# Constraining model biases in a global general circulation model with ensemble data assimilation methods

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## Overview

## Introduction

## Method formulation

- 3 Lorenz '96 Twin experiment
  - Single assimilation
  - Iterative assimilation

## 4 NEMO-LIM2

## 5 Twin Experiment

- Monovariate assimilation
- Multivariate assimilation

## 6 Real Case

- Single assimilation
- Iterative assimilation

Perspectives and conclusions

Most **numerical models** suffer important errors due to poorly represented processes. This leads to a systematic error with a non-zero mean: **bias**.

Bias is considered to be the **main source** of errors in climatic model. It allows one only to study the variation of a model, not its absolute results (Zunz et al., 2013).

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**Objective**: Develop a method aiming at correcting and coming closer to numerical model bias.

## Model Bias

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#### Numerical model bias

Dee, 2005: **Systematic** error with regard to the notion of the **average** of a model or estimator.

- Spatially variable
- Time dependence
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Origin:

- Poor parametrisation and representation of physical processes
- Bias in boundary and initial conditions
- Bias on observations

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Understanding and predicting Antarctic sea ice variability at the decadal timescale.

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Characteristics and requirements of the project:

- Long term simulations
- Low resolution model: NEMO-LIM2
- Large uncertainties on the effects of small scale processes

Comparison between 5<sup>th</sup> Coupled Model Intercomparison Project (CMIP5) using the ORCA2 grid:

- CMCC-CM (Centro Euro-Mediterraneo sui Cambiamenti Climatici -Climate Model)
- CMCC-CMS
- NEMO-LIM2 Free run
- NEMO-LIM2 with data assimilation

and observational data from OSTIA (Operational SST and Sea Ice Analysis).

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#### Antarctic ice coverage RMSE (in fraction) for period 1985-2005.



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NEMO-LIM2 free run.



NEMO-LIM2 with assimilation of OSTIA observations.

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$$\mathbf{x}_m = M(\mathbf{x}_{m-1}) + \eta_m, \tag{1}$$

Variables	Descriptions
т	Time index subscript.
<b>x</b> , <b>y</b>	State vector, observations.
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## State of the art: Data Assimilation

#### Data assimilation methods



Data assimilation methods overview.

# State of the art: Original and Extended Kalman Filter

**Original Kalman Filter** from Kalman (1960), in the Bayesian framework. Computes background and error covariance matrix at each step. **Original Kalman Filter** from Kalman (1960), in the Bayesian framework. Computes background and error covariance matrix at each step.

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**Extended Kalman Filter**: Linearisation of non linear model and observation operators.

However, error covariance matrix not computable: Size  $N_x^2$ ,  $N_y^2$ , where  $N_x > 10^6$ ,  $N_y > 10^4$  in realistic models.

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**Ensemble Transform Kalman filter**: No perturbations on the observations.

# State of the art: Bias correction

Data assimilation schemes:

- **Bias blind**: Ignores bias on the observations and in the model background estimate.
- **Bias aware**: Estimate model state and bias. Bias is isolated from other state vector variables.

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- Offline: Bias estimated beforehand, from preliminary model run
- **Online**: Bias estimated and updated during the assimilation scheme.

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#### Data assimilation with numerical model bias



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- Reference and **unbiased** data set from which an estimation can be provided.
- Bias model or characterisation with parameters.

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Crucial, for that an erroneous correction deteriorates the model even more than bias blind assimilation.

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One wants to correct the bias term **b** at every time step.

Start from the free and biased model trajectory:  $\mathbf{x}=$ 

$$\begin{bmatrix} \mathbf{x}_1 \\ \mathbf{x}_2 \\ \vdots \\ \mathbf{x}_{m_{\text{max}}} \end{bmatrix}$$

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Generate an ensemble of bias correction estimators:  $\hat{\mathbf{b}}^{(i)}$ .

Variables	Descriptions
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Generate an ensemble of bias correction estimators:  $\hat{\mathbf{b}}^{(i)}$ .

Run that ensemble to obtain an ensemble of forced runs:  $\mathbf{x}^{(i)}$ .

