1D geological imaging of the subsurface from geophysical data with Bayesian Evidential Learning.

<u>Authors¹:</u>

Hadrien MICHEL^{1,2,3}, Frédéric NGUYEN¹, Thomas KREMER¹, Ann ELEN⁴, Thomas HERMANS²

(1) University of Liège, Urban and Environmental Engineering Department, Faculty of Applied Sciences, Liège, Belgium, (2) Ghent University, Department of Geology, Ghent, Belgium, (3) F.R.S.-FNRS (Fonds de la Recherche Scientifique), Brussels, Belgium, (4) KU Leuven, Department of Earth and Environmental Sciences, Leuven, Belgium

Corresponding author:

Hadrien MICHEL

University of Liège

Allée de la découverte, 9

B-4000, Liège (Belgium)

E-mail: <u>hadrien.michel@uliege.be</u>

Declaration of interest: none

Link to the code: https://github.com/hadrienmichel/BEL1D

Highlights:

- A new approach for 1D imaging from geophysical data.
- A set of MATLAB graphical user interfaces is provided.
- Estimation of the uncertainty on the model parameters.
- Consistent results for SNMR.
- BEL1D can be generalized to any type of 1D geophysical data.

¹ <u>Authorship statement:</u> Hadrien Michel developed the BEL1D codes and produced the results presented in this paper. Thomas Kremer has significantly contributed and improved all the SNMR related elements developed in this paper. The Mont Rigi SNMR dataset was acquired and pre-processed by Ann Elen for her Master Thesis. Frédéric Nguyen and Thomas Hermans both co-supervised the whole work and are at the basis of methodological development.

Abstract:

Imaging the subsurface of the Earth is of prime concern in geosciences. In this scope, geophysics offers a wide range of methods that are able to produce models of the subsurface, classically through inversion processes. Deterministic inversions lack the ability to produce satisfactory quantifications of uncertainty, whereas stochastic inversions are often computationally demanding. In this paper, a new method to interpret geophysical data is proposed in order to produce 1D imaging of the subsurface along with the uncertainty on the associated parameters. This new approach called Bayesian Evidential Learning 1D imaging (BEL1D) relies on the constitution of statistics-based relationships between simulated data and associated model parameters. The method is applied to surface nuclear magnetic resonance for both a numerical example and field data. The obtained results are compared to the solutions provided by other approaches for the interpretation of these datasets, to demonstrate the robustness of BEL1D. Although this contribution demonstrates the framework for surface nuclear magnetic resonance geophysical data, it is not restricted to this type of data but can be applied to any 1D inverse problem.

<u>Keywords:</u> 1D imaging, uncertainty, Bayesian Evidential Learning, machine learning, SNMR, MATLAB.

1 1. Introduction

2 Imaging the subsurface of the Earth through geophysical methods allows reducing 3 costly invasive investigations or is sometimes the only way to investigate the Earth. Behind the 4 geophysical images are models conceptualized in terms of physical properties (density, seismic 5 velocities or resistivity) which may be converted to properties of interests (hydraulic 6 conductivity, porosity or temperature) using petrophysical relationships. Decision-making, 7 whether planning additional investigations for a scientific project, exploiting a geo-resource, or 8 constructing a building, is often a driver for geophysical methods (Lawyer et al., 2001). 9 Uncertainty assessment on those geophysical models is therefore a key concern to 10 geophysicists, as it provides a much more consistent information in decision-making processes 11 (Scheidt et al., 2018).

12 Computation methods to process geophysical data into models (other than simple 13 mapping of the data) are either deterministic, often using the regularization approach, or 14 stochastic, often using the Bayesian framework. Deterministic methods are limited in terms of 15 uncertainty quantification and often rely on linear data error propagation based on a linearized 16 approximation if needed (Aster et al., 2013; Kemna et al., 2007). On the other hand, stochastic 17 methods may compute sets of models based on numerous inversions of perturbed datasets, or 18 on the exploration of a prior model space using Markov chain Monte Carlos (McMC) methods 19 (Aster et al., 2013; Sambridge, 2002).

Recently, due to the rise of computer performances, stochastic methods have been applied to geophysics and are able to quantify more properly uncertainty. Most algorithms rely on Markov chains Monte Carlos (McMC) methods in order to explore the prior model space (Sambridge, 2002). Those methods offer the advantage to provide a much more complete appraisal of the ensemble of possible solutions. Nevertheless, stochastic methods are often

3

25 linked to high computation costs, since they require a large number of forward model runs to 26 sample the posterior distribution in the prior model space, explaining why they are still often 27 limited to scientific or numerical studies (e.g., Irving and Singha, 2010; Trainor-Guitton and 28 Hoversten, 2011). Approximation of the posterior can be obtained (under some assumptions) 29 through ensemble-based inversion (e.g., Bobe et al., 2019). Even more recently, machine 30 learning also emerged as a potential contender to replace the inversion process, typically 31 yielding to deterministic-like results due to the nature of neural networks and such (e.g., Laloy 32 et al., 2019, 2017; Yang and Ma, 2019).

In this paper, we present Bayesian Evidential Learning (BEL) (Scheidt et al., 2018) as an innovative method for the unidimensional geophysical imaging of the subsurface. Bayesian Evidential Learning (BEL) is a general framework to handle data in geosciences, from the optimization of the data acquisition to the data interpretation, prior falsification, global sensitivity analysis, and the prediction of relevant model responses (Scheidt et al., 2018).

38 We will refer to the approach presented here as Bayesian Evidential Learning 1D 39 imaging (BEL1D). It uses statistics-based relationships between forecast variables and data 40 learned from realizations sampled from a prior distribution. As opposed to inversion methods, 41 this approach does not require the time-consuming inversion of the dataset, but rather numerous 42 (and parallelized) runs of the faster forward model. So far, BEL has been used to predict the 43 dynamical response of subsurface models from well monitoring data (Satija and Caers, 2015), 44 push-pull tests (Hermans et al., 2019) or time-lapse geophysical experiments (Hermans et al., 45 2018), i.e. model predictions mostly characterized by smooth variations in time and/or in space. BEL was also applied to show the non-uniqueness of the gravity inverse problem (Phelps et al., 46 47 2018).

In this contribution, we extend BEL to 1D geophysical imaging. Such 1D imaging arises
in many geophysical methods where forward models assume a succession of homogeneous

50	horizontal layers such as surface nuclear magnetic resonance (SNMR e.g., Behroozmand et al.,
51	2015), surface seismic waves (e.g., Socco et al., 2010), electromagnetic surveys (e.g., Hanssens
52	et al., 2019; Li et al., 2018), or vertical electrical sounding (e.g., Jha et al., 2008).
53	The main contributions of this paper compared to previous works are:
54	1) The application of BEL1D for static geophysical imaging. Here, we predict directly
55	the posterior distribution of spatially distributed geophysical model parameters.
56	2) The estimation of the data-prediction relationship using kernel density estimation
57	instead of Gaussian regression, because the former can be applied when the
58	relationship deviates from linearity and Gaussian assumptions.
59	3) The validation of the approach by comparing it with state-of-the-art McMC
60	4) The link made between distance-based global sensitivity analysis and the canonical
61	correlation analysis, yielding similar results.
62	5) The analysis of the number of samples in the prior.
63	This paper first introduces our implementation of BEL: Bayesian Evidential Learning
64	1D imaging (BEL1D). The method is then validated within a numerical example. To
65	demonstrate the broad applicability of our approach, we also provide a validation within a non-
66	geophysical related inverse problem (supplementary material: "Testing BEL1D on a synthetic
67	mind experiment: oscillations of a pendulum"). Then, we present an example of BEL1D applied
68	to field data, providing evidence that the method is already mature. Finally, we discuss on
69	elements governing the accuracy of BEL1D, such as the number of models sampled in the prior
70	and the choice of the bandwidth for kernel density estimation.

