

# Modelling the excavation damaged zone in Callovo-Oxfordian claystone using shear strain localisation

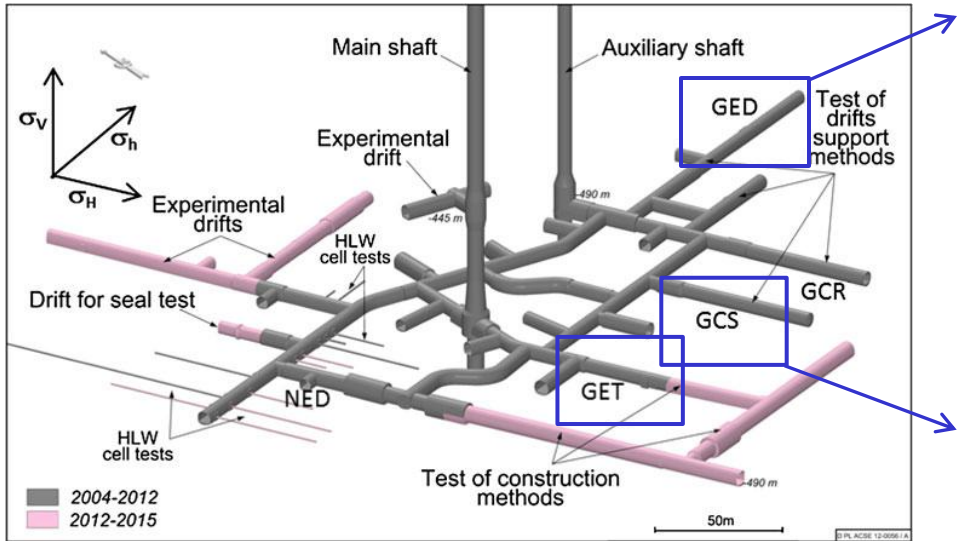
**B. Pardoën - S. Levasseur - F. Collin - D. Seyedi**



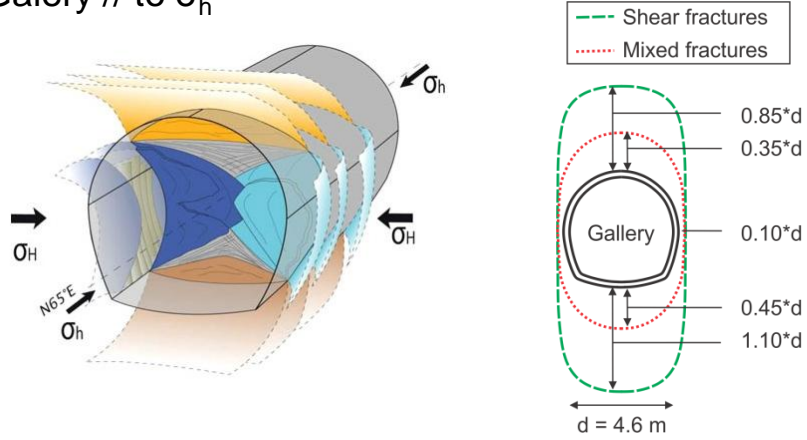
# Excavation damaged zone

**In situ evidences (Andra) :** (Armand et al. 2014)

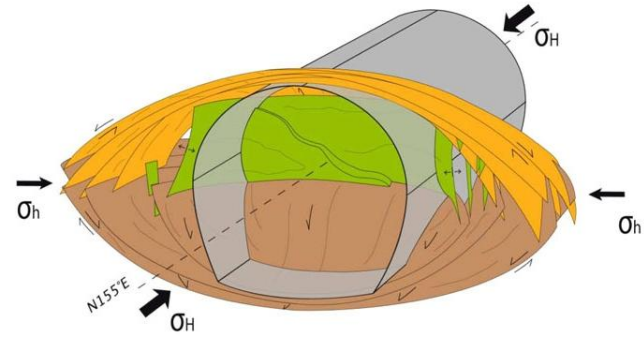
Anisotropy: - stress :  $\sigma_H > \sigma_h \sim \sigma_v$   
 - material : cross-anisotropy



Galery // to  $\sigma_h$



Galery // to  $\sigma_H$



Major issues : prediction of the extension, fracturing structure and properties modifications.

Study :  
 - fractures modelling with shear strain localisation  
 - influence of permeability variation

1. CONSTITUTIVE MODELS
  
2. FRACTURES MODELLING
  - GALLERY // TO  $\sigma_h$
  
  - GALLERY // TO  $\sigma_H$
  
3. PERMEABILITY EVOLUTION

# 1. Constitutive models

## 1.1 Strain localisation with regularization - Coupled 2<sup>d</sup> gradient model : (Chambon *et al.*, 1998 and 2001)

The continuum is enriched with microstructure effects. The kinematics include the classical one (macro) and the microkinematics (Toupin 1962, Mindlin 1964, Germain 1973).