Variables	Descriptions		
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#### Bias correction: A new method

State vector **augmentation** of the model trajectory with the bias correction estimator  $\hat{\mathbf{b}}^{(i)}$ , **ensemble mean**, and assimilation scheme:

(9)

$$\mathbf{x}^{\prime(i)} = \begin{bmatrix} \mathbf{x}_{1}^{(i)} \\ \mathbf{x}_{2}^{(i)} \\ \vdots \\ \mathbf{x}_{mmax}^{(i)} \\ \widehat{\mathbf{b}}^{(i)} \end{bmatrix}, \quad (6)$$

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Variables	Descriptions				
fa	Forecast and analysis superscripts.				
Р	Covariance matrix of state vector.				
R	Observation error covariance matrix.				
K	Kalman gain.				
N <sub>e</sub>	Ensemble size.				

Observations are already time averages. To limit the model trajectory and state vector size, one can take the **time average** as follow

$$\mathbf{H}'\mathbf{x}' = \frac{1}{m_{max}} \sum_{m=1}^{m_{max}} \mathbf{H}\mathbf{x}_m = \mathbf{H}\overline{\mathbf{x}}.$$
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Augmented state vector and observation operator become

$$\mathbf{x}'' = \begin{bmatrix} \overline{\mathbf{x}} \\ \widehat{\mathbf{b}} \end{bmatrix},$$
 (11)  $\mathbf{H}'' \mathbf{x}'' = \mathbf{H} \overline{\mathbf{x}}.$  (12)

Analysis with the average model state is equivalent to full trajectory, and is expressed as

$$\mathbf{x}^{\prime\prime a} = \mathbf{x}^{\prime\prime f} + \mathbf{K}^{\prime\prime} \left( \mathbf{y}^{o} - \mathbf{H}^{\prime\prime} \mathbf{x}^{\prime\prime f} \right).$$
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The bias correction term  $\hat{\mathbf{b}}^a$  is provided by the analysis  $\mathbf{x}''^a$ . One obtains the **corrected model rerun** with

$$\mathbf{x}_{m}^{r} = M_{m} \left( \mathbf{x}_{m-1}^{r} \right) + \widehat{\mathbf{b}}^{a}.$$
(14)



General schematic of the bias correction method, from the initial model run  $\mathbf{x}_m$  to the corrected model run  $\mathbf{x}_m^r$ .

## Lorenz '96 model: introduction

Formulated in 1996 (K = 40 variables), exhibits advection, diffusion, periodicity, and chaotic behaviour. With k = 1, ..., K, it is described by

$$\frac{dX_k}{dt} = -X_{k-2}X_{k-1} + X_{k-1}X_{k+1} - X_k + F.$$
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Creates a spatially variable constant in time forcing term  $F_k$  with a mean **F** and a perturbation depending on covariance **P** defined by

$$P_{i,j} = 0.3 \ e^{\frac{-(i-j)^2}{15}}.$$
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Model retains the same characteristics, but spatial  $F_k$  is more challenging to recover.



#### Lorenz '96 model mean state

Parameters:

- Model mean  $\overline{F^t} = 4$
- Initial conditions  $I_{max} = 15$
- Time span  $m_{max} = 1000$
- Ensemble size  $N_e = 100$

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Model mean obtained from average model trajectory by

$$X_{k}^{t} = \frac{1}{I_{max}} \sum_{l=1}^{I_{max}} \frac{1}{m_{max} - 199} \sum_{m=200}^{m_{max}} X_{k,l,m}^{t}.$$
(18)

trajectory by

Parameters:

- Model mean  $\overline{F^t} = 4$
- Initial conditions  $I_{max} = 15$
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Model mean obtained from average model

Observations are created by adding noise to the reference model run.



Reference parameter from twin experiment.



Reference parameter and corresponding model output

Forcing parameter



Output

Ensemble of forcings is generated.

#### Forcing parameter

Output



Ensemble is run to provide an ensemble of model outputs.

#### Forcing parameter

Output



Observations are assimilated.



Output



Ensemble of forcing parameters is corrected.