71 2. BEL for 1D imaging: methodology

72 In this work, the focus is on computing geophysical models and assessing the 73 uncertainty of the inferred models (Phelps et al., 2018). This interpretation part of BEL is also 74 known as the prediction-focused approach (PFA), introduced by Scheidt et al. (2015). Contrary 75 to deterministic approaches, this method does not rely on the stabilization of the ill-posed 76 inverse problem, such as regularization for example, that imposes non-realistic constraint to the 77 solution. It rather relies on the constitution of statistical relationships between predictors (the 78 set of parameters of interest to the end-user) and data. These relationships are originating from 79 models and (numerical) simulations of the geophysical problem that reflect the available prior knowledge. 80

BEL has already been applied previously to geophysical data – time-lapse electrical
resistivity – by Hermans et al. (2016). Hermans et al. (2016) demonstrated the applicability of
BEL to predict variations of subsurface physical properties with time-lapse electrical resistivity
data. More broadly, they also demonstrated that BEL was a possible tool for geophysicists.

85 However, Hermans et al. (2016) presented a scheme that required numerical simulations 86 of groundwater flow and transport and petrophysical relationships specific to hydrogeophysical 87 monitoring (i.e. time-lapse data), limiting the extend of the geophysical prior model. Solving 88 the time-lapse problem in geophysics is easier because we can invert for changes in the model 89 (e.g., Kemna et al., 2002; Nguyen et al., 2016). In Hermans et al., (2016), a petrophysical 90 relationship and the background distribution of resistivity were taken as known, which 91 simplified the inference of the posterior distribution. Static data are also characterized by a 92 higher noise level and thus more challenging to process (LaBrecque et al., 1996; Lesparre et 93 al., 2017). This is also the first application where BEL is used to directly estimate spatially 94 distributed model parameters. In our case, the forecast variables are static geophysical models,

95	directly linked to the geophysical experiment, hence no petrophysical relationship is required.		
96	Such situations generally bear more prior uncertainty and wider variability. Moreover, our aim		
97	is to propose a fast generation of subsurface models along with uncertainties. The adapted		
98	BEL1D method consists of six steps (Figure 1):		
99	1. Definition of the prior uncertainty based on prior field knowledge, generation of		
100	prior models (yellow box in Figure 1) and associated data (blue box in Figure 1)		
101	by forward modeling		
102	2. Reduction of the dimensionality of the data using principal component analysis		
103	(PCA)		
104	3. Constitution of statistical relationships between the model parameters and the		
105	reduced data (green cloud in Figure 1), using canonical correlation analysis		
106	(CCA)		
107	4. Generation of posterior distributions for the model parameters in reduced space		
108	by constraining the bivariate distributions to field data (red box in Figure 1)		
109	using kernel density estimators (contrary to previously used Gaussian process		
110	regression (Hermans et al., 2016))		
111	5. Sampling of the constituted (non-Gaussian) distributions		
112	6. Back-transformation of the samples into the original space, delivering a set of		
113	1D models of the subsurface that are constrained to the knowledge of the		
114	geophysical data (purple box in Figure 1).		
115	Each step is explained in detail in the following sections.		





Figure 1: Illustration of the BEL1D process. See text for abbreviation.

118 Although BEL1D can be applied to any 1D inverse problem, for the clarity of the paper, 119 we briefly introduce SNMR, the specific method chosen to illustrate the BEL1D method. 120 SNMR is a geophysical technique dedicated to the detection and characterization of aquifers, 121 whose main advantage and particularity is that the measured signal (an exponential decay) is 122 directly linked to the amount of water present in the subsurface. For more details about the 123 nature and origin of the SNMR signal, we refer to Behroozmand et al. (2015). The main goal 124 of SNMR experiments is to retrieve the distribution of the water content and the relaxation time 125 with depth. The latter depends on the pore space geometry and can be related to hydrodynamic 126 parameters such as the porosity or the permeability. The acquisition of SNMR data on the field 127 consists in the injection of an alternative current in a transmission loop (typically circular or in 128 an eight-shape) directly followed by the measurement of the induced current in the reception 129 loop (often the same as the transmission loop). The injection is tuned at the resonance frequency of water (called the Larmor frequency), perturbing the quantum state of protons contained in water molecules. Once the injection is stopped, the protons are relaxed back to their original state, inducing a current with a decreasing amplitude with time in the receiver loop. This response is called the free-induction decay (FID) and constitutes the collected data. The injection is repeated with different pulse moments, sounding different zones of the subsurface.

135

2.1. Step 1: prior realizations and associated data

136 In BEL1D, the models are described in a classic manner with a finite number of layers 137 N_L , the last one being a half-space. For each layer, the N_P physical properties related to the 138 geophysical experiment and layer thickness are assigned. This means that a model can be described with a low number of parameters: $(N_L \times (N_P + 1)) - 1 = q$. Defining the prior 139 model space $f_H(\boldsymbol{m})$, i.e. the space describing the prior knowledge of the subsurface physical 140 141 properties before data acquisition, consists in assigning a distribution to the values of each of 142 those q parameters. The shapes of those distributions are unconstrained and must best represent 143 the prior knowledge of the survey site, either originating from previous experiments or from 144 general geological and geophysical considerations. Once the prior model space is defined, we 145 generate random models within the prior boundaries, hereafter called prior realization: 146 $\boldsymbol{m}(q \times 1) \in f_H(\boldsymbol{m}).$

For example, the prior model space for SNMR experiments is defined using both the water content (in %) and relaxation time (in ms) distributions along depth. This leads to a finite number of layers, of unknown thickness, and of unknown water content and relaxation time, thus to the description of a model with only five parameters for a 2-layer model.

151 For all prior realizations, we compute their associated response using the same 152 acquisition parameters as the experimental conditions. Doing so, we produce the synthetic data 153 vector: d ($l \times 1$). To perform this operation, we use the (non-linear) forward operators (K) of 154 the geophysical problem of interest such that

$$\boldsymbol{d} = \boldsymbol{K}(\boldsymbol{m}) \tag{1}$$

For more details about the forward operator associated with the SNMR method, we refer to Hertrich et al. (2007). Often, the dimensionality of the data (l) is relatively large. For SNMR, the data space is composed of tens of time-decaying signals (one for each pulse moment) with sampling rates on the order of 100 Hz measured for about 1 seconds, thus resulting in several thousands of dimensions.

