

# **Geophysical Research Letters**

UVS

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#### **Key Points:**

- Juno Ultraviolet Spectrograph recorded transient blackbody emission from a point source in Jupiter's atmosphere
- The emission is consistent with a fireball produced by a 250-5,000 kg impactor in Jupiter's upper atmosphere
- We estimate an impact flux on Jupiter of 24,000 per year for masses greater than 250-5,000 kg

**Supporting Information:** 

• Supporting Information S1

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consistent with a 9600-K blackbody located 225 km above the 1-bar level and the duration of the emission was between 17 ms and 150 s. These characteristics are consistent with a bolide in Jupiter's atmosphere. Based on the energy emitted, we estimate that the impactor had a mass of 250–5,000 kg, which corresponds to a diameter of 1-4 m. By considering all observations made with Juno UVS over the first 27 perijoves of the mission, we estimate an impact flux rate of 24,000 per year for impactors with masses greater than 250-5,000 kg.

Plain Language Summary UVS is an ultraviolet spectrograph on NASA's Juno spacecraft, which has been in orbit around Jupiter since 2016. In April 2020, UVS observed short-lived bright emission from a point source in Jupiter's atmosphere. The emission was consistent with a blackbody of temperature 9,600 K. We suggest that this was a fireball produced by a 250-5,000 kg meteoroid entering Jupiter's atmosphere.

Detection of a Bolide in Jupiter's Atmosphere With Juno

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**Abstract** The Ultraviolet Spectrograph (UVS) instrument on the Juno mission recorded transient

bright emission from a point source in Jupiter's atmosphere. The spectrum shows that the emission is

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# 1. Introduction

As the largest and most massive planet in the Solar System, Jupiter undergoes a heavy bombardment of objects, ranging from tiny dust grains to kilometer-sized comets. The largest impacts, such as comet Shoemaker-Levy 9 (Asphaug & Benz, 1996) and the 2009 impactor (Sánchez-Lavega et al., 2010), occur rarely but leave scars on the planet that can persist for several months (Sánchez-Lavega et al., 1998, 2011) and can affect the distribution of trace species in Jupiter's upper atmosphere for decades (Benmahi et al., 2020; Cavalié et al., 2013; Lellouch et al., 2006). Impacts from objects in the 5-20 m diameter range occur more frequently and produce short bright flashes of light that can be observed by amateur astronomers on Earth but no observable debris (Hueso et al., 2018). Over a period of 8 years (2010-2017), amateur astronomers observed five of these smaller impacts, leading to an estimated impact rate of 10-65 per year (Hueso et al., 2018), an estimate that was not significantly modified by the observation of a sixth impact in 2019 (Sankar et al., 2020). On the smallest scales, the constant influx of dust particles equates to hundreds of thousands of tons of material per year (Poppe, 2016; Sremčević et al., 2005).

While larger (>5 m) impacts can be observed from Earth, observations from orbiting spacecraft have the advantage of being able to detect the fainter flashes that are produced from the more frequent smaller impacts. One example of this is the small fireball observed by the camera on the Voyager 1 spacecraft, which Cook and Duxbury (1981) estimated was caused by an 11 kg (<0.5 m) meteoroid. In this paper, we present observations of another fireball in Jupiter's atmosphere, this time observed by the Ultraviolet Spectrograph (UVS) instrument on the Juno spacecraft. In Section 2, we describe the instrument and the method of observation. In Section 3, we discuss the properties of the bright blackbody-emission flash observed. In Section 4, we conclude that this flash was caused by a meteoroid of mass 250–5,000 kg (diameter 1–4 m) entering Jupiter's atmosphere, and we use the sum of all of our observations to estimate an impact flux rate.

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**Figure 1.** A spatial map of Juno UVS measurements during a single spacecraft spin on April 10, 2020. Photons with wavelengths 130–160 nm are shown in green and photons with wavelengths 170–200 nm are shown in red. The inset image magnifies the area around the bright spot. UVS, Ultraviolet Spectrograph.

# 2. Observations

#### 2.1. Juno UVS

The UVS is an instrument on the Juno spacecraft, which has been in orbit around Jupiter since July 2016 (Bolton et al., 2017). UVS is a photon-counting imaging spectrograph, covering wavelengths of 68–210 nm in the far-ultraviolet (Gladstone, Persyn, et al., 2017). The main purpose of UVS is to study the morphology, brightness, and spectral characteristics of Jupiter's auroras, and the instrument's spectral range covers important H and  $H_2$  auroral emissions (Bonfond et al., 2017; Gladstone, Versteeg, et al., 2017).