Forcing parameter



Output

Ensemble is rerun to obtain rerun model outputs.

To reduce nonlinear behaviour, one can **iterate** the assimilation procedure. Observation batches creation as follows

$$\mathbf{y}_2^o = \begin{pmatrix} \mathbf{y}^o \\ \mathbf{y}^o \end{pmatrix}, (19) \qquad \mathbf{R}_2 = \begin{pmatrix} 2\mathbf{R} & \mathbf{0} \\ \mathbf{0} & 2\mathbf{R} \end{pmatrix}, (20) \qquad \mathbf{H}_2 = \begin{pmatrix} \mathbf{H} \\ \mathbf{H} \end{pmatrix}. (21)$$

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In absence of correlation between subsets of data, iterative assimilation is **equivalent** to single assimilation.

$$(\mathbf{P}^{a})^{-1}\mathbf{x}^{a} = (\mathbf{P}^{f})^{-1}\mathbf{x}^{f} + \mathbf{H}^{T}\mathbf{R}^{-1}\mathbf{y}^{o} = (\mathbf{P}^{f})^{-1}\mathbf{x}^{f} + \mathbf{H}_{2}^{T}\mathbf{R}_{2}^{-1}\mathbf{y}_{2}^{o}.$$
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However, for a **nonlinear** observation operator, iterations allow smaller steps by model rerun.

Suggested in Annan et al., 2005, and similar to running in place (RIP) algorithm (Kalnay and Yang, 2010).

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Experiment parameters:

- $\overline{F^t} = 5$
- $\overline{F^f} = 6$
- *I<sub>max</sub>* = 10
- $m_{max} = 1000$
- $N_e = 50$
- $n_{iter}^{max} = 1, 4$
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Different  $\overline{F^t}$  and  $\overline{F^f}$  for readability. Spread is sufficient.

Same **initial condition** for every iteration experiment.

Increase in computational cost proportional to  $n_{iter}^{max}$ .

## Lorenz '96 model: iterative assimilation





(b) Corresponding time averaged model state.

## Lorenz '96 model: iterative assimilation



(a)  $F_k$  for an iterated assimilation.

(b) Corresponding time averaged model state.

Variable index

-Iter 3

-Forecast

10

—Truth

Iter 2

Model output for 4 iterations

30

Iter 1

Iter 4

Exact RMSE values on the forcing parameter:

Backg	round RMSE	Analysed RMSE			
n <sub>iter</sub>	Background	$n_{iter} = 1$	$n_{iter} = 2$	$n_{iter} = 3$	$n_{iter} = 4$
1	1.270	0.726			
2	1.270	0.915	0.663		
3	1.270	1.007	0.799	0.639	
4	1.270	1.060	0.887	0.737	0.619

Table : RMSE on  $F_k$
Exact RMSE values of the corrected model state rerun ensemble mean:

Backg	round RMSE		Rerun	ın RMSE			
n <sup>max</sup>	Background	$n_{iter} = 1$	$n_{iter} = 2$	$n_{iter} = 3$	$n_{iter} = 4$		
1	0.304	0.187					
2	0.304	0.233	0.170				
3	0.304	0.254	0.203	0.163			
4	0.304	0.263	0.227	0.195	0.160		

Table : RMSE on the time average of the model state

### Conclusion:

- Modified Lorenz '96 model application.
- Bias correction term is estimated.
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Time to test the method with a realistic model.

**Realistic model**: NEMO (Nucleus for European Modelling of the Ocean), coupled to the LIM2 (Louvain-la-Neuve Sea Ice Model) sea ice model.

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Constraints: Model stability, physical restrictions on currents.

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**Method implementation**: Correct currents by adding a random forcing into the ocean dynamics equations.

Constraints: Model stability, physical restrictions on currents.

**Assimilation Scheme**: Local assimilation with ETKF scheme from OAK (Ocean Assimilation Kit).

Forcing term generation from **random field**  $\Psi = \Psi(x, y)$  and from the Hessian of the cost function  $J(\Psi)$  by

$$J(\Psi) = \int_{\Omega} L_h^4 (\nabla^2 \Psi)^2 + 2L_h^2 (\nabla \Psi)^2 + \Psi dx.$$
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This provides a random stream function tendency with constraints on correlation length  $L_h$ .