Biphasic porous media : solid + fluid (Collin *et al.*, 2006)

Balance equations for biphasic porous media :

$$\int_{\Omega} \left( \sigma_{ij} \frac{\partial u_i^*}{\partial x_j} + \underline{\Sigma_{ijk} \frac{\partial^2 u_i^*}{\partial x_j \partial x_k}} \right) d\Omega = \int_{\Omega} G_i u_i^* d\Omega + \int_{\Gamma_\sigma} (\bar{t}_i u_i^* + \underline{\bar{T}_i D u_i^*}) d\Gamma$$

$$\int_{\Omega} \left( \frac{\partial M}{\partial t} p_w^* - m_{w,i} \frac{\partial p_w^*}{\partial x_i} \right) d\Omega = \int_{\Omega} Q p_w^* d\Omega + \int_{\Gamma_q} \bar{q} p_w^* d\Gamma$$

Bishop's effective stress :  $\sigma_{ij} = \sigma'_{ij} + b_{ij} S_{rw} p_w$       Double stress :  $\tilde{\Sigma}_{ijk} = f \left( B, \frac{\partial^2 u_i^*}{\partial x_j \partial x_k} \right)$

# 1. Constitutive models

## 1.2 Mechanical model :

Linear elasticity : Cross-anisotropic (5 param.) + Biot's coefficient

$$d\varepsilon_{ij}^e = D_{ijkl}^e d\sigma_{kl}' \quad E_{//}, E_{\perp}, \nu_{////}, \nu_{//\perp}, G_{//\perp}$$

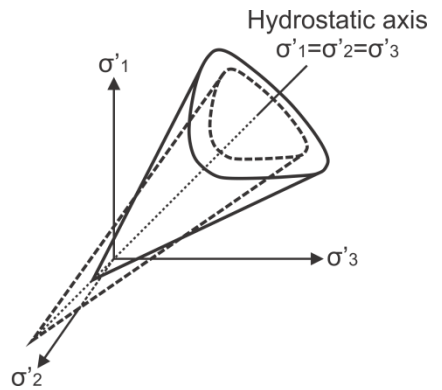
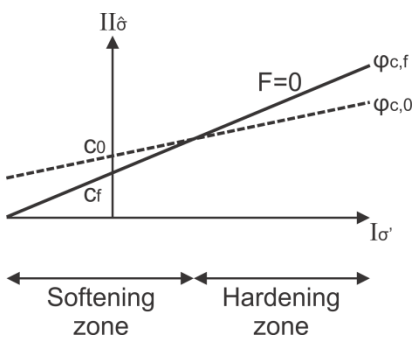
$$b_{ij} = \delta_{ij} - \frac{C_{ijkk}^e}{3K_s} \quad b_{ij} = \begin{bmatrix} b_{//} & & \\ & b_{//} & \\ & & b_{\perp} \end{bmatrix}$$

$$D_{ijkl}^e = \begin{bmatrix} \frac{1}{E_{//}} & -\frac{\nu_{////}}{E_{//}} & -\frac{\nu_{\perp//}}{E_{\perp}} \\ -\frac{\nu_{////}}{E_{//}} & \frac{1}{E_{//}} & -\frac{\nu_{\perp//}}{E_{\perp}} \\ -\frac{\nu_{\perp//}}{E_{//}} & -\frac{\nu_{\perp//}}{E_{//}} & \frac{1}{E_{\perp}} \\ & & & \frac{1+\nu_{////}}{E_{//}} & & \\ & & & & \frac{1}{2G_{//\perp}} & \\ & & & & & \frac{1}{2G_{//\perp}} \end{bmatrix}$$

Plasticity :

Van Eeckelen yield surface  
Hardening/softening of  $\phi/c$  :

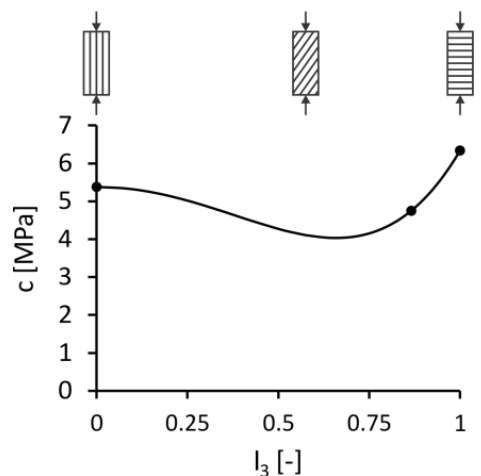
$$F \equiv \Pi_{\hat{\sigma}} - m \left( I_{\sigma'} + \frac{3c}{\tan \phi_c} \right) = 0$$



Cohesion anisotropy :

$$c = a_{ij} l_i l_j = \bar{c} \left( 1 + A_{11}(1 - 3I_3^2) + b_1 A_{11}^2 (1 - 3I_3^2)^2 + \dots \right)$$

$$l_i = \sqrt{\frac{\sigma_{i1}'^2 + \sigma_{i2}'^2 + \sigma_{i3}'^2}{\sigma_{ij}' \sigma_{ij}'}}$$



