W. Bader (1), A. Perrin (2), D. Jacquemart (3), K. Sudo (4), H. Yashiro (5), O. A. Søvde (6), P. Demoulin (1), C. Servais (1), and E. Mahieu (1).

> (1) University of Liège, Liège, Belgium (w.bader@ulg.ac.be), (2) Laboratoire Interuniversitaire des Systèmes Atmosphériques, Paris, France, (3) Laboratoire de Dynamique, Interactions et Réactivité, Université Pierre et Marie Curie, Paris, France, (4) Graduate School of Environmental Studies, Nagoya University, Nagoya, Aichi, Japan,

(5) Japan Agency for Marine-Earth Science and Technology (JAMSTEC), Yokohama, Japan, (6) Center for International Climate and Environmental Research Oslo (CICERO), Oslo, Norway.

de Liège



http://orbi.ulg.ac.be/handle/2268/124230

Université

Groupe InfraRouge de Physique Atmosphérique et Solaire

1. Introduction

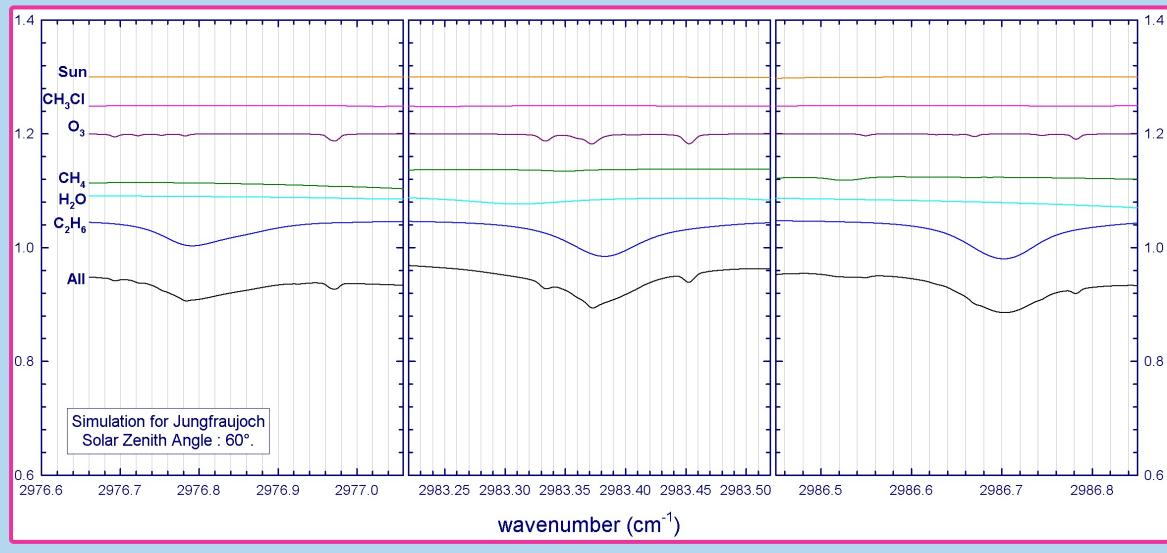
• The Ethane (C_2H_6) is the most abundant Non-Methane HydroCarbon (NMHC) in the Earth's atmosphere, with a lifetime of approximately 2 months. C_2H_6 main sources are biomass burning, natural gas loss and biofuel consumption. Oxidation by the hydroxyl radical is by far the major C_2H_6 sink as the seasonally changing OH concentration controls the strong modulation of the ethane abundance throughout the year. Ethane reduces CI atom concentrations in the lower stratosphere and is a major source of peroxyacetyl nitrate (PAN) and carbon monoxide (by reaction with OH). Involved in the formation of tropospheric ozone and in the destruction of atmospheric methane through changes in OH, C₂H₆ is an indirect greenhouse gas with a netglobal warming potential of 5.5 (100-yr horizon).

• All the spectra analyzed here have been recorded at the International Scientific Station of the Jungfraujoch (46.5°N, 8°E, 3580 m asl) with a Bruker IFS-120HR Fourier Transform Infrared (FTIR) spectrometer. It has been put into regular operation since 1984 allowing to record wide-band high-resolution IR solar spectra either with a MCT or InSb detector. Since 1991, the FTIR instrument is affiliated to the framework of the Network for the Detection of Atmospheric Composition Change (NDACC, visit http://www.ndacc.org).

2. Retrieval Strategy

Selected Micro-windows

Parameters have been settled down on the basis of tests on a full year minimizing residuals and maximizing DOFS.