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This provides a random stream function tendency with constraints on correlation length  $L_h$ .

One avoids perpendicular currents with  $\nabla \Psi \bullet \mathbf{t} = 0$ , and uses spatial filtering to increase model stability.

# NEMO-LIM2: Bias correction generation

#### Average mixed layer depth



Yearly average of the mixed layer depth from a NEMO-LIM2 free run, in m.

### Zonal and meridional forcings from stream function

$$F_u(x, y, z) = -\frac{\partial \Psi'(x, y, z)}{\partial y}, \quad (24) \qquad F_v(x, y, z) = \frac{\partial \Psi'(x, y, z)}{\partial x}. \quad (25)$$

Zonal and meridional forcings from stream function

$$F_u(x,y,z) = -\frac{\partial \Psi'(x,y,z)}{\partial y}, \quad (24) \qquad F_v(x,y,z) = \frac{\partial \Psi'(x,y,z)}{\partial x}.$$
 (25)

Modified ocean dynamics equations:

$$\frac{du}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial x} + fv + \frac{1}{\rho} \frac{\partial \tau_x}{\partial z} + F_u,$$
(26)
$$\frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial y} - fu + \frac{1}{\rho} \frac{\partial \tau_y}{\partial z} + F_v.$$
(27)

#### **Experiment parameters**

- Correlation length: 5000km
- SSH spatial variability: 30*cm*
- Ensemble size: 100 members

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### Procedure

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- Assimilate SSH observations
- Extract optimal forcing
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### Procedure

- Run ensemble with random forcing
- Assimilate SSH observations
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State vector is augmented with the observations:  $\mathbf{x}'' =$ 

$$= \begin{bmatrix} \overline{SSH} \\ \widehat{F}_u \\ \widehat{F}_v \end{bmatrix}.$$

Ensemble of forcings is run. Observations are taken from the reference run.

Yearly mean sea surface height (in m)



(a) Average of the ensemble.



(b) Twin experiment true run.

# NEMO-LIM2 Twin Experiment: Zonal forcing

After SSH assimilation, the analysis provides estimated bias correction terms.

Zonal forcing (in ms<sup>-2</sup>)



# NEMO-LIM2 Twin Experiment: Meridional forcing

After SSH assimilation, the analysis provides estimated bias correction terms.

Meridional forcing (in ms<sup>-2</sup>)



# NEMO-LIM2 Twin Experiment: SSH rerun

The optimal forcing is rerun, providing a bias corrected run.

Model rerun SSH (in m)



(a) Rerun with optimal forcing.



(b) Reference run.

## NEMO-LIM2 Twin Experiment: SSH RMSE

#### Twin experiment SSH RMSE



RMSE on SSH from Ensemble Mean before and after analysis, and Rerun, with True Run (in m)

# NEMO-LIM2 Twin Experiment: SST Validation





(b) RMSE on SST from Ensemble Mean after analysis, and Rerun, with True Run (in C°).

# NEMO-LIM2 Twin Experiment: SST Validation

### SST validation (in $C^{\circ}$ )



(a) Rerun with optimal forcing.



(b) Reference run.

Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>					
$\widehat{F}_{v}$ in ms <sup>-2</sup>					
SSH in m					
SST in $C^{\circ}$					
SSS in PSU					

Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$				
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$				
SSH in m					
SST in C°					
SSS in PSU					

Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$				
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$				
SSH in m	0.220				
SST in $C^{\circ}$	0.999				
SSS in PSU	0.268				

Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$	$6.27 imes10^{-7}$			
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$	$5.81 imes10^{-7}$			
SSH in m	0.220				
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Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$	$6.27 imes10^{-7}$			
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$	$5.81 imes10^{-7}$			
SSH in m	0.220		0.068		
SST in C°	0.999		0.539		
SSS in PSU	0.268		0.197		

Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$	$6.27 imes10^{-7}$		$5.96 imes10^{-7}$	
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$	$5.81 imes10^{-7}$		$5.45 imes10^{-7}$	
SSH in m	0.220		0.068		
SST in C°	0.999		0.539		
SSS in PSU	0.268		0.197		

Variable	Forecast	Monovariate		Multivariate	
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$	$6.27 imes10^{-7}$		$5.96 imes10^{-7}$	
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$	$5.81 imes10^{-7}$		$5.45 imes10^{-7}$	
SSH in m	0.220		0.068		0.061
SST in C°	0.999		0.539		0.509
SSS in PSU	0.268		0.197		0.150

# NEMO-LIM2 Real Case

SSH observations come from the **mean dynamic topography** (MDT) of CNES-CLS09 (Centre National d'Etudes Spatiales, Collecte Localisation Satellites).



#### SSH RMSE

RMSE on SSH with CNES-MDT observations (in m).

Strong currents are too weak in NEMO.

Yearly mean SSH (in m)



(a) CNES-MDT observations.



(b) Model free run.

Corrected rerun shows small scale structure, but also stronger currents.

Yearly mean SSH (in m)



(a) Average of the ensemble.



(b) Optimal forcing rerun.

## NEMO-LIM2 Real Case: SSH average error

#### Yearly mean SSH average error with the observations (in m)



(a) Average of the ensemble.

(b) Model free run.

## NEMO-LIM2 Real Case: SSH average error

Yearly mean SSH average error with the observations (in m)



(a) Analysis.

(b) Optimal forcing rerun.

## NEMO-LIM2 Real Case: final forcing

**Final forcing** 



Analysed forcing from CNES-MDT observations, used to rerun the model (in  $\rm ms^{-2}).$ 

# NEMO-LIM2 Real Case: final forcing

#### Real global current map



Real global average current map of the oceans. Adapted from http://www.georgemaps.com/  $\,$
## NEMO-LIM2 Real Case: SST Validation

**SST climatology** from NODC-WOA13V2 data provided by the National Oceanic and Atmospheric Administration (NOAA).

Yearly mean SST average error (in C°)





(a) Optimal forcing.

(b) Model free run.

SSH RMSE



Exact RMSE values on the SSH for single and iterative experiments.

ARMSE on SSH (in m)		RMSE (in m)			
		Background	Single Assim	Iterativ	e Assim
R	2 <b>R</b>			iter 1	iter 2
0.0215	0.0431				
0.0464	0.0928				
0.1000	0.2000				
0.2154	0.4308				
0.4642	0.9284				

Exact RMSE values on the SSH for single and iterative experiments.

ARMSE on SSH (in m)		RMSE (in m)			
		Background	Single Assim	Iterativ	e Assim
R	2 <b>R</b>			iter 1	iter 2
0.0215	0.0431	0.1965			
0.0464	0.0928	0.1965			
0.1000	0.2000	0.1965			
0.2154	0.4308	0.1965			
0.4642	0.9284	0.1965			

Exact RMSE values on the SSH for single and iterative experiments.

ARMSE on SSH (in m)		RMSE (in m)			
		Background	Single Assim	Iterative Assi	
R	2 <b>R</b>			iter 1	iter 2
0.0215	0.0431	0.1965	0.1604		
0.0464	0.0928	0.1965	0.1592		
0.1000	0.2000	0.1965	0.1571		
0.2154	0.4308	0.1965	0.1554		
0.4642	0.9284	0.1965	0.1589		

Exact RMSE values on the SSH for single and iterative experiments.

ARMSE	RMSE on SSH (in m)		RMSE (in m)		
		Background	Single Assim	Iterative Assim	
R	2 <b>R</b>			iter 1	iter 2
0.0215	0.0431	0.1965	0.1604	0.1604	
0.0464	0.0928	0.1965	0.1592	0.1579	
0.1000	0.2000	0.1965	0.1571	0.1554	
0.2154	0.4308	0.1965	0.1554	0.1574	
0.4642	0.9284	0.1965	0.1589	0.1640	

Exact RMSE values on the SSH for single and iterative experiments.

