

Biological controls of the CO₂ air-sea flux in a high-latitude shelf sea

A. E. Friederike Prowe^{1,2&}, Helmuth Thomas^{1,3}, Johannes Pätsch⁴, Wilfried Kühn⁴, Yann Bozec^{3,5}, Laure-Sophie Schiettecatte⁶, Alberto V. Borges⁶

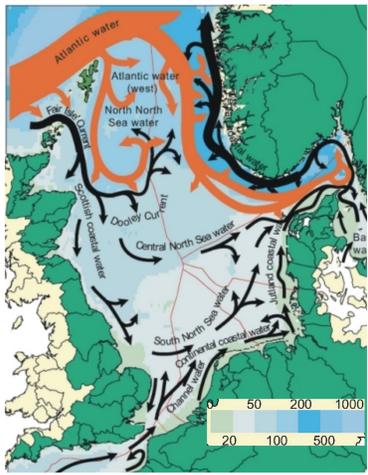


Fig. 1: General circulation and depth [m] of the North Sea [2]. Red arrows indicate water masses with dominant Atlantic signature.

The North Sea is a highly productive shelf sea receiving significant terrestrial nutrient and carbon input from Northwest Europe. Nonetheless, it takes up 1.38 mol C m⁻² yr⁻¹ (8·10⁻³ Pg C yr⁻¹) atmospheric CO₂ per year [1]. Observational studies [1] suggest that this uptake is linked to an oceanic export of carbon into the North Atlantic following the prevailing circulation (Fig. 1). They characterize two different hydrographic regimes in the southern and northern North Sea, respectively, enabling carbon export via the continental shelf pump mechanism. Here, a 3d coupled ecosystem model is used to investigate the controls of the CO₂ air-sea flux in the different regions.

The continental shelf pump (Fig. 2)

southern North Sea:

shallow, mixed water column: biological processes in a single compartment

- primary production: CO₂ uptake by phytoplankton => net CO₂ drawdown from atmosphere
- respiration: CO₂ release by heterotrophic food web => net CO₂ release to atmosphere

primary production and respiration occur in the same compartment of the water column, minor net CO₂ air-sea flux

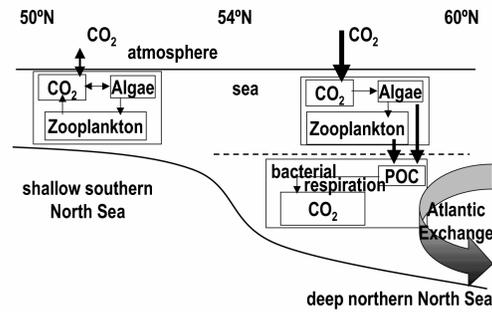


Fig. 2: The continental shelf pump mechanism [1].

Model: ECOHAM3 - 3d coupled physical-biogeochemical ecosystem model (IfM Hamburg), simulating years 2001 and 2002; state variables: DIC, TA, NO₃⁻, NH₄⁺, 1xPhytoplankton, 1xZooplankton, 2xDetritus, Bacteria, DOM, O₂; full C-, N- and O-cycles; resolution: 20km horizontal, 5-25m vertical (about 12 z-levels); fixed C/N ratio for Phyto-, Zooplankton, Bacteria, and related processes except for primary production: C fixation is decoupled from N uptake via simple parameterization allowing C uptake and excretion as semi-labile organic matter under nutrient-limited conditions, e.g. representing photoinhibition.

Observations: Data sets for the full carbon system (DIC, TA, pCO₂) from four North Sea cruises in Aug/Sep 2001, Nov 2001, Feb 2002 and May 2002, covering 90 identical stations.

northern North Sea: deep, seasonally stratified: surface mixed layer (SML) and deeper layer

- primary production: CO₂ uptake in SML, organic matter sinks down into deeper layer
- respiration of organic matter in deeper layer, i.e. not in contact with atmosphere by stratification during productive season => DIC enriched in deeper layer, depleted in SML (Fig. 3). Currents (Fig. 1) remove DIC enriched water into North Atlantic, while CO₂ is drawn down from the atmosphere