Juno is a spin-stabilized spacecraft with a rotation period of 30 s; as the spacecraft rotates, the UVS instrument slit sweeps across Jupiter. The wavelength, slit position and precise time of UV photon detection events are recorded, and this information is then used to build up spatial maps of the ultraviolet radiation (see Figure 1 for an example from a single spin). The UVS instrument slit consists of two wide segments on either side of a narrow segment. The wide parts of the slit have a width of  $0.2^{\circ}$  and a spectral resolution of 2.0-3.0 nm, while the narrow part of the slit has a width of  $\sim 0.025^{\circ}$  and a spectral resolution of  $\sim 1.3$  nm (Greathouse et al., 2013). The spatial resolution along the slit is  $0.1-0.3^{\circ}$ . The observations described in this paper were made at the top edge of the upper wide slit, where the spatial resolution is  $\sim 0.3^{\circ}$ .

#### 2.2. Bright Flash

On occasion, the Juno UVS instrument has observed short-lived, localized ultraviolet emission outside of the auroral zone. In Giles et al. (2020), we described a set of 11 bright transient flashes that were observed with Juno UVS and shared similar characteristics; they lasted ~1.4 ms, they were dominated by  $H_2$  emission and the source region was located ~260 km above the 1-bar level. Based on these characteristics, we

concluded that these were likely to be elves, sprites or sprite halos, forms of Transient Luminous Events (TLEs) that occur in the upper atmosphere in response to tropospheric lightning.

In this paper, we focus on another transient bright flash observed by UVS, but this observation has very different characteristics than the aforementioned TLEs. The observation was made on April 10, 2020 at 12:57:10 UTC and the bright spot was located at a planetocentric latitude of  $53^{\circ}$ N and a System III longitude of 200°W. It was observed at an emission angle of 43° and a solar incidence angle of 125° (i.e., on the planet's nightside). At the time of the observation, the spacecraft was at an altitude of 80,000 km above the 1-bar level of the planet and the subspacecraft latitude and longitude were  $63^{\circ}$ N and  $243^{\circ}$ W, respectively. This location was not observed by Juno's infrared instrument (JIRAM, Adriani et al., 2008) or the visible light camera (JunoCam, Hansen et al., 2017) during this time period or afterward. Figure 1 shows the UVS map from the spin in which the bright spot was recorded and includes a magnified image of the spot itself. This magnified image has dimensions  $3 \times 3^{\circ}$  on the sky, with each pixel being  $0.1 \times 0.1^{\circ}$ . The spot was only observed during a single spin; it was not seen two spins earlier or three spins later, which were the closest times when the same latitude and longitude was observed.

The colors used in Figure 1 show the number of photon detections made in two different spectral regions. Green is used to represent photons with wavelengths 130–160 nm, and red is used to represent photons with wavelengths 170–200 nm. The UVS swath includes a segment of Jupiter's northern auroral oval and the auroral emission appears purely green in the image; this is because the main auroral H<sub>2</sub> emission bands are at <170 nm. In contrast, the bright spot appears mostly yellow, indicating that there is significant emission at longer wavelengths as well.

Although this bright spot is located close to the northern auroral region, the long-wavelength emission marks it out as unique when compared with auroral features. It is also unique compared to the transient bright flashes described in Giles et al. (2020), which were all dominated by  $H_2$  emission, like the aurora. Despite these different spectral characteristics, the bright spot does clearly originate from within Jupiter's atmosphere; Figure 1 shows that it is located far from the limb of the planet so there is no possibility of confusion with a star close to Jupiter, and the bright spot spectrum shows  $CH_4$  absorption from Jupiter's atmosphere (see Section 3.2) so it cannot originate from an object located between the spacecraft and the planet.

We conducted a search for other such events, by filtering for occasions when there was a sharp increase and then decrease in the number of long-wavelength photon counts recorded, but no similar events were found. The spot is analyzed further in Section 3, and the results are discussed in Section 4.

#### 3. Results

#### 3.1. Spatial Extent and Duration

Because of the way in which the UVS maps are built up as the spacecraft spins, Figure 1 contains information about both the spatial extent and the duration of the bright spot. By fitting a two-dimensional Gaussian to the total photon counts, we find that the spot has a FWHM of 0.29° in the along slit direction and 0.23° in the across slit direction. This is consistent with the shape that we would expect for a point source that has a constant brightness on the 17-ms timescale it takes the slit to pass over the source (Greathouse et al., 2013). Based on the distance of the spacecraft at the time of the observations (80,000 km), the upper limit for the spot diameter is 400 km.

Because the shape of the bright spot is consistent with the FWHM of a point source, the bright spot must also be approximately stationary in latitude and longitude over the 17-ms duration of the observation. In order to observe motion, the spot would have had to move  $>0.1^{\circ}$  in 17 ms, equating to a horizontal velocity of  $>8,000 \text{ kms}^{-1}$ . Since we do not observe any elongation, any horizontal movement must be below this speed.