160 2.2. Step 2: reduction of the dimensionality of the data using principal component161 analysis

162 At this stage, we have generated the "model space" which contains the prior realizations, 163 and the "data space" which contains the synthetic geophysical data sets associated with each 164 model. To explore the statistical relationships that exist between these two spaces, one must 165 first ensure that their dimensionality is sufficiently low (typically about 10 dimensions per space) to allow for reasonable computational costs. In the case of SNMR, as in most other cases 166 167 if the prior models space remains simple, the model space will generally have a reasonable number of dimensions, thus no dimension reduction is required ($m^f = m (q \times 1)$). As stated 168 169 previously, the data space is generally large and cannot be handled properly under reasonable 170 computational costs. In this case, the Principle Component Analysis (PCA) technique can be 171 used to transform the data space and reduce its dimensionality while preserving its original 172 variability.

The PCA method seeks for linear combinations of variables that maximize the variability in the first dimensions using eigenvalue decomposition (Krzanowski, 2000). This operation results in PCA scores, the values of the transformed data. This way, the first dimensions are the most informative about the dataset, and the others only represents a small 177 amount of the variability. Those later can be discarded, to gain memory and ease following 178 computations. In the SNMR context, the same process is applied to the original data space and 179 performs very well since the dimensions are often highly correlated. The obtained reduced data space has a dimensionality $k \ll l$ and is noted d^{f} (k \times 1). Note that in other geophysical 180 181 problems, this PCA step may not be necessary if the dimensions of both the model and data 182 space are low or may also be necessary for both the model and data spaces if their dimensions 183 are too high. Noise analysis using Monte Carlo simulations (Hermans et al., 2016) provides covariance between the different PCA components representing the data (C^{f}) . This translates 184 185 in uncertainties on the PCA scores of the test data. The detailed procedure for noise propagation 186 is presented in subsection 2.7.

187 2.3. Step 3: constitution of statistical relationships between the models parameters188 and the reduced data, using canonical correlation analysis

189 BEL1D relies on the constitution of statistics-based relationships between earth models 190 and simulated data. This step is the core of BEL1D. Several methods are suitable to derive such 191 a relationship, but we choose canonical correlation analysis (CCA) for its simplicity. In essence, CCA transforms the (possibly PCA-reduced) model space (m^{f}) and the PCA data space (d^{f}). 192 193 so that the canonical correlation (an approach formula of the correlation) between the resulting CCA spaces (m^c – Equation 2 – and d^c – Equation 3) is maximized (Krzanowski, 2000). Each 194 195 of the dimensions of the CCA spaces are orthogonal, meaning that there is no redundancy 196 between the dimensions. We obtain:

$$d^c = d^f A^T \tag{2}$$

$$\boldsymbol{m}^{c} = \boldsymbol{m}^{f} \boldsymbol{B}^{T} \tag{3}$$

where **A** is a matrix of dimensions $m \times k$ and **B** is a matrix of dimensions $m \times q$ (*m* is the minimum between *k* and *q*). Graphically, we can observe the statistical relationship between d^c and m^c as both spaces have the same dimensions $(1 \times m)$. As long as *k* is larger than *q*, it is possible to back-transform the models from the CCA model space to the original model space
using the inverse of *B* (Hermans et al., 2016).

202 2.4. Step 4: generation of posterior distributions for the model parameters in CCA
 203 space using kernel density estimators

204 In canonically correlated space, the correlation between dimensions is maximized. 205 Hence, it is possible to produce a meaningful statistical description of the bi-variate distribution. 206 This enables to determine a posterior distribution in the CCA model space constrained by the 207 knowledge of field data. This posterior distribution represents the probability density function 208 for the model parameters in the CCA model space. To approximate the posterior distribution corresponding to the observed data $f_H(\boldsymbol{m}^c | \boldsymbol{d}_{obs}^c)$, we use a kernel density estimator (KDE) with 209 210 a Gaussian kernel (Wand and Jones, 1993) similar to Hermans et al. (2019) where they faced 211 non-linearity issues when validating BEL for hydrogeology. Gaussian regression (Hermans et 212 al., 2016; Satija and Caers, 2015) was tested; but in many cases, CCA did not yield perfectly 213 linearly correlated relationships with Gaussian prior distributions, two necessary conditions to 214 apply Gaussian regression. We thus implement KDE for its robustness to estimate posterior 215 distribution in a large variety of situations. KDE computes, for a given location in space, the 216 sum of the contribution of each point from the CCA space. We use a multi-Gaussian kernel 217 centered on the points with bandwidths chosen accordingly to the point density (see subsection 218 5.2). The resulting distributions in the CCA model space are not constrained to any given 219 distribution of known shape. The process is illustrated in the center of Figure 1.

220

2.5. Step 5: sampling of the constituted distributions

The posterior distribution is not constrained to a known distribution. Therefore, sampling is done through the inverse transform sampling method (Devroye, 1986). This sampling procedure benefits from the properties of the cumulative distribution function that links a uniformly distributed variable to any distribution. The sampler thus transforms a value sampled in a uniform distribution to the corresponding value in the original distribution via the inverse of the cumulative distribution function. We can now easily generate a set of samples from the posterior distributions in the reduced space:

$$\boldsymbol{m_{post}^c} \in f_H(\boldsymbol{m^c}|\boldsymbol{d_{obs}^c}) \tag{4}$$

228 2.6. Step 6: back-transformation of the samples into original space

The set of samples in CCA model space can be back-transformed into the original model space, using the inverse of the *B* matrix ($q \times q$). We thus have the posterior in the original physical space:

$$\boldsymbol{m_{post}} = \boldsymbol{m_{post}}^f = \boldsymbol{B}^{-1} \boldsymbol{m_{post}}^c \tag{5}$$

Here, in the SNMR context, since the model space was not reduced using principle component analysis, the models in PCA space are the same as the ones in original space. Because Gaussian regression or KDE with Gaussian kernel has no limit on the value of sampled parameters, a few samples might occasionally be located outside the prior model space, in particular when the model is at the extremity of the prior. Those models are removed from the sampled space as they do not correspond to the definition of the problem.

238 2.7. Noise propagation in BEL1D

In order to investigate the effect of noise on BEL1D, one can analyze the variation of the PCA scores to the estimated field noise level (if the latter can be estimated). If the scores of the first *k* components are not varying in a significant way, then, one can assume that noise has no impact on the imaging process. On the other hand, if the noise impacts the scores significantly, BEL1D considers this impact by propagating uncertainty (represented by C^f in Equation 6) on the reduced observed data in the canonically correlated space, leading potentially to more complex computation of the posterior distributions (Hermans et al., 2016). This uncertainty is propagated in the CCA dimension according to linear error propagation fromequation (2):

$$C^c = A C^f A^T \tag{6}$$

To estimate C^{f} , we explore the PCA scores and their modifications when noise is present in the dataset. To do so, we take the data associated to each of the prior realizations as a basis and compute the PCA scores when one model is perturbed with a noise level similar to the data. The operation is repeated for a significant sub-sample (in the latter example, we used 50 randomly selected models) of the prior samples allowing to derive the PCA score covariance matrix C^{f} .