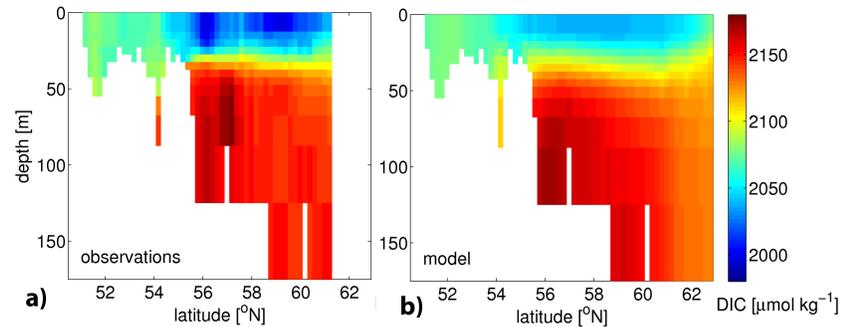


Fig. 3: (a) Observed and (b) simulated DIC [μmol kg⁻¹] along a North-South section at 2°E through the North Sea in Aug 2001.

What controls the CO₂ air-sea flux in the two regimes?

Northern North Sea (Fig. 4a):

water column: deep, seasonally stratified

Primary production: spring bloom supported by stratification (line a); bloom collapses when

NO₃⁻ is depleted (line b); ongoing PP until decrease with mixing in fall (line c)

=> non-Redfield production!

Respiration: increases in spring/fall due to increased PP/mixing; does not balance PP in upper 30 m (SML) during productive season: organic matter sinks out of SML

=> **net community production (NCP = GPP - R) positive** in SML over most of the year:

more CO₂ fixation than CO₂ release.

ΔpCO₂: strong decrease during spring bloom: strongly undersaturated surface pCO₂.

Subsequent increase caused by increasing temperatures counteracting biological uptake.

Mixing in fall brings up DIC rich waters, increasing pCO₂ to supersaturated values.

=> **net CO₂ uptake** (positive air-sea flux), CO₂ release only in winter.

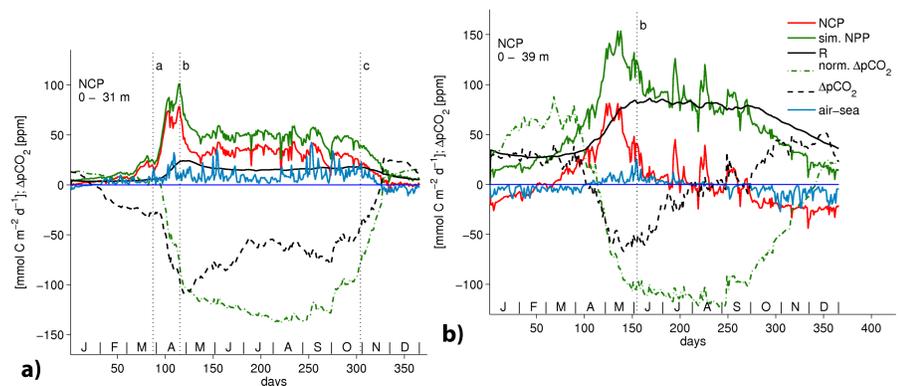


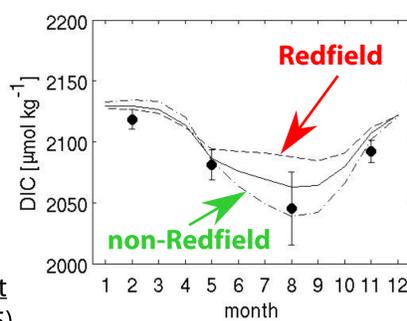
Fig. 4: Simulated net community production (NCP), primary production (PP), respiration (R), ΔpCO₂ (in situ; normalized to temperature at Jan 1, 2001), CO₂ air-sea flux. NCP, PP and R integrated over (a) the total depth at the southern location (Fig. 6a) and (b) 0-30 m at the northern location. Lines: (a) onset of stratification, (b) NO₃⁻ depletion, (c) onset of mixing.

Fig. 5: Average simulated DIC [μmol kg⁻¹] of the upper 30 m for the central North Sea (ICES box 4) in 2001. • observations; -- Redfield production; —, · - non-Redfield production.

non-Redfield production: C fixation under nutrient-depleted conditions!

