Preliminary Jiram Results from Juno Polar Observations: 1 -Methodology and Analysis Applied to the Jovian Northern Polar Region

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24 Key Points:

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25	• First global maps of H_3^+ intensity, column density and temperature for the Jupiter
26	Northern aurora with high spatial resolution
27	• One side of the auroral oval shows higher H_3^+ column density and lower tempera-
28	tures in comparison with the other side
29	· Column densities main oval and temperature main oval do not superimpose

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30 Abstract

31	During the first orbit around Jupiter of the NASA/JUNO mission, the Jovian Auro-
32	ral InfraRed Mapper (JIRAM) instrument observed the auroral regions with a large num-
33	ber of measurements. The measured spectra show both the emission of the H_3^+ ion and
34	of methane in the 3-4 μ m spectral region. In this paper we describe the analysis method
35	developed to retrieve temperature and Column Density (CD) of the H_3^+ ion from JIRAM
36	spectra in the Northern auroral region. The high spatial resolution of JIRAM shows an
37	asymmetric aurora, with CD and temperature ovals not superimposed and not exactly lo-
38	cated where models and previous observations suggested. On the main oval averaged H_3^+
39	CDs span between 1.8×10^{12} cm ⁻² and 2.8×10^{12} cm ⁻² , while the retrieved temperatures
40	show values between 800 and 950 K. JIRAM indicates a complex relationship among H_3^+
41	CDs and temperatures on the Jupiter Northern aurora.

42 **1 Introduction**

In the Infra-Red (IR) spectral range Jupiter's aurora can be mapped thanks to the 43 thermal emissions of the H₃⁺ molecular ion, first detected by Drossart et al. [1989]. H₃⁺ 44 forms at altitudes mainly above the Jovian homopause. At polar latitudes, accelerated en-45 ergetic electrons that flow downward along magnetic field lines from the magnetosphere 46 drive ionization of both atomic and molecular hydrogen [Atreya, 1986]. The ionised molec-47 ular hydrogen almost instantaneously react with H₂ itself and create H₃⁺ (H₂⁺ + H₂ \rightarrow H₃⁺ 48 + H). This is by far the major creation pathway for H_3^+ in planetary atmospheres; other 49 minor creation routes involving H⁺ and H₂ are present but their effect is negligible with 50 respect to the main process [Grodent et al., 2001]. In the upper atmosphere, where the 51 density of free electrons is large, the H₃⁺ ion is converted back to neutral hydrogen mainly 52 by dissociative recombination (H₃⁺ + e⁻ \rightarrow H₂ + HÂă or H₃⁺ + e⁻ \rightarrow 3 H). At lower al-53 titudes, the main destruction pathway for H₃⁺ is the ion-neutral charge exchange reactions 54 with hydrocarbons (mainly CH₄ and C₂H₂) producing molecular hydrogen and hydrocar-55 bon ions. 56

⁵⁷ The IR spectral range around 3.5 μ m is particularly suitable to study the H₃⁺ emis-⁵⁸ sion as, in this spectral region, methane absorbs most of the light from the lower atmo-⁵⁹ sphere of Jupiter and H₃⁺ lines can be detected with a high contrast with respect to the ⁶⁰ dark planetary disk below. Auroral morphologies mapped through H₂ UltraViolet (UV)

and H₃⁺ IR emissions are very similar on a global scale [e.g. Radioti et al., 2013] and three 61 main components are usually identified [e.g. Clarke et al., 2004; Grodent, 2015]: the main 62 oval, the polar emissions (poleward of the main emission), and the satellites footprints 63 (equatorward of the main emission). However, unlike the auroral UV emissions that are 64 a tracer of instantaneous energy inputs of the impacting electrons, the equatorward IR 65 aurora also provides the information on the atmospheric response to the inputs. The sur-66 rounding neutral atmosphere quickly thermalises the H_{2}^{+} ions after their formation. For 67 this reason, H_2^+ IR emission lines can be used to derive the atmospheric temperature [e.g. 68 Lam et al., 1997; Stallard et al., 2002], while integrated column densities retrieved using 69 the intensities of the emission lines allow mapping the ion distribution. 70