254 Contrary to Gaussian regression where the addition of covariance to the observed 255 reduced data only slightly increases the complexity of the formula for the sampling parameters 256 (Hermans et al., 2016), the addition of uncertainty on the reduced observed data highly hinders 257 performances of the Kernel density estimation. Assigning uncertainty to the observed data (X in 258 Figure 2) requires the computation of the kernel density estimation of the reduced posterior distributions for multiple sets of randomly sampled $d_{obs}^{c} + \epsilon$, then processing to sampling and 259 260 back transformation from those sets of distributions to constitute models of the posterior distributions. This can be avoided by enlarging the bandwidth of the kernel density estimator 261 (Bowman and Azzalini, 1997) in the X direction (in the CCA data space) according to C^{c} . This 262 263 accounts for the uncertainty on the exact position of the data in the reduced space and therefore 264 account for noise, without adding complexity to the computations, as is shown in Figure 2.



Figure 2 : Example of the effect of noise on the computation of the kernel density estimator for the probability density functions. On the left, a random dataset that is highly correlated and the associated true value of X. On the right, the estimated distributions for the values of Y constrained to the known value of X and its associated error of 1. The computation time t is given for the different cases.

270 3. Numerical example

In this section we will present results obtained by applying BEL1D to numerical examples of SNMR. The model that we will use corresponds to a classical context with two layers: a first unsaturated layer and a saturated half-space (Table 1). A more complex case is presented in section 4, and a non-geophysical related inverse problem is presented in supplementary material ("Testing BEL1D on a synthetic mind experiment: oscillations of a pendulum").

	Thickness (m)	Water content (%)	Relaxation time (ms)
Layer 1	5	5	100
Half-space	/	25	200

277

Table 1: Characteristics of the synthetic model

3.1. Computation of the model space and data space

279 The data are simulated for SNMR using the MRSMatlab toolbox (Müller-Petke et al.,
280 2016). For the simulated experiment, the circular transmitter is the same as the receiver (the

281 loops have a 50 m diameter with 2 turns). To simplify the computations, the earth is set as 282 resistive, neglecting the effect of electrical conductivity in signal attenuation and phase shift. 283 Therefore, the signal is purely real. However, the process presented in this paper is not limited 284 to real-valued data. SNMR is known to be impacted significantly by electromagnetic (EM) 285 noise (e.g., Behroozmand et al., 2016). Different types of EM noise can affect the SNMR data. 286 However, when most of these components have been filtered out, the SNMR signal will remain 287 contaminated by a Gaussian noise distribution, whose magnitude cannot be reduced further. To 288 represent this, we chose to add a Gaussian noise level to the dataset, corresponding to a mean 289 signal-to-noise ratio of 10 (35 nV), which is a relatively high value in usual SNMR surveys.



290

291 Figure 3 : Simulated data for the SNMR experiment. The black curves represent the noisy dataset and the red
292 curves the noise-free one.

As can be observed in Figure 3, the initial amplitude of the signal is around 500 nV, and signal-to-noise ratio of 10 produces quite noisy signals, but the decaying behavior remains observable.

The prior model space reflects the prior knowledge of the study site. In our case, we will assume that it is known that the subsurface is composed of one layer resting over an infinite half-space. Then, the rest of the properties will be defined as uniformly distributed variables varying in the intervals described in Table 2. The prior intervals are not necessarily centered on the actual value of the parameter. This is to better simulate real situations, where the user will only have rough information about the subsurface constitution. 5000 prior realizations are generated using a Latin-hypercube sampler (McKay et al., 1979). The forward model that was used to simulate the test data is reused here to generate the synthetic data set.

304

	Thickness [m]		Water content [%]		Relaxation time [ms]	
	Minimum	Maximum	Minimum	Maximum	Minimum	Maximum
Layer 1	2.5	7.5	3.5	10	5	350
Half-space	/	/	10	30	5	350

305

Table 2: Prior model space for the SNMR experiment

306 3.2. Global sensitivity analysis and dimension reduction

307 We performed a distance-based global sensitivity analysis (Park et al., 2016) whose

308 results are presented in Figure 4 (Left). The parameters corresponding to the first layer (W_1

and $T_{2,1}^*$) are not sensitive to the SNMR signal in this specific experimental configuration and should therefore lead to a poorly reduced uncertainty on those parameters.



312 Figure 4 : <u>Left:</u> multivariate sensitivity analysis performed on the simulated data (the red bars correspond to 313 parameters that have an effect on the data and the blue ones to parameters that have a negligible effect). <u>Right:</u> 314 combinations of parameters representing the canonically correlated model space dimensions (the values are 315 scaled for visual purposes)



316

Figure 5 : SNMR - Data space versus model space in the canonically correlated space. The blue dots represent
the prior models and data spaces and the orange line represents the exact value of the test data in the reduced
space.

320 From the initial 11.904 dimensions in the dataspace, PCA enabled a reduction to only 5 321 dimensions, while managing to keep more than 90% of the variability. The canonical 322 correlation analysis is shown in Figure 5. It is observed that (1) the correlation between the models and the data is larger for the first 3 CCA dimensions, (2) the correlation is not linear 323 from the 2nd dimension and (3) from the 4th dimension, the points are scattered and not 324 325 correlated. The fact that the last dimensions are scattered (note that each dimension is a linear 326 combination of the model parameters) informs us on the inability of the data to efficiently 327 predict all the parameters simultaneously. This is confirmed by the linear combinations that 328 constitute the CCA model space dimensions (Figure 4, Right): we observe that the three last 329 reduced dimensions are mainly linear combinations of the insensitive values (Figure 4, Left). 330 Note that non-linear statistical techniques (e.g., Gorban et al., 2008; Lawrence, 2012) could 331 potentially lead to better characterization of the data-model relationship.



332 3.3 Results of BEL1D on the numerical example

333

Figure 6 : Results of the BEL1D process on the synthetic noise-free data from the 2 layers model. The obtained
distributions are the posterior distributions estimated from the noisy dataset in red. The sampled prior model
space is represented by the blue distributions. The test values are represented with dashed red lines in the
histogram plots. Results of the McMC algorithm DREAM are presented in yellow.

BEL1D is applied to the simulated noisy dataset. Figure 6 shows the prior and posterior distributions of parameters obtained through BEL1D. We observe that the uncertainty on the half-space parameters is dramatically reduced through BEL1D. On the other hand, the insensitive parameters $T_{2,1}^*$ and W_1 are poorly reduced. This insensitivity to some of the parameters is also clearly identified in the pendulum example (Supplementary Material). As 343 can be observed, the test value of the parameter is always contained in the posterior distribution. 344 The multivariate analysis of the parameters shows that most of the parameters are independent. However, some correlations are observed for parameters that are related to the same layer $(T_{2,2}^*)$ 345 346 and W_2 for example). Figure 7 (Left) presents another representation of the posterior results, 347 where the data RMS of posterior samples are color-coded, and where each model is a set of 348 linked parameters. The presented results show that a large part of the posterior models has a 349 very low RMS error. The trend in the water content set through the prior model space is 350 respected. In contrast, the trend for the relaxation time is not resolved by BEL1D. The analysis 351 of the RMS and the comparison with DREAM (Figure 7, Right) shows that is related to the 352 insensitivity of the data and not to the method itself.