# 1. Constitutive models

## 1.3 Flow model :

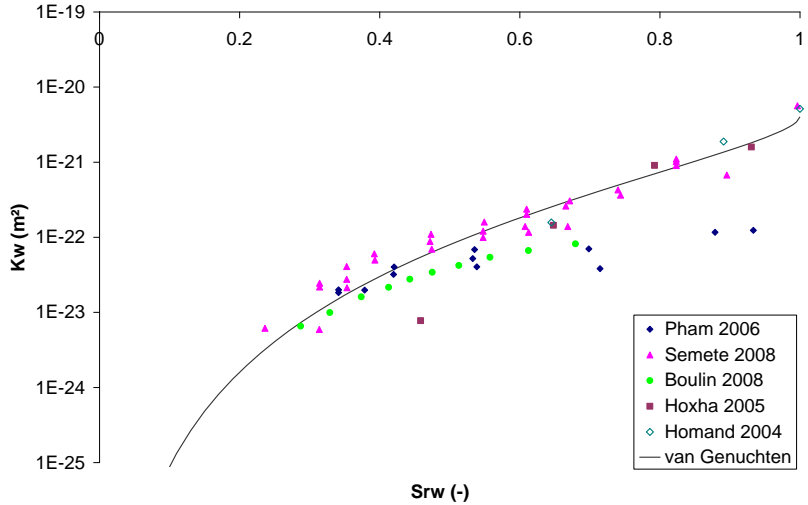
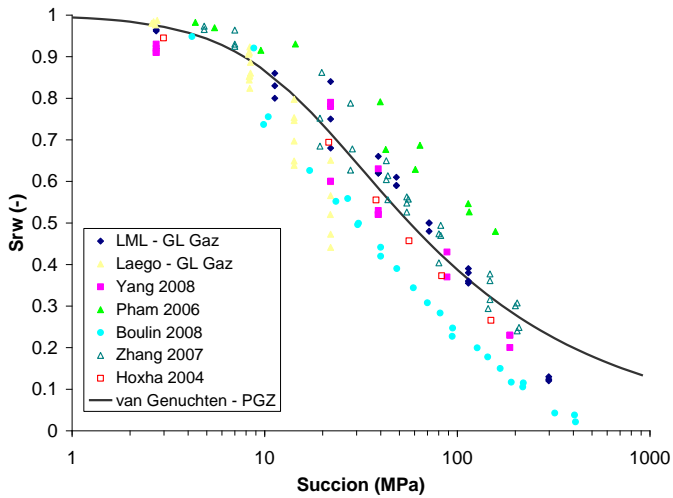
Advection of liquid phase (Darcy's flow) :

$$m_{w,i} = -\rho_w \frac{k_{ij} k_{r,w}}{\mu_w} \frac{\partial p_w}{\partial x_j}$$

Water retention and permeability curves (Van Genuchten's model) :

$$S_{r,w} = S_{res} + (S_{max} - S_{res}) \left[ 1 + \left( \frac{p_c}{P_r} \right)^n \right]^{-m}$$

$$k_{r,w} = \sqrt{S_{r,w}} \left[ 1 - (1 - S_{r,w}^{1/m})^m \right]^2$$



# 2. Fractures modelling

## 2.1 Gallery // to $\sigma_h$ :

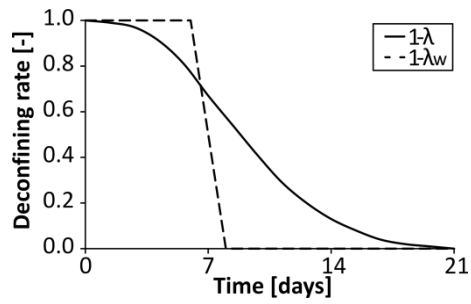
Anisotropic stress state, isotropic model

Initial anisotropic stress state(Andra URL) :

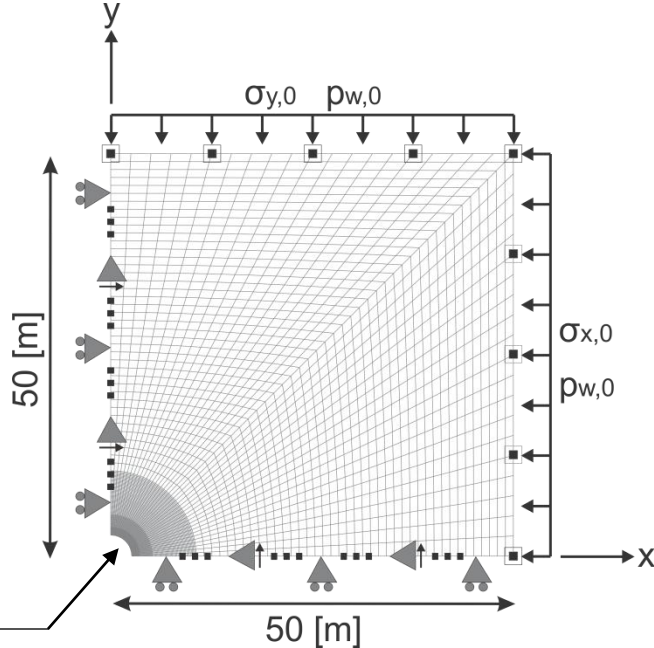
$$\begin{aligned}
 p_{w,0} &= 4.5 \text{ [MPa]} \\
 \sigma_{v,0} &= \sigma_{h,0} = 12 \text{ [MPa]} \\
 \sigma_{H,0} &= 1.3 \sigma_{v,0} = 15.6 \text{ [MPa]}
 \end{aligned}$$

HM modelling in 2D plane strain state (LAGAMINE-Ulg)

