



Evolution of Antarctic Surface Mass Balance by high-resolution downscaling of LMDZ4 AGCM and contribution to sea-level change

Cécile AGOSTA

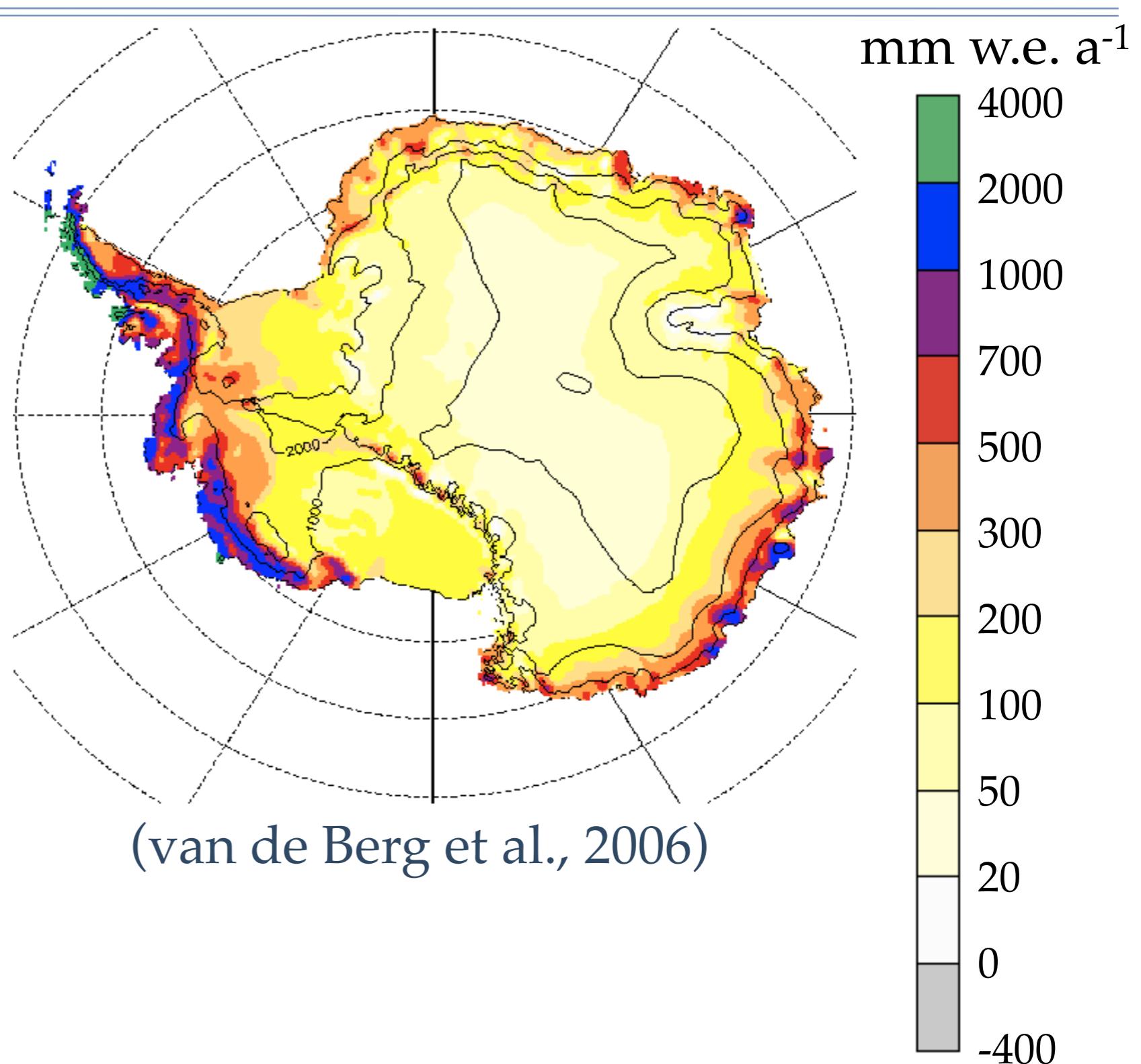


Laboratoire de Glaciologie et Géophysique de l'Environnement

Vincent Favier, Christophe Genthon, Gerhard Krinner, Hubert Gallée

Antarctic Surface Mass Balance

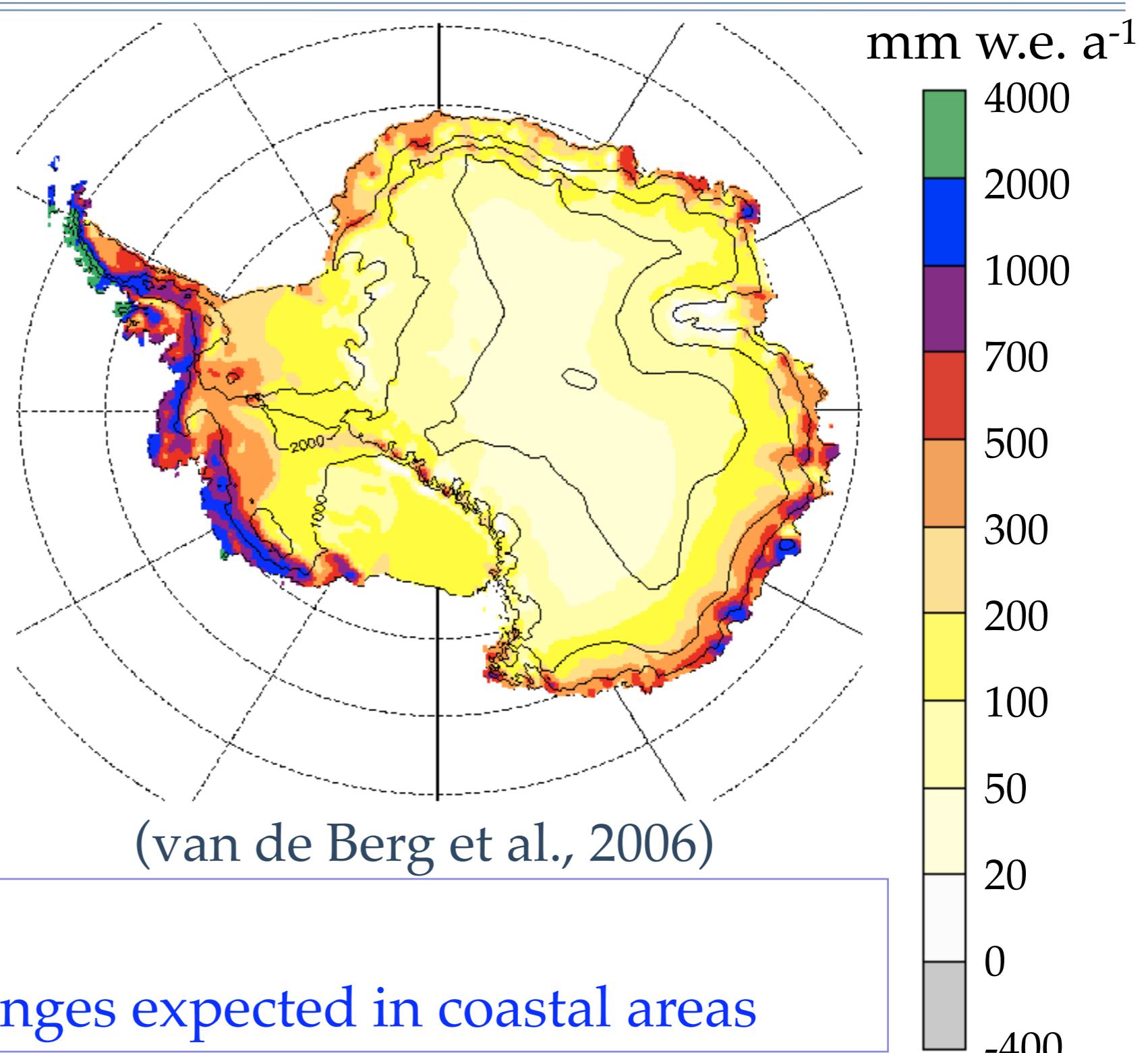
- ▶ Coastal areas :
snowy and windy
- ▶ Antarctic Plateau :
cold and dry



Antarctic Surface Mass Balance

► Coastal areas :
snowy and windy

► Antarctic Plateau :
cold and dry



► Climate models:
major SMB changes expected in coastal areas

SMB Downscaling : Why ?

SMB estimation

Precipitation, Sublimation, Melting, Refreezing, Blowing snow

Spatial extent

Antarctica (5600 km x 5600 km)

Time extent

~800 yrs (1980-2200 * 2 Scenarios * 2 Boundary conditions)

GCM resolution : ~ 60 km



Required resolution : ≤ 15 km

→ Reduced computation time needed

The HiDEP model

High-Resolution Downscaling of surface Energy balance and Precipitation

INPUTS (~50 KM RESOLUTION)

GCM Outputs :
 P, T, Qv, U, V, W

3D Fields

Time step : 6H

Surface Fields

Time step : 3H

High-resolution topography

HiDEP

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Rain_{HiDEP}

Snow_{HiDEP}

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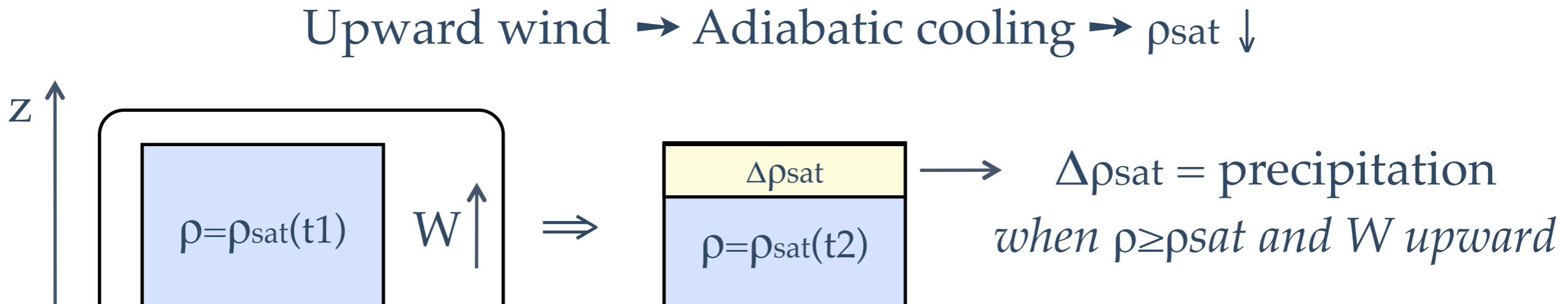
$Snow_{\text{HiDEP}}$

$Sublimation_{\text{HiDEP}}$

$Melting_{\text{HiDEP}}$

$Refreezing_{\text{HiDEP}}$ 3

Precipitation downscaling : an orographic precipitation model

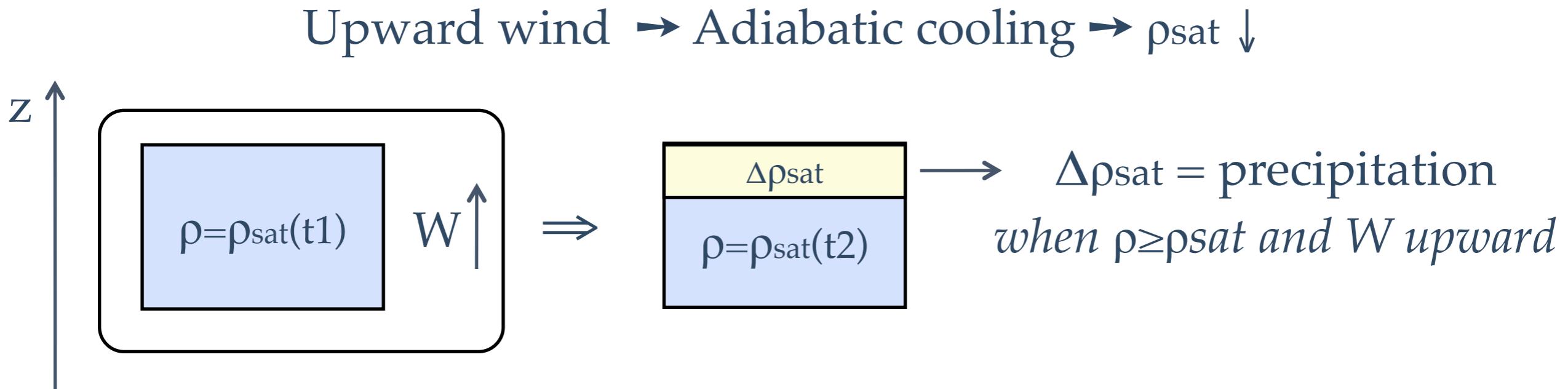


Integration of the Clausius-Clapeyron equation at saturation :

$$\Delta\rho_{\text{sat}} / \Delta t = F(\rho_{\text{sat}}, T, P) \times W$$

when $\rho \geq \rho_{\text{sat}}$ and W upward

Precipitation downscaling : an orographic precipitation model



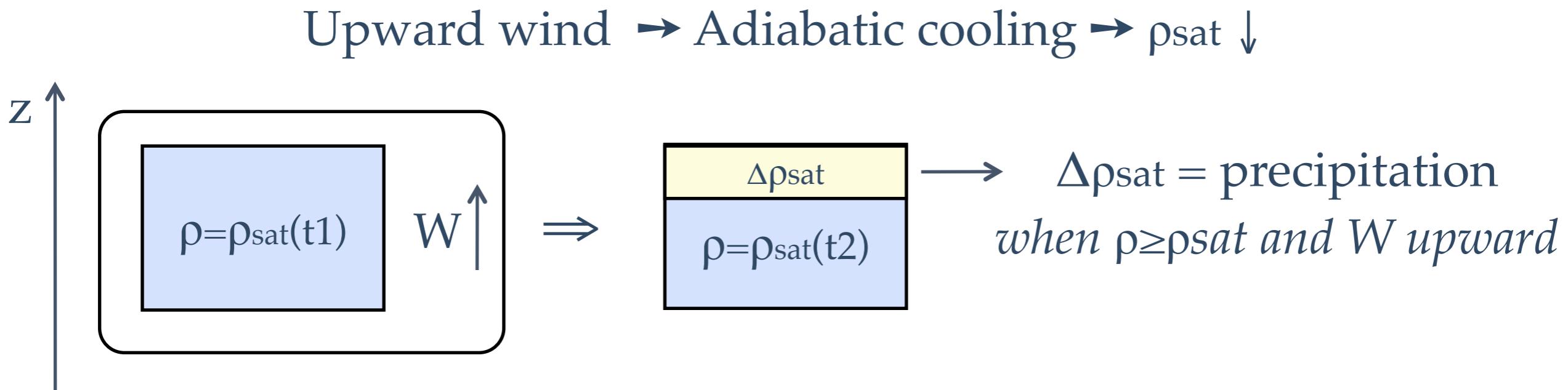
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- Feedback on Q_{sat}, T, P
- Time-delay for hydrometeor formation
- Hydrometeor advection