While the brightness appears approximately constant on the timescale, it takes the slit to pass over the source, the bright spot was only observed during a single spin, so it must be transient on slightly longer timescales. UVS data was recorded from the same latitude and longitude two spins (60 s) earlier and three spins (90 s) later, and the bright spot was not present in either of these maps. This places an upper bound of 150 s on the duration of the bright spot, giving a duration of between 17 ms and 150 s.





**Figure 2.** The bright spot spectrum (black) compared with various model spectra. (a) The colored lines show the shape of blackbodies with different temperatures. There is no atmospheric absorption. (b) The colored lines show the spectral shape of a 9,000 K blackbody placed at different altitudes and therefore with different amounts of atmospheric absorption. (c) The colored lines show two different best-fit model spectra.

Because the bright spot was observed on the night side of the planet, it is not possible to use UVS data to search for any local atmospheric changes in its aftermath, such as changes in the aerosol abundance. Any such changes could therefore have lasted longer than 150 s.

#### 3.2. Spectral Analysis

The bright spot spectrum was calculated by summing all photon detections within a  $1 \times 1^{\circ}$  box centered on the bright spot. The spectrum is shown by the black data points in Figure 2 and is presented in terms of spectral irradiance measured at the spacecraft. As suggested earlier by the color of the bright spot in Figure 1, the bright spot emission has an unusual spectral shape compared to both the auroral emission and the TLE emission described in Giles et al. (2020). The auroral and TLE spectra are dominated by H<sub>2</sub> emission, and the H<sub>2</sub> Lyman band can be recognized by a double peak at 160 nm. Beyond 170 nm, the H<sub>2</sub> emission drops to low levels. In contrast, the spectrum shown in black in Figure 2 increases smoothly between 150 and 190 nm, suggesting blackbody emission. The rapid drop-off at wavelengths shorter than 140 nm suggests strong atmospheric absorption by  $CH_4$ .

In order to model a blackbody spectrum attenuated by atmospheric absorption, we used the atmospheric composition Model C from Moses et al. (2005). In this model atmosphere, the  $CH_4$  homopause is located at a pressure of  $2.9 \times 10^{-4}$  mbar, or 360 km above the 1 bar level. For a given altitude and emission angle, we used the atmospheric composition model to calculate the column density for each atmospheric gas. The three stratospheric gases that contribute significantly to absorption in the 125–210 nm range are  $CH_4$ ,  $C_2H_2$ , and  $C_2H_6$ . Absorption cross-sections for these three gases were obtained from Chen and Wu (2004), Lee et al. (2001), Smith et al. (1991), and Bénilan et al. (2000). The atmospheric transmission was then calculated using the column density for each gas and the absorption cross-sections, and this was multiplied by a blackbody spectrum to obtain a modeled top-of-atmosphere spectrum. The colored lines in Figures 2a and 2b show these modeled spectra for a range of blackbody temperatures and emission altitudes.

Figure 2a shows modeled spectra for blackbodies with five different temperatures. For these spectra, no atmospheric absorption is included and the spectra have been scaled to match the average irradiance of the bright spot spectrum. Within the 125–210 nm spectral range, increasing the temperature from 7,000 K to 12,000 K flattens the spectrum. At shorter wavelengths, the data diverges sharply from all of the blackbody spectra due to the lack of atmospheric absorption. We note that the data also diverges from the blackbody spectra at ~195 nm. As demonstrated by the larger error bars, the UVS sensitivity is much lower at these longer wavelengths than at shorter wavelengths; by 195 nm, the instrument's effective area is an order of magnitude smaller than at 160 nm (Hue et al., 2018). We therefore attribute this apparent small dip in the irradiance to limitations in the radiometric calibration.

Figure 2b shows modeled spectra obtained using different emission altitudes (i.e., different amounts of atmospheric absorption). Each spectrum uses the same 9,000 K blackbody as the source emission and the altitude is defined as the height above the 1-bar level. The red lines in Figures 2a and 2b are identical, as there is negligible absorption above an altitude of 400 km. As the source is moved deeper in the atmosphere, the amount of absorption increases. By 250 km, there is significant  $CH_4$  absorption shortwards of 140 nm. By 200 km, there is also a significant  $C_2H_2$  feature at 152 nm. By 150 km, the atmosphere is opaque shortwards of 155 nm due to the combination of  $C_2H_2$ ,  $C_2H_6$ , and  $CH_4$ .