Spectroscopic linelist

As the current state of ethane parameters in HITRAN (e.g. : Rothman et al., 2009, http://www.hitran.com) was rather unsatisfactory in the 3 µm region, new ethane absorption cross sections recorded at the Molecular Spectroscopy Facility of the

Rutherford Appleton Laboratory (Harrison et al., 2010) were combined with HITRAN 2004 line parameters (including all 2006 updates) and therefore used in our retrievals, after conversion into pseudolines by G. C. Toon (personnal communication, 2011). These cross sections were calibrated in intensity by using reference low-resolution spectra from the Pacific Northwest National Laboratory (PNNL, Washington, USA, http://www.pnl.gov/) IR database.

We quantified the impact of two updates of the spectroscopic parameters for both Hitran 2004 and 2008 on spectral residuals (see Table I):

1. The update of two O₃ lines (encompassed in the 1-PQ₃ μ-window) corrected by P. Chelin (LPMA, Paris, France) in the framework of the UFTIR project.

2. The improvement brought by the update of the line positions and intensities of methyl chloride (CH₃CI) in the 3.4 μm region (Bray et al., 2011). 2983.20 2983.25 2983.30 2983.35 2983.40 2983.45 2983.50 Improvements brought by Hitran-08 over the 2004 edition are illustrated on Figure 3.

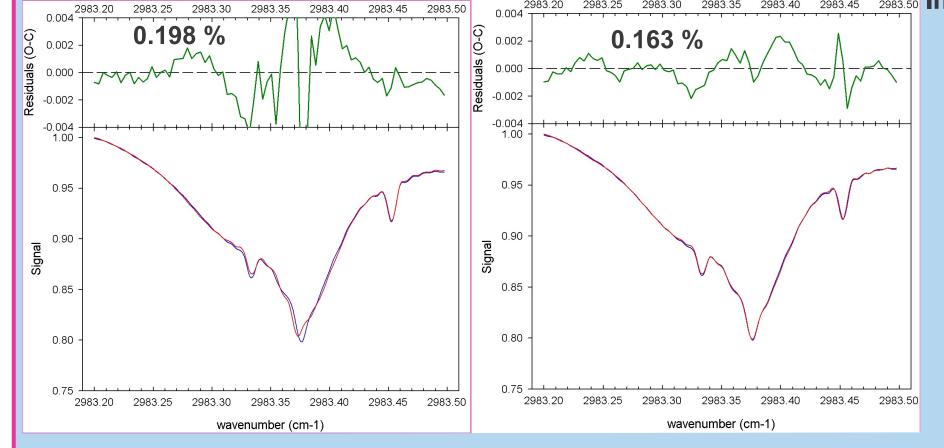


Figure 3 - Mean observed (blue curve)/simulated (red curve) spectra and residuals (green curve) for 187 spectral fits (4.96 mK), using Hitran-04 (left panel) and Hitran-08 (right panel). Harrison's, Chelin's and Bray's updates are used for both.

Table I - RMS Values (in %) for each μ-window (see Table II) according to each linelist tested.

Smoothing Error

Natural Variability

Measurement Error

	2- Harrison + 1						
	3- Chelin + 2		0.169	0.158	0.	172	0.
	4- Bray + 3		0.153	0.151	0.152		0.
n ⁻¹)		Interfering species					2
77.059		C ₂ H ₆ , H ₂ O, CH ₄ , O ₃ , CH ₃ CI					W re

1- Hitran-08 | 0.431 | 0.206 | 0.471

able II - List of microwindows used for our C2H6 inversions,							
3	2986.43 - 2986.85	C ₂ H ₆ , H ₂ O, CH ₄ , O ₃ , CH ₃ Cl					
2	2983.2 - 2983.5	C ₂ H ₆ , H ₂ O, CH ₄ , O ₃ , CH ₃ Cl					
1 - 「Q ₃	2976.66 - 2977.059	C_2H_6 , H_2O , CH_4 , O_3 , CH_3CI					

for each of them, the third column provides interfering gases adjusted during the retrieval.

Figure 1 - Synthetic spectra of our three fitted μ-windows (see limits on Table I) for all gases (in black) as well as for the individual contributors (C₂H₆ in blue, H₂O in cyan, CH₄ in green, O₃ in dark pink, CH₃Cl in pink and solar lines in orange) to the absorption in this spectral interval. For clárity, the contributions of each species have been vertically displaced. Simulated solar zenith angle: 60°.