⁷¹ On the NASA mission Juno, orbiting around Jupiter starting in early July 2016, the ⁷² IR spectral range from 3 to 4 μ m was covered by the Jovian Infrared Auroral Mapper, JI-⁷³ RAM [*Adriani et al.*, 2014]. This paper is the first of three papers where we report the ob-⁷⁴ servations of the Jupiter auroras made with the spectrometer of JIRAM: here we describe ⁷⁵ the methodology used to analyse JIRAM spectra of the auroral regions, and we report the ⁷⁶ results obtained for the Jupiter Northern polar region. Further results obtained for both the ⁷⁷ auroral regions are reported in *Adriani et al.* [this issue] and *Moriconi et al.* [this issue].

78 2 JIRAM Instrument and Observations

JIRAM [Adriani et al., 2014] is an imager/spectrometer designed to study the Jovian aurorae, as well as the planet's atmospheric structure, dynamics and composition. It 80 is composed of two IR imager channels (M centered at 4.78 μ m and L centered at 3.45 81 μ m) and by a spectrometer. On the rows of the bi-dimensional spectrometer sensor, the 82 entrance slit (with a Field Of View, FOV, of 3.5°) is covered by 256 pixels. The Instan-83 taneous Field Of View (IFOV) of each pixel is about $250 \times 250 \ \mu$ rad. On the sensor 84 column, the spectrum is sampled in 336 spectral channels in the 2 - 5 μ m range (mean 85 spectral resolution about 9 nm) with integration time of 30s. Thanks to the motion of the 86 spacecraft and the JIRAM de-spinning mirror, each slit of the spectrometer is combined 87 to provide a hyperspectral image (so called "image cube"), being the X-dimension of the 88 cube provided by the slit (corresponding to the cross-track direction), the Y-dimension col-89 lected along-track, and the λ -dimension provided by the spectrometer spectral range. Data 90 acquired during the Moon flyby of October 2013 and during the first perijove demon-9

strated the JIRAM radiometric performances [Adriani et al., 2016]. As already mentioned,

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in this work we focus on the spectrometer data in the 3 - 4 μ m sub-range.

Figure 1 shows a collage of all the analyzed measurements of the northern auroral 94 region. As described in section 3.2 only measurements with emission angle lower than 95 75° have been included in the figure. Panels **a**, **b** and **c** of Figure 1 show Jupiter ortho-96 graphic map in planetocentric coordinates focussed on the North auroral region with su-97 perimposed spots at the geo-location of the intercept of the Line Of Sight (LOS) of each 98 spectrum acquired by JIRAM with the surface located 500 km above Jupiter 1 bar surface. 99 The colour of each dot indicates the value of the represented quantity, while the size of 100 the spot represents its spatial resolution. Panel a shows the emission angle of the measure-101 ments, that is the angle made by the LOS with the vertical to the 500 km altitude surface. 102 It indicates that the measurements have been made observing the same part of Jupiter 103 from different directions, and the measurements with higher spatial resolution are also 104 the ones with lower emission angle. Panel **b** shows the solar incidence angle, that is the 105 Solar Zenith Angle (SZA) at the intercept of the LOS with the 500 km surface. We see 106 that measurements of both the dark and illuminated regions are available. Panel \mathbf{c} shows 107 the integrated intensities of the analysed measurements, obtained integrating the recorded 108 signal over the 3.35-3.75 μ m spectral region (where most of H₃⁺ emission is located and 109 no interferences with other molecules are present) and multiplying it by the cosine of the 110 emission angle, to correct for the slant optical path. The orthographic surface shown in 111 the panels has been divided in squared bins, obtained dividing each axis in regular inter-112 vals. The single intensities have been averaged over each bin, and bins containing less 113 than 3 measurements have not been included in the final dataset. Panel d of Figure 1 rep-114 resents the contour plot of the binned distribution. In all panels, the dashed line represents 115 the geolocation of the auroral oval from existing models [Connerney et al., 1998] and the 116 solid line is the statistical geolocation of the aurora [Bagenal et al., 2014]. In panel \mathbf{d} we 117 identify a region where the signal is maximum (that indicates a strong H_3^+ emission) and 118 that falls close to the two auroral ovals. 119