When the blackbody temperature, emission altitude and a scaling factor were all allowed to vary together, the best fit was obtained with a temperature of 9,600  $\pm$  600 K and an altitude of 225  $\pm$  5 km above the 1-bar level (assuming an error of 20% on the abundances of the absorbers in the model). This best-fit spectrum is shown by the red line in Figure 2c. While this modeled spectrum provides a good fit at most wavelengths, the C<sub>2</sub>H<sub>2</sub> absorption feature at 152 nm is clearly too strong; an altitude of 225 km provides a good fit for the CH<sub>4</sub> and C<sub>2</sub>H<sub>6</sub> column densities, but the C<sub>2</sub>H<sub>2</sub> column density (2.1 × 10<sup>16</sup> cm<sup>-2</sup>) is too high. Instead, if we allow the C<sub>2</sub>H<sub>2</sub> column density to vary independently of the other gases, we retrieve a value of 1.1 × 10<sup>16</sup> cm<sup>-2</sup>, approximately half of the value at 225 km. The spectrum obtained with this value is shown by the blue line in Figure 2c. Compared to the Moses et al. (2005) model atmosphere, this suggests that the atmosphere near the bright spot has lower C<sub>2</sub>H<sub>2</sub> volume mixing ratio. This is unsurprizing as the Moses et al. (2005) model describes equatorial latitudes, and Nixon et al. (2007) used Cassini CIRS observations to show that the C<sub>2</sub>H<sub>2</sub> volume mixing ratio decreases at higher latitudes. The blue line in Figure 2c is therefore consistent with an altitude of 225 km at a latitude of 53°N.

# 4. Discussion

The bright spot described in Section 3 has the following properties:

- 1. Its duration is longer than 17 ms and less than 150 s
- 2. The shape is consistent with a point source, which equates to a maximum diameter of 400 km
- 3. Its spectral shape is consistent with a blackbody of temperature 9,600 K located at an altitude of 225 km above the 1-bar level (a pressure of 0.04 mbar)

In Giles et al. (2020), we presented several instances of transient bright flashes observed in the Juno UVS data, and we concluded that these were consistent with TLEs known as elves, sprites, or sprite halos, upper





**Figure 3.** A comparison of the time series and spectra of the bolide (this paper) and the TLEs (Giles et al., 2020). (a) and (b) show the number of observed UV photons as a function of time, relative to the peak event time. The time series shown in (b) is for the TLE observed on April 10, 2020 at 17:24:35 UTC. (c) shows the same spectrum presented in Figure 2, and (d) shows the averaged TLE spectrum from Giles et al. (2020). TLE, Transient Luminous Event.

atmospheric responses to tropospheric lightning. The bright spot that we present in this paper differs from the previous bright flashes in two main ways, highlighted in Figure 3. First, it has a significantly longer duration. The TLEs had a duration of ~1.4 ms, which is short enough that UVS could observe their decay during the time it took the instrument slit to pass across the point source. In contrast, this bright spot is approximately constant as the slit passes across it, giving it a minimum duration of ~17 ms. This difference in duration is shown in Figures 3a and 3b; the Gaussian shape in Figure 3a is governed by the width of the slit, while the rapid exponential decay of Figure 3b takes place on a much shorter timescale. Second, the spectra of the TLEs were dominated by H<sub>2</sub> Lyman band emission, giving them a very similar spectral shape to auroral emission. In contrast, this bright spot has a blackbody-like emission spectrum. The spectra are compared in Figures 3c and 3d. As discussed in Section 3, the shape of Figure 3c is consistent with blackbody emission, while the double peak at 160 nm seen in Figure 3d is indicative of H<sub>2</sub> emission.

Because of both the duration and the spectral shape, the bright spot described in this paper is unlikely to be a TLE. Lightning and associated TLEs do not have blackbody spectra; on earth, they have emission spectra that are dominated by nitrogen and oxygen lines (Rodger, 1999; Walker & Christian, 2017) and on Jupiter their emission spectra are dominated by hydrogen (Borucki et al., 1996; Yair et al., 2009). Elves have submillisecond timescales, and while sprites can last for several tens of milliseconds, the brighter ones are often shorter (Lyons, 2006). There are other longer-lasting TLEs, such as blue jets, but these would still be  $H_2$  dominated and they emerge directly from the thunderstorm anvils (Rodger, 1999), so we would expect them to be deeper in the atmosphere. The shape of the spectrum also rules out the possibility that this is a transient auroral event, as auroral emission is  $H_2$  dominated.

Given the spectrum and the duration, we suggest that the bright spot we observe may be a bolide/fireball in Jupiter's atmosphere. Borovička and Spurnỳ (1996) studied the lightcurves and spectra of two of the brightest bolides observed in the Earth's atmosphere and found that the spectra were a mixture of blackbody continuum and emission lines. The emission lines come from ablated meteoroid material and the blackbody continuum comes from thermal bremsstrahlung (free-free) emission. In the beginning phase of the bolide, Borovička and Spurnỳ (1996) found that the line emission dominated, but during the brightest part of the lightcurve, the blackbody continuum dominated. A temperature of 9,600 K is approximately consistent with the Planck temperatures measured in Earth bolides. Borovička (1994) found that fireball spectra in the Earth's atmosphere exhibit two distinct characteristic temperatures: a primary component with a brightness temperature of 4,000 K and a secondary component with a brightness temperature of 10,000 K. The lower temperature is thought to be from the thermal radiation of the meteoroid itself, while the second part is associated with the shock wave. The greater the speed of the meteoroid, the more dominant this second part is. We do not observe any movement of the bright spot during our 17-ms observation timescale, and this is consistent with a typical impact velocity of tens of kms<sup>-1</sup> (Crawford et al., 1994); a horizontal velocity of >8,000 kms<sup>-1</sup> would be required to be detectable by UVS.