Precipitation downscaling : an orographic precipitation model



Integration of the Clausius-Clapeyron equation at saturation :

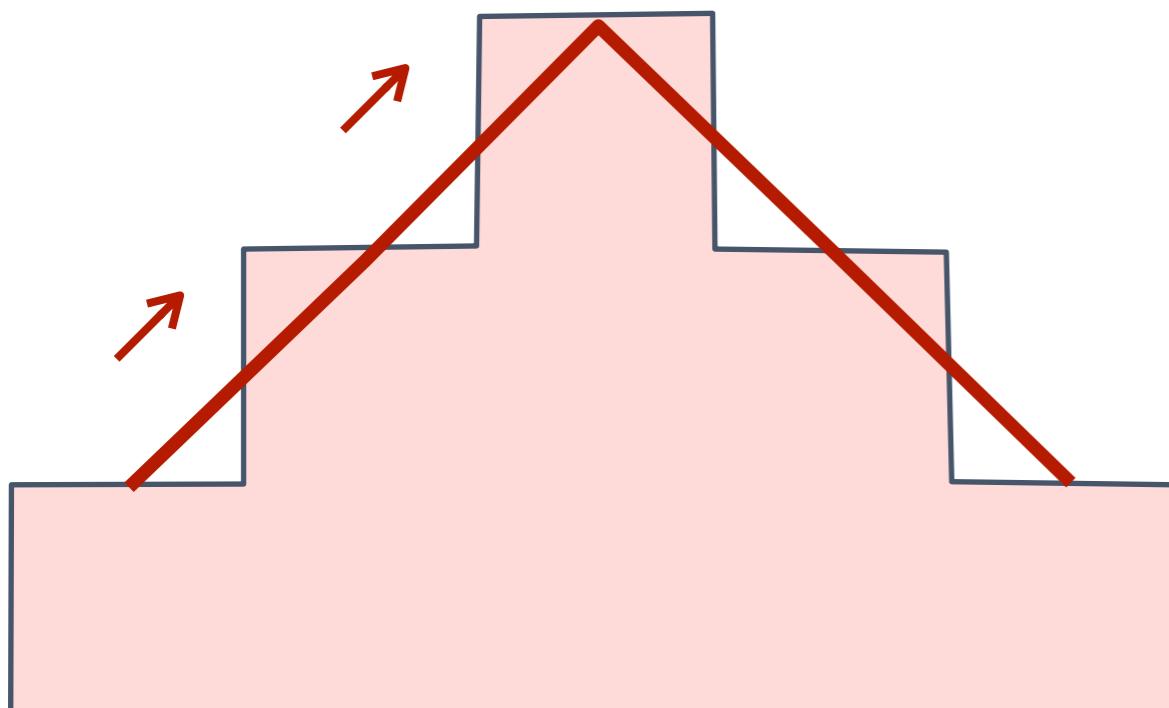
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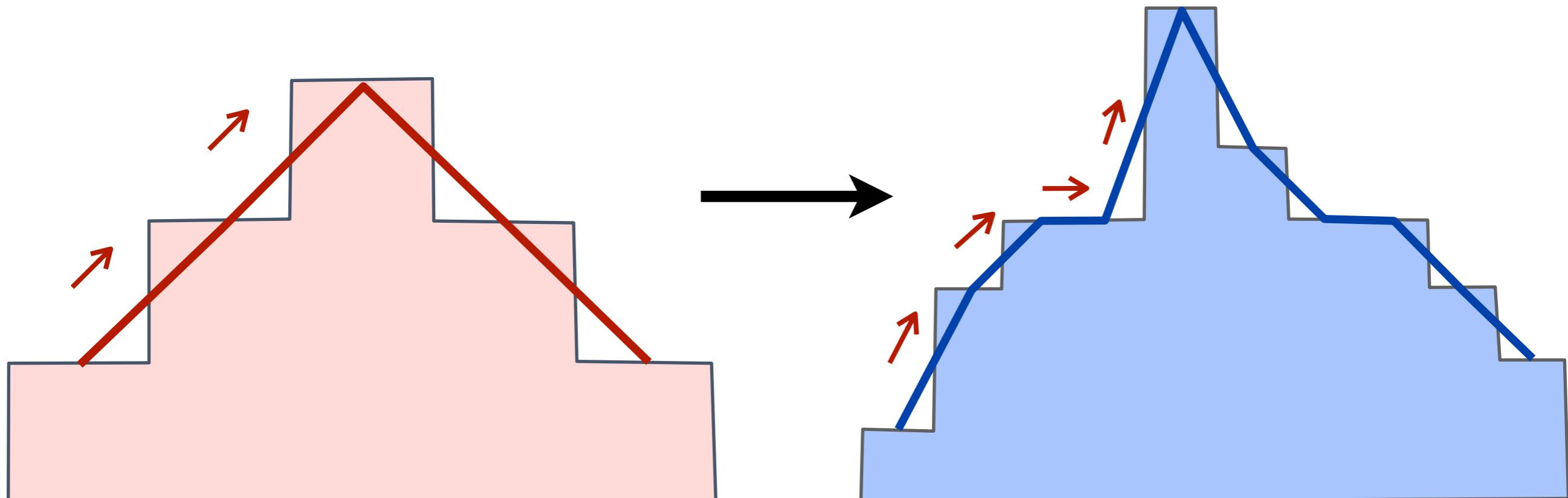
Orographic precipitation : Determination of the vertical wind W

At the surface : the wind is tangent to the topography



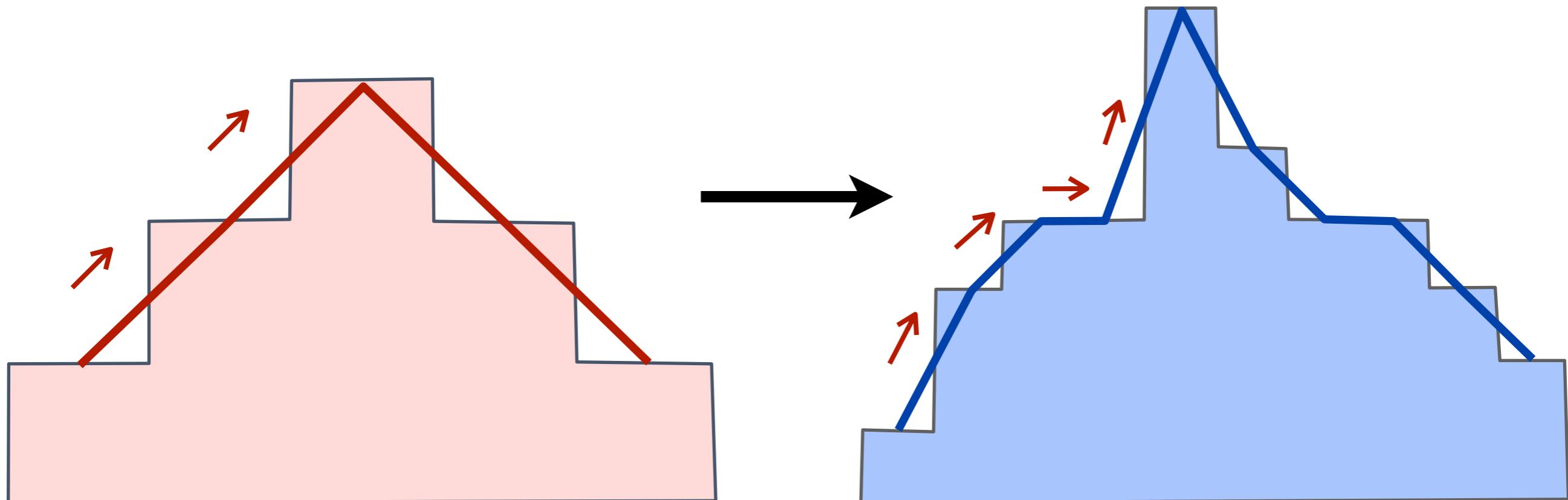
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Orographic precipitation : Determination of the vertical wind W

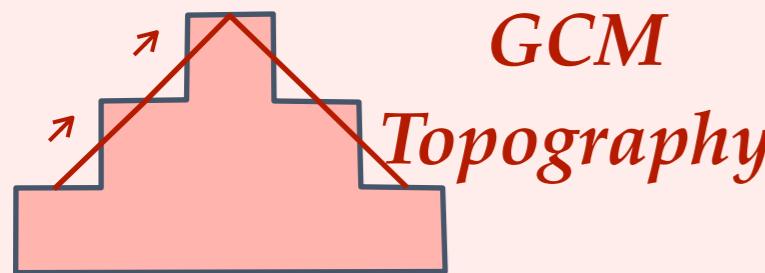
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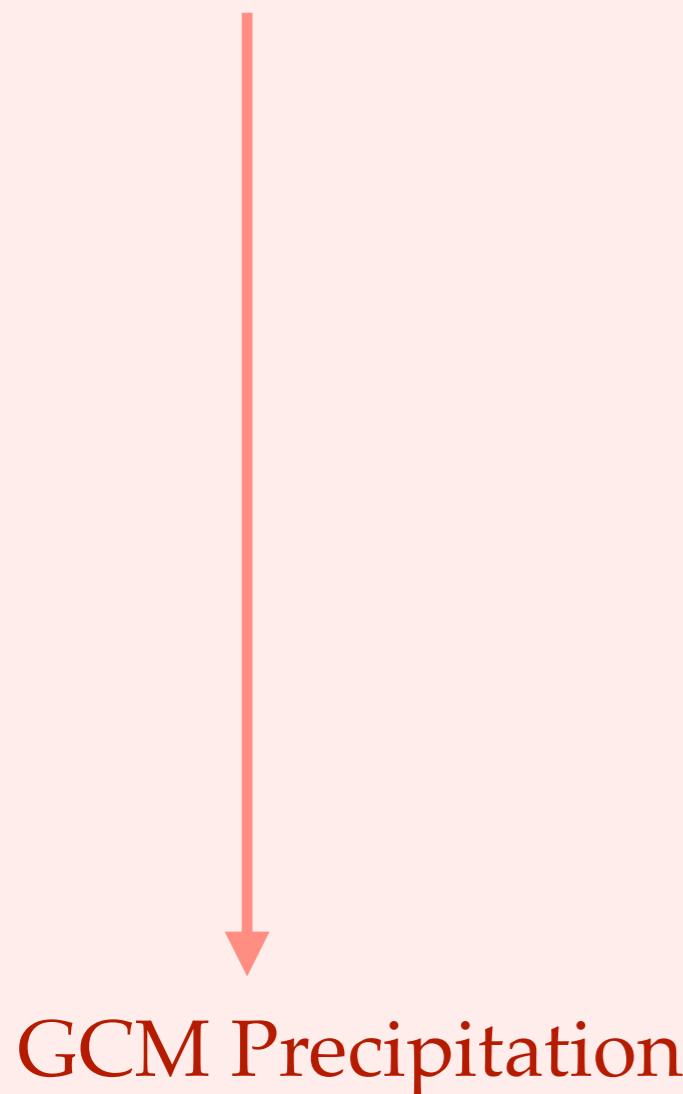
→ Computation for W : resolution of mountain gravity wave

Total precipitation : Orographic + Non-Orographic

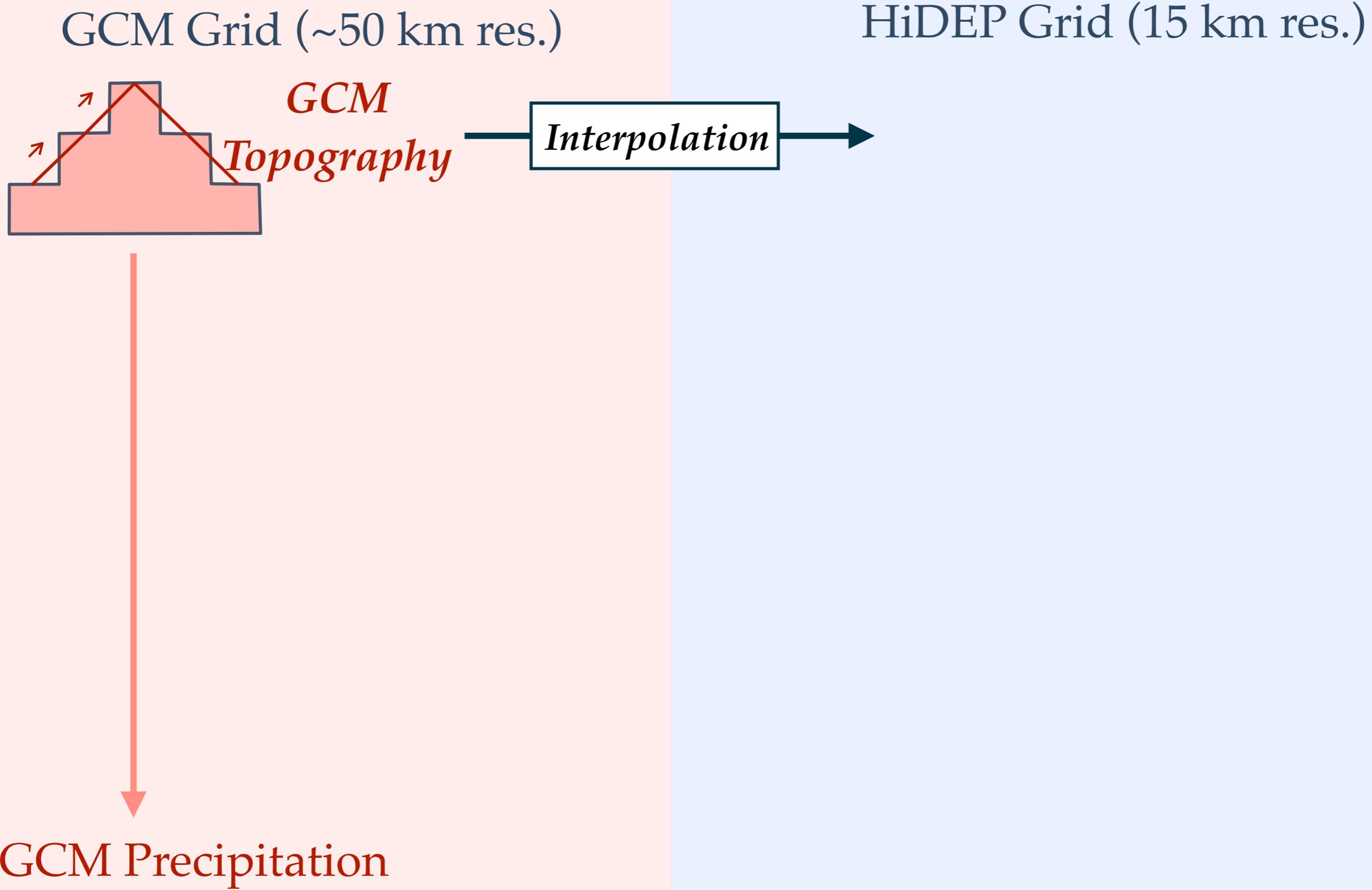
GCM Grid (~50 km res.)



