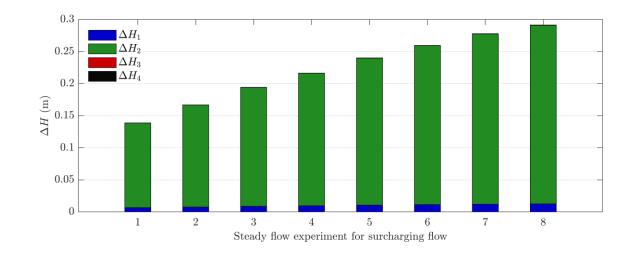
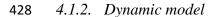


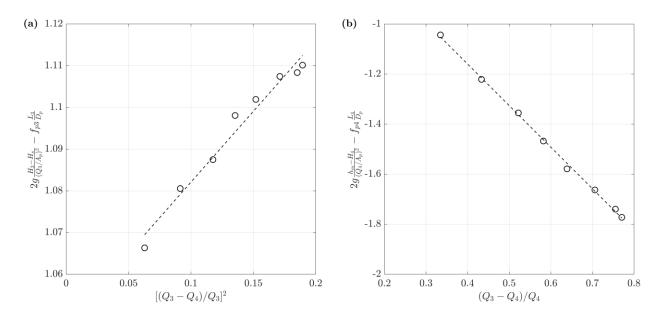
Figure 4. Distributions of head losses for surcharging sewer under steady flow conditions basedon experimental data of Rubinato et al. (2017).

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Based on the measurements of Rubinato et al. (2017) for steady surcharging flow, the observed values of the left-hand-side of Eq. (12) can be plotted as a function of the measured values of $(Q_3 - Q_4)/Q_3$, as shown in Figure 5a. A linear regression with $[(Q_3 - Q_4)/Q_3]^2$ being the independent variable, leads to $A_p^2/(C_3 A_m)^2 = 0.340$, from which it can be deduced that the discharge coefficient C_3 for the dynamic model is equal to 0.168. The discharge coefficient that was generated with this method is remarkably similar to that estimated by Rubinato et al. (2017) (Table 1), despite the fact that the two methods have notable differences. It should be noted that in 436 Eq. (12), k_2'' was considered to be independent of $(Q_3 - Q_4)/Q_3$, because otherwise a linear 437 dependency would lead to an unrealistic value of C_3 .

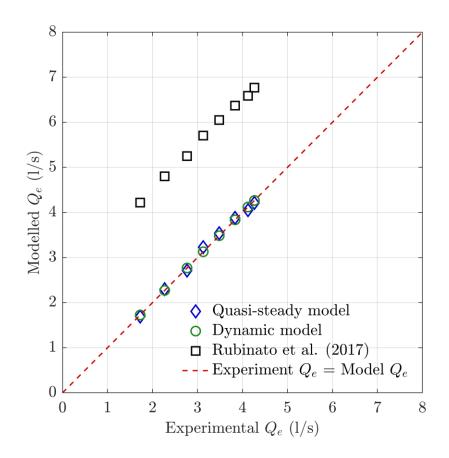


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Figure 5. Calibration of the dynamic model with (a) Linear regression between $[(Q_3 - Q_4)/Q_3]^2$ and the dimensionless head loss from Eq. (12) for the determination of the discharge coefficient C_3 and (b) Linear regression between $(Q_3 - Q_4)/Q_4$ and the dimensionless head loss from Eq. (10) for the determination of parameters α' and β' .

Subsequently, h_m is calculated from Eq. (9), which allows the application of linear regression in Eq. (10) for the determination of parameters α' and β' . Figure 5b shows that the linear regression fits the data well with $\alpha' = -1.660$ and $\beta' = -0.496$. In case Eq. (11) was used for the calculation of h_m , the parameters α' and β' would differ by less than 1%.