ARMSE on SSH (in m)		RMSE (in m)				
		Background	Single Assim	Iterative Assim		
R	2 <b>R</b>			iter 1	iter 2	
0.0215	0.0431	0.1965	0.1604	0.1604	0.1315	
0.0464	0.0928	0.1965	0.1592	0.1579	0.1305	
0.1000	0.2000	0.1965	0.1571	0.1554	0.1341	
0.2154	0.4308	0.1965	0.1554	0.1574	0.1416	
0.4642	0.9284	0.1965	0.1589	0.1640	0.1511	

# NEMO-LIM2 Real Case: Iterative Analysis SSH average error

### Yearly mean SSH average errors (in m)





(a) First iteration rerun.

(b) Second iteration rerun.

The twin experiment shows a **significant** reduction on the SSH bias through adequate model forcing. The real case shows that the bias **generation** is important, and the strong **seasonal** cycle in the Antarctic ocean.

The twin experiment shows a **significant** reduction on the SSH bias through adequate model forcing. The real case shows that the bias **generation** is important, and the strong **seasonal** cycle in the Antarctic ocean.

Possible development options for the **future**:

- Real 3D forcing
- Time dependent forcing, with seasonal variations
- Validation against other bias correction methods
- Parametrisation of the final forcing
- Extract local optimal forcing

## Thesis: Conclusions

Theoretical formulation of the bias correction method.

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The application on the Lorenz '96 shows:

- Estimation of the bias correction term.
- Reduced model bias on the model rerun.

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NEMO-LIM2 implementation:

- Correction term forcing is stable on the model.
- Bias correction works on complex models.
- Iterative assimilation is more accurate.

Theoretical formulation of the bias correction method.

The application on the Lorenz '96 shows:

- Estimation of the bias correction term.
- Reduced model bias on the model rerun.

NEMO-LIM2 implementation:

- Correction term forcing is stable on the model.
- Bias correction works on complex models.
- Iterative assimilation is more accurate.

Thesis objectives:

- Model correction can be used for future projections.
- Applicable to other models.

Thank you for your attention.

Multiple notations through time. Unified notation in Ide et al. (1997).

Variables	Descriptions
т	Time index subscript.
fa	Forecast and analysis superscripts.
$\mathbf{x}, \mathbf{y}$	State vector, observations.
$\mathbf{M}, \mathbf{H}$	Forward model and observation operators.
Р	Covariance matrix of state vector.
$\mathbf{Q}, \mathbf{R}$	Model and observation error covariance matrix.
K	Kalman gain.
i	Ensemble index superscripts.

SSH and SST can be assimilated together, for a better estimation of the bias correction forcing.

Variable	Forecast	Monovariate		Multivari	ate
name		analysis	rerun	analysis	rerun
$\widehat{F}_u$ in ms <sup>-2</sup>	$1.66 imes10^{-6}$	$6.27 imes10^{-7}$		$5.96 imes10^{-7}$	
$\widehat{F}_{v}$ in ms <sup>-2</sup>	$1.24 imes10^{-6}$	$5.81 imes10^{-7}$		$5.45 imes10^{-7}$	
SSH in m	0.220	0.039	0.068	0.0457	0.061
SST in C°	0.999		0.539	0.453	0.509
SSS in PSU	0.268		0.197		0.150

Table : RMSE values of the multivariate rerun for a  $ARMSE = 1 \text{ C}^{\circ}$  value, compared to the monovariate assimilation. Empty values are not relevant.

	D.4.	The	program	structur	e
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#### D.4 The program structure

To be done ....

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## NEMO-LIM2 Real Case: final forcing

**Final forcing** 



Analysed forcing from CNES-MDT observations, used to rerun the model (in  $\rm ms^{-2}).$ 

## NEMO-LIM2 Real Case: Gulf Stream

#### Zonal current Gulf Stream (in ms)



## NEMO-LIM2 Real Case: Gulf Stream

#### Zonal current Gulf Stream, latitude cut (in m/s)



Free model range: -0.2921 to 0.25035 (m/s). Forced model: -0.57111 to 0.53793 (m/s).