Figure 7 : <u>Left:</u> results of the BEL1D process on the synthetic noisy data from the 2 layers model. Each displayed model is presented with its associated RMS error value. In order to enhance visualisation, each colour is associated to the same number of models. The dashed white lines represent the test values and the dashed black lines represent the extent of the prior model space. The similar graph is also presented for DREAM results (<u>Right</u>).

Finally, we applied the DREAM McMC algorithm (Vrugt, 2016) to the same dataset. We applied a Gaussian likelihood with 15 generations and 10000 chains and a burn-in period of 5000 chains. However, to achieve convergence towards reasonable uncertainties, changes to the definition of the dataset are required. We had to use gate integration, similarly to what is 363 applied in MRSMatlab (Müller-Petke et al., 2016) in order to attribute higher loads to the first 364 part of the decay curves, thus lowering the impact of noise on the estimation. For the presented 365 results, we use 5 gates for each pulse moment. However, we must note that the results of 366 DREAM are highly sensitive to the number of gates: with 1 or 2 gates, nearly no reduction of 367 uncertainty is observed, whereas with a too large number of gates (to the limit with no gate 368 integration) false posterior distributions are observed (the test values are outside of the predicted 369 posterior). The results are presented alongside the BEL1D results in Figure 6. They show that 370 BEL1D tends to have a slightly lower performance than DREAM. This is related to the 371 approximation of the problem in a lower dimension space where the actual data misfit is never 372 computed. Applying a cutoff on the RMS values observed in BEL1D would result in the same 373 distribution (Figure 7).



Figure 8 : Impact of noise on the PCA scores (the first 5, kept after dimensionality reduction), covariance matrix
for the PCA scores impacted by noise (left) and the corresponding covariance in CCA space (right). This first
matrix is estimated through the mean of the covariance matrices observed between the noise-free scores and the
noisy ones.

To estimate the posterior distributions, we needed to account for the impact of noise in the process. The matrix C^f presented in Figure 8 shows very low values for the covariance. However, once transformed into CCA space (C^c , Equation 6), we observe that the effect of noise, even if limited, is still observable. From a default bandwidth of 0.01 in the X direction
for the kernel density estimation, we computed bandwidths ranging from 0.017 to 0.059.

Those increased bandwidths explain the presence of some outliers in the dataspace (Figure 7, Left**Erreur ! Source du renvoi introuvable.**). When observing the CCA space (Figure 5), it is noticeable that the second dimension (mainly representing the decay time of the second layer) shows models that are outside the main distribution but close to the position of the reduced test data. When using the bandwidth that takes into account the impact of noise, those models are englobed into the posterior. Nonetheless, those models correspond to large RMS error and are easily identified (Figure 7).

391 An example with more than two layers is presented in supplementary material392 ("Application of BEL1D to a four-layer model").

393 4. Case study: Mont Rigi (Belgium)

The Mont Rigi is located in the Belgian Fagnes region, in the Eastern part of the country.
This site presents an ideal case study for SNMR as it is remote and far from any electromagnetic
noise sources, due to its natural reserve classification.

397 Geologically, the site is characterized by a metric peat layer on top of a Cambrian 398 bedrock (La Venne formation) known as an aquiclude with a very low water content (Gilson et 399 al., 2017). In contrast, peat is known to present very high water contents with observed total 400 porosities around 90% (Wastiaux, 2008). According to previous GPR exploration (Wastiaux 401 and Schumacker, 2003), the peat layer at the experimental site should have a thickness between 402 2.5 and 4.5 meters. The SNMR response should therefore allow to easily distinguish the two 403 layers with a properly designed experiment. The designed on-site experiment consisted of a 404 classical coincident loops transmitter/receiver couple with a diameter of 20 meters.

405 The raw data have been preprocessed to lower the impact of noise with despiking, 406 harmonic modelling, reference noise cancellation and despiking again (Müller-Petke et al., 407 2016). The signals have been further truncated to 0.2 seconds to decrease the impact of noise, 408 mostly present at large timeframes. The resulting noise amplitude is quite low for SNMR (18 409 nV). Then, the signals have been inverted using the QT Inversion (Müller-Petke and Yaramanci, 410 2010), to constitute a benchmark to compare to the results of BEL1D. We used a smooth-mono 411 model description (smooth models with a single relaxation time for a given depth) and a 412 regularization parameter of 6000 for both water content and relaxation time using the L-curve 413 criteria (Aster et al., 2013). Then, BEL1D was applied with the uniformly distributed prior 414 model space described in Table 3. We sampled 5000 prior realizations and produced 5000 415 models for the posterior.

	Thickness [m]		Water content [%]		Relaxation time [ms]	
	Minimum	Maximum	Minimum	Maximum	Minimum	Maximum
Layer 1	0	7.5	30	80	0	200
Layer 2	/	/	0	15	100	400

416

Table 3: Prior model space for the Mont Rigi SNMR experiment

The CCA space analysis (Figure 9) shows that we could expect a significant reduction of uncertainty for some parameters, as the first three dimensions show high correlations. The first dimension, dominated by the water content of the half-space shows an especially narrow relationship between the reduced data and model space. On the other hand, the last dimension, mostly represented by the water content of the first layer, shows no specific correlation. We should therefore expect a much better reduction for the parameters of the half-space than for the first layer.



424

425 Figure 9: Mont Rigi – Data versus model space in the CCA space. The blue dots represent each prior
426 realizations and the orange line the position of the field data.

The noise level (after denoising) impact on BEL1D was assessed and resulted in
significant changes in bandwidths for the kernel density estimation, between 0.04 and 0.12
instead of the default 0.01 value.



430

431

Figure 10 : SNMR at Mont Rigi - Prior (blue) and posterior (orange) distributions of the parameters.

The analysis of the posterior parameters model space (Figure 10) shows that the reduction of uncertainty from the prior model space to the posterior is significant and produces consistent results, when compared to the geological context and the QT inversion solution (dashed gray lines in Figure 11). The thickness of the peat layer tends to be smaller than the maximum estimate (4.5 m) from to the GPR interpretation (Wastiaux and Schumacker, 2003). The use of a narrower prior taking into account the GPR information would prevent such a behavior. This is illustrated by the decrease of the maximum value of the thickness for the first layer when the water content of the aforementioned layer increases. This behavior arises fromthe non-uniqueness of the SNMR response stated above.

441 In the posterior models produced by BEL1D (Figure 11), a large majority shows a root 442 mean square error lower than the noise level in the dataset, implying that the set of probable 443 models provides an efficient estimation of the site parameters. The RMS errors of the posterior 444 samples show that they all explain the data to a similar level; the uncertainty is thus intrinsic to 445 the data set and the non-unicity of the solution. The non-uniqueness of the solution is easily 446 observable, with two distinct behaviors: one with a median water content and a thick first layer 447 and another with a very high water content but a thinner first layer. Those two cases show 448 similar RMS errors, hence are explaining the data at the same level. The probability of each 449 posterior model is shown in the RMS error colorbar for which each color represents the same 450 number of models. Therefore, it is observed that the low RMS error models also correspond 451 with the most probable ones.