The resulting modelled exchange discharges, Q_e , obtained from the quasi-steady and the dynamic models agree well with the experimental data of Rubinato et al. (2017), as shown in Figure 6. The results are also compared to the results obtained with the orifice equation calibrated experimentally by Rubinato et al. (2017). Note that the perfect agreement of the dynamic model is owed to the fact that h_m was calculated from Eq. (9). Rubinato et al. (2017) estimated expected values and upper and lower values of the exchange discharge, based on an error parameter associated with the instrumentation error. Our models are compared to the expected values of
Rubinato et al. (2017), which overestimate the exchange discharge by approximately 2.5 l/s
(Figure 6). This bias corresponds to the intercept visible in Fig. 8 in Rubinato et al. (2017), which
is indeed of the order of 2.5 l/s.



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Figure 6. Comparison of experimental and modelled exchange discharge, Q_e , for surcharging sewer under steady flow conditions. The model tested in Rubinato et al. (2017) was used with its calibrated expected discharge coefficient, with respect to its experimental measurement uncertainty. The data are from the experiments of Rubinato et al. (2017).

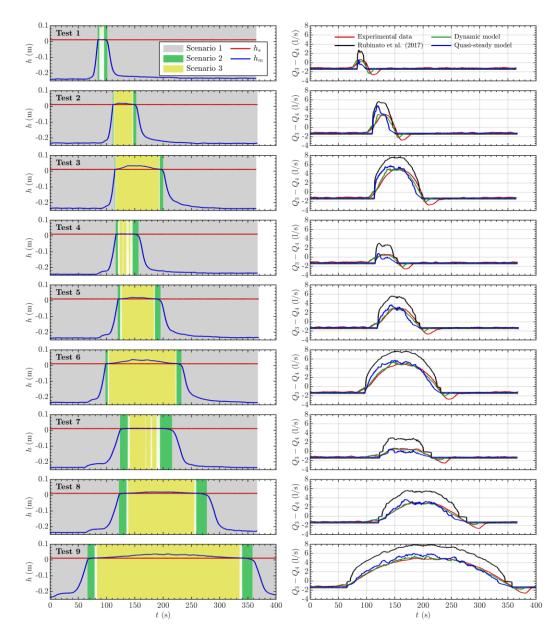
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463 4.2. Validation of models with unsteady flow data

Figure 7 presents a comparison of the results of the new quasi-steady and dynamic models to the experimental observations and the computations presented in Rubinato et al. (2017), for the 466 nine unsteady experiments reported by Rubinato et al. (2017). The numerical results of Rubinato et al. (2017) displayed in this figure were obtained by using the values of the discharge coefficients 467 C_1 , C_2 , and C_3 calibrated with steady flow experiments ("expected values" in Table 3 of Rubinato 468 et al. 2017) and the observed values of h_{p3} , while h_s was calculated with the equation of Manning. 469 The evolution of the depth h_m in the manhole is computed only by the dynamic model (left column 470 471 in Figure 7). In all cases, the dynamic model exhibits a better agreement with the measured data compared to the quasi-steady model and to the model tested in Rubinato et al. (2017) (Table 2), 472 both at the rising and the falling limbs of the hydrographs (right column in Figure 7). The quasi-473 474 steady model performs generally better than the model tested in Rubinato et al. (2017), which overestimates the exchange discharge. This is consistent with the overestimation of the exchange 475 discharge by the model tested in Rubinato et al. (2017) when there is surcharge under steady flow 476 conditions, as highlighted in Figure 6. 477

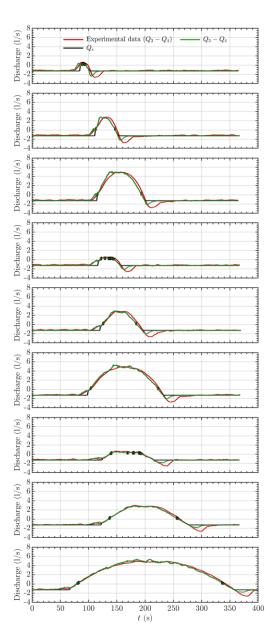
478 The flow in the drainage system is mostly classified as Scenarios 1 and 3 (Figure 2), with 479 Scenario 2 occurring only for brief transitional periods of time (Figure 7 and Table 2). While the transition between Scenarios 1 and 2 is smooth, the transition between Scenarios 2 and 3 can be 480 abrupt, as shown by the dynamic model in Figure 7. Although the raw pressure input data 481 /measurements in the sewer pipes were filtered to smoothen the time-series, the dynamic model is 482 still sensitive to the unsteadiness of the flow, which leads to rapid fluctuations between Scenarios 483 3 and 2, i.e., manhole surcharge or not. This sensitivity is particularly evident in tests 1, 4, and 7, 484 where the peaks of the hydrographs are the lowest and the quasi-steady and dynamic models, 485 486 particularly the former, are not always able to classify correctly when the flow enters Scenario 3, (in which $Q_3 - Q_4 > 0$, Table 2). 487

