The structure and variability of Mars upper atmosphere as seen in MAVEN/IUVS dayglow observations

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Abstract

We report a comprehensive study of Mars dayglow observations focusing on upper atmospheric structure and seasonal variability. We analyzed 744 vertical brightness profiles comprised of ~109,300 spectra obtained with the Imaging Ultraviolet Spectrograph (IUVS) aboard the Mars Atmosphere and Volatile Evolution (MAVEN) satellite. The dayglow emission spectra show features similar to previous UV measurements at Mars. We find a significant drop in thermospheric scale height and temperature between \( L_S = 218\)° and \( L_S = 337-352\)°, attributed primarily to the decrease in solar activity and increase in heliocentric distance. We report the detection of a second, low-altitude peak in the emission profile of OI 297.2 nm, confirmation of the prediction that the absorption of solar Lyman alpha emission is an important energy source there. The CO⁺ UV doublet peak intensity is well correlated with simultaneous observations of solar 17–22 nm irradiance at Mars.

1. Introduction

Mars has been studied extensively at ultraviolet wavelengths starting from Mariner 6 and 7 [Barth et al., 1969, 1971; Stewart, 1972; Strickland et al., 1972], Mariner 9 [Stewart et al., 1972, 1973], and more recently by Spectroscopy for Investigation of Characteristics of the Atmosphere of Mars (SPICAM) aboard Mars Express [Bertaux et al., 2006; Leblanc et al., 2006; Simon et al., 2009; Cox et al., 2010; Stiepen et al., 2015]. The results from these measurements reveal a large variability in the composition and structure of the Martian upper atmosphere. However, due to the lack of simultaneous measurements of energy input in the atmosphere, such as solar electromagnetic and particle flux as well as limitations in the observation geometry and data itself, this variability is still not fully understood. The Imaging Ultraviolet Spectrograph (IUVS) is one of the remote sensing instruments on the Mars Atmosphere and Volatile Evolution (MAVEN) spacecraft. MAVEN's observations allow IUVS to conduct limb scans near periapse, usually at every other orbit, largely without interference from other instruments aboard MAVEN, and can thus observe the Martian atmospheric variability over an extended period with a wide spatial and temporal coverage. The detailed science goals of IUVS are provided by McClintock et al. [2014].

In this paper, we present ultraviolet limb scan dayglow observations made by IUVS and study the major characteristic of Martian dayglow and its seasonal variations.

2. Observation Geometry and Data

The MAVEN spacecraft is in an elliptical orbit with apoapsis near 6000 km altitude and periapsis near 160 km [Jakosky et al., 2015]. IUVS carries two detectors: a far ultraviolet (FUV) detector (115–190 nm) and a middle ultraviolet (MUV) detector (180–340 nm) with a spectral resolution of ~0.6 and 1.2 nm, respectively, [McClintock et al., 2014]. The instrument is mounted on an Articulated Payload Platform (APP) that can orient IUVS's field of view relative to Mars depending on spacecraft location, orientation, and desired viewing geometry. IUVS uses a long, narrow slit (11.3° × 0.06°) in the telescope focal plane to provide entrance to
Table 1. Observations of Martian Dayglow From IUVS and Prior Missions