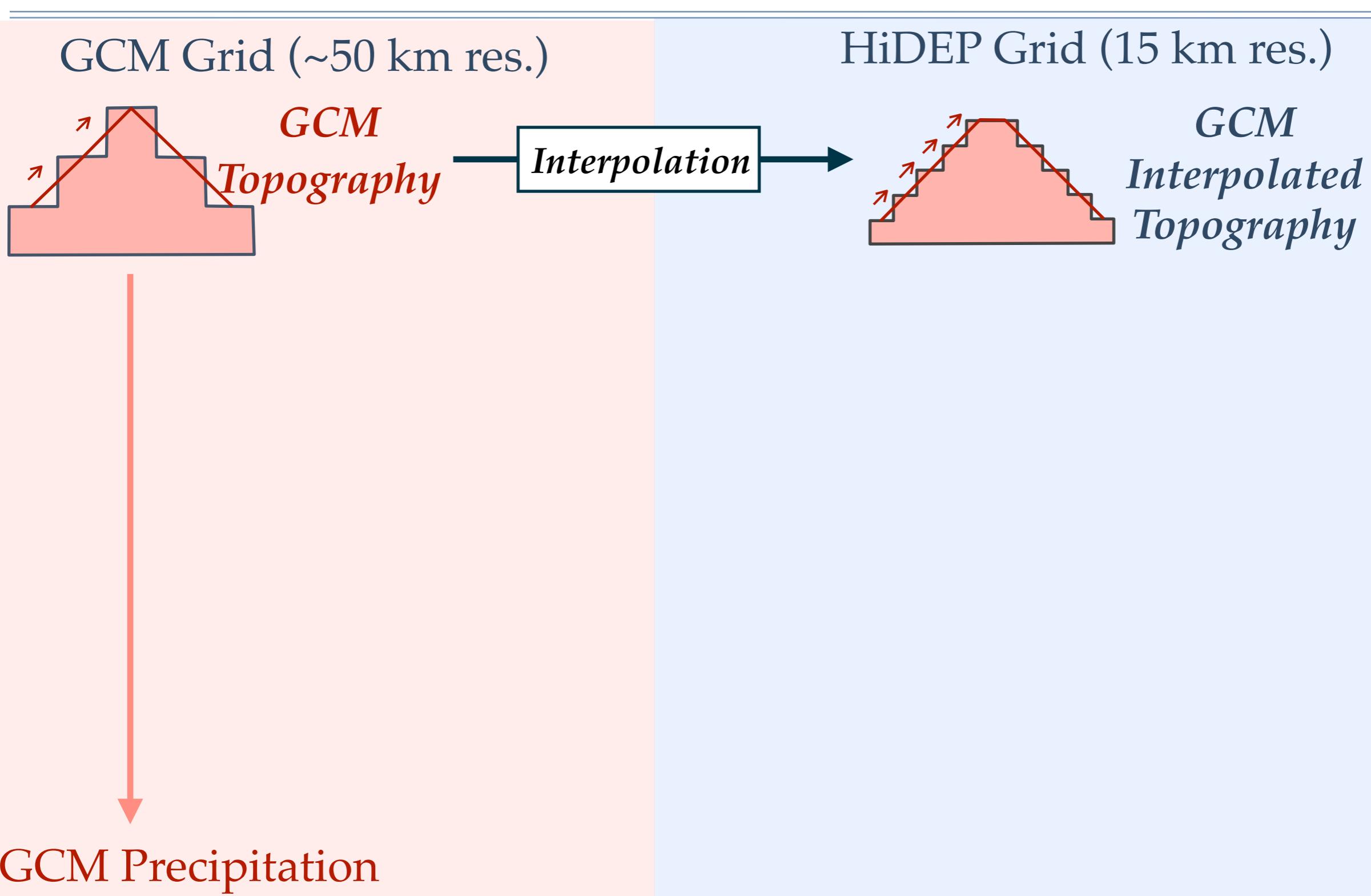
HiDEP Grid (15 km res.)



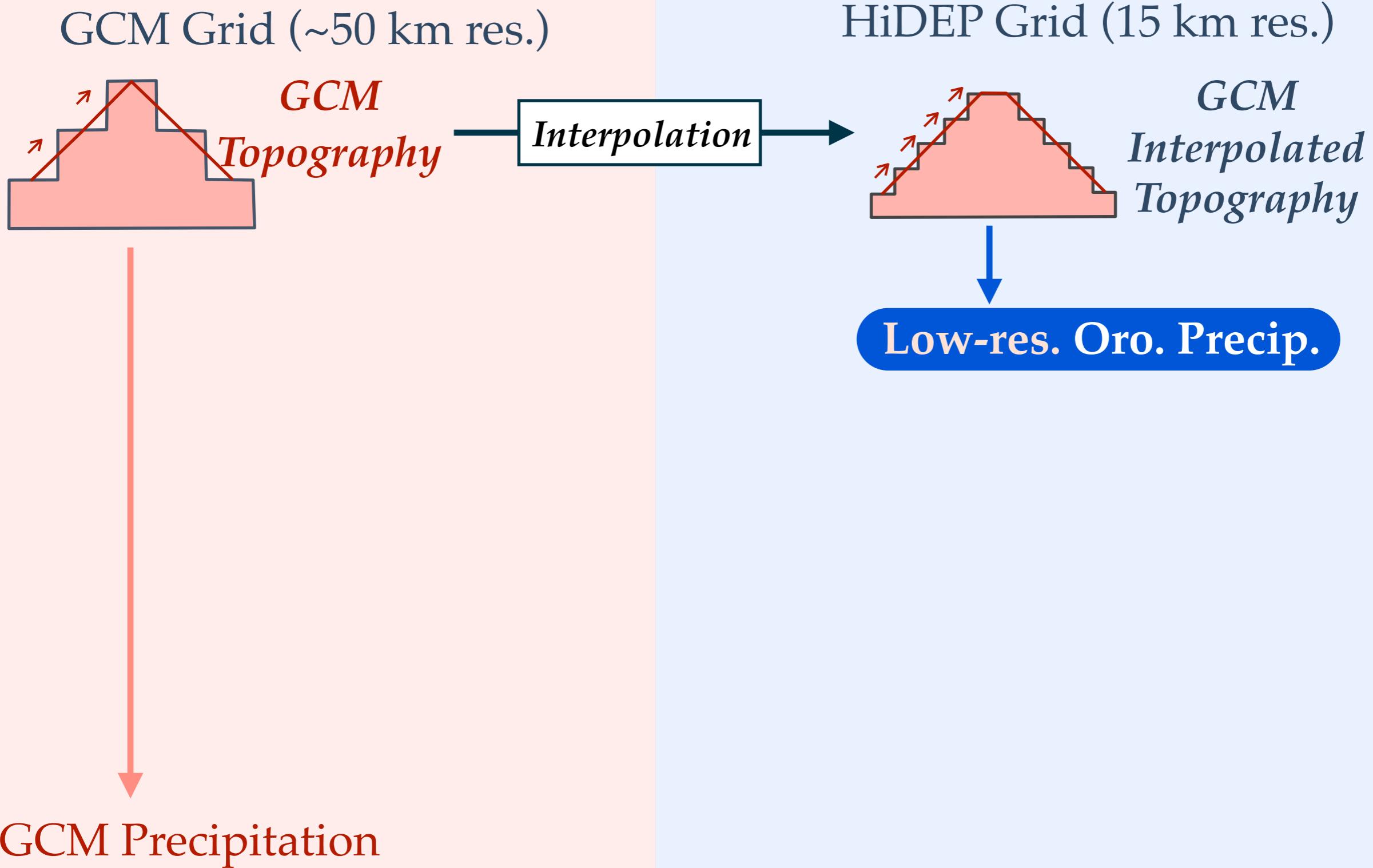
Total precipitation : Orographic + Non-Orographic