Our observations are somewhat similar to the observations made by the UVS instrument on the Galileo spacecraft when comet Shoemaker-Levy 9 collided with Jupiter in 1994 (Hord et al., 1995). They observed a brightening during one spatial scan only (no brightening in scans obtained 5 s beforehand or 5 s afterward) and when combined with simultaneous observations from the Photopolarimeter Radiometer, they concluded that the brightness temperature was  $7800^{+500}_{-600}$  K. Our observations are also similar to several observations made by amateur observers of bolides on Jupiter (Hueso et al., 2018). These observations were recorded in the visible and the observed bright flashes lasted 1–2 s in each case. One of these events was simultaneously recorded in two different filters and fitting a blackbody to these two measurements gave a temperature of 6,500–8,500 K (Hueso et al., 2013).

Hueso et al. (2010) used the Earth-based observation of a Jovian bolide to estimate the size of the impacting object and we follow their analysis approach here. If we correct for atmospheric absorption, extend the blackbody shape to all wavelengths and integrate, the total irradiance observed at the spacecraft would be  $1.2-2.3 \times 10^{-10}$  Wcm<sup>-2</sup> (taking into account the 600 K error on the brightness temperature). Assuming isotropic radiation, the total power emitted is  $1.0-1.9 \times 10^{11}$  W. As we only observe the bright flash for a very short period of time as the spacecraft spins, we do not know the total duration of the emission and this adds significant uncertainty to the calculation. Here, we assume a duration of 1-2 s, based on Hueso et al. (2018). There is additional uncertainty from the fact that we do not know which stage of the light curve we are observing, it could be the peak of the emission, or the peak could occur shortly before/after our observations. Including a 50% uncertainty for this factor, we obtain a total optical energy of  $5 \times 10^{10}$ –6 ×  $10^{11}$  J.

Brown et al. (2002) found that the efficiency,  $\mu$ , with which the kinetic energy of an impactor is converted into optical energy can be empirically described by

$$u = 0.121 \times E_0^{0.115} \tag{1}$$

where  $E_0$  is the optical energy in terms of kilotons of TNT (1 kton =  $4.185 \times 10^{12}$  J). For our observed total optical energy range, the efficiency is 7%–10% and the kinetic energy of the impactor is therefore  $7 \times 10^{11}$ –  $6 \times 10^{12}$  J. Assuming an impactor velocity of 60 kms<sup>-1</sup>, the velocity at which fragments of Shoemaker-Levy 9 collided with Jupiter (Crawford et al., 1994), and a velocity error of 20% leads to a mass estimate of 250–5,000 kg. For densities of of 250–2,000 kgm<sup>-3</sup>, this equates to a diameter of 1–4 m.

By considering the total power emitted and the blackbody temperature, we can use the Stefan-Boltzmann law to calculate the area of the emitting region. Assuming an emissivity of 1, the effective diameter of the emitting region is 9 m. This is consistent with an impactor diameter of 1-4 m; Borovička and Spurný (1996) found that the maximum diameter of the radiating region is on the order of 10 times larger than the initial diameter of the body.

A mass estimate of 250–5,000 kg seems broadly consistent with the altitude at which we observe the bright flash: 225 km above the 1-bar level, or a pressure of 0.04 mbar. Numerical impact simulations have been used to estimate the altitudes of bolides in Jupiter's atmosphere, and have found that impactors with masses of  $\sim 1 \times 10^6$  kg reach their peak brightness at 60–150 km above the 1 bar level (Hueso et al., 2013; Sankar et al., 2020). Smaller impactors, such as our observation, would not penetrate as deeply into the atmosphere. Borovička and Spurný (1996) studied the entry of a 5,000-kg impactor in the Earth's atmosphere, equivalent to the upper end of our mass estimate. They found that the peak brightness occurred at an altitude of 67 km, which corresponds to a pressure of 0.07 mbar (ISO, 1975). This is slightly deeper in the atmosphere than our observed bright flash pressure level of 0.04 mbar, which is consistent with 5,000 kg being the upper bound of our mass estimate.