<table>
<thead>
<tr>
<th>Spacecraft</th>
<th>$L_\odot$ (deg)</th>
<th>$D_{\odot}$ (AU)</th>
<th>SZA Range (deg)</th>
<th>$F_{10.7}$</th>
<th>$T$ (K)</th>
<th>Peak Brightness (kR)</th>
<th>Peak Altitude (km)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Mariner 6</td>
<td>200</td>
<td>1.42</td>
<td>44</td>
<td>180</td>
<td>350</td>
<td>509</td>
<td>131</td>
</tr>
<tr>
<td>Mariner 7</td>
<td>200</td>
<td>1.42</td>
<td>27</td>
<td>180</td>
<td>350</td>
<td>555</td>
<td>129</td>
</tr>
<tr>
<td>Mariner 9</td>
<td>312</td>
<td>1.45</td>
<td>7–23</td>
<td>127</td>
<td>325</td>
<td>300</td>
<td>138</td>
</tr>
<tr>
<td>SPICAM</td>
<td>100–130</td>
<td>1.63</td>
<td>30–60</td>
<td>105</td>
<td>200</td>
<td>155</td>
<td>118</td>
</tr>
<tr>
<td>SPICAM</td>
<td>138–171</td>
<td>1.52</td>
<td>30–60</td>
<td>105</td>
<td>200</td>
<td>153</td>
<td>125</td>
</tr>
<tr>
<td>IUVS</td>
<td>218</td>
<td>1.45</td>
<td>30–72</td>
<td>110</td>
<td>250</td>
<td>583</td>
<td>127</td>
</tr>
<tr>
<td>IUVS</td>
<td>337–352</td>
<td>1.5</td>
<td>10–73</td>
<td>110</td>
<td>250</td>
<td>440</td>
<td>122.5</td>
</tr>
</tbody>
</table>

$^a$Sun-Mars distance.
$^b$F$_{10.7}$ is the solar EUV flux index (in units of $10^{-22}$ W/m$^2$/Hz), measured at 1 AU, and corrected for Sun-Earth-Mars angle.
$^c$Temperature values retrieved from CO$_2$ UVD profile. References are same as provided for instrument.
$^d$Studies on Mariner 6, 7, and Mariner 9 did not report CO$_2$ UVD brightness near the airglow peak.
$^e$Barth et al. [1971] and Stewart [1972].
$^f$Stewart et al. [1972].
$^g$Leblanc et al. [2006]; values are for SZA = 15–37°.

We have divided the data set into two solar longitude sets $L_\odot \sim 218$° (18–22 October 2014) and $L_\odot = 337–352$° (5 May to 2 June 2015). For $L_\odot \sim 218$°, our analysis includes 19 orbits (109–128) except for orbit 115, when the spacecraft was in safe mode during the maximum predicted dust fluence associated with the comet C/2013 A1 Siding Spring’s close flyby of Mars [Schneider et al., 2015, and references therein]. For $L_\odot$ between 337° and 352°, 43 orbits are used; every other orbit between 1160 and 1220, and between 1275 and 1309 (except orbits 1198, 1216, 1289, and 1295). A total of 109,368 individual spectra are used to generate 744 altitude profiles (228 for $L_\odot = 218$° and 516 for $L_\odot = 337–352$°). The occasional gaps in between the nominal periapse observation mode are due to sharing pointing control with other instruments, communication passes, solar Keep-Out Zone constraints, calibration activities, and other irregularities. The lighting and geometry for the IUVS observations are given in Table 1.

Figure 1 shows the color composite of MUV and FUV detectors assembled from a full limb scan (orbit 110) of the Martian airglow layer. In the MUV channel, the CO Cameron (a$^3$Π − X$^1$Σ+) band system between 175 and 270 nm, CO$_2$ ultraviolet doublet (UVD, B$^2$Σ$^+$ − X$^2$Π$_G^+$) near 289 nm, CO$_2$ Fox-Duffendack-Barker (FDB, A$^2$Π$_u$ − X$^2$Π$_g$) bands near 310–330 nm, and OI (1$^3$P − 1$^1$S) 297.2 nm emissions are prominent as expected from previous observations by Mariner 6 and 7 [Barth et al., 1971], Mariner 9 [Stewart et al., 1972], and SPICAM/Mars Express [Leblanc et al., 2006]. The prominent features in FUV consist of H Ly α, OI 130.4 and 135.6 nm, CI 156.1 nm, and CO fourth positive (a$^3$Π − X$^1$Σ+) band system. Figure 1 (bottom) shows the averaged spectra in the FUV and MUV observed at altitudes between 120 and 130 km.