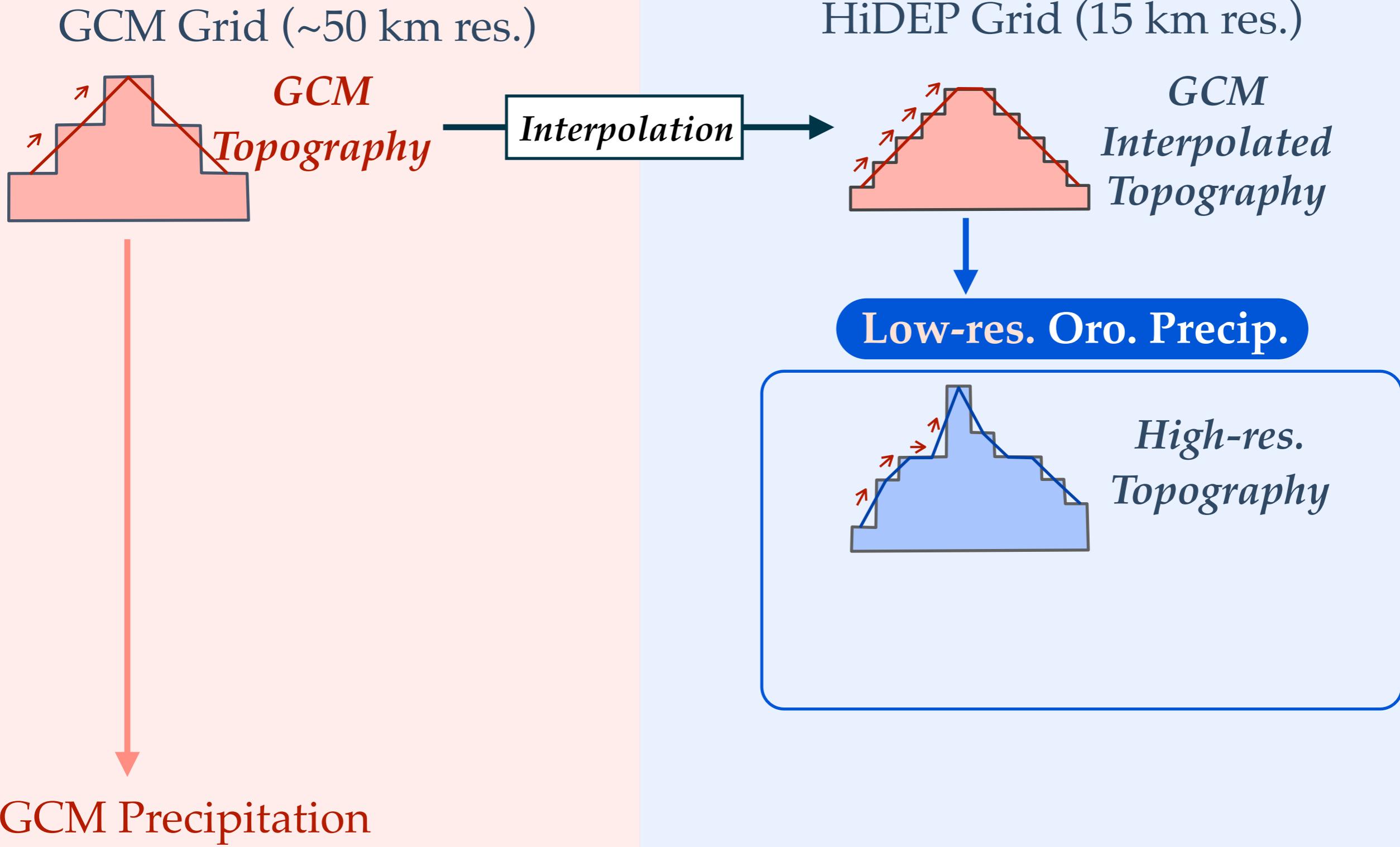
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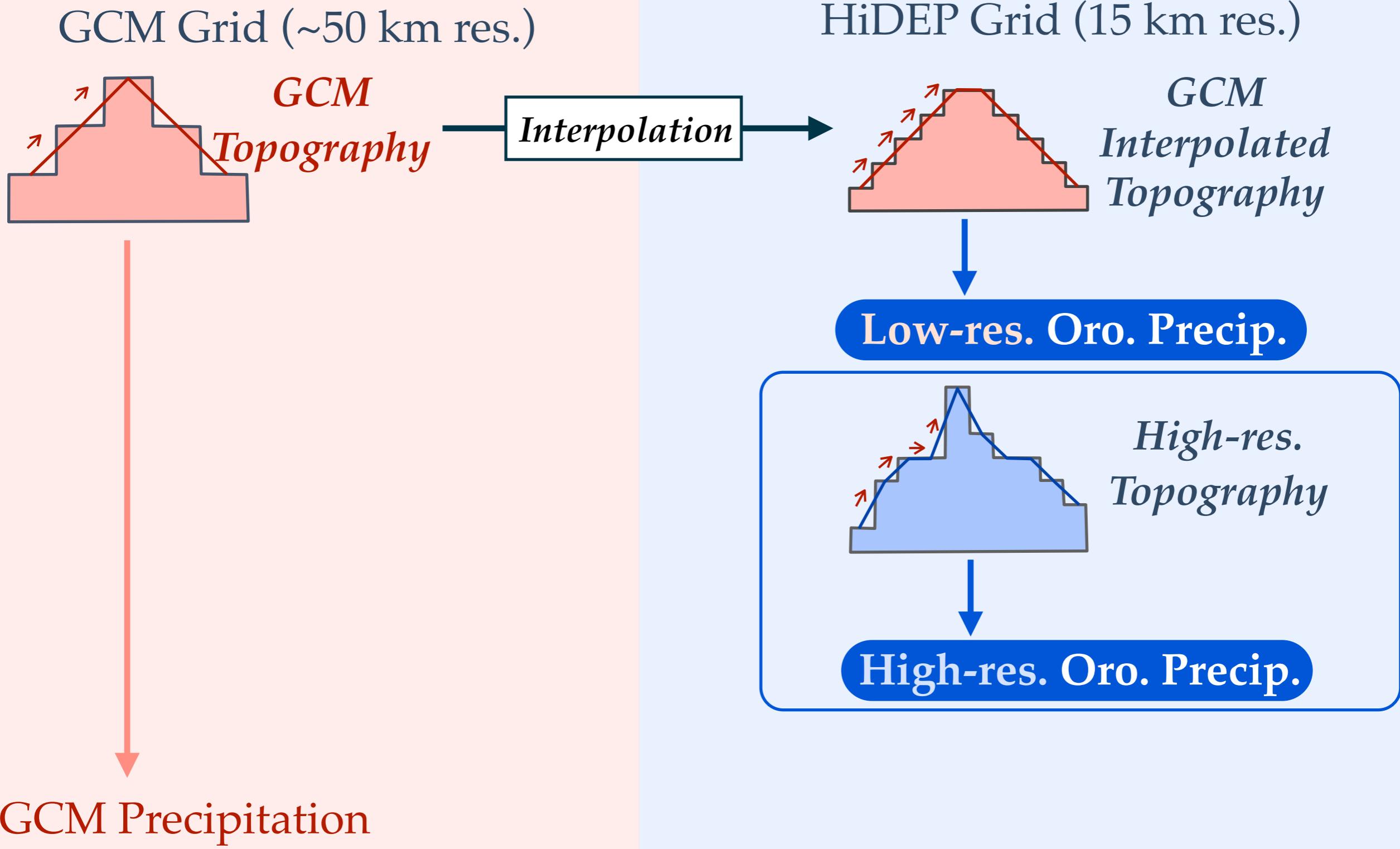
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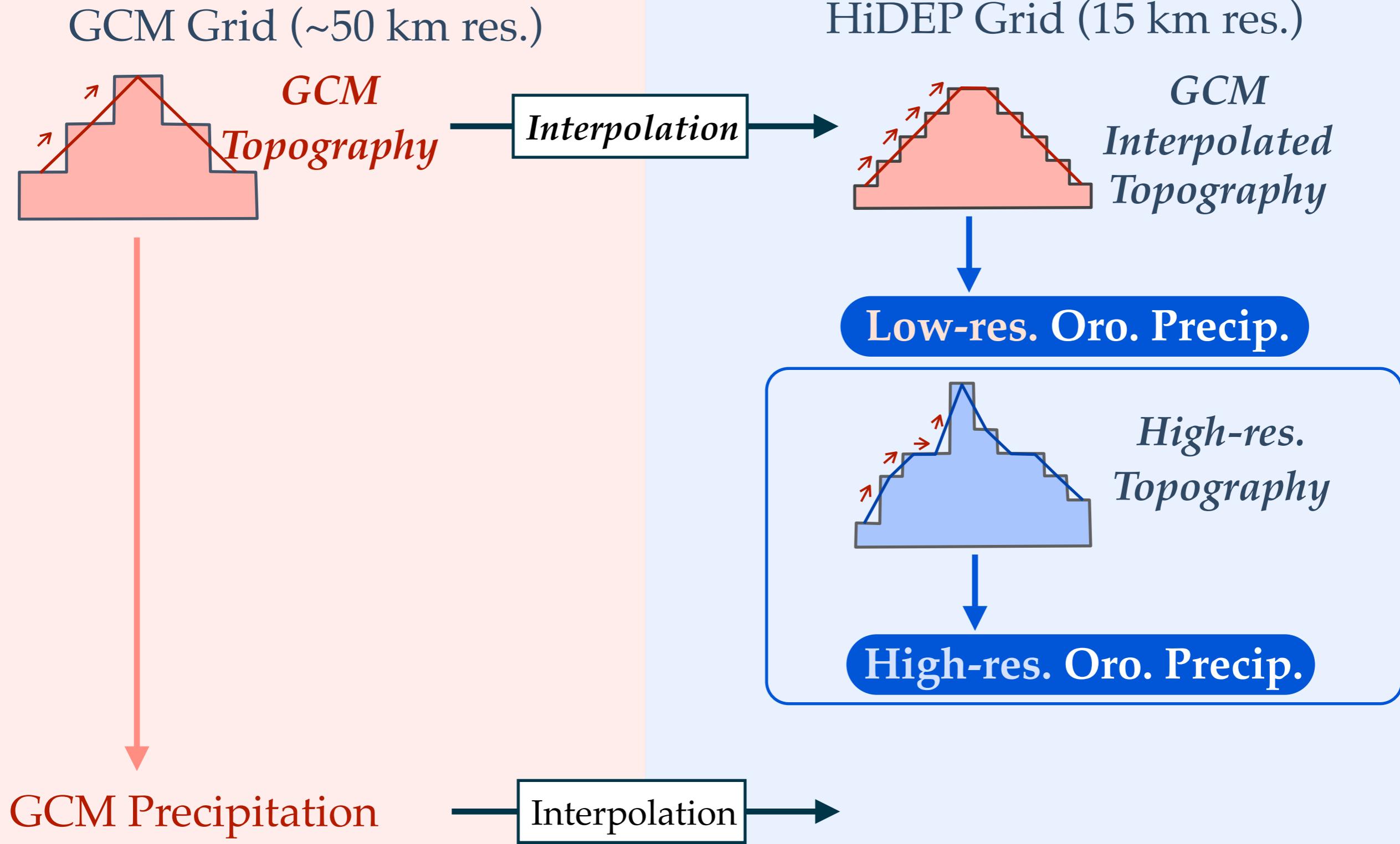
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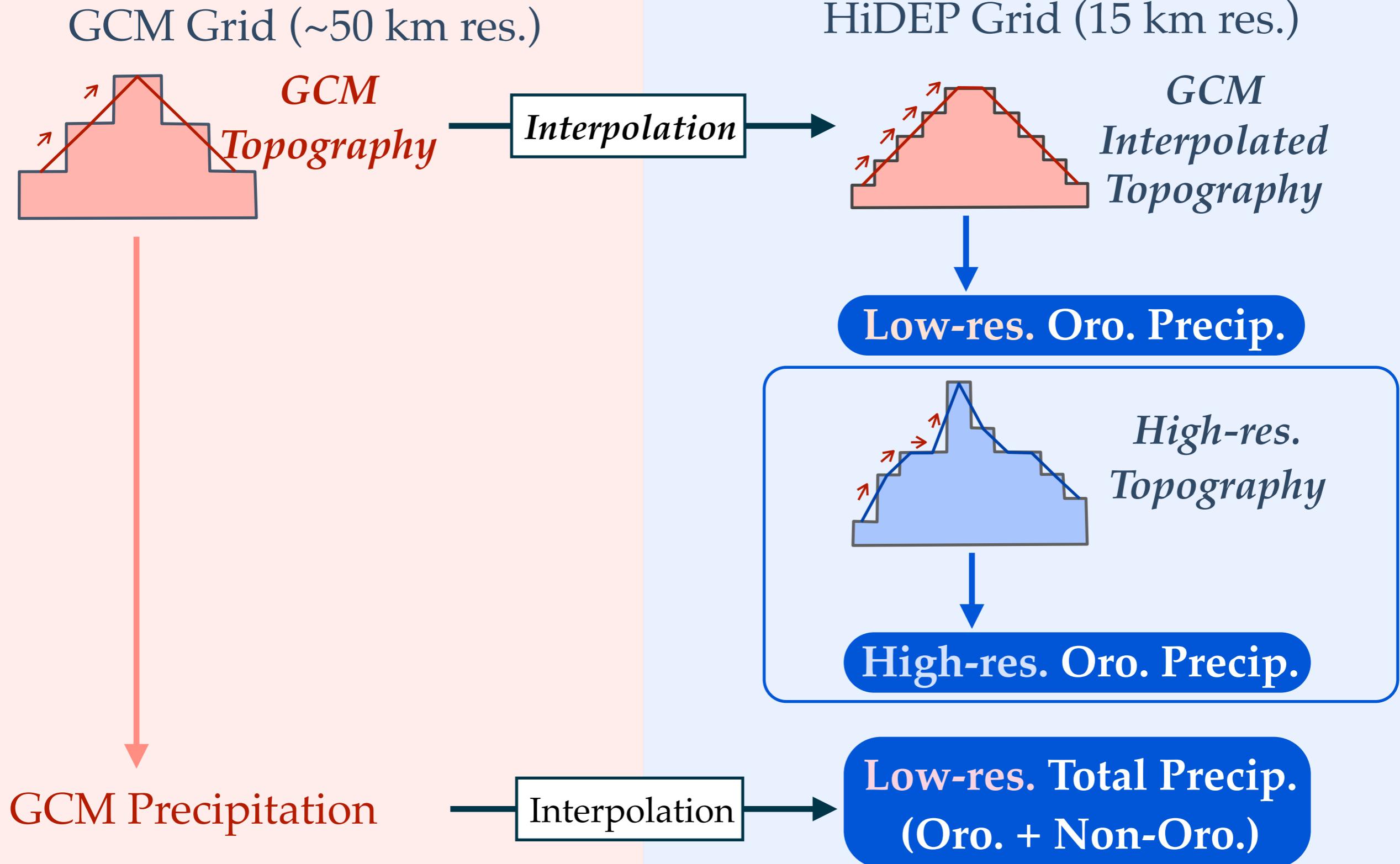
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Low-res. NON-Orographic Precipitation

Low-res. Total Precip. (*Interpolated from GCM*)

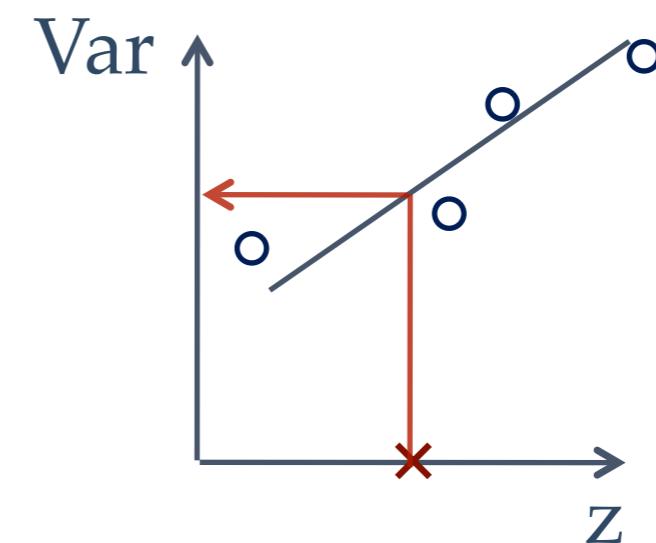
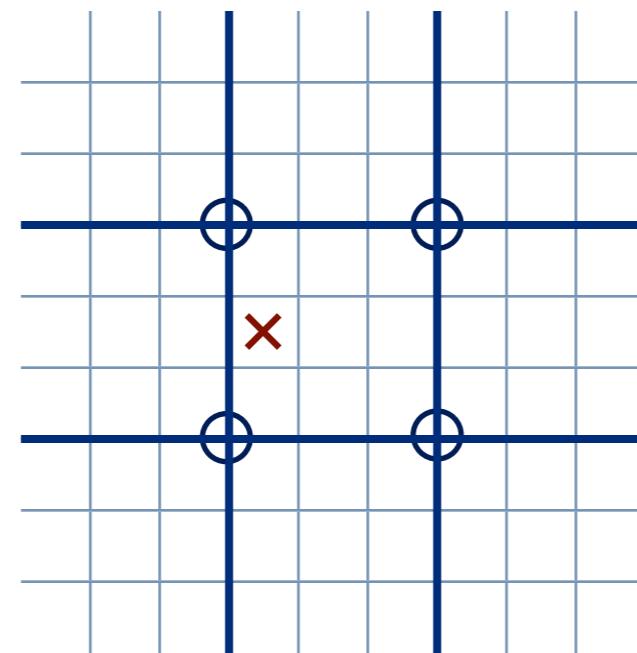
– Low-resolution Orographic Precip.

+ High-resolution Orographic Precip.

High-resolution Total Precip.