488 For the computations performed with the quasi-steady model and by Rubinato et al. (2017), the calculated exchange discharge Q_e is considered equal to the value of $Q_3 - Q_4$, according to 489 Eq. (1). In contrast, in reality, the exchange discharge Q_e differs from the value of $Q_3 - Q_4$ during 490 transient phases, as a result of variations in the storage in the manhole as expressed by Eq. (2). 491 Here, the experimental dataset reports only $Q_3 - Q_4$ due to the infeasibility of measuring 492 493 continuously the evolution of Q_e in the laboratory setup (Rubinato 2015). Only the dynamic model gives access to both Q_e and $Q_3 - Q_4$, as these quantities are computed separately by this model. 494 Figure 8 shows a comparison between these two discharges obtained from the dynamic model, 495 496 from which significant overlap can be observed for the largest part of the hydrograph, besides the start and the end of the unsteady sections. Particularly at the end of the unsteady section of the 497 hydrograph, the suction that is observed in the manhole as the water depth decreases in the 498 transition from Scenario 2 (submerged weir) to Scenario 1 (free weir) is partially captured only by 499 the $Q_3 - Q_4$ results. This is owed to the transient nature of the dynamic model and its ability to 500 represent the evolution of storage in the manhole. The abrupt changes between Scenarios 2 and 3 501 are also evident in the hydrographs of Q_e , where the exchange discharge fluctuates between 502 positive and negative values before it stabilizes. These abrupt transitions in the exchange discharge 503 504 correspond to the white areas between Scenarios 2 and 3 in Figure 7 (left column). Despite this sensitivity in the computation of the exchange discharge, the dynamic model results of $Q_3 - Q_4$ 505 506 agree well with the experimental data.



508

Figure 7. Evolution of water level in the manhole (left column) and comparison of modelling results for the discharge in the manhole, $Q_3 - Q_4$, (right column) with data from unsteady flow experiments from Rubinato et al. (2017). For the quasi-steady model and the model tested in Rubinato et al. (2017), Q_e was used as a proxy for $Q_3 - Q_4$. In the left column, *h* denotes the water level with respect to the surface, which is located at h = 0, and h_m was computed with the dynamic model. The different scenarios in the left column were determined with the dynamic model with the white areas in between Scenarios 2 and 3 denoting rapid transitions between these scenarios.



517

Figure 8. Comparison of Q_e and $Q_3 - Q_4$ predicted by the dynamic model. The data are from the experiments of Rubinato et al. (2017).

A quantitative evaluation of the unsteady modelling results for the discharge in the manhole is provided in Table 2, which shows the Nash–Sutcliffe efficiency (NSE) coefficient between the results of each model and the corresponding measurements for the unsteady part of each hydrograph. The NSE coefficient is consistently higher for the dynamic model followed by the 524 quasi-steady model. The performance of both models improves as the surcharge becomes more

525 intense, while the difference between the two models is the lowest for tests 3, 5, 6, 8, and 9, which

are cases with high hydrograph peaks and long duration of the unsteady section (Figure 7).