IUVS has also detected N$_2$ Vegard-Kaplan bands in the MUV and for the first time, the N$_2$ Lyman-Birge-Hopfield bands in Mars’s FUV dayglow; details are reported in Stevens et al. [2015].
3. Analysis of IUVS Data and Results

As shown in Figure 1, the Martian ultraviolet dayglow spectrum is a blended composite of many emission features, each of which is identified from its expected location and shape in the spectrum. We have used a multiple linear regression (MLR) method to fit the observed spectrum. Aside from atomic and molecular emissions described above, the fit templates include a reference-reflected solar spectrum measured by IUVS during disk observations. The inclusion of this spectrum is important at altitudes below 100 km, where solar scattered light can appear in the MUV spectra both from high-altitude clouds and dust, and from stray light within the instrument. Other templates use the IUVS line spread function (LSF), obtained from an IUVS observation of interplanetary and Mars coronal hydrogen Lyman-\(\alpha\). More details about MLR, the LSF, and the synthetic spectrum are given in Stevens et al. [2015] and Evans et al. [2015].

The observed raw data numbers (DN) are corrected for detector dark current and then converted to physical brightness in Rayleigh using the sensitivity derived from UV bright stellar observations made during the MAVEN cruise phase [McClintock et al., 2014]. The MUV and FUV systematic uncertainty estimated from these stellar calibrations are \(\pm 30\%\) and \(\pm 25\\%\), respectively. The flatfield errors have not been corrected in the data, which can introduce additional 10% uncertainty. A detailed study on IUVS calibrations and instrument performance is currently underway.

3.1. Altitude Profiles of Thermospheric Emission

To derive the brightness of the emissions, the fitted spectra of targeted emissions from each section of the slit are binned at 5 km altitude intervals. Figure 2 shows the altitude profiles of CO Cameron, CO\(_2\) UVD, OI 297.2 nm, and CO\(_2\) FDB bands, for the two seasons and solar zenith angles 30–45\(^\circ\) (top) and 50–70\(^\circ\) (bottom).

Some of the spectra used in this analysis (for \(L_s = 345^\circ\)) captured only wavelength \(\geq 198\) nm (affects total CO Cameron band intensity). Comparison with the complete spectra shown that these contained 83% of the total CO Cameron band intensity. A normalization factor of 1.21 has therefore been applied to CO Cameron band intensity observed during \(L_s = 345^\circ\).

The peak intensities of the CO\(_2\) UVD and CO Cameron bands are about 35% higher for \(L_s = 218^\circ\) than for \(L_s \approx 345^\circ\). This difference is readily accounted for by changes in heliocentric distance and solar activity (see Table 1). As expected from Chapman theory, we find that the variation of peak intensities of CO Cameron band and CO\(_2\) doublet emission vary as the cosine of solar zenith angle (SZA) [Schunk and Nagy, 2000]. The altitude of the maximum dayglow intensity is \(\sim 127\) km and \(122\) km (for CO Cameron and CO\(_2\) UVD) for \(L_s = 218^\circ\) and \(L_s \sim 345^\circ\), respectively. Large dust storms on Mars also affect the peak altitude [Stewart et al., 1972; Leblanc et al., 2006]. There was regional dust storm activity near \(L_s \sim 217\), but it was minor compared to the storms of 1971 and 2007 (S. W. Bougher, personal communication, 2015). We did not find any signifi-
Figure 2. Mean altitude profiles for the CO$_2$ ultraviolet doublet, CO Cameron band, OI 297.2 nm, and CO$_2$ Fox-Duffendack-Barker bands for two solar longitudes and (top) SZA = 30–45° and (bottom) SZA = 50–70°. Solid curves show the brightness for $L_s = 218°$, and dash-dotted curves shows them for $L_s \sim 345°$. The intensity of CO$_2$ FDB band is divided by a factor of 10. Horizontal bars show the standard deviation of the profiles. The intensities decrease by $\sim 50\%$ for SZA = 50–70°.