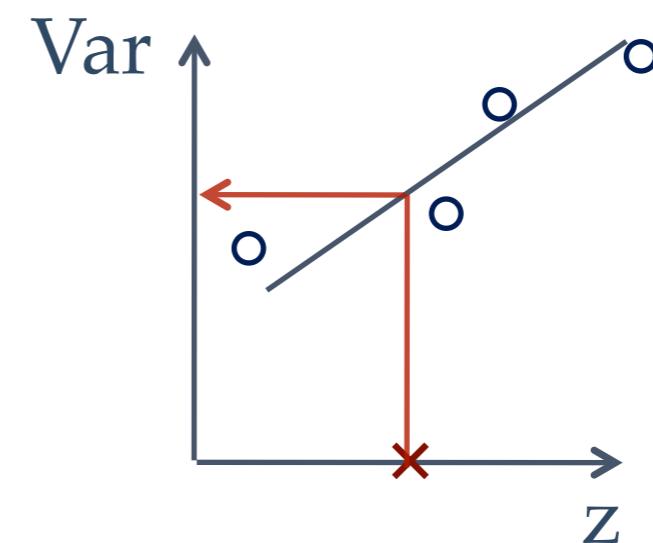
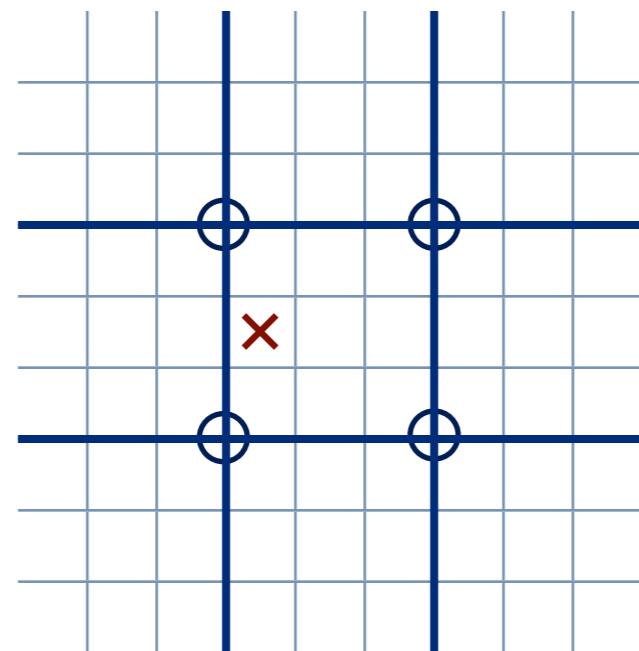
Surface Energy Balance

Extrapolation of GCM surface fields against the topography



Surface Energy Balance

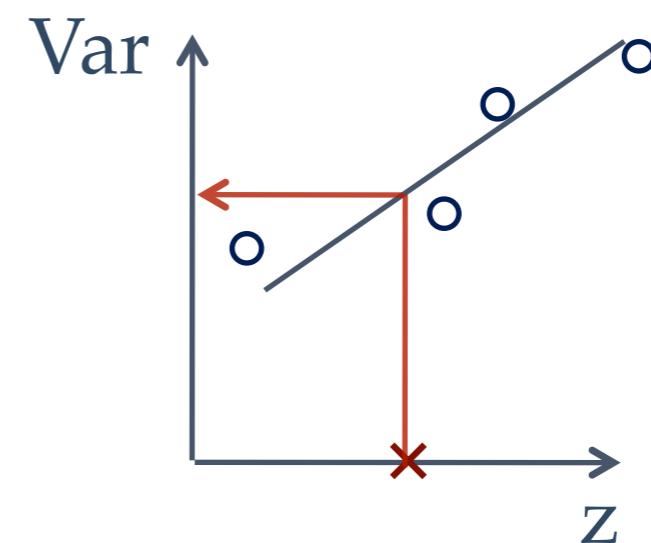
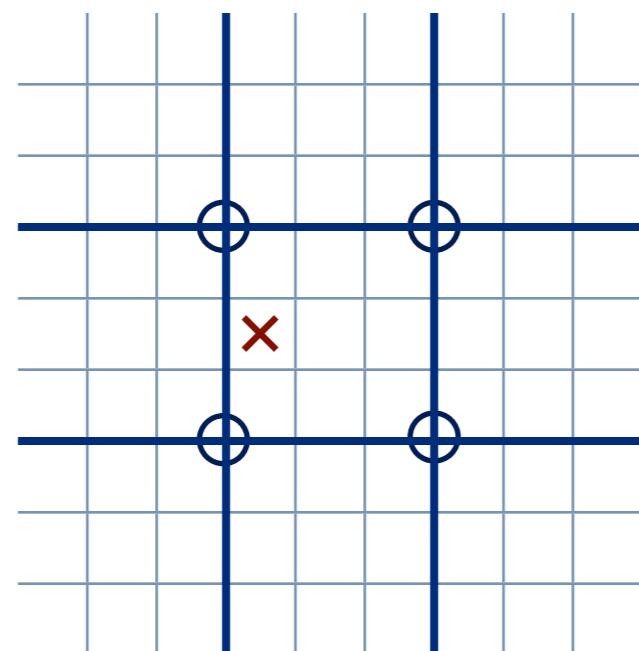
Extrapolation of GCM surface fields against the topography



Surface Scheme

Surface Energy Balance

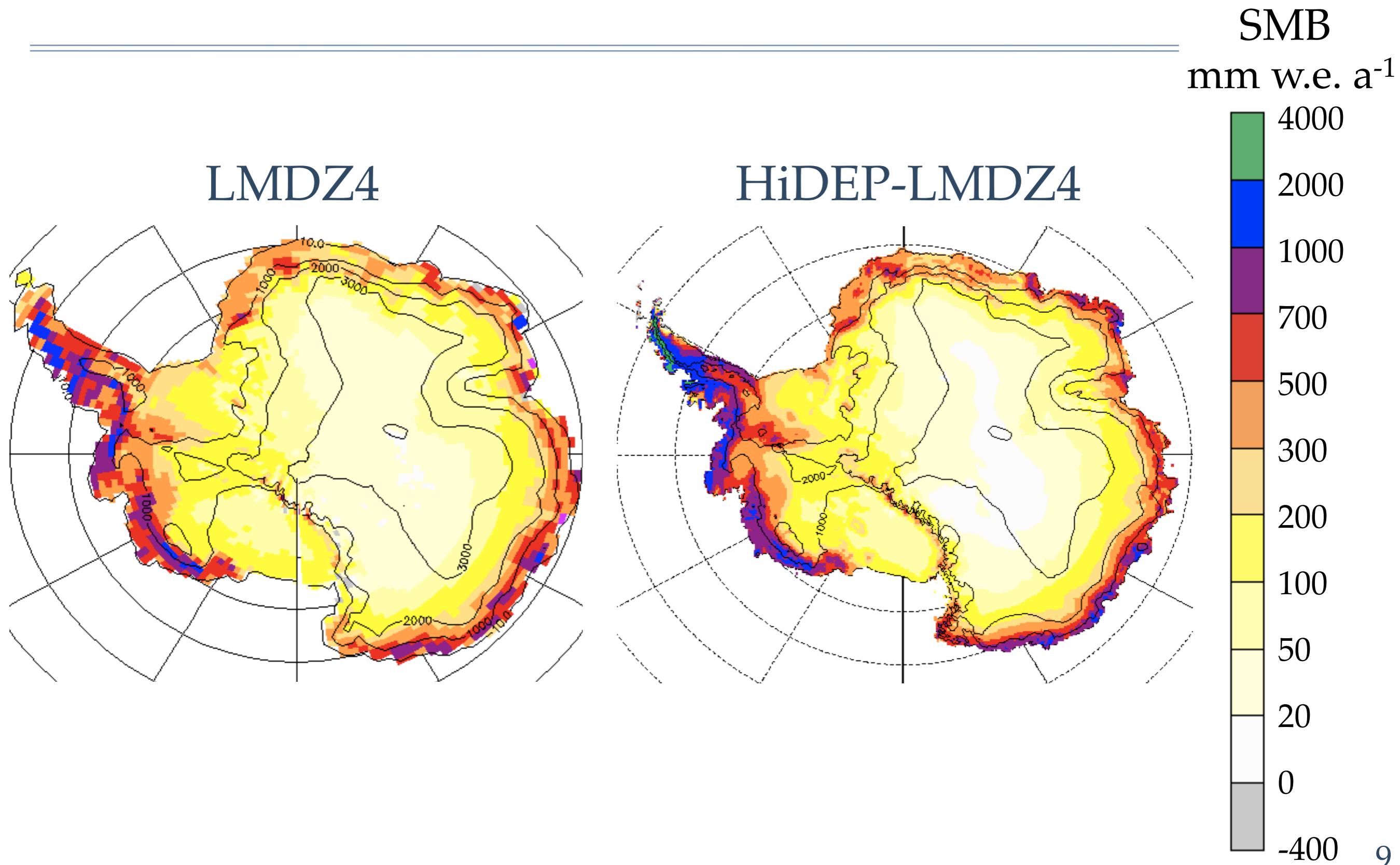
Extrapolation of GCM surface fields against the topography