By considering all observations made with Juno UVS over the first 27 perijoves of the mission, we find that UVS obtained a total effective coverage of  $8.2 \times 10^{13}$  km<sup>2</sup>s. In that time, there was a single bolide observation. There are clear limitations in calculating an impact flux rate from a single observation, but a maximum likelihood estimation calculation leads to an impact rate of ~24,000 per year (See Text S1 in the supporting information). By considering how impact rates vary as a function of meteoroid size, we can compare this impact rate with the rates deduced from previous studies. We find that our rate is lower than the flux rate estimated by Cook and Duxbury (1981) and higher than the flux rate of Hueso et al. (2018) (See Text S2 in the supporting information). This comparison is dependent on the assumed mass of our observed meteor, and the lower limit of 250 kg is significantly more consistent with Hueso et al. (2018) than the upper limit of 5,000 kg. We note that there are considerable uncertainties in our estimated rate, as we have thus far only observed a single bolide with Juno UVS.

### 5. Conclusions

Juno UVS observations recorded transient blackbody emission from a point source in Jupiter's atmosphere. Spectral modeling showed that the emission is consistent with a 9,600 K source located 225 km above the 1-bar level, and the emission lasted between 17 ms and 150 s. The blackbody nature of the spectrum, the temperature and the duration of the emission are all consistent with a bolide in Jupiter's atmosphere. Based on the energy emitted, we estimate that the impactor had a mass of 250–5,000 kg, which would correspond to a diameter of 1–4 m. This impactor size is larger than the small fireball observed by Cook and Duxbury (1981) and smaller than the superbolides described by (Hueso et al., 2018). We estimate an impact flux rate of 24,000 per year, for masses of 250–5,000 kg or greater; if we use the lower limit of our mass estimate, this is consistent with the rate determined by Hueso et al. (2018). However, we note that there are large uncertainties associated with this value due to the fact that we have thus far only observed one bolide with Juno UVS. As the Juno mission continues, our integration time on the planet will continue to steadily increase and our estimate of the impact flux rate will improve.

# Data Availability Statement

The Juno UVS data used in this paper are archived in NASA's Planetary Data System Atmospheres Node: https://pds-atmospheres.nmsu.edu/PDS/data/jnouvs\_3001 (Trantham, 2014). The data used to produce the figures in this paper are available in Giles (2021).

# References

- Adriani, A., Coradini, A., Filacchione, G., Lunine, J. I., Bini, A., Pasqui, C., et al. (2008). JIRAM, the image spectrometer in the near infrared on board the Juno mission to Jupiter. *Astrobiology*, *8*, 613–622. https://doi.org/10.1089/ast.2007.0167
- Asphaug, E., & Benz, W. (1996). Size, density, and structure of comet Shoemaker–Levy 9 inferred from the physics of tidal breakup. *Icarus*, 121, 225–248. https://doi.org/10.1006/icar.1996.0083
  - Bénilan, Y., Smith, N., Jolly, A., & Raulin, F. (2000). The long wavelength range temperature variations of the mid-UV acetylene absorption coefficient. Planetary and Space Science, 48(5), 463–471. https://doi.org/10.1016/S0032-0633(00)00019-2
  - Benmahi, B., Cavalié, T., Dobrijevic, M., Biver, N., Bermudez-Diaz, K., Sandqvist, A., et al. (2020). Monitoring of the evolution of H<sub>2</sub>O vapor in the stratosphere of Jupiter over an 18-yr period with the Odin space telescope. Astronomy & Astrophysics, 641, A140. https://doi.org/10.1051/0004-6361/202038188
  - Bolton, S. J., Lunine, J., Stevenson, D., Connerney, J. E. P., Levin, S., Owen, T. C., et al. (2017). The Juno Mission. Space Science Reviews, 213, 5–37. https://doi.org/10.1007/s11214-017-0429-6
- Bonfond, B., Gladstone, G. R., Grodent, D., Greathouse, T. K., Versteeg, M. H., Hue, V., et al. (2017). Morphology of the UV aurorae Jupiter during Juno's first perijove observations. *Geophysical Research Letters*, 44(10), 4463–4471. https://doi.org/10.1002/2017GL073114
- Borovička, J. (1994). Two components in meteor spectra. Planetary and Space Science, 42(2), 145–150. https://doi. org/10.1016/0032-0633(94)90025-6

Borovička, J., & Spurný, P. (1996). Radiation study of two very bright terrestrial bolides and an application to the comet S–L 9 collision with Jupiter. *Icarus*, 121(2), 484–510. https://doi.org/10.1006/icar.1996.0104