Significant dependence of the peak altitude on SZA. For the present observations, SZAs are highly correlated with latitude. The Chapman theory for a given temperature profile predicts increase in peak altitude for higher SZA, but this might be offset by a decrease in lower thermospheric temperature at higher latitudes and for SZAs. Compared to the SPICAM observations, the maximum intensities of CO Cameron band and CO$_2$ UV observed by IUVS is $\sim 3$ times higher. Most of the differences in the intensity are due to solar activity and Sun-Mars distance (see Table 1). IUVS observations conditions are more similar to Mariner 6, 7, and 9 than SPICAM because they were taken near perihelion and during moderate to high solar activity period. Nonetheless, the overall variation in Martian dayglow is not fully understood.

Figure 2 presents CO$_2$ FDB and OI 297 nm emission profiles. At Mars, the major CO$_2$ FDB excitation sources include photoionization of CO$_2$ and solar fluorescent scattering from CO$_2$. Fluorescent scattering becomes the dominant source at higher altitudes (>160 km), which accounts for the increase in their scale height at higher altitudes [Fox and Dalgarno, 1979; Fox, 1992]. These bands may be used to infer the ion densities of CO$_2$ [Evans et al., 2015].

The OI 297.2 nm emission originates from the O($^1S \rightarrow ^3P$) transition. This emission has been observed by Mariner 6 and 7 [Stewart, 1972] and SPICAM [Gronoff et al., 2012]. The major production source of O($^1S$) in the Martian atmosphere is photodissociation of CO$_2$ [Fox and Dalgarno, 1979; Gronoff et al., 2012; Bhardwaj and Jain, 2013], and its production rate has two distinct peaks: one below 100 km due to solar
Lyman \(\alpha\) and another peak at 120–130 km due to EUV absorption [Fox and Dalgarno, 1979; Simon et al., 2009; Gronoff et al., 2012; Bhardwaj and Jain, 2013]. None of the previous observations reported a lower peak of OI 297.2 nm, because of scattered solar light below 100 km. Due to low amount of stray light in the IUVS data, we have fitted and removed the solar continuum from the periapse data below 100 km and find evidence of a second peak in the OI 297.2 nm brightness profile. The detection of lower peak of OI 297.2 nm emission is important in light of simultaneous observations of solar Lyman alpha irradiance by extreme ultraviolet monitor (EUVM)/MAVEN. This will enable us to constrain the yield of O(\(^1S\)) atoms in photodissociation of CO\(_2\) at 121.6 nm wavelength [Fox and Dalgarno, 1979; Gronoff et al., 2012; Bhardwaj and Jain, 2013]. A clear IUVS detection of OI 297.2 nm emission lower peak might also help understand/reinterpret SPICAM observation of OI 297.2 nm brightness below 100 km.

The upper peak intensity of OI 297.2 nm (for SZA = 30–45°) is \(\sim 12\) KR (8 KR) for \(L_s = 218°\) (\(L_s \sim 345°\)). Stewart [1972] reported a value of \(\sim 20\) KR at 120 km, whereas Gronoff et al. [2012] reported a value of 4.5 KR (SZA = 40°, \(F_{10.7} = 115, L_s = 148°\)) near 125 km (G. Gronoff, personal communication, 2014). The upper peak intensity of OI 297 nm emission observed by Mariner and IUVS agree but are about a factor 4 higher than that reported by Gronoff et al. [2012]. This difference can partly be explained due to solar activity and the heliocentric distance of Mars between IUVS and Mariner, and SPICAM observations.

Figure 3 shows the ratio of CO\(_2^+\) UVD to the CO Cameron band. The ratio is \(\sim 0.14\) below 135 km and \(\sim 0.10\) above 160 km. The ratio decreases rapidly between 135 and 160 km, and it shows that the CO Cameron band brightness does not arise solely from dissociative excitation of CO\(_2\), and contributions from dissociative recombination of CO\(_2^+\) and electron impact on CO become important at higher altitudes [Fox and Dalgarno, 1979; Jain and Bhardwaj, 2012; Stiepen et al., 2015]. The ratio does not change significantly with the SZA.

A detailed model calculation for IUVS dayglow data is underway to quantify the role of various production sources of CO(a\(^3\)\(\Pi\)) and CO\(_2^+\) (B\(^2\)\(\Sigma^+\)).