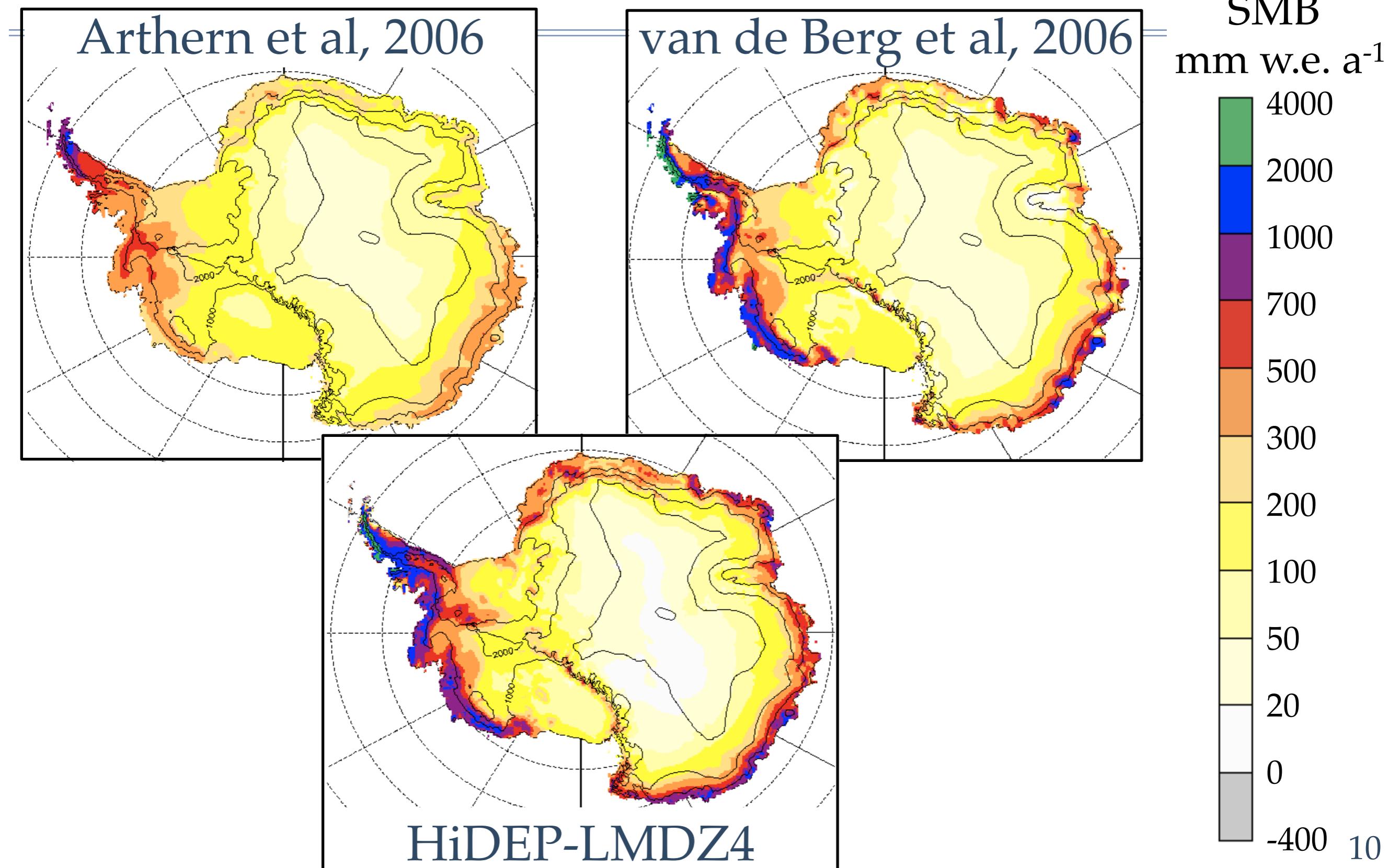
Surface Scheme

Sublimation
Melting
Refreezing

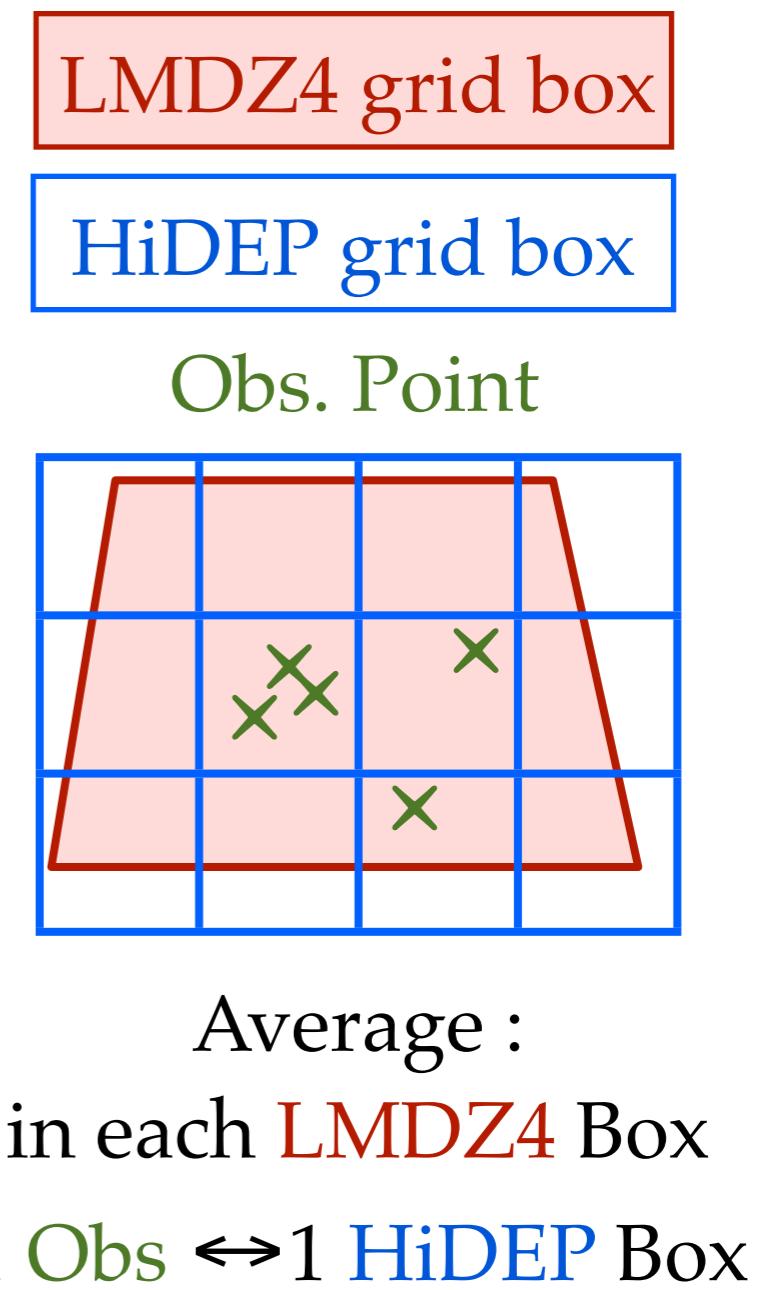
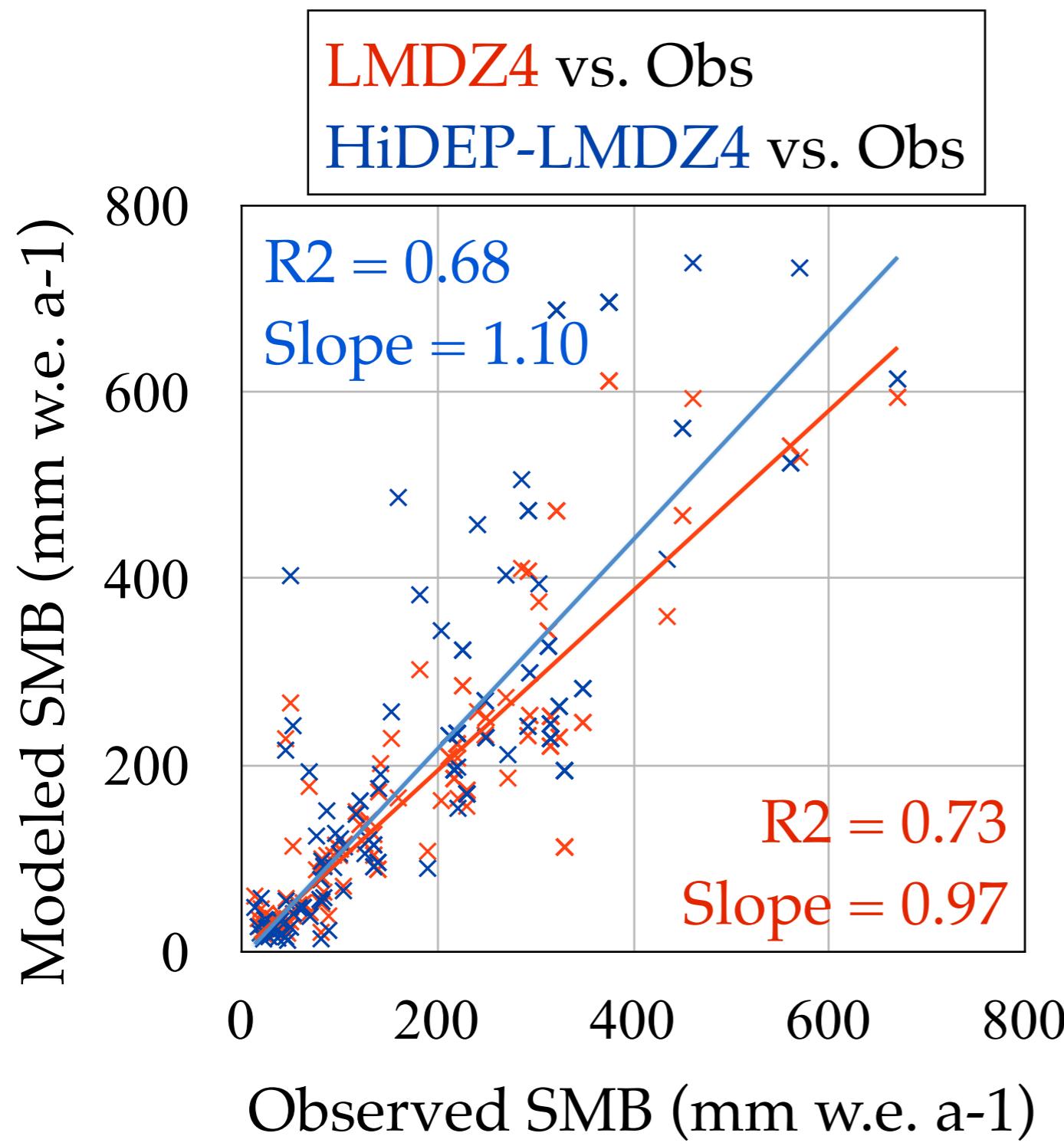
Application to LMDZ4 : 1980-2007



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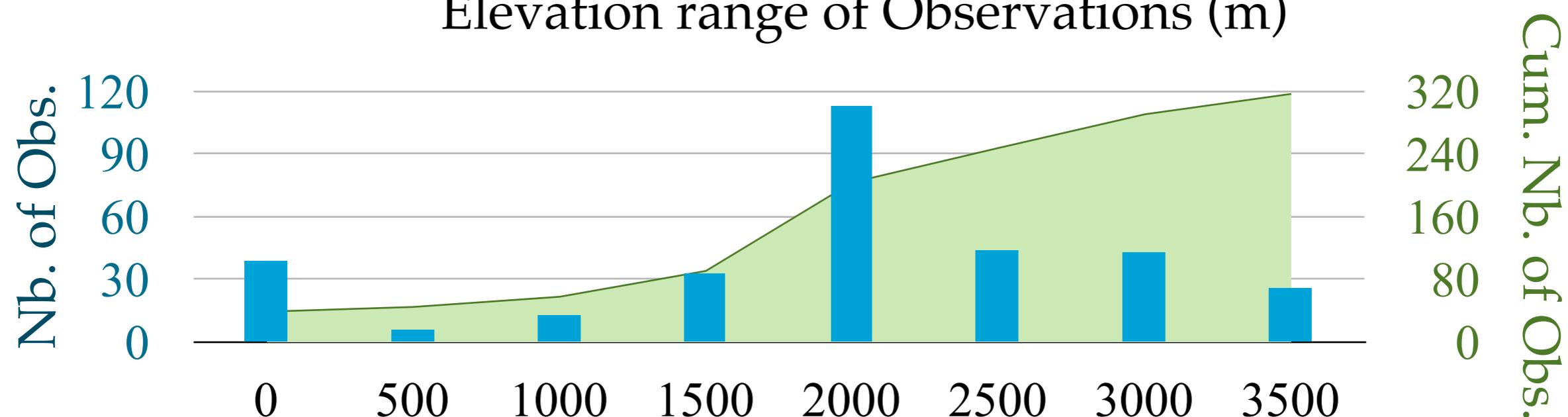
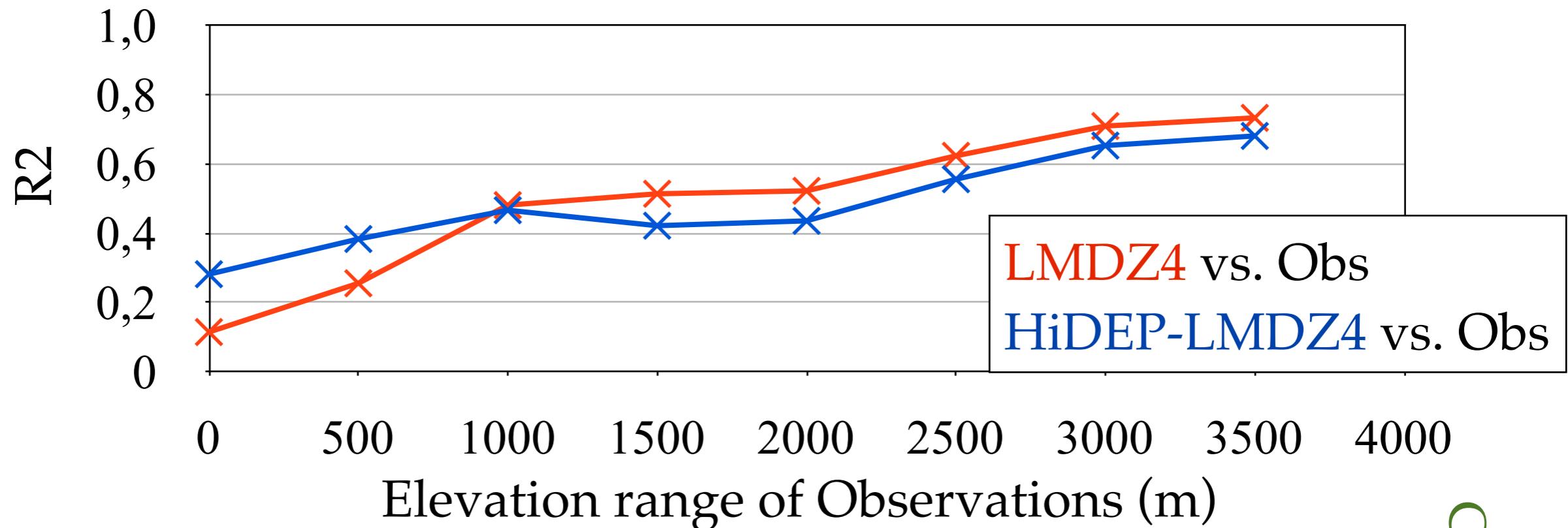


Validation with a quality-controlled SMB data-set (Magand et al., 2007) : $90^{\circ} - 180^{\circ}$ E

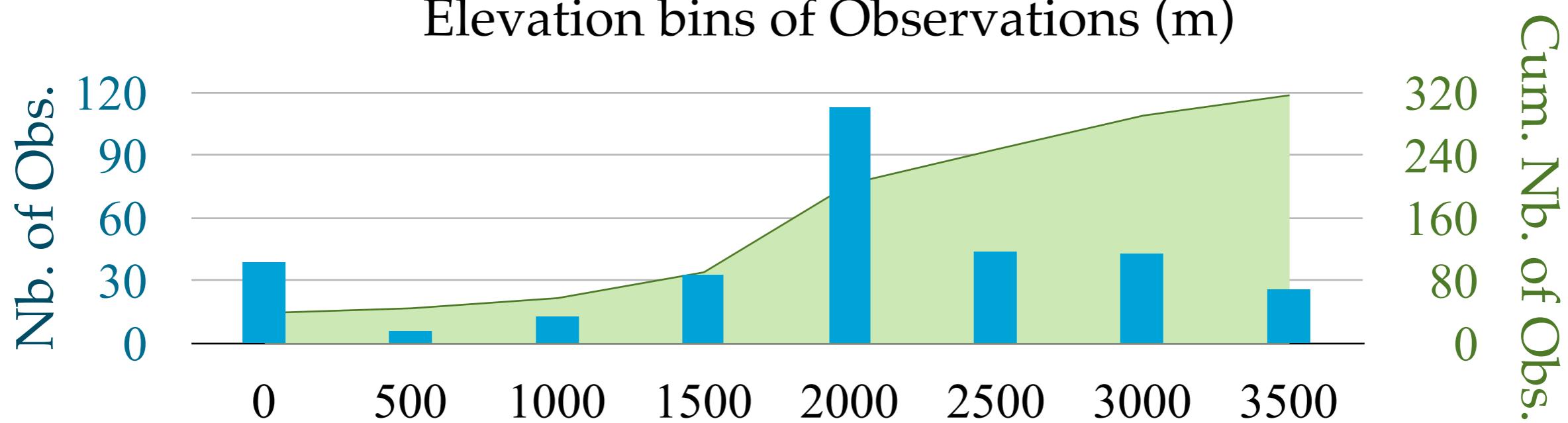
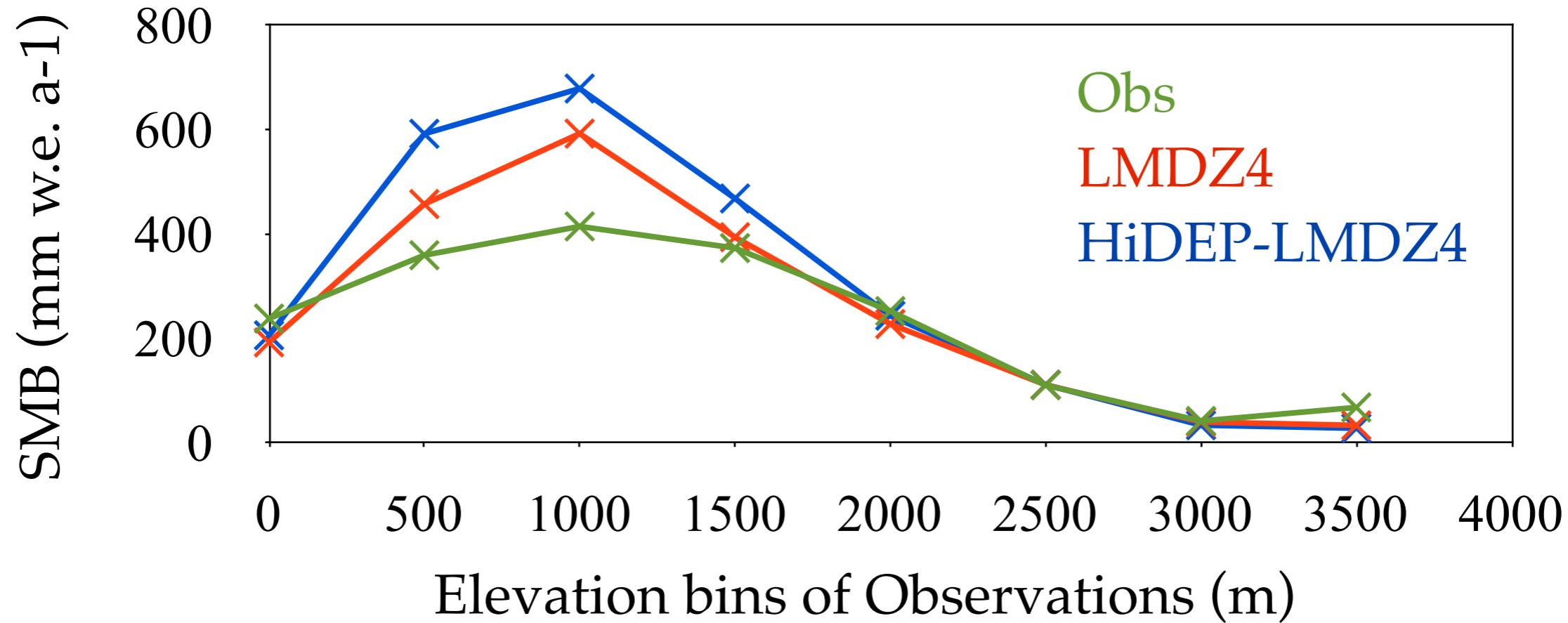