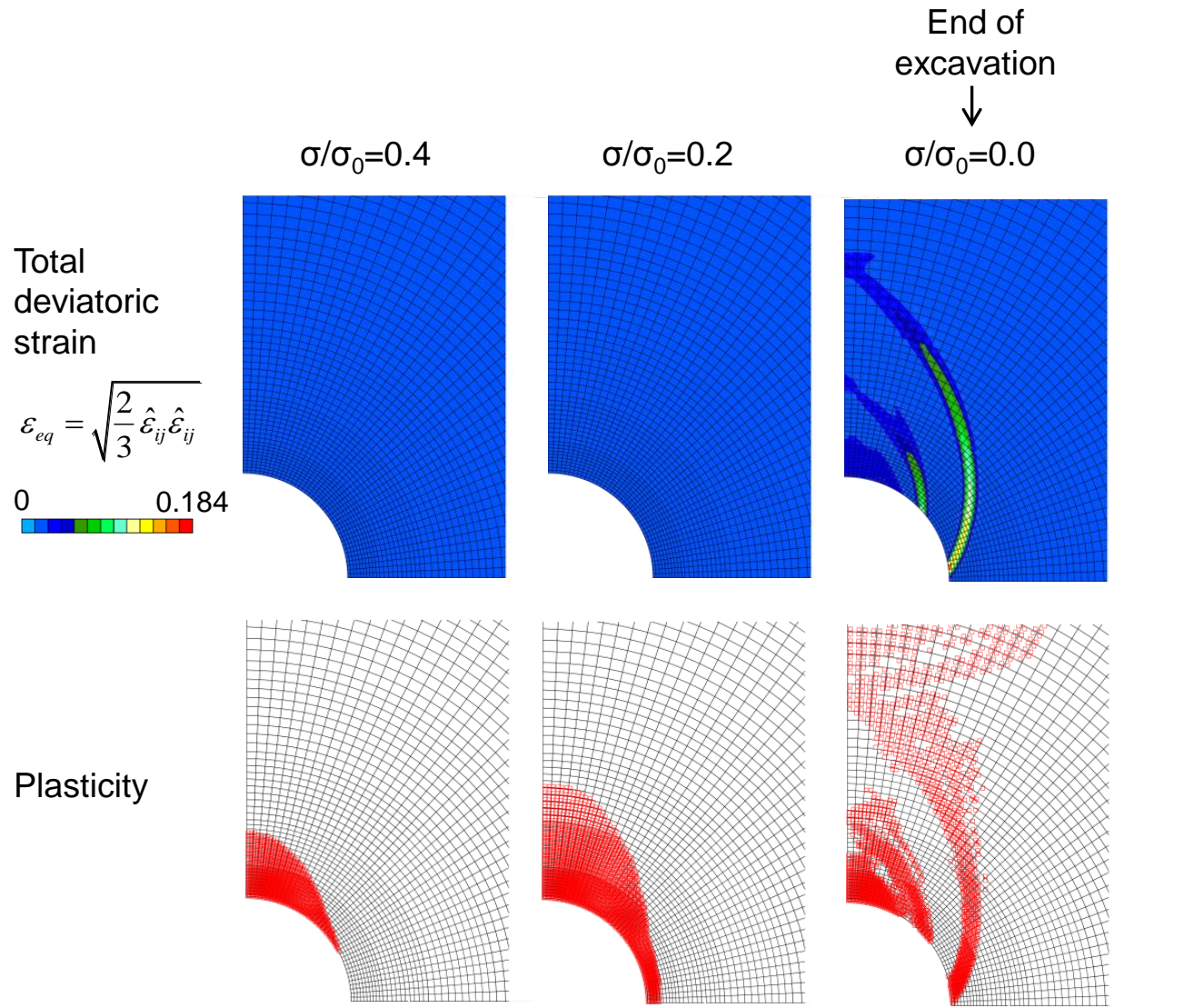
Excavation :  $\sigma_r^\Gamma = (1 - \lambda) \sigma_{r,0}^\Gamma$   
 $p_w^\Gamma = (1 - \lambda_w) p_{w,0}^\Gamma$



- ▣ Constant pore water pressure ( $p_{w,0}$ )
- ← Constant total stress ( $\sigma_{y,0} / \sigma_{x,0}$ )
- ▶ Constrained displacement perpendicular to the boundary
- ▲ Constrained normal derivative of the radial displacement
- ⋯ Impervious boundary