² A priori Volume Mixing Ratio (VMR) profiles

We selected the a priori VMR profile among four ethane profiles (see Figure 2) giving retrieved profiles with the least oscillations and least negative VMR values. We also tested those profiles adjusted on EMEP measurements made at the Rigi station (47°N, 8°E, 1031 m a.s.l.).

Figure 2 - The a priori profiles tested (left panel) and their associated relative standard deviation (right panel) are illustrated. The first adopted *a priori* C₂H₆ profile (grey crosses) is a zonal mean (for the 41-51°N latitude band) of 771 occultations recorded by the ACE-FTS instrument between the 2nd of November in 2004 and the 8th of February in 2011 extending from 8.5 to 20 km. The profile extension down to 3.58 km is based on EMEP in situ measurements (bottom panel) while the upper extension to 100 km is based on the WACCM model climatology (Whole Atmosphere Community Climate, http://waccm.acd.ucar.edu). Pink crosses is the *a priori* profile issued from the CHASER 3-D Chemical Transport Model developed at the Center for Climate System Research (CCSR), University of Tokyo/National Institute for Environmental Studies (NIES). Chemical Transport Model v.2 and v.3 from the University of Oslo (Berntsen et al., 1997) are plotted in dark and light green crosses, respectively.

C₂H₆ VMR

We selected the a priori VMR profile and their uncertainties issued from the Chemical AGCM for Study of atmospheric Environment and Radiative Forcing (CHASÉR) Model (Sudo et al., 2002). Residuals were reduced by 7.6 % while DOFS is improved by 17.4 and 11.1 %, in comparison to the Oslo's CTM v2 and v3 a priori profiles, respectively.

µ-windows

--- ACE + WACCM --- Oslo_CTM 2_Corrected -- Oslo_CTM 3 E) 1000

Global

0.424

 All retrievals have been performed with the SFIT-2 algorithm (v 3.91) (Rinsland et al., 1998) in order to retrieve three ethane vertical columns (seé next frame) on a series of about 13 000 spectra recorded between 1994 and

2011. The adopted settings are: (i) the fitting has been narrowed down to 3 microwindows described on Table II, (ii) the adopted C₂H₆ a priori VMR profile in all our retrievals, as well as its uncertainties, are based on synthetic profiles produced by the CHASER model for the 2007-2009 time period. A priori profiles for the interfering gases are based on the WACCM model climatology (iii) HITRAN 2008 line parameters including Harrison's pseudo lines, the two ozone lines update provided by P. Chelin (Personal Communication, 2004) and the updated CH₃Cl lines (Bray et al., 2011) as well as the solar line compilation provided by F. Hase (KIT) have been assumed for target and interfering absorptions, (iv) adopted temperature and geopotential height data sets are provided by the National Centers for Environmental Prediction (NCEP, Washington,

3. Information Content and Error Budget Information content and error budget

have been carefully evaluated. Figure 4 displays typical results computed for VMR. The first eigen vector and corresponding eigenvalues (see left 18 18 frame, in red) show that information on both selected C₂H₆ partial columns, namely 3.58-6.79 km (lowtropospheric) and 8.45-14.3 km (Upper Tropospheric-Lower Stratospheric, 14 UTLS), is mainly coming from the retrieval (99 %). 4 - Information content

calculated for typical C₂H₆ retrievals at the Jungfraujoch station. The three first eigenvectors are reproduced in the left frame. Right frame gives the corresponding error budget, with identification of the main error components, together with the assumed variability (see color codes).

error budget is still ongoing. However, preliminary error analysis is displayed on Table III.