452

Figure 11: SNMR at Mont Rigi - Posterior models distributions with their associated misfit and results of the QT
inversion (dashed gray line). The black line in the RMS error graph represents the level of noise in the original
dataset.

456 5. Discussion

457

5.1. Impact of the number of sampled models in the prior model space

458 The optimum number of models in the prior model space sample is a tradeoff between 459 precision of the posterior models distributions and ease of computation: the larger the number 460 of samples, the larger the memory requirement and the harder the computations. We ran BEL1D 461 on the previous synthetic SNMR dataset with 10, 50, 100, 500, 1000, 5000, 10 000, 50 000 and 462 100 000 models in the prior samples. For each case, we ran BEL1D 10 times and aggregate the 463 results. It is observed that the RMS error distribution tends to stabilize above 1000 models, 464 suggesting that this value is optimal in our case (Figure 12), especially for the minimum and 5th 465 percentile values. This observation is confirmed by the evolution of the parameter distribution 466 characteristics (Figure 12). In those graphs, the mean values of the parameters are normalized 467 with the test values, and the standard deviations are normalized with the standard deviation in 468 the prior model space. We see that the parameter distributions are stabilized above 1000 models, 469 even if the values are not minimal.



471 Figure 12 : Impact of the number of prior samples on the posterior distribution. The top graph presents RMS
472 evolution, while the two bottom graphs show (from left to right) the normalized mean value and the normalized
473 standard deviation.

474 5.2. Impact of the chosen kernel density bandwidth on the posterior distributions

To assess the impact of the kernel density bandwidth on the posterior distribution, we ran 10 times 10 different values ranging from 10⁻⁵ and 1. In each case, 5000 models were sampled in the prior (Latin-hypercube sampler) and 5000 models were predicted, to ensure that no side effect is linked to other parameters. A similar analysis to the one performed in the previous subsection is then performed (Figure 13).

We observe that the bandwidth has a strong impact on the resulting distributions. Too low values result in erratic behaviors of the posterior distributions. On the other end, too large bandwidths result in the smoothing of the posterior distributions, as is observed in the evolution 483 of the standard deviation where this value increases sharply for the sensitive parameters. 484 Therefore, the optimal bandwidth is a value located in-between the two extremes, where the 485 reduction of uncertainty is significant for the sensitive parameters and the other parameters are 486 stable. This occurs around 10^{-2} in our case.

487 In our implementation of the kernel density estimator, the effect of a too small 488 bandwidth is avoided by verifying that enough points (at least 1% of the sampled prior) are present in the direct surrounding of the reduced data ($d_{obs}^{c} \pm 3 \times bandwidth$). To perform the 489 490 test presented in this bandwidth analysis, this safeguarding measure was disabled. In the case 491 the default bandwidth is too small, then, the bandwidth is multiplied by 2 and the verification 492 is made again until the criterion is satisfied and only then, the kernel density estimation is performed. We ran BEL1D with a bandwidth set to 10^{-5} and the safeguard algorithm enabled. 493 494 The obtained result is here similar to the one obtained with a suitable bandwidth as is shown by 495 the diamonds in the graph presenting the evolution of the normalized standard deviation in 496 Figure 13.



Figure 13 : Impact of the bandwidth in the kernel density estimation on the posterior distribution. The top graph
presents RMS evolution, while the two bottom graphs show (from left to right) the normalized mean value and
the normalized standard deviation. This latter also presents results obtained with the safeguarding measure
enabled (diamonds) for the most extreme case: a bandwidth of 1E-5.

502 6. Conclusion

We propose an adaptation of the Bayesian Evidential Learning framework as a new tool for the unidimensional interpretation of geophysical data called BEL1D. This approach offers an alternative to stochastic inversions, as it requires fewer and simpler computations, thus lowering the CPU cost. This approach is also presented as an alternative to deterministic inversions as its behavior is superior both in terms of ease of implementation and completeness of the results (handling properly non-uniqueness, quantifying the uncertainty, etc.). We therefore developed this method theoretically, as a new general algorithm. 510 The framework does not require the use of regularization parameters or any other form 511 of bias, but rather requires the proper definition of the prior model space. Contrary to other 512 Bayesian methods, the use of a large prior model space (hence, introducing less bias in the 513 posterior information) does not involve heavier computations to converge as the algorithm is 514 always building in one step the posterior model space that corresponds to the observed data 515 (see, 2.4 Step 4: generation of posterior distributions for the model parameters in CCA space 516 using kernel density estimators). This method can therefore be considered unbiased if the prior 517 model space is sufficiently large. BEL1D is thus a computationally efficient alternative to 518 stochastic inversions as it only requires *n* forward models corresponding to the prior samples. 519 Moreover, in BEL1D noise is considered via an imbedded theoretically based method, contrary 520 to classically used high regularization parameters that lack of theoretical basis to be properly 521 assessed and are therefore difficult to estimate.

522 When compared to McMC methods (such as DREAM), one of the key advantages of 523 BEL1D is the tractability of all operations. When McMC classically applies random changes to 524 the models and select them randomly, BEL1D constitutes a correlation between models 525 parameters and the data. This means that one could seek for the origin of all the posterior models 526 in the prior model space. This latter enables, for example, to efficiently sample multi-modal 527 distribution or to understand the origin of odd models in the posterior distribution (see the 528 numerical example), impossible to track back in McMC algorithms. The other key advantage 529 is the simpler tuning to convergence of BEL1D compared to DREAM (McMC). In DREAM, 530 it was required to change the dataspace to enable convergence towards reasonable (or even 531 correct) posterior distributions. In BEL1D, as long as the prior is correctly assumed, the only 532 parameter that the user needs to take care of is the number of models to sample in this space. 533 As shown in the discussion, this parameter can be selected by observing the stabilization of the 534 posterior distributions. In DREAM, the application without changes in the dataset representation lead to either nearly no reduction of uncertainty (using the RMS as a proxy of the probability) or to a reduction towards level of uncertainty close to deterministic values but without the true model inside the posterior (using all data points as proxy of the probability). Therefore, applying this method to an unknown dataset might lead to erroneous interpretations, whereas BEL1D, even if producing slightly larger uncertainties should always predict consistent posterior spaces.

541 We demonstrated that BEL1D is efficient for the interpretation of SNMR data. We 542 presented results for noisy datasets originating from a numerical model and field. In both cases, 543 we validate the results from BEL1D, either against the test values from the model and McMC 544 methods - DREAM (Vrugt, 2016) - for the numerical example or other geophysical 545 experiments and classical state-of-the art inversion technics for the field case. On the other end, 546 the careful analysis of the RMS error of the proposed models indicates that the range of 547 uncertainty delivered by BEL1D is intrinsic to the non-unicity of the solution for geophysical 548 inverse problems: many models explain the data to the same level.