$_{120}$ 3 H⁺₃ temperature and column density estimation

3.1 Retrieval Code

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Temperature and column density of H_{1}^{+} in the auroral regions have been determined 122 by analysing JIRAM data in the 3.5 micron region. For this purpose, we have used an 123 update of the code previously developed for the analysis of Galileo/NIMS spectra [Al-124 tieri et al., 2016]. Here we recall the basics of this system. The code is divided into two 125 modules: the Forward Model (FM) and the Retrieval Module (RM). The FM is used to 126 simulate the spectra measured by JIRAM. The spectra are simulated by assuming that the 127 emission of the auroral region is optically thin, which is supported by the fact that the H_{2}^{+} 128 layer is located in the highest part of Jupiter's atmosphere. We also assume that the pres-129 sure broadening of the spectral lines is negligible and that we can use the same tempera-130 ture for all the lines of each gas. The latter hypothesis is justified by the fact that most of 131 the auroral emission originates at altitudes where vibrational local thermal equilibrium can 132 be safely assumed [Melin et al., 2005], and may produce a maximum error of 5%. The 133 spectrum is simulated by first computing the intensity of each transition of the gases that 134 we want to include into the simulation. The intensity of the transition k of the molecule m135 can be computed using the expression reported by Altieri et al. [2016] taken from Stallard 136 et al. [2002, and references therein]. 137

The intensities are computed for all the spectral lines of the considered spectral re-138 gion. To reproduce the measurements, the computed intensities are then convolved with 139 the instrumental spectral response. Since part of the JIRAM data were acquired on the 140 dayside, to take into account a variable background emission introduced by the scattered 141 sunlight, we have introduced in the FM the possibility to add a radiometric offset, constant 142 over the whole analysed spectral region, to the simulated spectra. The FM also includes 143 the possibility to evaluate analytically the derivatives of the spectra with respect to temper-144 ature and column density of each gas. 145

The RM is the part that takes care of the determination of the required parameters. We invert the measured spectra using an iterative Bayesian approach. The parameters that can be retrieved with the RM are the column density along the instrument LOS and the effective temperature of each considered gas. Moreover, to account for spectral calibration problems, we can also retrieve a wavelength shift, the width of the instrumental response function, and the radiometric offset value. At each iteration, we evaluate the weighted χ^2

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(χ -test = $n^T S^{-1}n$) and the loop is stopped when two consecutive iterations do not yield values that differ for more than 1% percent.

3.2 Analysis

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The targets of our analysis are the effective temperature of H_3^+ and its column density along the Line Of Sight (LOS) of each observation. Since in recent studies [see *Altieri et al.*, 2016, and reference therein] methane emissions have been detected in the auroral region, along with the H_3^+ data we may simultaneously retrieve temperature and column density of methane along the LOS.

H₃⁺ transitions and spectral properties have been downloaded from the web site 160 http://www.tampa.phys.ucl.ac.uk/ftp/astrodata/H3+/ [Neale et al., 1996] and the parti-16 tion function has been computed using the expression of Miller et al. [2013]. CH₄ spectro-162 scopic data have been taken from the HITRAN 2012 database [Rothman et al., 2013] and 163 its partition function is evaluated using the routine of Gamache described in Laraia et al. 164 [2011]. In this work, JIRAM instrumental response function is assumed to be a Gaus-165 sian function whose width has been evaluated during the on-ground calibration campaign 166 [Adriani et al., 2014], therefore we did not include in the retrieval the corresponding pa-167 rameter. 168

Figure 2 shows two JIRAM spectra in the H_3^+ emission spectral region, one acquired in the sunlit part and one in the dark part of the northern aurora. As can be seen in the Figure, the spectral region below 3.2 μ m is affected by a continuum due to the scattering of sunlight by the lower atmospheric layers, that cannot be corrected using a simple expression. The spectral region above 3.8 μ m is contaminated by both instrumental effects and scattered radiation. Therefore, to avoid systematic errors due to the poor representation of the measured spectrum, we have restricted our analysis to the 3.2-3.8 μ m region.