=> sustains high PP during summer in stratified regime (Fig. 4a)

Redfield production: DIC drawdown in summer not strong enough compared to observations (Fig. 5)



net air-sea flux (Fig. 6):

northern North Sea: net uptake => C transfer from atmosphere via sinking and export to North Atlantic: continental shelf pump!

southern North Sea: net release

model performance: captures DIC and air-sea flux characteristics in both regions; overestimates CO₂ release in

southern Bight: missing sediment resuspension => overestimates DIC; hydrography

off British Coast: hydrography, primary production?

off Norwegian Coast: Alkalinity too low

Summary:

high primary production ≠ atmospheric CO₂ uptake

high primary production ≠ high NCP

NCP governs biological component of CO₂ air-sea flux

stratification, sinking and DIC export into North Atlantic enable net CO₂ uptake

effective C continental shelf pump

C overconsumption needed!

Southern North Sea (Fig. 4b):

water column: shallow, predominantly mixed

Primary production: stronger spring bloom than in northern North Sea; high PP sustained by NH₄⁺ after NO₃⁻ is depleted (line b) => regenerated production;

Respiration: builds up to high level following spring bloom; balances PP in summer

=> **NCP neutral or negative** from summer to winter, **positive only in spring**

ΔpCO₂: undersaturated during spring bloom, subsequently rising temperatures lead to neutral state; supersaturation in winter despite decreasing temperature due to negative NCP

=> positive CO₂ air-sea flux only following spring bloom, **CO₂ release dominates**

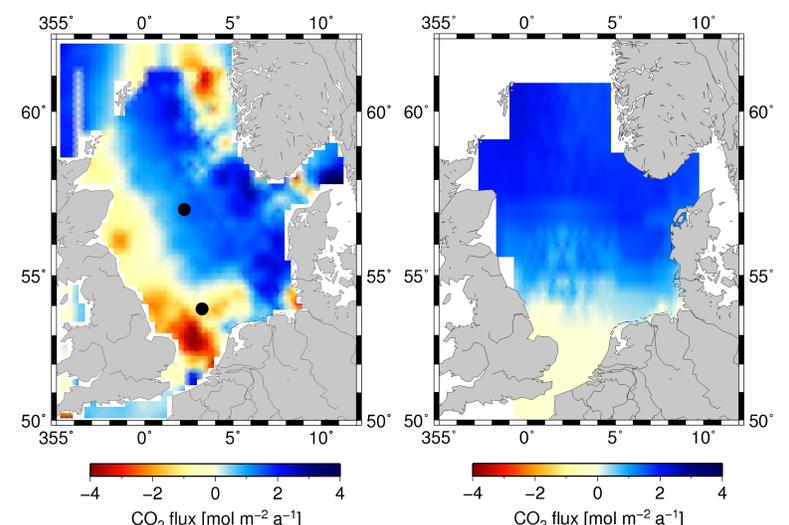


Fig. 6: Annual CO₂ air-sea flux [mol m⁻² yr⁻¹] for 2001 (a) simulated and (b) deduced from observations [1].

References:

[1] Thomas, H. et al. (2004): Enhanced open ocean storage of CO₂ from shelf sea pumping, Science 304, 1005-1008;

[2] mod. from Ospar Commission (2000): Quality Status Report, Region II - Greater North Sea; Turrell, W.R. (1992): New hypotheses concerning the circulation of the northern North Sea and its relation to North Sea fish stock recruitment, ICES Journal of Marine Science 49, 107-123;

¹ Dalhousie University, Department of Oceanography, Halifax, NS, Canada; ² Institut für Chemie und Biologie des Meeres, Carl von Ossietzky Universität Oldenburg, Oldenburg, Germany; ³ Royal Netherlands Institute of Sea Research, Den Burg, Texel, The Netherlands;

⁴ Institut für Meereskunde der Universität Hamburg, Hamburg, Germany; ⁵ now at: Scripps Institution of Oceanography, San Diego, CA, USA; ⁶ Chemical Oceanography Unit, University of Liège, Liège, Belgium

& contact: friederike.prowe@dal.ca