549 In term of performance, the computation cost of BEL1D is directly proportional to the 550 efficiency of the forward model and the number of samples in the prior model space. It must be 551 noted that the prior samples being independent, the simulation of synthetic data can be fully 552 parallelized. The performance of BEL1D can thus be easily estimated from the available 553 computing facilities. The cost is therefore much smaller than other stochastic methods requiring 554 ten to hundreds thousands of runs to converge towards a solution. Moreover, computation 555 timing could be dramatically reduced by pre-field forward modelling operations for large 556 geophysical campaigns where the prior would be similar. Prior realizations and the associated 557 datasets would then be reused (PCA and CCA operations also), letting only the kernel density 558 estimation, sampling and back-transformation operations to be applied after field acquisition, leading to extremely rapid imaging. This is an extreme advantage compared to other stochasticmethods that rely on the data misfit (such as McMC) to sample the posterior.

561 This application, along with the pendulum case presented in supplementary material, 562 demonstrate the versatility of BEL1D and its applicability to multiple cases with very few 563 adaptations. Provided that a forward model already exists, the adaptation of the codes to any 564 other 1D geophysical method would be straightforward. For example, the forward operator 565 from HV-Inv (García-Jerez et al., 2016) could be used for the interpretation of H/V data. System 566 calls may also help further extend the applicability of BEL1D. This may be the case for the use 567 of all the 1D forward operators implemented in pyGIMLi (Rücker et al., 2017) - vertical 568 electrical sounding, frequency- and time-domain electromagnetic and magnetotelluric.

In its present version, BEL1D is dedicated to the 1D inversion of geophysical data in layered Earth model where the contrast between layers is supposed to be sharp. In the future, we plan to extend the method to smoothly varying systems accounting for a large number of thin, correlated layers.

573 7. Acknowledgment

574 The authors would like to thank the contributors to the development of MRSmatlab, the 575 cornerstone of SNMR data interpretation in BEL1D. Hadrien Michel is a fellow funded by the 576 F.R.S.-FNRS.

577 8. Computer codes availability

578 The different MATLAB codes that have been used to produce the results presented in 579 this paper are available at <u>github.com/hadrienmichel/BEL1D</u>. They are released with this paper 580 under the BSD-2-Clause license. The codes consist in a series of graphical user interfaces for 581 the interpretation of SNMR data (BEL1DSNMR), among others. All the functions are

582	standalone (no specific toolbox required) in MATLAB but can run faster when the Parallel
583	Computing toolbox is available.
584	The main codes have been developed by H. Michel (ULiège, Belgium):
585	Hadrien MICHEL
586	Urban and Environmental Engineering Department
587	Allée de la découverte, 9
588	B-4000 Liège
589	Belgium
590	The forward models that is used is an Open Source MATLAB toolboxes for SNMR
591	(MRSMatlab: Müller-Petke et al., 2016).
592	In order to run the codes, a MATLAB license is required. Minimum requirements for
593	the computations are the ones for your MATLAB license, at the exception of RAM, for which
594	a minimum of 8 Go is recommended and more could be required for more complex
595	computations. If using the parallel capabilities of MATLAB (parallel computing toolbox), a
596	large number of cores is preferred and, globally, a 4-core (or more) processor is advised.

597 9. References

- Aster, R., Borchers, B., Thurber, C., 2013. Parameters estimation and inverse problems, 2nd
 ed. Academic Press.
- Behroozmand, A.A., Auken, E., Fiandaca, G., Rejkjaer, S., 2016. Increasing the resolution and
 the signal-to-noise ratio of magnetic resonance sounding data using a central loop
 configuration. Geophysical Journal International 205, 243–256.
 https://doi.org/10.1093/gji/ggw004
- Behroozmand, A.A., Keating, K., Auken, E., 2015. A Review of the Principles and Applications
 of the NMR Technique for Near-Surface Characterization. Surveys in Geophysics 36,
 27–85. https://doi.org/10.1007/s10712-014-9304-0
- Bobe, C., Van De Vijver, E., Keller, J., Hanssens, D., Van Meirvenne, M., De Smedt, P., 2019.
 Probabilistic One-Dimensional Inversion of Frequency-Domain Electromagnetic Data Using a Kalman Ensemble Generator. arXiv:1911.03275 [physics].

- Bowman, A.W., Azzalini, A., 1997. Applied Smoothing Techniques for Data Analysis: The
 Kernel Approach with S-Plus Illustrations, Oxford Statistical Science Series. Oxford
 University Press, Oxford, New York.
- 613 Devroye, L., 1986. Non-uniform random variate generation. Springer, New York.
- García-Jerez, A., Piña-Flores, J., Sánchez-Sesma, F.J., Luzón, F., Perton, M., 2016. A computer
 code for forward calculation and inversion of the H/V spectral ratio under the diffuse
 field assumption. Computers & Geosciences 97, 67–78.
 https://doi.org/10.1016/j.cageo.2016.06.016
- 618 Gilson, M., Briers, P., Ruthy, I., Dassargues, A., 2017. Elsenborn Langert Dreiherrenwald.
- Gorban, A.N., Kégl, B., Wunsch, D.C., Zinovyev, A.Y. (Eds.), 2008. Principal Manifolds for
 Data Visualization and Dimension Reduction, Lecture Notes in Computational Science
 and Enginee. Springer Berlin Heidelberg, Berlin, Heidelberg.
 https://doi.org/10.1007/978-3-540-73750-6
- Hanssens, D., Delefortrie, S., Bobe, C., Hermans, T., De Smedt, P., 2019. Improving the
 reliability of soil EC-mapping: Robust apparent electrical conductivity (rECa)
 estimation in ground-based frequency domain electromagnetics. Geoderma 337, 1155–
 1163. https://doi.org/10.1016/j.geoderma.2018.11.030
- Hermans, T., Lesparre, N., De Schepper, G., Robert, T., 2019. Bayesian evidential learning : a
 field validation using push-pull tests. Hydrogeology journal in press.
- Hermans, T., Nguyen, F., Klepikova, M., Dassargues, A., Caers, J., 2018. Uncertainty
 Quantification of Medium-Term Heat Storage From Short-Term Geophysical
 Experiments Using Bayesian Evidential Learning. Water Resources Research 54, 2931–
 2948. https://doi.org/10.1002/2017WR022135
- Hermans, T., Oware, E., Caers, J., 2016. Direct prediction of spatially and temporally varying
 physical properties from time-lapse electrical resistance data. Water Resources
 Research 52, 7262–7283. https://doi.org/10.1002/2016WR019126
- Hertrich, M., Braun, M., Gunther, T., Green, A.G., Yaramanci, U., 2007. Surface Nuclear
 Magnetic Resonance Tomography. IEEE Transactions on Geoscience and Remote
 Sensing 45, 3752–3759. https://doi.org/10.1109/TGRS.2007.903829
- Irving, J., Singha, K., 2010. Stochastic inversion of tracer test and electrical geophysical data
 to estimate hydraulic conductivities. Water Resources Research 46.
 https://doi.org/10.1029/2009WR008340
- Jha, M.K., Kumar, S., Chowdhury, A., 2008. Vertical electrical sounding survey and resistivity
 inversion using genetic algorithm optimization technique. Journal of Hydrology 359,
 71–87. https://doi.org/10.1016/j.jhydrol.2008.06.018
- Kemna, A., Kulessa, B., Vereecken, H., 2002. Imaging and characterisation of subsurface solute
 transport using electrical resistivity tomography (ERT) and equivalent transport models.
 Journal of Hydrology 267, 125–146.
- Kemna, A., Nguyen, F., Gossen, S., 2007. On linear model uncertainty computation in electrical
 imaging. Presented at the SIAM Conferance on Mathematical and Computational Issues
 in Geosciences, Santa Fe.
- Krzanowski, W.J., 2000. Principles of multivariate analysis: A user's perspective, Revised
 Edition. ed, Oxford Statistical Series. Oxford University Press, New York.
- LaBrecque, D.J., Miletto, M., Daily, W., Ramirez, A., Owen, E., 1996. The effects of noise on
 Occam's inversion of resistivity tomography data. Geophysics 61, 538–548.
- Laloy, E., Hérault, R., Lee, J., Jacques, D., Linde, N., 2017. Inversion using a new low-dimensional representation of complex binary geological media based on a deep neural network. Advances in Water Resources 110, 387–405.
 https://doi.org/10.1016/j.advwatres.2017.09.029