### 3.2. Scale Height and Temperature

The primary sources of both CO(a\(^3\)\(\Pi\)) and CO\(_2^+\) (B\(^2\)\(\Sigma^+\)) involve dissociation/ionization of CO\(_2\), making them a useful tool to infer the neutral atmosphere scale height. As discussed in the previous section, the ratio of CO\(_2^+\) doublet to CO Cameron band does not vary significantly between 150 and 180 km; thus, we have chosen this region of atmosphere to derive the scale heights from CO Cameron band and CO\(_2^+\) doublet emission.
Figure 4. Variation of CO$_2^+$ ultraviolet doublet scale height for two seasons. Symbol shows the derived scale height for each individual scan in an orbit. Thick colored lines show the mean scale height and thin lines show the standard deviation for each orbit. The mean scale height for $L_s = 218^\circ$ and $L_s \sim 345^\circ$ are $16.2 \pm 0.12$ km and $14 \pm 0.1$ km, respectively. The error on scale height determination is smaller than the observed variations within an orbit.

profiles assuming an isothermal atmosphere. The scale height is calculated by fitting an exponential function to topside portion of brightness profile using a Levenber-Marquardt least squares minimization algorithm.

Figure 4 shows the scale height derived from CO$_2^+$ doublet emission. The scale heights derived from our measurements show statistically significant variations even within an orbit. The minimum and maximum values of scale heights derived for $L_s = 218^\circ$ ($L_s = 337–352^\circ$) are $12.7$ km ($10.2$ km) and $21$ km ($18.5$ km), respectively. Such variability in scale height was also reported by Stewart et al. [1972] for Mariner 9 and by Leblanc et al. [2006] and Stiepen et al. [2015] for SPICAM data. This large variability in scale height might be due to strong coupling between the lower and upper atmosphere via wave and tidal activities [Stewart et al., 1972; Bougher et al., 2001; Forbes et al., 2008; Stiepen et al., 2015]. A study on atmospheric tides in Martian atmosphere using IUVS dayglow data is presented by Lo et al. [2015].

The mean scale height for the two seasons is $16.2 \pm 0.1$ km and $14 \pm 0.1$ km with a standard deviation of $1.6$ km for both seasons. The uncertainty in scale height is dominated by intrinsic variability as opposed to measured error. These scale heights correspond to a mean thermospheric temperature of $300 \pm 2$ K and $250.6 \pm 1.7$ K for $L_s = 218^\circ$ and $L_s = 337–352^\circ$, respectively, (standard deviation is $29$ K for both seasons). This indicates $\sim 50$ K cooling over this time period, consistent with the altitude profile shapes presented in Figure 2. Evans et al. [2015] have used the CO$_2^+$ UVD profiles to retrieve the density of CO$_2$. They also inferred the temperature from retrieved CO$_2$ densities between 150 and 220 km and reported a mean temperature of $325 \pm 25$ K for $L_s = 218^\circ$. Their calculated temperature is about $25$ K higher than that in our study for the same period. The difference in temperature could be due to their choice of altitude region to infer the temperature, uncertainty in the retrieved density, and quality of fit.

We did not find any significant dependence of scale height on the latitude or solar zenith angle, which is consistent with the studies of Stewart et al. [1972], Leblanc et al. [2006], and Stiepen et al. [2015]. Leblanc et al. [2006] reported a mean scale height of $\sim 11.2$ km and $\sim 14$ km derived from CO$_2^+$ doublet and CO Cameron bands, respectively, which resulted in a temperature difference of $\sim 50$ K estimated from the two emission profiles. The mean scale height derived from CO Cameron profiles in our study is $16.7 \pm 0.1$ km ($13.9 \pm 0.1$ km) for $L_s = 218^\circ$ ($L_s = 337–352^\circ$), consistent with the scale height derived from CO$_2^+$ doublet emission. Stiepen et al. [2015] have also not found any statistical difference between the scale heights derived from the CO Cameron band and CO$_2^+$ UVD emissions.