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R2 weighted by the number of observation in LMDZ4 grid boxes



Validation with a quality-controlled SMB data-set (Magand et al., 2007) : $90^{\circ} - 180^{\circ}$ E



Validation with a quality-controlled SMB data-set (Magand et al., 2007)

Extension of the data quality-control to the rest of Antarctica :
Work in progress at LGGE

For further information, you can contact **Soazig Parouty**

Grounded SMB 1980-2007

LMDZ4
P-E

$$\begin{aligned} 175.2 \text{ mm a}^{-1} &= \text{kg m}^{-2} \text{ a}^{-1} \\ \Leftrightarrow 2159 \text{ Gt a}^{-1} \\ \Leftrightarrow 6.0 \text{ mm a}^{-1} \text{ sea level equivalent} \end{aligned}$$

HiDEP-LMDZ4
SMB

$$\begin{aligned} 208 \text{ mm a}^{-1} &= \text{kg m}^{-2} \text{ a}^{-1} \\ \Leftrightarrow 2410 \text{ Gt a}^{-1} \\ \Leftrightarrow 6.7 \text{ mm a}^{-1} \text{ sea level equivalent} \end{aligned}$$

Present SMB (1950-2000) :

- Range : 1475 à 2331 Gt a⁻¹ [Monaghan et al., 2006]

Conclusion

- High-resolution SMB (15 km) obtained from LMDZ4 downscaling
Partial validation for present :
 - ♦ Downscaled SMB close to LMDZ4 SMB performance
 - ♦ Increase the wet bias of LMDZ4 in coastal areas

BUT lack of field data in (crucial) coastal areas

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Running :

- HiDEP-LMDZ4 for :

		Scenario		Just Finished
		A1B	E1	
LMDZ4 Boundary Conditions	HadCM3	21C	21C & 22C	
	ECHAM5	21C	21C	

Outlook

Further validation :

- Extended quality-controlled data set over all Antarctica
- Coastal-to-plateau transects

Model development in progress :

- MAR Surface Scheme
- Humidity advection



Thank you

Validation with a quality-controlled SMB data-set (Magand et al., 2007)

Quality criteria defined on the basis of :

- 1 - **Essential information** available : location, time period, method
- 2 - Quality rating of SMB **measurement methods**

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Applied to the **90° – 180°E Antarctic sector** :

Review of references compiled by Vaughan and Russell (1997)
+ New results from recent field campaigns

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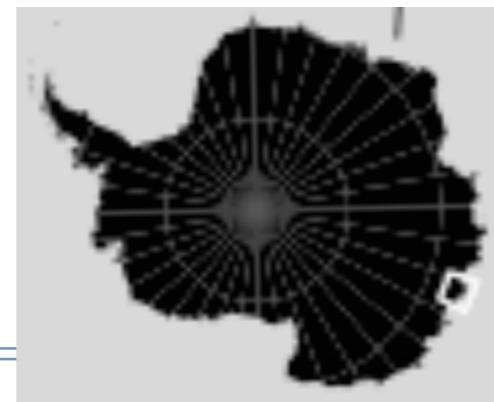
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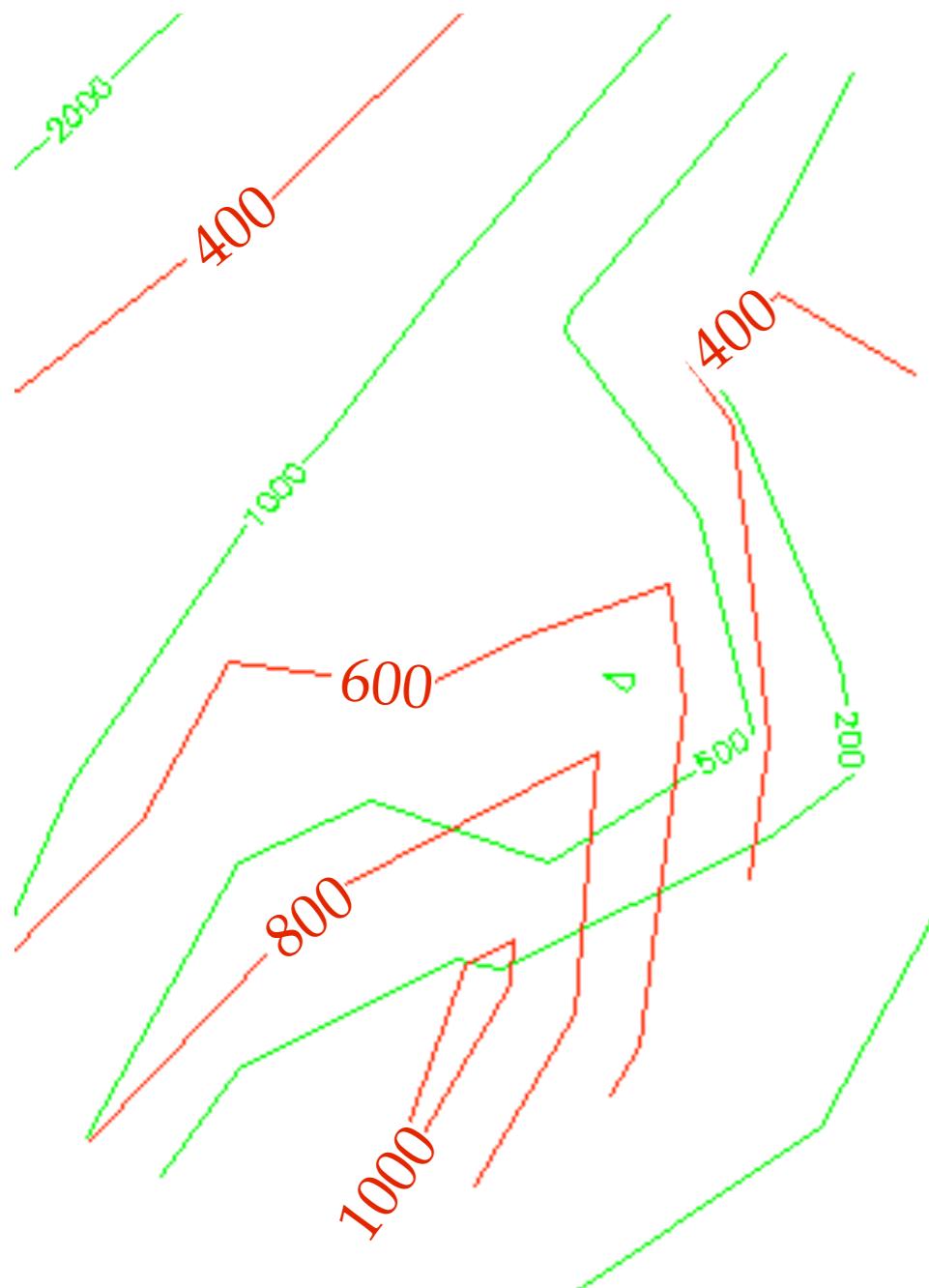
Review of references compiled by Vaughan and Russell (1997)
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- Quality-controlled data set for the **1950–2005 time period**
- Reduction in density and coverage
 - BUT reduced uncertainties compared to other compilations