- Laloy, E., Linde, N., Ruffino, C., Hérault, R., Gasso, G., Jacques, D., 2019. Gradient-based deterministic inversion of geophysical data with generative adversarial networks: Is it feasible? Computers & Geosciences 133, 104333.
 https://doi.org/10.1016/j.cageo.2019.104333
- Lawrence, N.D., 2012. A Unifying Probabilistic Perspective for Spectral Dimensionality
 Reduction: Insights and New Models 30.
- Lawyer, L.C., Bates, C., Rice, R., 2001. Geophysics in the affairs of mankind: a personalized
 history of exploration geophysics 2nd edition, Society of Exploration Geophysics. ed.
 Tulsa, Oklahoma.
- Lesparre, N., Nguyen, F., Kemna, A., Robert, T., Hermans, T., Daoudi, M., Flores-Orozco, A.,
 2017. A new approach for time-lapse data weighting in electrical resistivity
 tomography. Geophysics 82, E325–E333. https://doi.org/10.1190/geo2017-0024.1
- Li, J., Lu, X., Farquharson, C.G., Hu, X., 2018. A finite-element time-domain forward solver
 for electromagnetic methods with complex-shaped loop sources. GEOPHYSICS 83,
 E117–E132. https://doi.org/10.1190/geo2017-0216.1
- McKay, M.D., Beckman, R.J., Conover, W.J., 1979. A Comparison of Three Methods for
 Selecting Values of Input Variables in the Analysis of Output from a Computer Code.
 Technometrics 21, 239. https://doi.org/10.2307/1268522
- Müller-Petke, M., Braun, M., Hertrich, M., Costabel, S., Walbrecker, J., 2016. MRSmatlab —
 A software tool for processing, modeling, and inversion of magnetic resonance
 sounding data. GEOPHYSICS 81, WB9–WB21. https://doi.org/10.1190/geo20150461.1
- Müller-Petke, M., Yaramanci, U., 2010. QT inversion Comprehensive use of the complete
 surface NMR data set. GEOPHYSICS 75, WA199–WA209.
 https://doi.org/10.1190/1.3471523
- Nguyen, F., Kemna, A., Robert, T., Hermans, T., 2016. Data-driven selection of the minimum gradient support parameter in time-lapse focused electric imaging. GEOPHYSICS 81,
 A1–A5. https://doi.org/10.1190/geo2015-0226.1
- Park, J., Yang, G., Satija, A., Scheidt, C., Caers, J., 2016. DGSA: A Matlab toolbox for
 distance-based generalized sensitivity analysis of geoscientific computer experiments.
 Computers & Geosciences 97, 15–29. https://doi.org/10.1016/j.cageo.2016.08.021
- Phelps, G., Scheidt, C., Caers, J., 2018. Exploring viable geologic interpretations of gravity
 models using distance-based global sensitivity analysis and kernel methods.
 GEOPHYSICS 83, G79–G92. https://doi.org/10.1190/geo2017-0742.1
- Rücker, C., Günther, T., Wagner, F.M., 2017. pyGIMLi: An open-source library for modelling
 and inversion in geophysics. Computers & Geosciences 109, 106–123.
 https://doi.org/10.1016/j.cageo.2017.07.011
- 696 Sambridge, M., 2002. Monte Carlo methods in geophysical inverse problems. Reviews of
 697 Geophysics 40. https://doi.org/10.1029/2000RG000089
- Satija, A., Caers, J., 2015. Direct forecasting of subsurface flow response from non-linear dynamic data by linear least-squares in canonical functional principal component space.
 Advances in Water Resources 77, 69–81.
 https://doi.org/10.1016/j.advwatres.2015.01.002
- Scheidt, C., Li, L., Caers, J., 2018. Quantifying Uncertainty in Subsurface Systems, Wiley Blackwell. ed.
- Scheidt, C., Renard, P., Caers, J., 2015. Prediction-focused subsurface modeling: investigating
 the need for accuracy in flow-based inverse modeling. Mathematical Geosciences 47,
 173–191.

- Socco, L.V., Foti, S., Boiero, D., 2010. Surface-wave analysis for building near-surface velocity
 models Established approaches and new perspectives. GEOPHYSICS 75, 75A83 709 75A102. https://doi.org/10.1190/1.3479491
- Trainor-Guitton, W., Hoversten, G.M., 2011. Stochastic inversion for electromagnetic
 geophysics: Practical challenges and improving convergence efficiency.
 GEOPHYSICS 76, F373–F386. https://doi.org/10.1190/geo2010-0223.1
- Vrugt, J.A., 2016. Markov chain Monte Carlo simulation using the DREAM software package:
 Theory, concepts, and MATLAB implementation. Environmental Modelling &
 Software 75, 273–316. https://doi.org/10.1016/j.envsoft.2015.08.013
- Wand, M.P., Jones, M.C., 1993. Comparison of Smoothing Parameterizations in Bivariate
 Kernel Density Estimation. Journal of the American Statistical Association 88, 520.
 https://doi.org/10.2307/2290332
- Wastiaux, C., 2008. Les tourbières sont-elles des éponges régularisant l'écoulement ? Bulletin
 de la Société géographique de Liège 50.
- Wastiaux, C., Schumacker, R., 2003. Topographie de surface et de subsurface des zones tourbeuses des réserves naturelles domaniales des Hautes-Fagnes (No. C60/1-2-3).
 Université de Liège, Service de phytosociologie, flore et végétation des Hautes-Fagnes.
- Yang, F., Ma, J., 2019. Deep-learning inversion: A next-generation seismic velocity model
 building method. GEOPHYSICS 84, R583–R599. https://doi.org/10.1190/geo20180249.1

727