- Borucki, W. J., McKay, C. P., Jebens, D., Lakkaraju, H. S., & Vanajakshi, C. T. (1996). Spectral irradiance measurements of simulated lightning in planetary atmospheres. *Icarus*, 123(2), 336–344. https://doi.org/10.1006/icar.1996.0162
- Brown, P., Spalding, R. E., ReVelle, D. O., Tagliaferri, E., & Worden, S. P. (2002). The flux of small near-earth objects colliding with the earth. *Nature*, 420(6913), 294–296. https://doi.org/10.1038/nature01238
- Cavalié, T., Feuchtgruber, H., Lellouch, E., de Val-Borro, M., Jarchow, C., Moreno, R., et al. (2013). Spatial distribution of water in the stratosphere of Jupiter from Herschel HIFI and PACS observations. *Astronomy & Astrophysics*, 553. A21. https://doi.org/10.1051/0004-6361/201220797

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- Chen, F., & Wu, C. R. (2004). Temperature-dependent photoabsorption cross sections in the VUV-UV region. I. Methane and ethane. Journal of Quantitative Spectroscopy and Radiative Transfer, 85(2), 195–209. https://doi.org/10.1016/S0022-4073(03)00225-5
- Cook, A. F., & Duxbury, T. C. (1981). A fireball in Jupiter's atmosphere. *Journal of Geophysical Research*, 86(A10), 8815–8817. https://doi.org/10.1029/JA086iA10p08815
- Crawford, D. A., Boslough, M. B., Trucano, T. G., & Robinson, A. C. (1994). The impact of comet Shoemaker-Levy 9 on Jupiter. Shock Waves, 4(1), 47-50. https://doi.org/10.1007/BF01414632

Giles, R. S. (2021). Detection of a Bolide in Jupiter's Atmosphere with Juno UVS. Mendeley Data, V1. https://doi.org/10.17632/nhc5r85xv3.1