¹⁷⁶ During the first Juno orbit, on Aug. 27^{th} 2016 in total JIRAM acquired about 75000 ¹⁷⁷ spectra when observing the North Polar region of Jupiter. More than 16000 were recorded ¹⁷⁸ over the North auroral region from 08:24 to 11:51 UTC. Those spectra were acquired at ¹⁷⁹ different times, and both the emission angle and the pixel size at Jupiter span a wide vari-¹⁸⁰ ety of values. Therefore, in order to properly map the emission on Jupiter disk, only spec-¹⁸¹ tra with an emission angle smaller than 75° have been retained in the analysis. However ¹⁸² some of the selected spectra were affected by strong intensity spikes or showed an H⁺₃ sig-

nal too weak to produce a reliable retrieval. Given the number of measured spectra, an 183 automatic procedure has been designed to perform a pre-filtering of the measurements. 184 The first step of this procedure was the identification of the spikes produced by energetic 185 particles on the detector. For each spectrum, we have evaluated the maximum intensity 186 recorded at the wavelengths of the H_3^+ lines. All the spectral points outside the H_3^+ lines 187 whose intensity was larger than 1.5 times the maximum intensity were flagged as spikes, 188 and masked out from the retrieval. Spectra with 3 or more spikes were completely dis-189 carded. The second step, applied after the spike removal, was the removal of the spectra 190 where the H_3^+ signal was below the detection limit, set to 0.0001 W/m²/sr. The final set of 19 measurements included a total of 14131 spectra. 192

Each spectrum in the final set of measurements has been analysed with the retrieval code described in section 3.1. In the first run we considered H_3^+ only in the simulated spectra and the target parameters were: H_3^+ effective temperature (T) and column density (CD), a wavelength shift and an offset value for each spectrum. The uncorrelated a-priori errors (diagonal S_a) used in the retrieval were chosen to ensure a very small constraint on the retrieval results. For all the spectra we have assumed the same Noise Equivalent Spectral Radiance (NESR) of 1.5×10^{-7} W/(m² nm sr), evaluated from deep space spectra.

A first inspection of the retrieval results highlighted that the spectra acquired over 200 the region inside the auroral oval showed higher χ -test values than the other spectra and 201 anomalous H₃⁺ temperatures. A visual inspection of some of these spectra showed that the 202 intensity of the H₃⁺ line at 3.32 μ m was always too high in comparison to other H₃⁺ diag-203 nostic lines. Considering that methane has already been observed in Jupiter auroral region 204 [Altieri et al., 2016] and its v_3 Q-branch lies in the same spectral region, we simulated the 205 analysed spectral region adding the methane emission around 3.3 μ m. A quick compari-206 son of the simulated spectra with our measurements showed that the recorded signal was 207 compatible with the CH₄ emission at 500 K superimposed to the H₃⁺ spectrum [see Mori-208 coni et al., this issue]. We therefore repeated the analysis of all the spectra including the 209 CH₄ column density among the target parameters, keeping its effective temperature fixed 210 at 500 K. Indeed this inclusion reduced the χ -test value of the retrieval. In the regions 211 where CH₄ emission was not a dominant feature, the inclusion of methane did not change 212 significantly the results of the fit for the H_3^+ parameters. Where the methane emission was 213 significant, the new H₃⁺ temperatures assumed values in the expected range (700-1100 K). 214 Finally, since the wavelength calibration of the measured spectra is expected to be depen-215

dent only from the position of the pixel on the spectrometer slit, we used the results of the first run to fit a second order polynomial function to the set of retrieved wavelength shift versus position of the pixel. We then used that polynomial to compute the real wavelength scale and repeated the analysis (final analysis) using as free parameters just T, CD, CH₄ column density and the offset. This was done to prevent the retrieval code to use the frequency shift to partially correct for other instrumental problems and therefore producing a bias in the retrieved parameters.

The results of the final analysis were further filtered by retaining only the retrievals of the spectra for which the final χ -test was smaller than 20, and the obtained T had a retrieval error lower than 100 K. No filter was applied to the size of the error on the H⁺₃ CDs. The final number of obtained results is 13198.