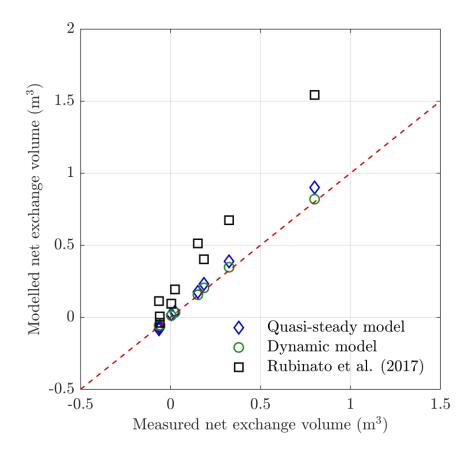
527

528	Table 2. Percentages of the duration of occurrence of $Q_3 - Q_4 > 0$ for the experimental data and
529	of the occurrence of the three different scenarios for the quasi-steady model, the dynamic model,
530	and the model tested in Rubinato et al. (2017) for the whole duration of each unsteady test. NSE
531	denotes the Nash-Sutcliffe efficiency coefficient for the modelled and measured discharge in the
532	manhole, $Q_3 - Q_4$, for the unsteady part of each hydrograph from Figure 7. For the quasi-steady
533	model and the model tested in Rubinato et al. (2017), Q_e was used as a proxy for $Q_3 - Q_4$. The
534	model tested in Rubinato et al. (2017) was used with its expected discharge coefficients, with
535	respect to their calibration experimental uncertainty, while the surface flow depth was estimated
536	with the equation of Manning.

		Test 1	Test 2	Test 3	Test 4	Test 5	Test 6	Test 7	Test 8	Test 9
	Data: $Q_3 - Q_4 > 0$	4.3%	11.0%	22.6%	7.9%	17.4%	33.6%	16.4%	34.5%	57.4%
Quasi-steady model	Scenario 1	96.1%	89.4%	77.3%	89.8%	80.1%	64.1%	75.5%	57.9%	36.8%
	Scenario 2	3.4%	2.8%	2.8%	8.5%	4.6%	4.3%	18.8%	10.2%	8.2%
lasi- mo	Scenario 3	0.5%	7.8%	19.9%	1.7%	15.2%	31.6%	5.7%	31.8%	55.0%
ð	NSE	0.338	0.390	0.774	0.427	0.757	0.856	0.634	0.878	0.916
	Scenario 1	95.3%	88.6%	76.5%	89.0%	79.5%	63.4%	74.9%	57.6%	36.4%
Dynamic model	Scenario 2	3.5%	2.2%	2.3%	5.7%	4.5%	4.0%	11.3%	9.6%	7.6%
)yn: mo	Scenario 3	1.2%	9.2%	21.2%	5.3%	16.0%	32.6%	13.8%	32.8%	55.9%
Γ	NSE	0.667	0.874	0.948	0.842	0.938	0.966	0.907	0.967	0.982
) et	Scenario 1	95.7%	89.0%	77.0%	89.4%	79.7%	63.6%	75.0%	57.5%	36.4%
Rubinato e al. (2017)	Scenario 2	1.2%	1.3%	1.2%	2.1%	2.2%	2.1%	5.5%	5.5%	4.0%
	Scenario 3	3.2%	9.7%	21.8%	8.5%	18.0%	34.3%	19.5%	37.0%	59.6%
	NSE	0.144	0.059	0.412	-0.604	0.093	0.221	-1.265	-0.268	-0.037

537

The modelled and measured net water volumes that are exchanged between the sewer and the floodplain are compared in Figure 9. A very good agreement is obtained for the dynamic model. In some cases (e.g. Tests 1 and 2), the quasi-steady model seems to predict the exchanged volume as well as the dynamic model, despite the fact that the overall evolution of the exchange discharge is less accurate than the dynamic model (Figure 7). In reality, the dynamic model is more reliable since it captures better the governing physics. Nonetheless, considering that the quasi-steady model exhibits a good agreement with the experiments in cases where the flow unsteady hydrograph is long and the suction effect becomes small, it can be inferred that this model remains also valuable, especially given the fact that it does not require a downstream boundary condition when compared to the dynamic model.



548

Figure 9. Comparison of the modelled and measured net exchange volumes of water between

the sewer and the floodplain for the unsteady part of each hydrograph from Figure 7.

552 5. Concluding remarks

In light of climate change and with the anticipation of an increase in the frequency of extreme rainfall events, the accurate design of drainage systems and accurate evaluation of flood risk is of paramount importance for the resilience of urban areas.