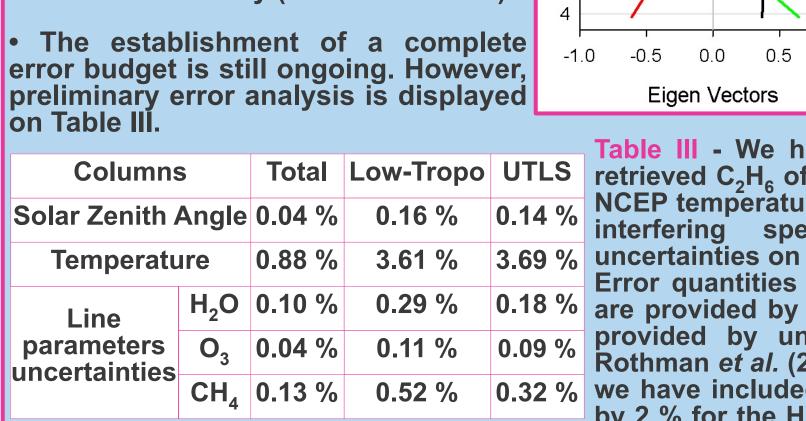


Table III - We have already quantified the impact on our Total Low-Tropo UTLS retrieved C₂H₆ of the error on the solar zenith angle, on the NCEP temperature profiles and on the line intensities of the interfering species through sensitivity tests. The 3.61 % 3.69 % uncertainties on the solar zenith angle are estimated at 0.2°. Error quantities associated with temperature uncertainties 0.29 % are provided by the NCEP while error on line intensities is provided by uncertainty indices reported on Hitran-04, Rothman et al. (2005). Therefore, according to these indices CH₄ 0.13 % 0.52 % we have included line intensities in our retrievals modified by 2 % for the H₂O lines, by 2 or 5 % for the O₃ line and by

0.001 0.01 0.1 1

Relative Errors (%)

10 % for the CH₄ line. We observe a seasonnal variation of the error as for the error on temperature with a maximum during Spring (10 %) and a minimum (< 0.6 %) during Autumn.

Eigen Vectors

 $\lambda_4 = 0.9997 \ \lambda_2 = 0.9613 \ \lambda_3 = 0.1432$

6. Conclusion

 Harrison's new ethane parameters coupled to Hitran 2008 compilation improve the retrieval of ethane in terms of spectral residuals and information content; as well as Chelin's O3 and Bray's CH₂Cl updates.

• The selected a priori VMR profiles issued from the CHASER Model gives the least negative profiles with best residuals and DOFS.

 Concerning the long-term trend of C₂H₆, we determined a significant decrease in its concentration over the 1994-2011 time period. We also characterized a seasonnal change in total and tropospheric columns of ethane greater than the previously mentionned decrease. In the UTLS the long-term decrease is as the same order of magnitude than the seasonnal modulation.

 We expect to put more investigations into the study of the interannual variations as well as of the seasonnal change of ethane in the lower-troposphere and in the UTLS region.

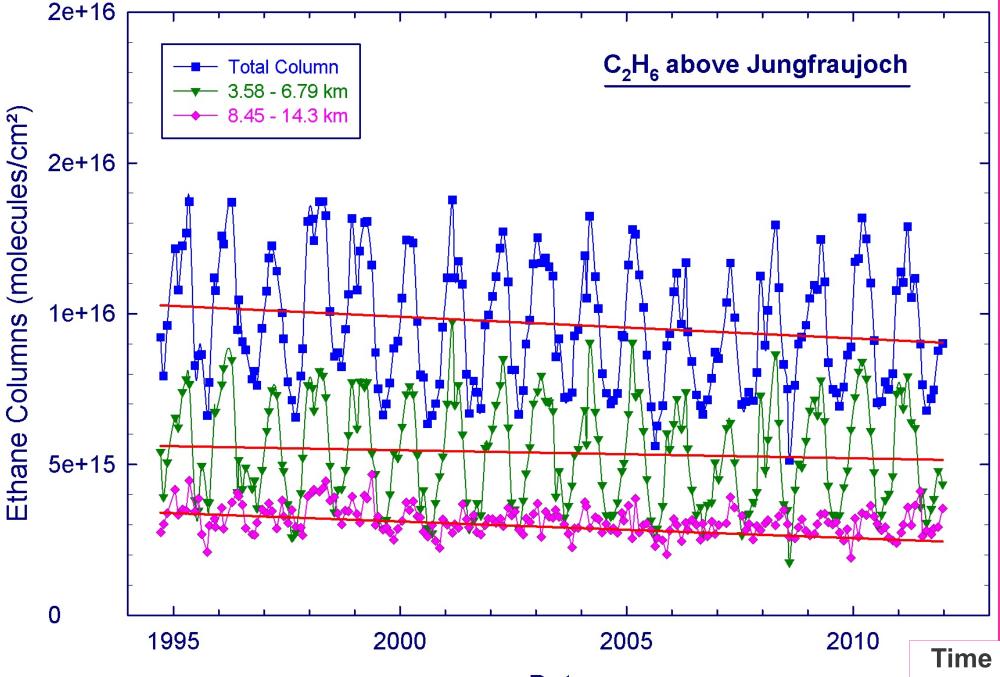
• We still have to evaluate the impact of C₂H₆ and CH₃Cl spectroscopy and instrumental line shape in

order to refine our error budget.