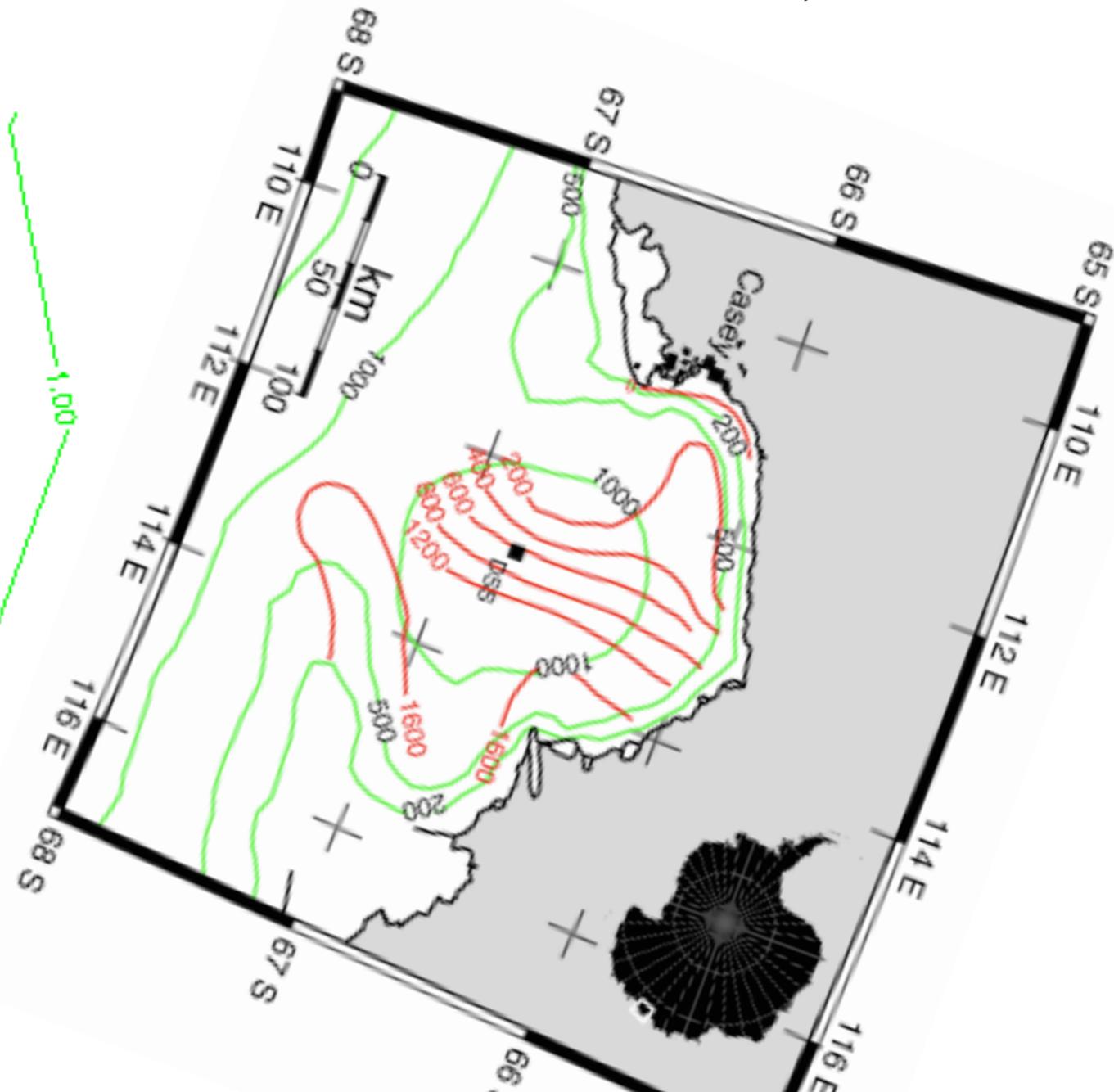
Law Dome



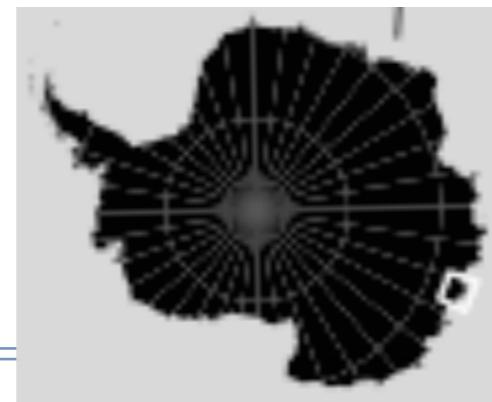
LMDZ4
SMB 1980-2007



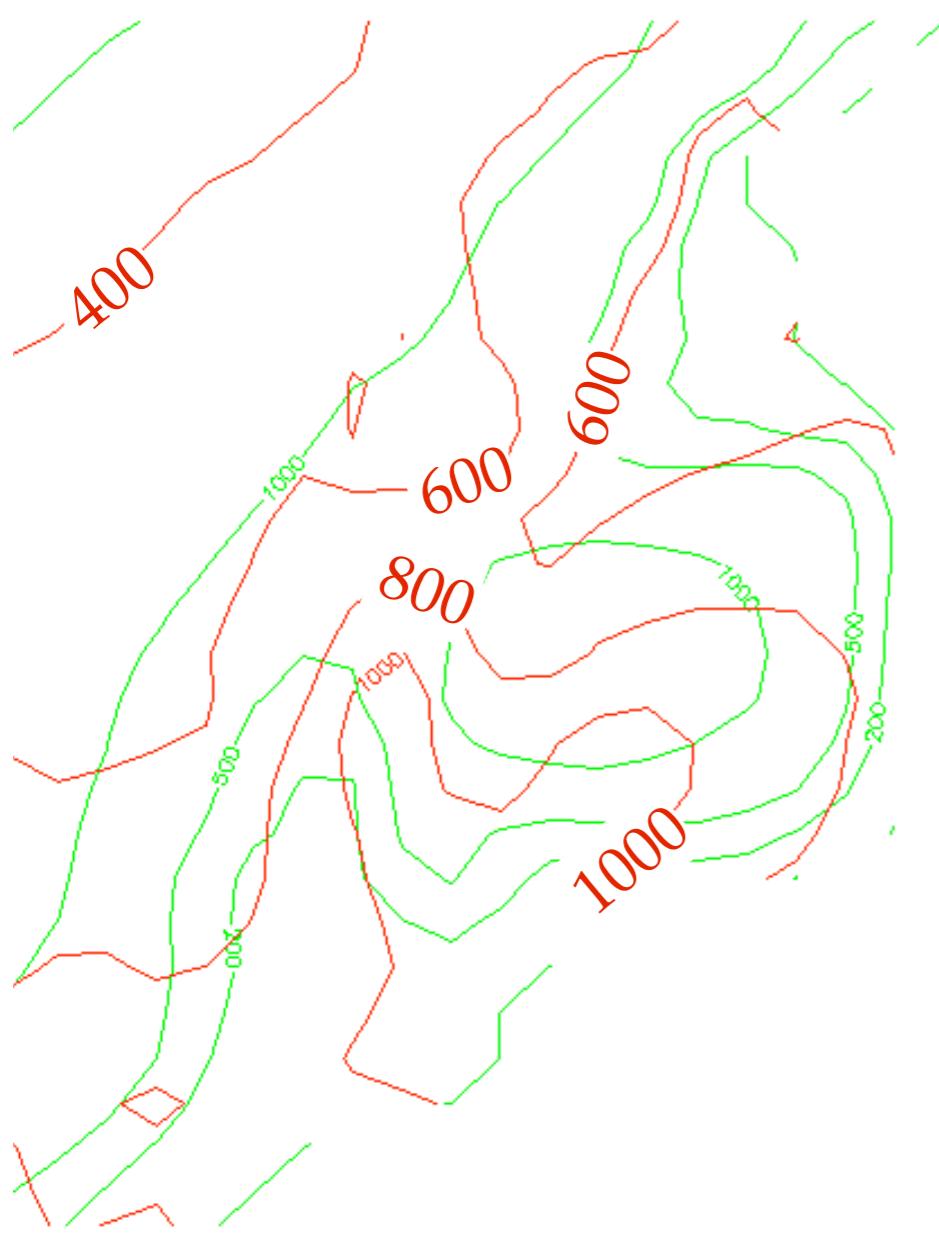
Climatologie SMB
van Ommen et al., 2004



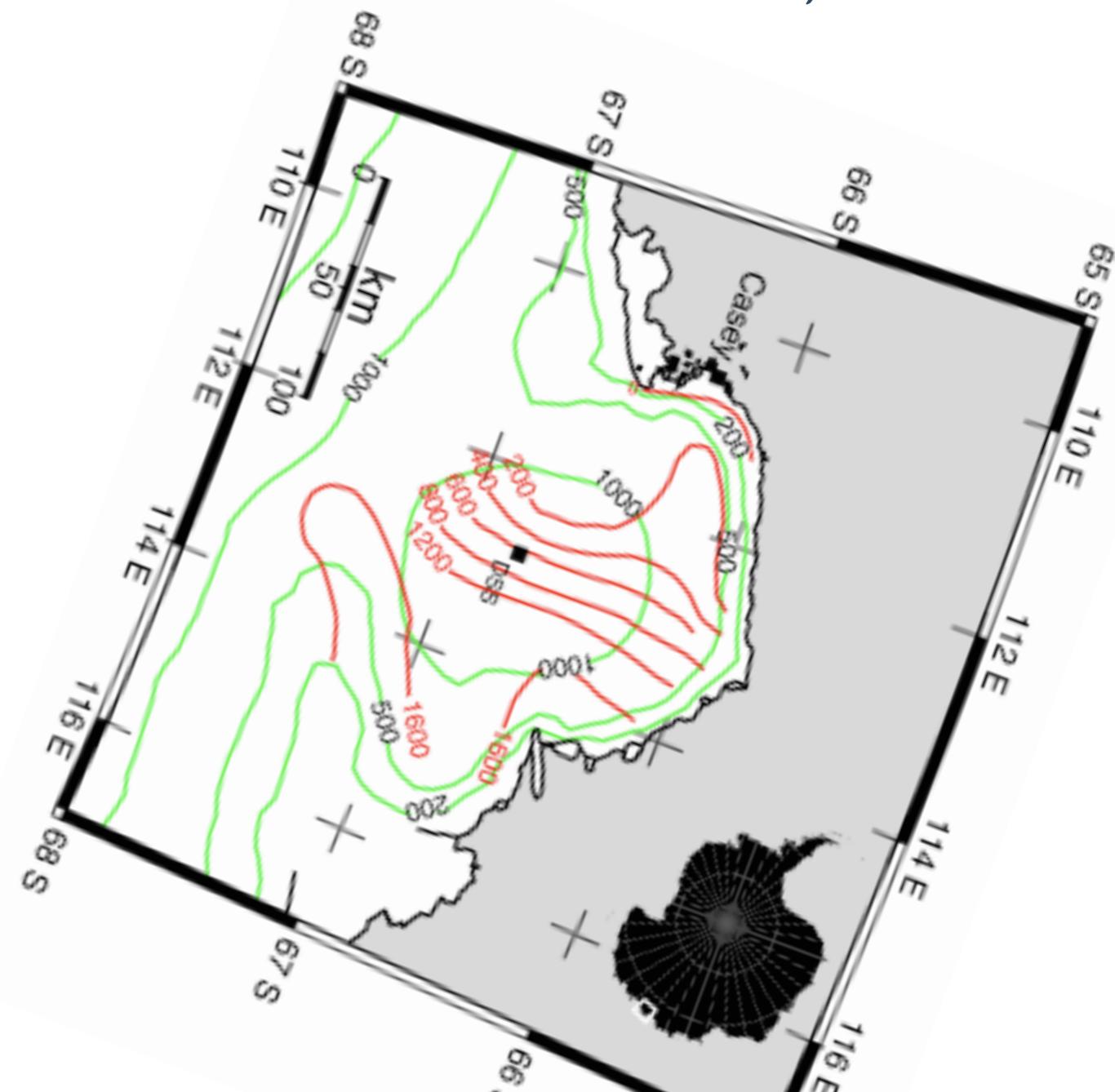
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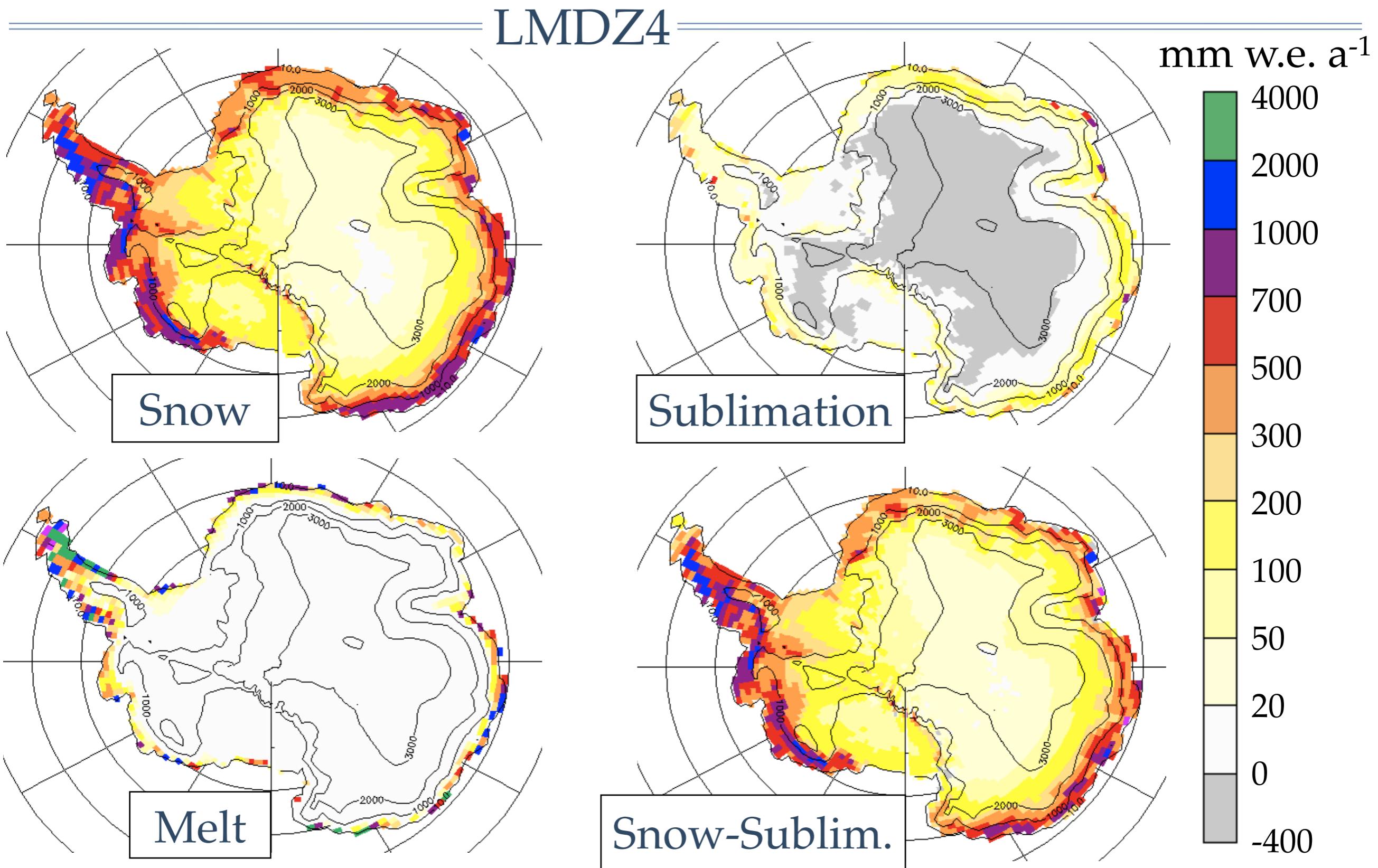
HiDEP-LMDZ4
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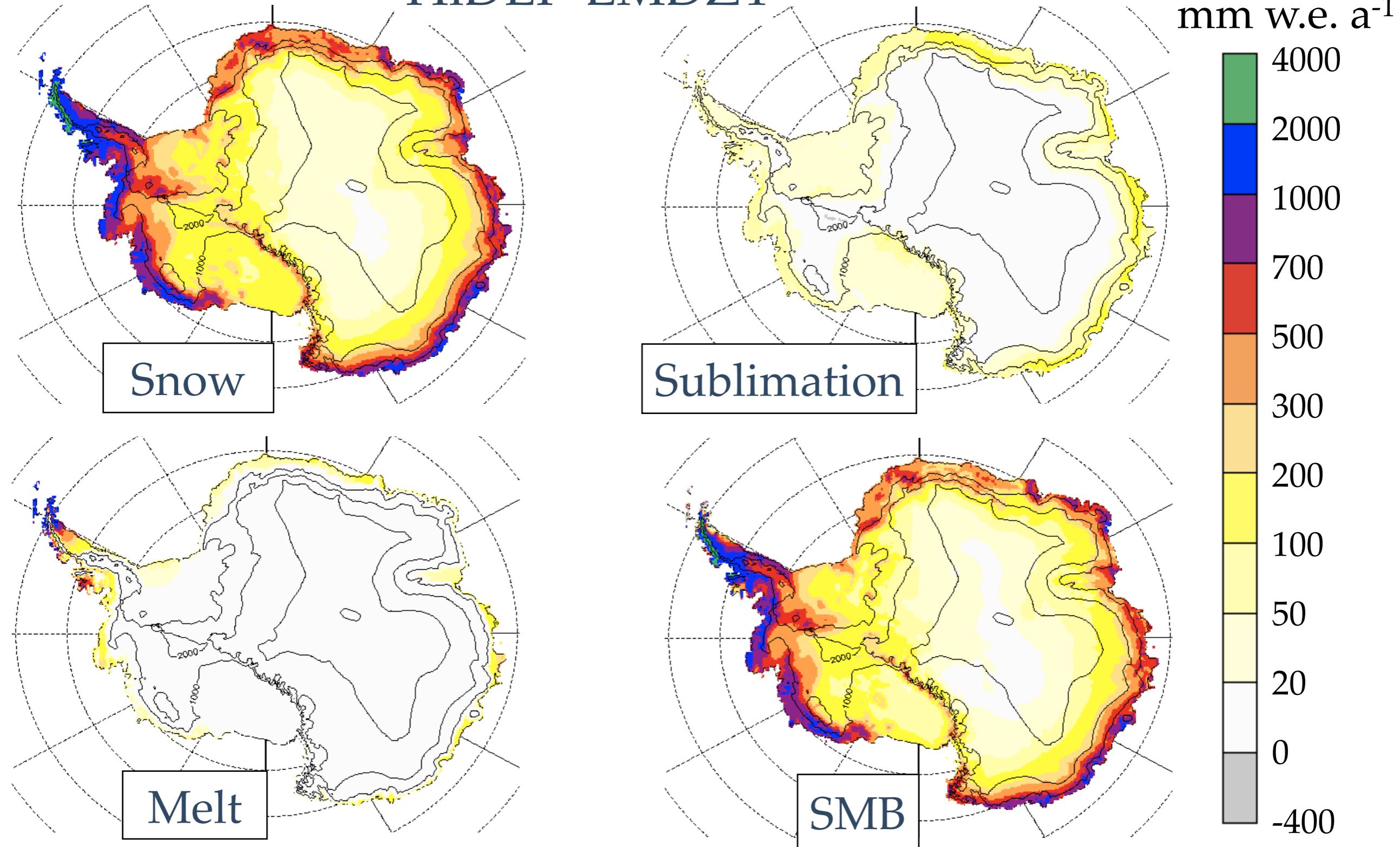


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HiDEP-LMDZ4



Désagrégateur de précipitation

Limites du désagrégateur de précipitation

Pas de dynamique (coûteux en temps de calculs)

- Pas de rétro-action de la physique sur la dynamique
- Pas de contournement du relief
- Pas d'effet de blocage

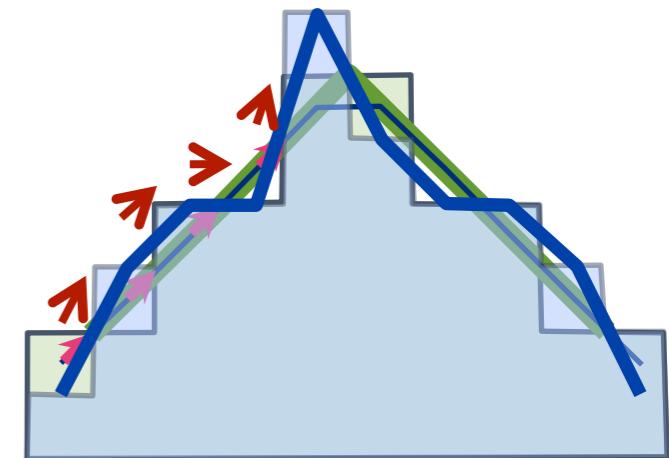
Orographic precipitation : Determination of the vertical wind W

Momentum equation

+ Continuity equation

+ Thermal dynamic equation

+ Steady adiabatic, inviscid, no-rotating flow



+ Hypothesis to simplify and linearize the equations :

- Small perturbations around the hydrostatic equilibrium
- Horizontal wind \gg Vertical wind
- Mean values slowly varying horizontally
- 2D
- Scorer parameter $l(z)=f(T(z),P(z))$ slowly varying

→ Wave equation on W (Mountain gravity wave) : WG