This study developed a quasi-steady and a dynamic model for the determination of the 556 exchange discharge between a sewer pipe and the surface floodplain through a manhole in a typical 557 setup of an urban drainage system. Both models can be utilized for a complete unsteady 558 hydrograph, ranging from inflow from the floodplain into non-surcharging sewer to overflow of 559 the surcharging sewer. When compared to the commonly utilized weir/orifice approach to 560 calculating exchange volumes (Nasello and Tucciarelli 2005, Seyoum et al. 2012), the quasi-steady 561 562 model explicitly accounts for the head losses along the flow path from the sewer pipe to the surface 563 and links the exchange flow to the total head in the sewer pipe minus the occurring head losses. The dynamic model also takes into account the head losses but is also able to estimate the evolution 564 565 of the water level in the manhole with the aid of one additional boundary condition at the downstream sewer pipe. 566

The models were calibrated with steady flow data from large-scale experiments from Rubinato et al. (2017) and were validated against unsteady flow conditions in the same experimental setup. Both models exhibited good agreement with the experimental measurements, with the dynamic model performing a little better with a Nash–Sutcliffe efficiency coefficient for the unsteady section of each tested hydrograph ranging between 0.667 and 0.982. The dynamic model captured better the physics of the problem since it was able to reproduce to a certain degree the suction in the manhole that was observed at the falling limb of the hydrograph. Both models performed

significantly better than the standard weir/orifice formulations for exchange volume as evaluated
in Rubinato et al. (2017). Past work suggests lumping head losses into a single coefficient, which
has resulted in a wide range of calibrated discharge coefficients. These existing methods may be
sensitive to the choice of boundary condition/measurement location as well as to the method of
calculation of pipe/surface hydraulic head (Rubinato et al. 2018a).

The utilization of the models at larger geometrical scales can be facilitated by using non-579 dimensional variables, such as the Froude number within the hypothetical surface and the 580 Reynolds numbers in the pipe and in the manhole. The values of these non-dimensional variables 581 are provided in Rubinato et al. (2017). The hydraulic conditions replicated include fully turbulent 582 583 pipe flows and subcritical flow conditions in reasonably flat floodplains. A topic of further work 584 would be to consider scale effects and transferability of energy loss parameters to full size systems, as well as the transferability of the findings and parameter sensitivity to different flow conditions 585 586 and geometrical configurations. Given an understanding of the relevant boundary conditions via 587 measurements or hydrodynamic modelling, the methodology of this study could be applied to systems with multiple interaction nodes and/or with lids covering the manholes. Further 588 experimental work could consider the calibration of energy loss parameters in such systems and 589 sensitivity of flood modelling predictions to these parameters. 590

Besides the development of the two models and the demonstration of their satisfactory predictive capabilities in unsteady flow conditions, this study also showed that the head losses that occur in the considered dual drainage system consist mostly of the head losses due to the flow expansion at the location where the sewer pipe meets the manhole. Frictional head losses in the sewer pipe are an order of magnitude smaller, while the frictional head losses in the manhole and the head losses where the flow exits the manhole at the surface are negligible, due to the significantly lower velocities involved. Therefore, in order to produce more transferable, standardized energy loss coefficients to describe flow exchange from sewer systems to surface flows, it is suggested that future work focuses on measuring sub-surface pipe/exchange structure hydraulic losses in flood/high flow conditions. It is noted that an extensive body of work already exists on head losses through such structures in non-surcharging/flooding conditions (e.g. Marsalek 1985, Pedersen and Mark 1990), and the feasibility of data from these studies to provide initial estimates of energy losses for use in flood conditions could be investigated.

604

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609 Appendix

610 The model of Idelchik (2007) considers a diverging tee and calculates the parameter k_2 with 611 the following equation:

$$k_2 = 1 + \left(\frac{Q_e A_p}{Q_3 A_m}\right)^2 \tag{13}$$

613 The parameter k_2 with the model of Hager (2010) is given by:

614
$$k_2 = 1 - 2\frac{Q_e}{Q_3} \cos\left(\frac{3}{4}\theta\right) + \left(\frac{Q_e}{Q_3}\right)^2$$
(14)

615 where θ is the angle between the manhole and the sewer pipe, which herein is equal to 90°.

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