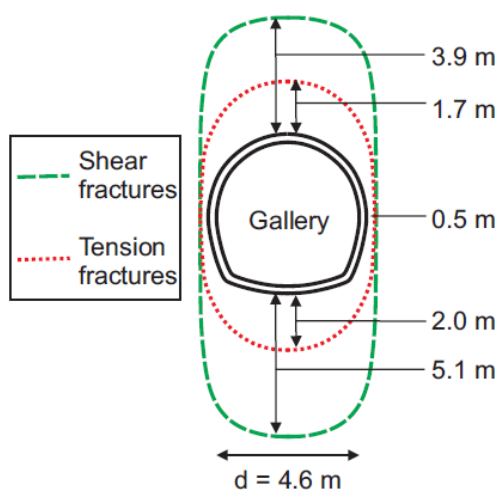
# 2. Fractures modelling



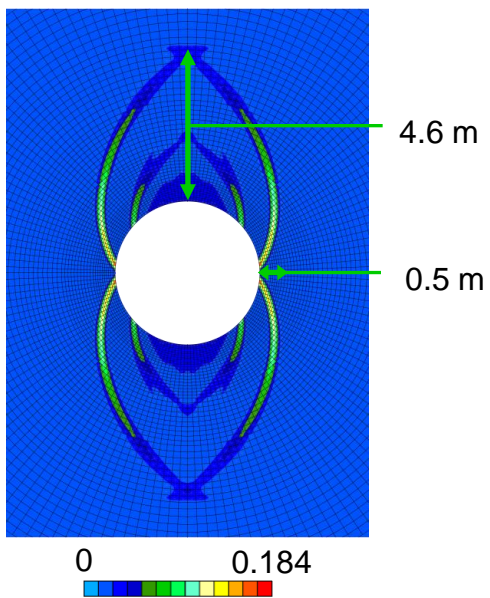


# 2. Fractures modelling

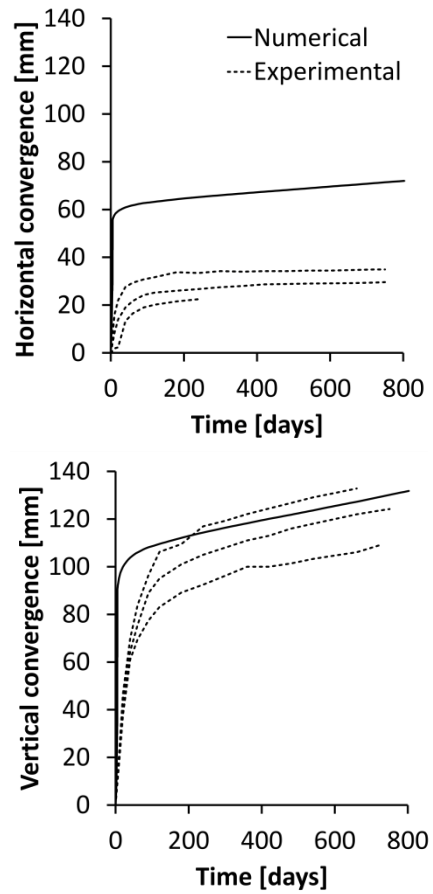
Fractures



Total deviatoric strain



Convergence



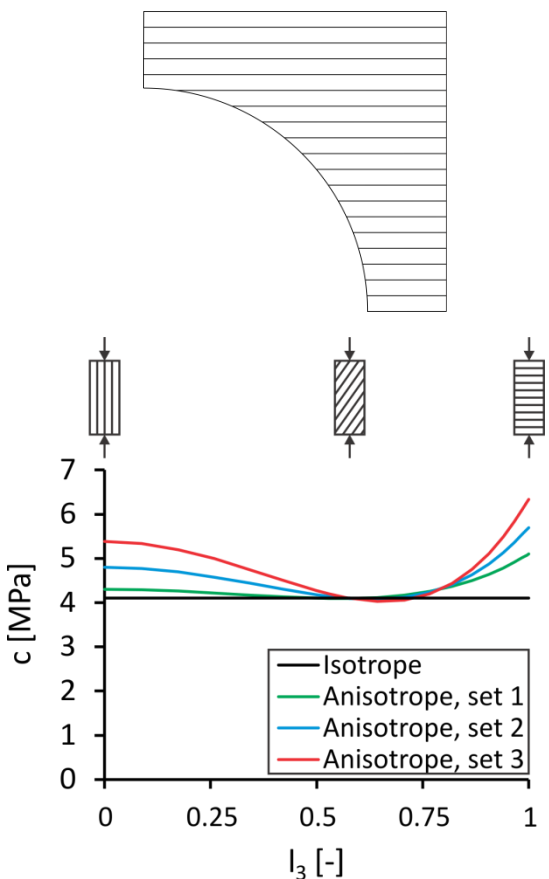
$\sigma$  anisotropy is the predominant factor leading to strain localisation and to the elliptical shape of the damaged zone.

# 2. Fractures modelling

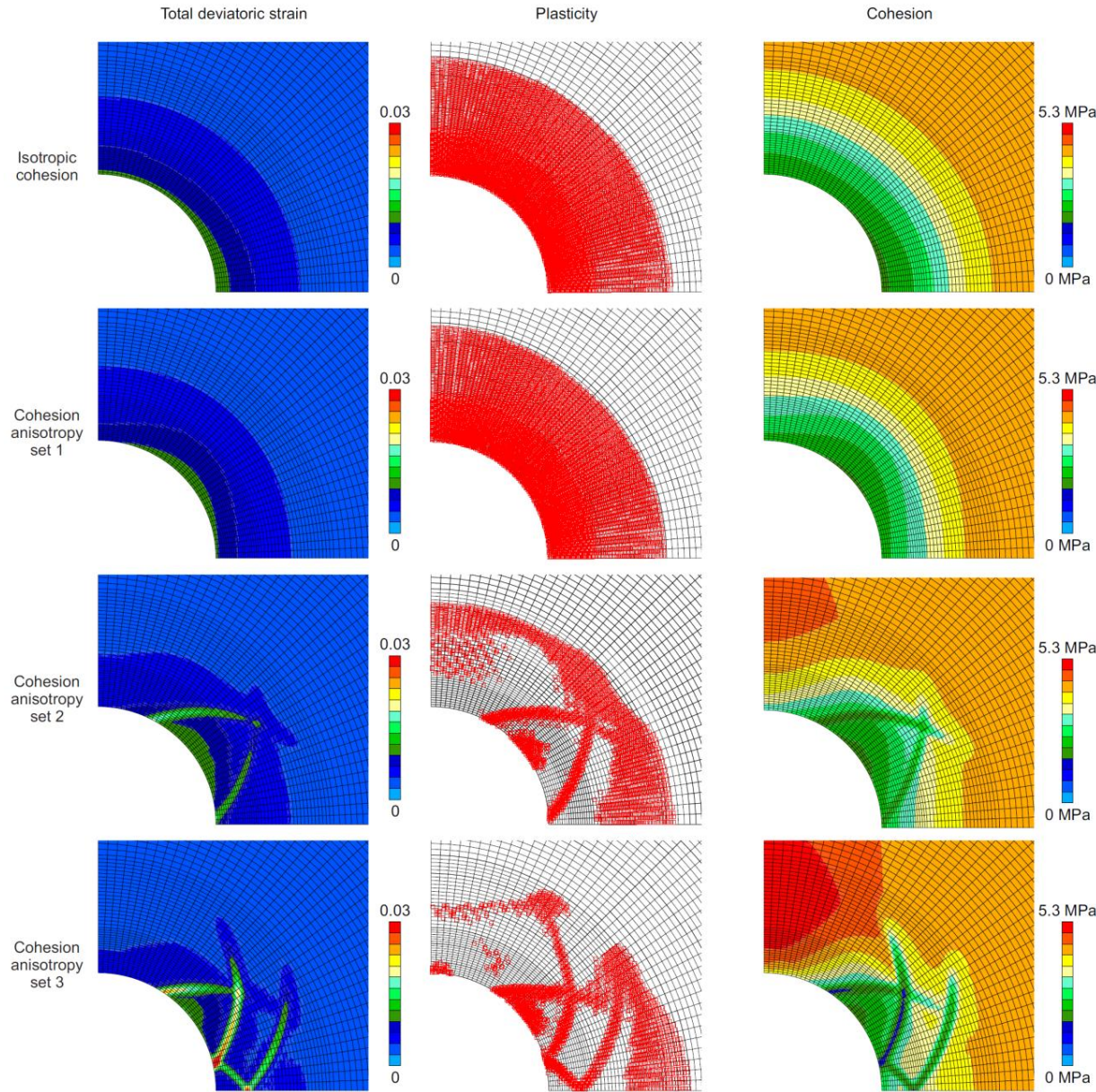
## 2.2 Gallery // to $\sigma_H$ :