$4 H_3^+$ results

As can be seen in panel **a** of Fig. 1, the emission angle of the analysed observations 228 spans a wide range of values. The retrieved CD is proportional to the length of the op-229 tical path inside the H₃⁺ layer (slant columns). These two facts make impossible a direct 230 comparison of the retrieved CDs. Therefore we have transformed all the retrieved slant 231 columns into vertical columns multiplying the retrieved values by the cosine of the emis-232 sion angles of the corresponding observation. Then, assuming that the vertical extent of 233 the H₃⁺ layer at a certain geo-location is constant, similarly to what we have already done 234 for the integrated intensities (see section 2), we have divided Jupiter surface into bins and 235 we have averaged all the retrieval results and their errors inside the bins. The results are 236 reported in Figure 3 for the column densities and in Figure 4 for the temperatures. 237

The right panel of Figure 3 shows that on average the retrieval error on the CD is 238 below 30% and that the error is lower where the CDs assume the highest values. The 239 left panel of Figure 3 shows that the peak of H_3^+ column densities lays in part above the 240 model oval (dashed line) and in part closer to the statistical oval (solid line). The right 241 panel of Figure 4 shows that the highest errors on T are located in the region inside the 242 auroral oval and in general where the H_3^+ signal is lower (see panel d of Figure 1 for 243 comparison). The left panel of Figure 4 shows that in general the highest temperatures 244 are located on the left side of the auroral region. 245

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Comparing the left panels of Figure 3 and 4 we have identified 3 regions of interest, highlighted in Fig. 5:

248	A:	Longitudes from 200° to 240° and latitudes from 90° to 65°N. This region of the
249		main auroral oval is characterised by high CDs inside (poleward of) the statistical
250		oval with a peak (CD larger than $3.0 x 10^{12} \mbox{ cm}^{-2})$ in the longitude range from 200°
251		to $210^\circ and \ 67^\circ N$ in latitude. The corresponding temperature is about 850 K on av-
252		erage. Higher values for the temperatures are retrieved equatorward, in the region
253		located between the model and the statistical oval. A high CD region (with values
254		about 2 times lower) was observed at the same latitudes by Miller et al. [1997] with
255		the United Kingdom Infrared Telescope. Similarly to what we find, the correspond-
256		ing H_3^+ temperatures were of the order of 880-900 K. The morphological analysis
257		in Mura et al. [this issue] made using the imager L channel of JIRAM (3.3-3.6 μ m)
258		shows that this is a region of broad emission with thin coherent features (arcs) that
259		are visible from the main oval to 10 degrees inward. Such region may be still mag-
260		netically connected to the equatorial plane; this is in agreement with our finding of
261		uniform temperature and CD.
262	B:	Longitudes from 60° to 95°, latitudes from 90° to 75° N. On the oval arc crossing
263		the pole, H_3^+ CDs show variation between 2 and 2.6x10 ¹² cm ⁻² , with a peak on
264		the North Pole. Temperatures show values between 800 and 850 K, with a peak

the North Pole. Temperatures show values between 800 and 850 K, with a peak eastward of the North Pole of about 900 K.

C: Longitudes from 90° to 160°, latitudes from 80° to 60° N. In this region higher 266 CD values (larger then 2.6x10¹² cm⁻²) are retrieved external to the statistical oval. 267 Temperatures show high variability between 800 and 950 K, with peaks on the 268 statistical oval. UV emission increases have been also found in previous UV data 269 [e.g. Clarke et al., 2004; Clarke, 2013], and have been associated to 'dawn storms'. 270 Moreover this side of the oval appears narrower than the other side. The shape of 271 the auroral oval in this region, as seen also in the images reported by [Mura et al., 272 this issue], appears extremely sharp. 273

We also notice that there are spots where the temperatures are very high. Spotty features have been identified also in the L-band images reported by [*Mura et al.*, this issue]. The first extensive mapping of H_3^+ temperatures and CDs in the Northern Aurora of Jupiter, reported in this paper, highlight many differences in their morfology. The