3.3. Effects of Solar EUV Flux

Figure 5 (top) shows the short-term variation of maximum intensity of CO$_2^+$ doublet ($L_s = 337–352^\circ$ and SZA = 30–40$^\circ$) with respect to the orbital average extreme ultraviolet channel irradiance (17–22 nm: solar irradiance mainly responsible for photoionization) of the extreme ultraviolet monitor (EUVM) on board MAVEN [Eparvier et al., 2015]. There is a positive correlation between peak intensities of CO$_2^+$ doublet bands
Figure 5. Short-term variations of (top) CO$_2^+$ doublet maximum intensity and (bottom) topside temperature with 17–22 nm irradiance measured by the EUVM instrument aboard MAVEN. Values are for $L_S \sim 337–343^\circ$ (orbits 1160–1220) and SZA between 30$^\circ$ and 40$^\circ$. $R$ is the linear correlation coefficient.

(Figure 5, top) and CO Cameron (not shown) and the EUV flux. Although this is expected, such simultaneous observations of the dayglow and the local solar forcing have heretofore not been made. Stewart et al. [1972] reported a dependence of CO Cameron band peak intensity with $F_{10.7}$, measured at Earth and corrected for Mars-Earth-Sun angle. Figure 5 (bottom) shows the short-term variation of derived temperature with 17–22 nm irradiance for $L_S \sim 345^\circ$ and SZA = 30–45$^\circ$. The temperatures show a little to no dependence on EUV flux (the slope is statistically insignificant). Stewart et al. [1972] also did not find any correlation between scale height and $F_{10.7}$ flux. The lack of correlation between solar EUV flux and temperature for small subset of data (5 – 18 May 2015) suggests that on short timescales, the temperature variability in the upper atmosphere depends not so much on solar activity but rather on wave and/or tidal activity from lower atmosphere [Stewart et al., 1972; Stiepen et al., 2015]. However, on longer timescales when changes in the solar EUV flux becomes dominant, the mean temperature in the upper atmosphere does respond to solar EUV flux as shown in temperature difference between the two seasons. Forbes et al. [2006] also found the response of Mars thermosphere to the quasiperiodic (27 day) variation of solar flux. In the future, we will use large IUVS dayglow data sets to study the overall impact of solar EUV flux on thermospheric temperatures. This also underscores the need for monitoring the EUV solar flux at Mars when assessing the Martian upper atmosphere response to solar forcing.

4. Conclusions

We report early IUVS/MAVEN observations of Martian UV dayglow spanning two seasons $L_S = 218^\circ$ (18–22 October 2014) and $L_S = 337–352^\circ$ (5 May to 2 June 2015). The altitude profiles for major atmospheric emissions have been generated by binning the spectra in altitude. The average peak emission intensity and peak altitude have decreased by about 30% and ~5 km, respectively, between October 2014 and May 2015, which contributes to a decrease in solar activity and an increase in heliocentric distance between the two seasons. We report the first indication of a lower peak (below 100 km) in OI 297.2 nm emission; this lower peak is produced solely by photodissociation of CO$_2$ by solar Lyman alpha.

We have retrieved the scale height and thermospheric temperatures from dayglow altitude profiles, which shows large variability even within an orbit, consistent with earlier findings [Stewart et al., 1972; Leblanc et al., 2006; Stiepen et al., 2015], indicating a strong coupling between lower and upper Martian atmosphere. The temperatures derived for IUVS measurements for $L_S = 218^\circ$ and $L_S \sim 345^\circ$ are 300 ± 2 K and 250 ± 1.7 K. We observed a positive linear correlation between the CO$_2^+$ UVD and CO Cameron bands peak intensity and simultaneous measured solar 17–22 nm irradiance by the EUV monitor aboard MAVEN. We noticed that solar
forcing is not a very important factor in determining upper atmosphere temperatures for short timescales, but it does affect the mean upper atmosphere temperature over long timescales. This simultaneous monitoring of both the UV dayglow and the local EUV flux represents a significant advance in understanding how planetary atmospheres respond to solar forcing.

References


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