Isotropic stress state ( $\sigma=12$  MPa),  
anisotropic model

HM modelling in 2D plane  
strain state



### End of excavation



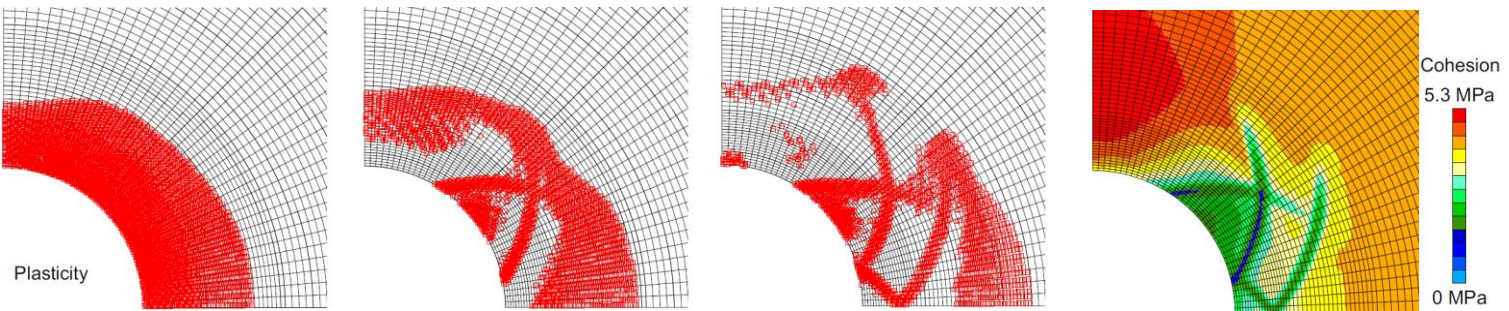
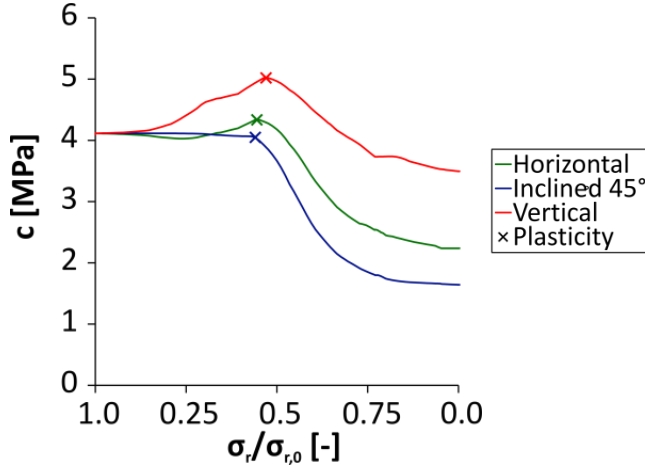
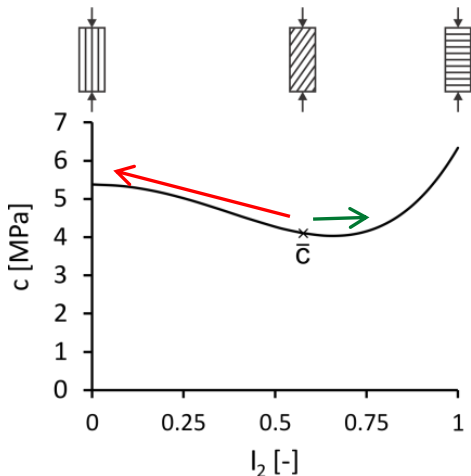
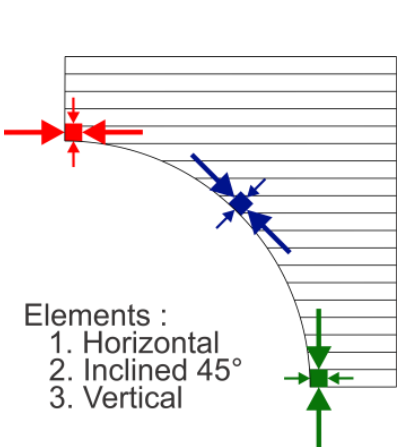
# 2. Fractures modelling