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T oval and the CD oval appear asymmetric and tilted one with respect to the other. In 278 general regions with high H_3^+ CDs do not coincide with regions with high T. Following 279 O'Donoghue et al. [2014], we can speculate that high energy electrons penetrate deep in 280 the atmosphere producing large H_3^+ quantities at low altitudes, where temperatures are low, 281 while less energetic electrons will stop at higher altitudes, generating larger H₃⁺ quantities 282 in the upper layers of the atmosphere, where the temperature is larger. Assuming the auro-283 ral temperature profile reported by Grodent et al. [2001], a 100 K difference (from 950 K 284 to 850 K) would mean a difference of about 150 km in the penetration depth of the elec-285 trons. If this is effectively the case, the temperature/CD maps would give some hint on the 286 variability of precipitating electron energies, as suggested by Hiraki and Tao [2008]. An-28 other possibility is that the lower temperatures observed in correspondence of large CDs 288 are a signature of the "H₃⁺ thermostat effect" [Miller et al., 2013]: larger H₃⁺ densities lead 289 to an increased infrared cooling of the atmosphere that lowers the local temperature. How-290 ever, we do not observe a strict anti-correlation between temperature and CD, suggesting 291 that the thermostat mechanism is not enough to explain the observed differences. 292

293 5 Conclusions

We have developed an analysis tool to invert JIRAM spectra in the 3 - 4 μ m spec-294 tral region to retrieve informations on H₃⁺ distribution and temperatures in Jupiter auroral 295 regions. The tool has been applied to JIRAM measurements acquired over the North Po-296 lar region during the first Juno orbit around Jupiter. Given the number of measurements 297 acquired by JIRAM, an automatic procedure to identify and discard problematic measure-298 ments has been developed. We have evaluated the distribution of the H_2^+ column densi-299 ties and temperatures in the Northern aurora. The analysis shows that the location of the 300 maximum H₃⁺ concentration is close to what models and previous observations suggest. 301 The high spatial resolution of the measurements suggest that the North aurora of Jupiter 302 is not uniform, with different distributions of the H₃⁺ abundance and temperature. On the 303 main oval averaged H_3^+ CDs span between 1.8×10^{12} cm⁻² and 2.8×10^{12} cm⁻², while the 304 retrieved temperatures show variation between 800 and 950 K. On the auroral region at 305 longitudes from 90° to 170° higher H₃⁺ column densities are observed equatorward of the 306 main oval, whereas the temperature is higher inside the statistical oval region. On the con-307 trary, in the auroral region from 200° to 210° in longitude higher temperatures are ob-308

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- served with increasing values equatorward. JIRAM first data confirm the complex rela-
- tionship among H_3^+ emission rates, CDs and temperatures on the Jupiter Northern aurora.

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270°

Figure 1. Orthographic map of Jupiter North Pole in planetocentric coordinates with superimposed spots at the geo-location of the intercept of JIRAM LOS with the surface located at 500 km above Jupiter 1-bar surface. The longitude scale is reported above the axis and dashed circles have been drawn every 10° of latitude. The colour of each dot indicates the value of the represented quantity, while the size of the spot represents its spatial resolution. The black lines represent the position of the aurora from models (dashed line) and from statistics (solid line) Panel **a** shows the emission angle of the measurements, panel **b** shows the solar incidence angle, panels **c** and **d** show the integrated intensities of the measurements.

- **Figure 2.** JIRAM spectra acquired in the sunlit auroral region (red line Jupiter solar time 6:36) and in the
- dark auroral region (blue line Jupiter solar time 9:50)

- Figure 3. Same region of Figure 1. Left panel: map of the retrieved H_3^+ column densities. Right panel: map
- ³⁹¹ of the average retrieval error on the retrieved CDs.

- Figure 4. Same region of Figure 1. Left panel: map of the retrieved H_3^+ effective temperatures. Right panel:
- ³⁹³ map of the average retrieval error on the retrieved Ts.

Figure 5. Comparison of the distribution of H_3^+ CD (left panel) and T (right panel).

394

'Figure 1'.



'Figure 2'.



'Figure 3'.





'Figure 4'.





'Figure 5'.