Acknowledgments

The University of Liège involvement has primarily been supported by the PRODEX program funded by the Belgian Federal Science Policy Office, Brussels and by the Swiss GAW-CH program. E. Mahieu is Research Associate with the F.R.S. – FNRS. The FRS-FNRS and the Fédération Wallonie-Bruxelles are further acknowledged for observational activities support. We thank the International Foundation High Altitude Research Stations Jungfraujoch and Gornergrat (HFSJG, Bern) for supporting the facilities needed to perform the observations. We further acknowledge the vital contribution from all the Belgian colleagues in performing the Jungfraujoch observations used here. We further thank G.C. Toon (NASA-JPL, Pasadena) for the conversion of the ethane cross sections into pseudolines which can be used by our retrieval algorithm. We finally thank K. Sudo and O. A. Søvde for providing synthetic data.

μ-windows Limits (cm



4. Time Series

• The Figure 5 displays our retrieved C₂H₆ total column and both partial columns (low-tropospheric and UTLS) above Jungfraujoch. We computed an overall decrease in ethane concentrations since 1994 of -14, -9 and -39 % resp. for our three columns. Trends have been determined using the boostrap resampling tool developed by Gardiner (2008) (see Table IV).

 Our measurements allow to characterize strong seasonal variations of C₂H₆ total and lowtropospheric columns with a maximum generally observed around mid-February. On average, the peak-to-peak amplitudes respectively amount to 50 and 76 % of the 1994 reference column.

Figure 5 - Time series of C₂H₆ total column (in blue), low-tropospheric (3.58-6.79 km, in green) and UTLS (8.45-14.3 km, in pink) partial columns above Jungfraujoch. Red lines are linear trends.

Total Column | 3.58 - 6.79 km | 8.45 - 14.3 km Time Series -0.47 ± 0.35 % | -0.92 ± 0.30 % | -0.69 ± 0.24 % Junfraujoch (1994-2011) (1994)(1994)(1994)Jungfraujoch -1.06 ± 0.31 % (1994-2008) (1994)CHASER 0.55 ± 0.18 % (1994-2008) Jungfraujoch | -1.58 ± 0.68 % | -0.79 ± 0.98 % | -2.59 ± 0.89 % (1998-2005) Oslo's CTM $|-0.89 \pm 0.56 \%|$ $|-0.72 \pm 0.57 \%|$ $|-1.26 \pm 1.23 \%|$ (1998-2005) (1998)(1998)(1998)

Model Comparison

Total Columns 8.0e+15 6.0e+15 3.58 - 6.79 km 8e+15 6e+15 4e+15 8.45 - 14.3 km 6e+15 4e+15 2e+15

Jungfraujoch Observations —— CHASER Model —— Oslo's CTM

Table IV - Annual Change (in %), its 2-σ uncertainties and its reference year for Jungfraujoch, CHASER Model and Oslo's CTM v.3 time series for our three columns.

• On Table IV, we notice a good agreement, significant within 2-σ between Oslo's CTM computed trends and the trends of our retrieved columns.

 On Figure 6, we notice an underestimation of the amount of ethane in the atmosphere by Oslo's model which may be explained by the use of

undervalued atmospheric pressures. • C₂H₆ seasonnal change has been evaluated for both CHASER and Oslo's CTM data with a peak-to-peak amplitude of 35 % (Ref. : 1994) and of 44 % (Ref. : 1998), respectively for ethane total columns. While our lowtropospheric peak-to-peak amplitude amounts to 67 % for the CHASER Model and for 51 % for Oslo's CTM.

Figure 6 - C₂H₆ three columns typical year for Jungfraujoch observations (in blue), results from CHASER Model (in pink) and from Oslo's CTM v.3 (in green). Total columns are respecitively averaged on the 1994-2011, 1970-2008 and 1998-2005 time periods while CHASER's partial columns are computed over the 2007-2009 time period.

eterences

- Berntsen T., et al. (1997), J. Geophys. Res., 102, 21239-21280. Bray C., et al. (2011), J.Q.S.R.T., 112, 2446–2462.
- Gardiner T., et al. (2008), Atmos. Chem. Phys., 8, 6719-6727. Harrison J. J., et al. (2010), J.Q.S.R.T., 111, 357-363.
- * Rinsland C. P., et al. (1998), J. Geophys. Res., 103 (D21), 28197. * Rothman L.S., et al. (2009), J.Q.S.R.T., 110, 533-572.
- * Sudo K., et al. (2002), J. Geophys. Res., 107, 4339.

Contact Information: w.bader@ulg.ac.be - http://girpas.astro.ulg.ac.be