## Cohesion evolution :

Anisotropy:  $c = a_{ij} l_i l_j = \bar{c} (1 + A_{11}(1-3l_2^2) + b_1 A_{11}^2(1-3l_2^2)^2 + \dots)$   $l_i = \sqrt{\frac{\sigma_{i1}^2 + \sigma_{i2}^2 + \sigma_{i3}^2}{\sigma'_{ij} \sigma'_{ij}}}$

Initially : isotropic  $\sigma_{ij} \rightarrow c = \bar{c}$   $l_2 = \sqrt{3}/3 = 0.58$

Excavation :  $\sigma_r \downarrow$  and  $\sigma_{ort} \uparrow$

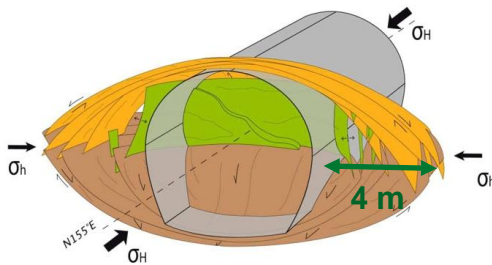


## 2. Fractures modelling

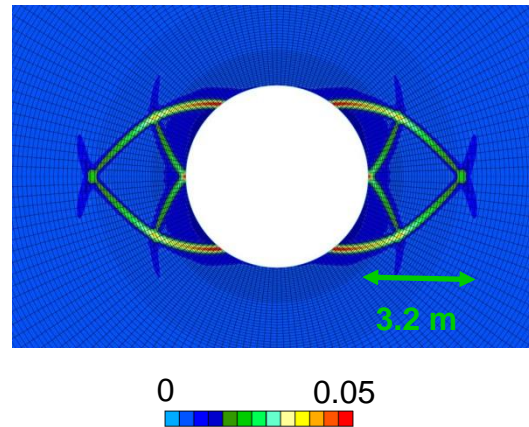
Anisotropic stress state, anisotropic model

$$\sigma_{H,0} = 1.3 \sigma_{v,0} > \sigma_{v,0} = \sigma_{h,0} = 12 \text{ [MPa]}$$

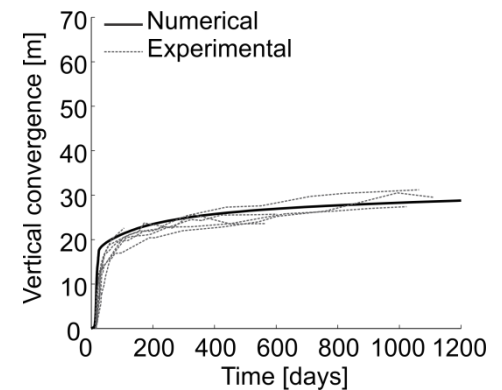
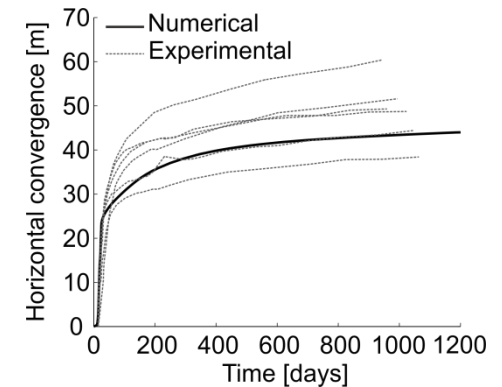
Fractures



Total deviatoric strain



Convergence



Isotropic stress state in the gallery section does not lead to shear strain localisation unless the material anisotropy is considered.

Material anisotropy seems to be the predominant factor leading to strain localisation and to the elliptical shape of the damaged zone.

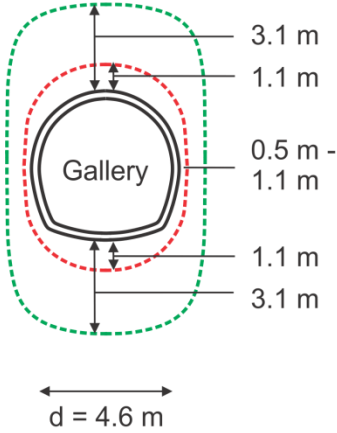
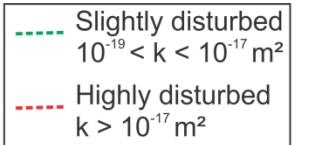
# 3. Permeability evolution

## 3.1 In situ evidences :

Hydraulic properties are not homogeneous in the damaged zone.

Influence of rock fracturing on intrinsic permeability.