- Giles, R. S., Greathouse, T. K., Bonfond, B., Gladstone, G. R., Kammer, J. A., Hue, V., Levin, et al. (2020). Possible Transient Luminous Events observed in Jupiter's upper atmosphere. *Journal of Geophysical Research: Planets*, 125(11), e2020JE006659. https://doi.org/10.1029/2020JE006659
- Gladstone, G. R., Persyn, S. C., Eterno, J. S., Walther, B. C., Slater, D. C., Davis, M. W., et al. (2017). The ultraviolet spectrograph on NASA's Juno mission. Space Science Reviews, 213(1-4), 447–473. https://doi.org/10.1007/s11214-014-0040-z
- Gladstone, G. R., Versteeg, M. H., Greathouse, T. K., Hue, V., Davis, M. W., Gérard, J.-C., et al. (2017). Juno-UVS approach observations of Jupiter's auroras. *Geophysical Research Letters*, 44(15), 7668–7675. https://doi.org/10.1002/2017GL073377
- Greathouse, T. K., Gladstone, G. R., Davis, M. W., Slater, D. C., Versteeg, M. H., Persson, K. B., et al. (2013). Performance results from in-flight commissioning of the Juno Ultraviolet Spectrograph (Juno-UVS). In UV, X-ray, and gamma-ray Space instrumentation for astronomy XVIII, 8859, 88590T. https://doi.org/10.1117/12.2024537
- Hansen, C. J., Caplinger, M. A., Ingersoll, A., Ravine, M. A., Jensen, E., Bolton, S., & Orton, G. (2017). Junocam: Juno's outreach camera. Space Science Reviews, 213(1-4), 475–506. https://doi.org/10.1007/s11214-014-0079-x
- Hord, C. W., Pryor, W. R., Stewart, A. I. F., Simmons, K. E., Gebben, J. J., Barth, C. A., et al. (1995). Direct observations of the comet Shoemaker-Levy 9 fragment G impact by Galileo UVS. *Geophysical Research Letters*, 22(12), 1565–1568. https://doi.org/10.1029/95GL01414
- Hue, V., Kammer, J., Gladstone, G. R., Greathouse, T. K., Davis, M. W., Bonfond, B., et al. (2018). In-flight characterization and calibration of the Juno-Ultraviolet Spectrograph (Juno-UVS). In Space Telescopes and Instrumentation 2018: Ultraviolet to Gamma Ray, SPIE 10699, (p. 1069931). https://doi.org/10.1117/12.2311563
- Hueso, R., Delcroix, M., Sánchez-Lavega, A., Pedranghelu, S., Kernbauer, G., McKeon, J., et al. (2018). Small impacts on the giant planet Jupiter. *Astronomy & Astrophysics*, 617, A68. https://doi.org/10.1051/0004-6361/201832689
- Hueso, R., Pérez-Hoyos, S., Sánchez-Lavega, A., Wesley, A., Hall, G., Go, C., et al. (2013). Impact flux on Jupiter: From superbolides to large-scale collisions. Astronomy & Astrophysics, 560, A55. https://doi.org/10.1051/0004-6361/201322216
- Hueso, R., Wesley, A., Go, C., Pérez-Hoyos, S., Wong, M., Fletcher, L., et al. (2010). First Earth-based detection of a superbolide on Jupiter. *The Astrophysical Journal Letters*, 721(2), L129. https://doi.org/10.1088/2041-8205/721/2/L129
- ISO (1975). ISO 2533:1975 Standard atmosphere. International Organization for Standardization.
- Lee, A. Y., Yung, Y. L., Cheng, B.-M., Bahou, M., Chung, C.-Y., & Lee, Y.-P. (2001). Enhancement of deuterated ethane on Jupiter. The Astrophysical Journal Letters, 551(1), L93. https://doi.org/10.1086/319827
- Lellouch, E., Bézard, B., Strobel, D. F., Bjoraker, G. L., Flasar, F. M., & Romani, P. N. (2006). On the HCN and CO<sub>2</sub> abundance and distribution in Jupiter's stratosphere. *Icarus*, 184(2), 478–497. https://doi.org/10.1016/j.icarus.2006.05.018
- Lyons, W. A. (2006). The meteorology of Transient Luminous Events An introduction and overview. Sprites, elves and intense lightning discharges (pp. 19–56). Springer.
- Moses, J. I., Fouchet, T., Bézard, B., Gladstone, G. R., Lellouch, E., & Feuchtgruber, H. (2005). Photochemistry and diffusion in Jupiter's stratosphere: Constraints from ISO observations and comparisons with other giant planets. *Journal of Geophysical Research*, 110(E8), E08001. https://doi.org/10.1029/2005JE002411
- Nixon, C. A., Achterberg, R. K., Conrath, B. J., Irwin, P. G. J., Teanby, N. A., Fouchet, T., et al. (2007). Meridional variations of C<sub>2</sub>H<sub>2</sub> and C<sub>2</sub>H<sub>6</sub> in Jupiter's atmosphere from Cassini CIRS infrared spectra. *Icarus*, *188*(1), 47–71. https://doi.org/10.1016/j.icarus.2006.11.016
- Poppe, A. R. (2016). An improved model for interplanetary dust fluxes in the outer solar system. *Icarus*, 264, 369–386. https://doi.org/10.1016/j.icarus.2015.10.001
- Rodger, C. J. (1999). Red sprites, upward lightning, and VLF perturbations. *Reviews of Geophysics*, 37(3), 317–336. https://doi.org/10.1029/1999RG900006
- Sánchez-Lavega, A., Gómez, J. M., Rojas, J. F., Acarreta, J. R., Lecacheux, J., Colas, F., et al. (1998). Long-term evolution of comet SL-9 impact features: July 1994–September 1996. *Icarus*, 131(2), 341–357. https://doi.org/10.1006/icar.1997.5871
- Sánchez-Lavega, A., Orton, G. S., Hueso, R., Pérez-Hoyos, S., Fletcher, L. N., García-Melendo, E., et al. (2011). Long-term evolution of the aerosol debris cloud produced by the 2009 impact on Jupiter. *Icarus*, 214(2), 462–476. https://doi.org/10.1016/j.icarus.2011.03.015
- Sánchez-Lavega, A., Wesley, A., Orton, G., Hueso, R., Perez-Hoyos, S., Fletcher, L., et al. (2010). The impact of a large object on Jupiter in 2009 July. *The Astrophysical Journal Letters*, 715(2). L155. https://doi.org/10.1088/2041-8205/715/2/L155
- Sankar, R., Palotai, C., Hueso, R., Delcroix, M., Chappel, E., & Sánchez-Lavega, A. (2020). Fragmentation modeling of the 2019 August impact on Jupiter. *Monthly Notices of the Royal Astronomical Society*, 493(4), 4622–4630. https://doi.org/10.1093/mnras/staa563
- Smith, P. L., Yoshino, K., Parkinson, W., Ito, K., & Stark, G. (1991). High-resolution, VUV (147–201 nm) photoabsorption cross sections for C<sub>2</sub>H<sub>2</sub> at 195 and 295 K. *Journal of Geophysical Research*, *96*(E2), 17529–17533. https://doi.org/10.1029/91JE01739
- Sremčević, M., Krivov, A. V., Krüger, H., & Spahn, F. (2005). Impact-generated dust clouds around planetary satellites: Model versus Galileo data. *Planetary and Space Science*, 53(6), 625–641. https://doi.org/10.1016/j.pss.2004.10.001
- Trantham, B. (2014). JNO-J-UVS-3-RDR-V1.0. NASA planetary data System.
- Walker, T. D., & Christian, H. J. (2017). Triggered lightning spectroscopy: Part 1. A qualitative analysis. Journal of Geophysical Research: Atmosphere, 122(15), 8000–8011. https://doi.org/10.1002/2016JD026419
- Yair, Y., Takahashi, Y., Yaniv, R., Ebert, U., & Goto, Y. (2009). A study of the possibility of sprites in the atmospheres of other planets. Journal of Geophysical Research, 114(E9), E09002. https://doi.org/10.1029/2008JE003311