*In situ* permeability in Callovo-Oxfordian claystone  
(Armand et al. 2014, Cruchaudet et al. 2010b)



## 3.2 Permeability variation :

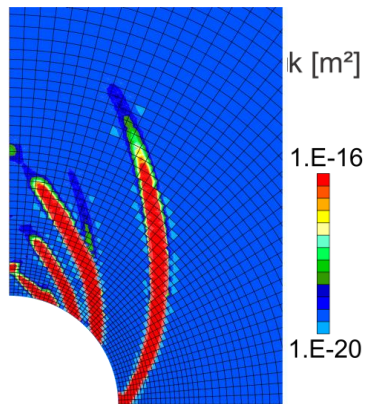
Total deviatoric strain (if  $\epsilon_{eq} > \epsilon_{eq}^{min}$ )

$$\frac{k_{ij}}{k_{ij,0}} = 1 + \alpha(\epsilon_{eq} - \epsilon_{eq}^{min})^\beta$$

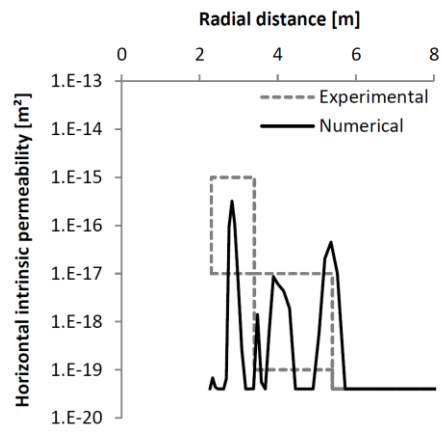
$$\epsilon_{eq} = \sqrt{\frac{2}{3} \hat{\epsilon}_{ij} \hat{\epsilon}_{ij}}$$

$$\alpha = 2 \times 10^8, \beta = 3$$

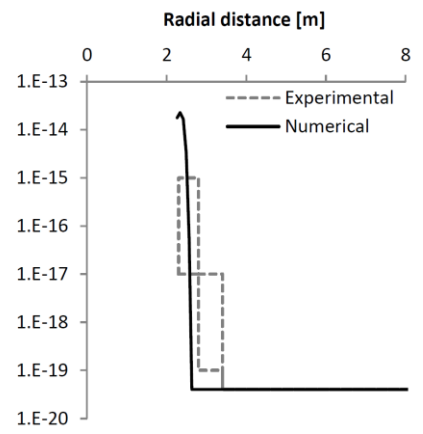
End of excavation



Vertical y-axis



Horizontal x-axis



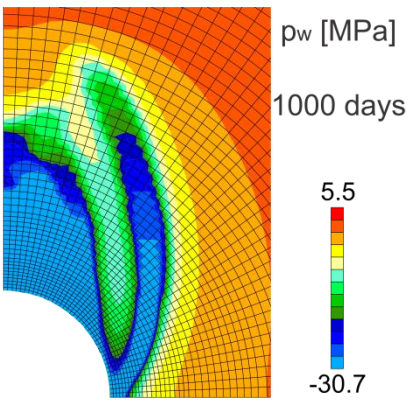
# 3. Permeability evolution

## 3.3 Rock-atmosphere interaction (gallery air ventilation) :

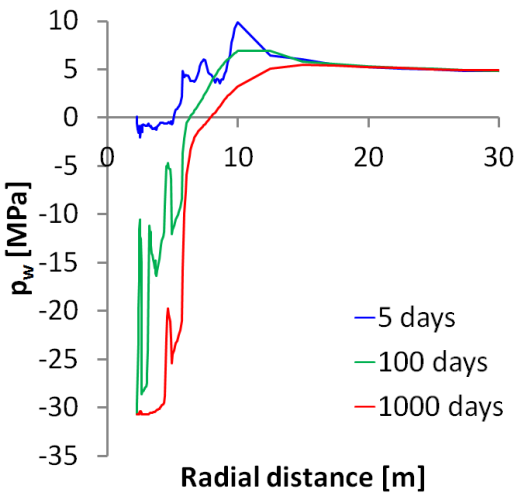
Water phases equilibrium at gallery wall, Kelvin's law : 
$$RH = \frac{p_v}{p_{v,0}} = \exp\left(\frac{-p_c M_v}{RT \rho_w}\right)$$

Air RH = 80%,  $p_w = -30.7$  [MPa]

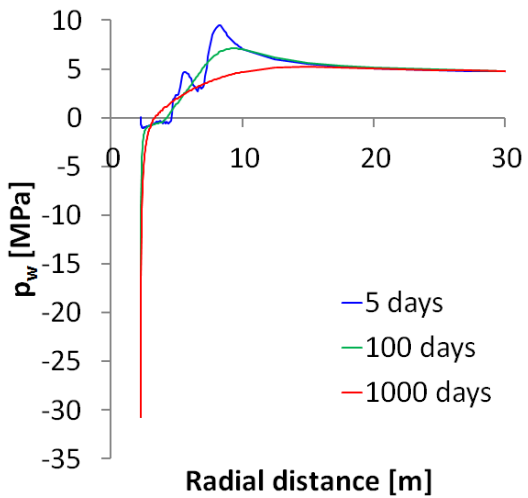
End of calculation  
 $k(\epsilon_{eq})$



Vertical y-axis  
 $k(\epsilon_{eq})$



Vertical y-axis  
 $k = cst$



# Conclusions

- Damaged zone → strain localisation zone similar to *in situ* measurements
- modelling provide information about the rock structure and evolution within this zone, as observed *in situ*.
- rock anisotropy and properties modification

